

2015 AOFD

Dynamic Partition on the Stratosphere-Troposphere Exchange (STE) of Air Mass along Isentropic Surface

Huang Yang¹, Gang Chen¹, Qi Tang^{2,3},
Peter Hess³ and David Plummer⁴

¹Earth and Atmospheric Sciences, Cornell University

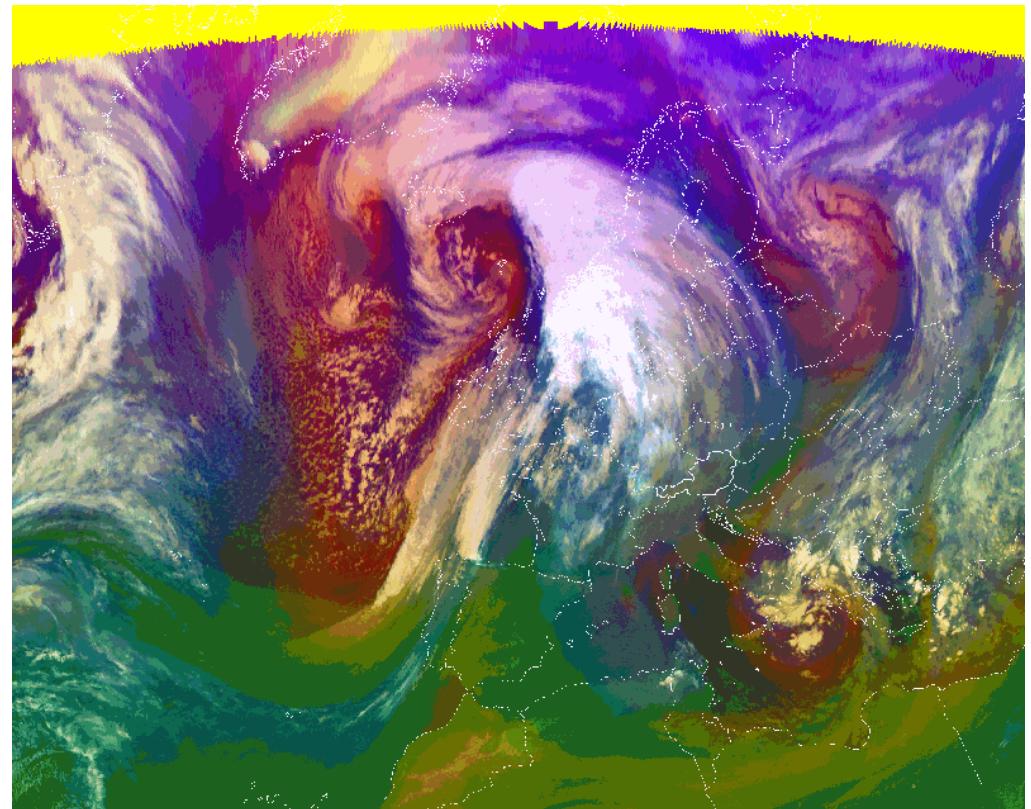
²Lawrence Livermore National Laboratory

³Biological and Environmental Engineering, Cornell University

⁴Canadian Centre for Climate Modelling and Analysis, Environment Canada

Introduction

Meteosat Second Generation (MSG) satellites (**EUMETRAIN**)



Air mass RGB loop from 03:00 UTC – 09:00 UTC, 29 Dec. 2012

Blue – moist polar air

Green – tropical air

Red – dry stratospheric air intrusion (STE)

Introduction

Impact one:

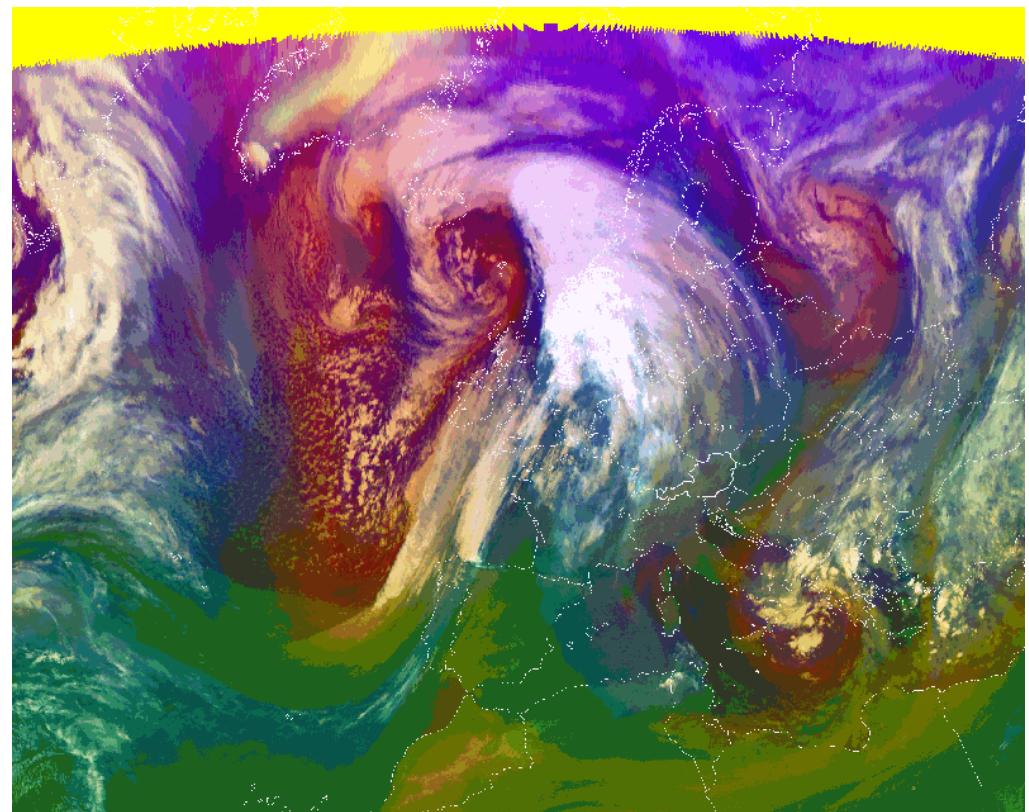
STE

stratos
column
ozone

surface
radiation

(Hegglin and Shepherd 2009)

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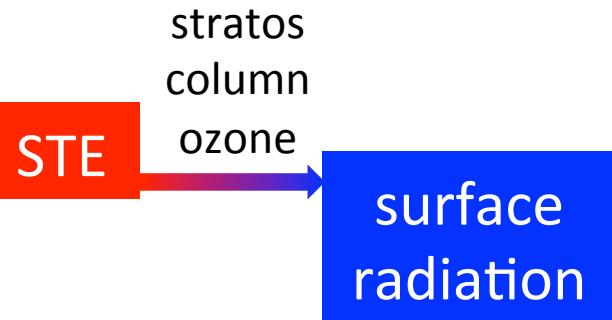
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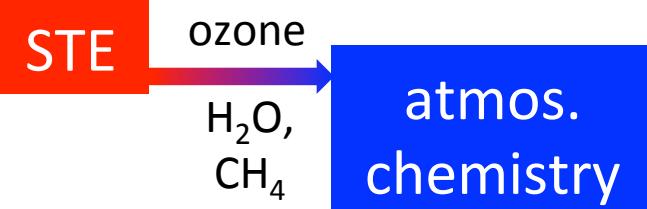
Introduction

Impact one:



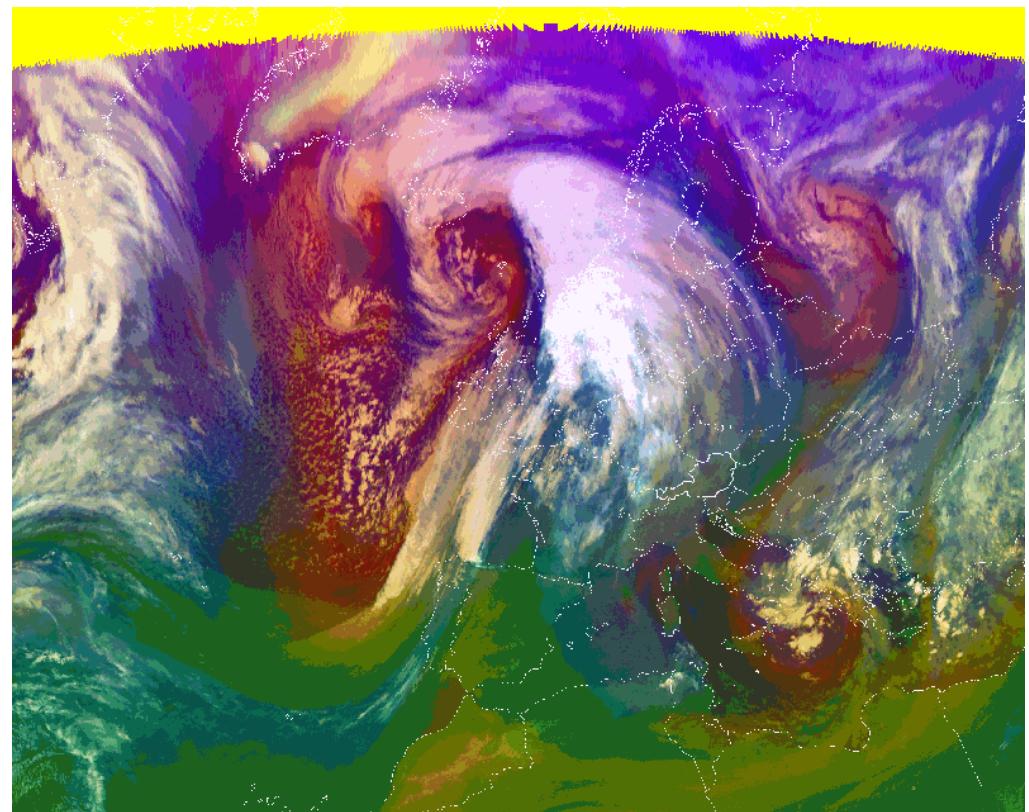
(Hegglin and Shepherd 2009)

Impact two:



(Kentarchos et al. 2003
Lin et al. 2012
Fiore et al. 2002)

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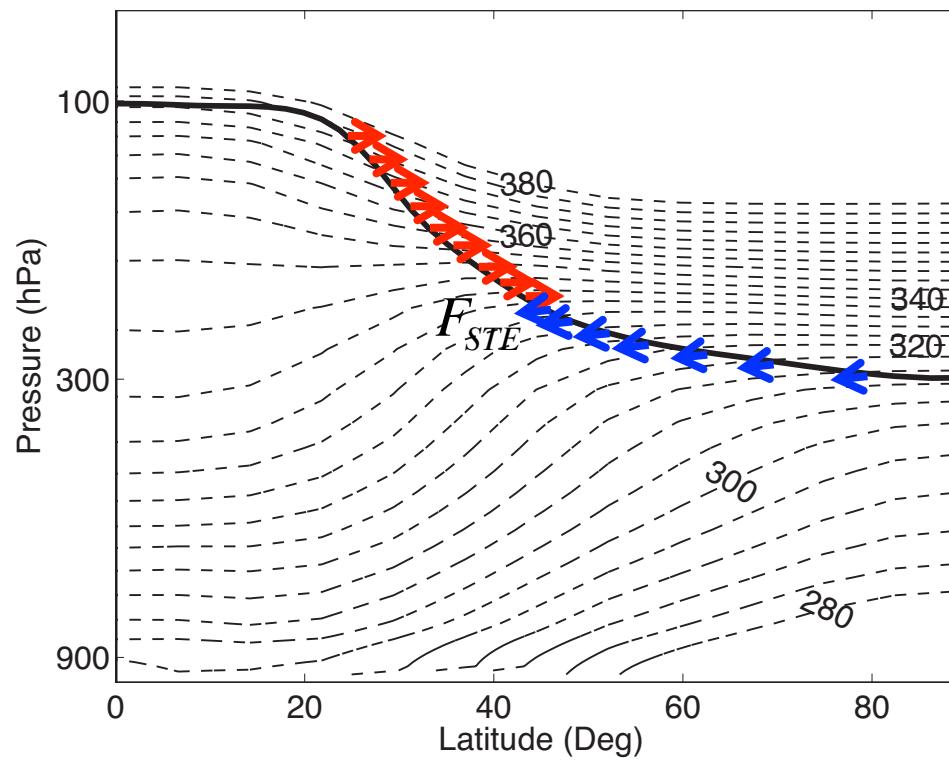
Green – tropical air

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Isentropic STE diagnostic

Nakamura, 2007

- diagnosing STE along individual isentropic surface (i.e. isentropic STE, F_{STE})
- a vertical series of isentropic STE approximately represents its meridional distribution



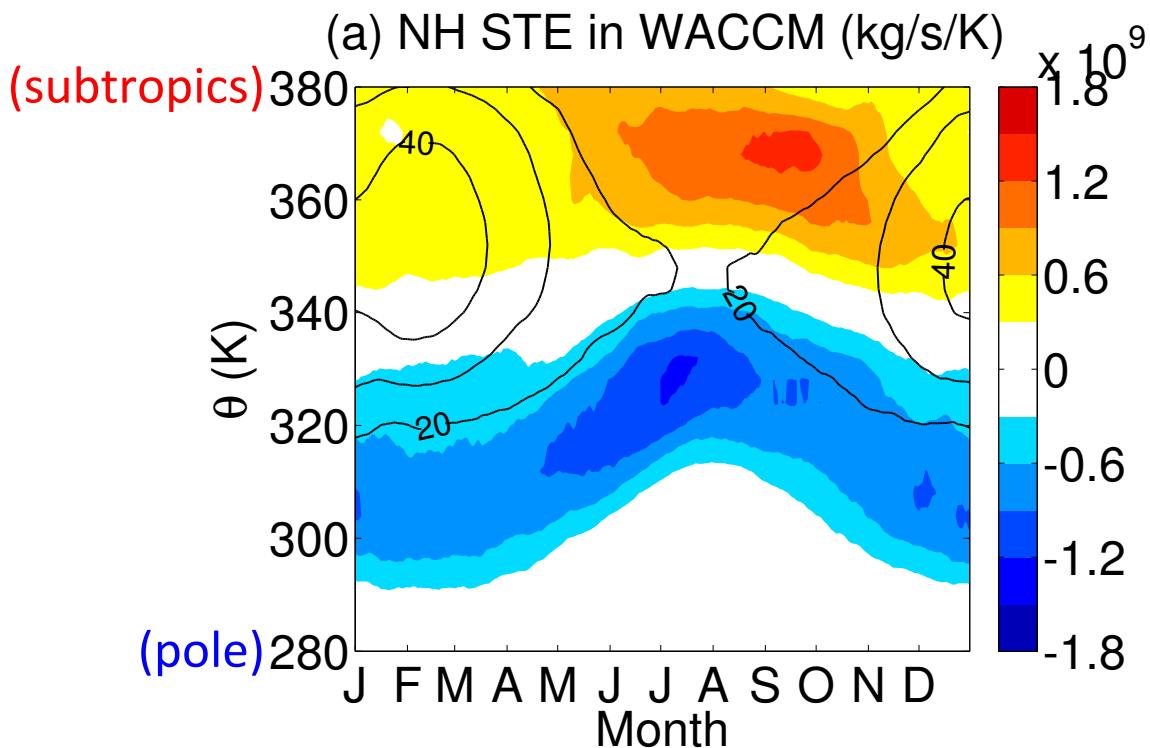
High isentropes <-> Low latitudes

Low isentropes <-> High latitudes

ISENTROPIC STE (COLORS) w/ ZONAL WIND (CONTOURS)

WACCM

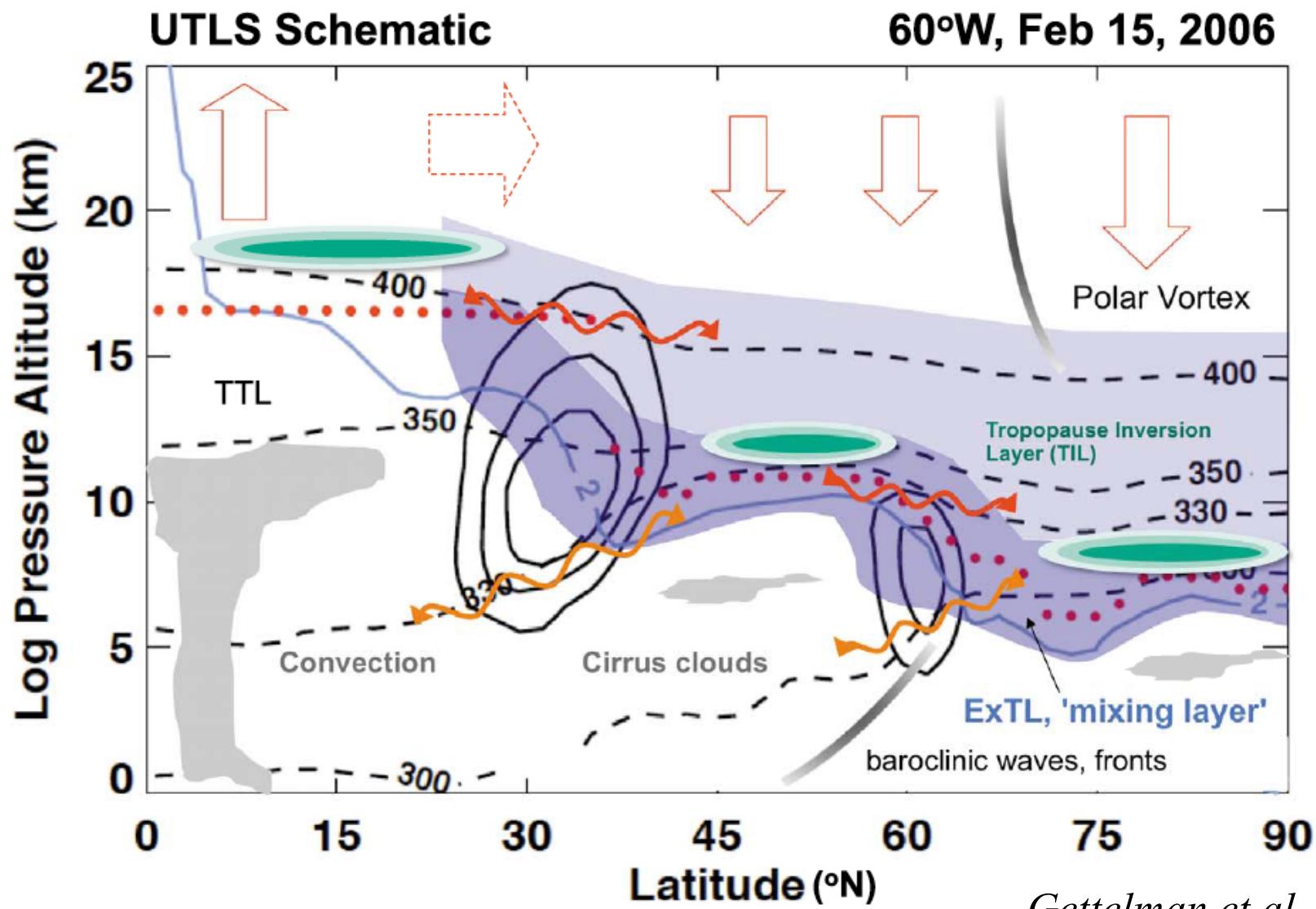
Whole Atmosphere
Community Climate
Model (1991-2009,
19-yr control run)
Focus on the **NH**, as
SH is similar.



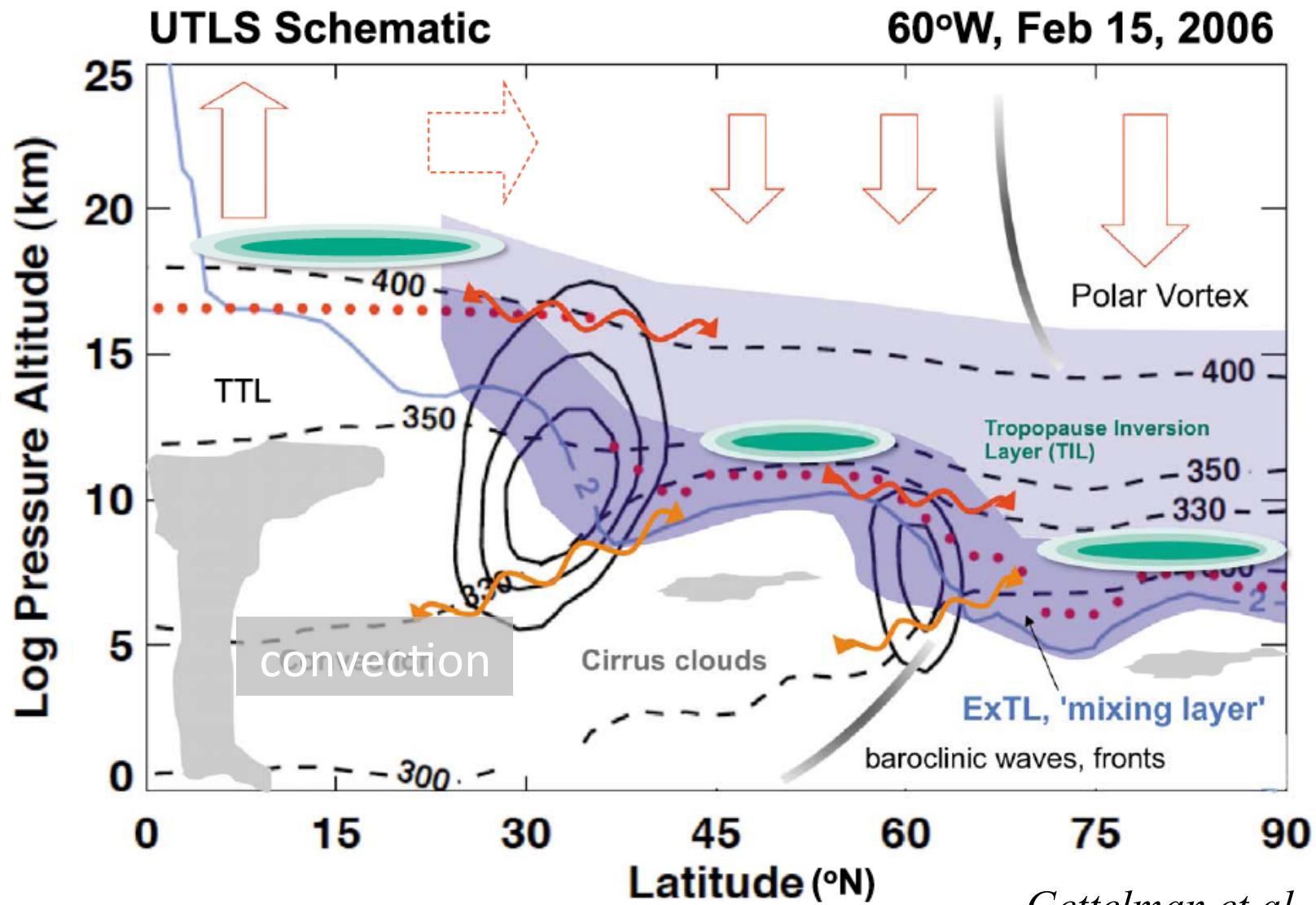
- Troposphere-to-stratosphere (upward) STE on higher isentropes (subtropics), stratosphere-to-troposphere (downward) STE on lower isentropes (extratropics)
- Maximum downward STE occurs on the poleward flank of the tropospheric jet, and moves seasonally with the jet

When and where STE occurs?

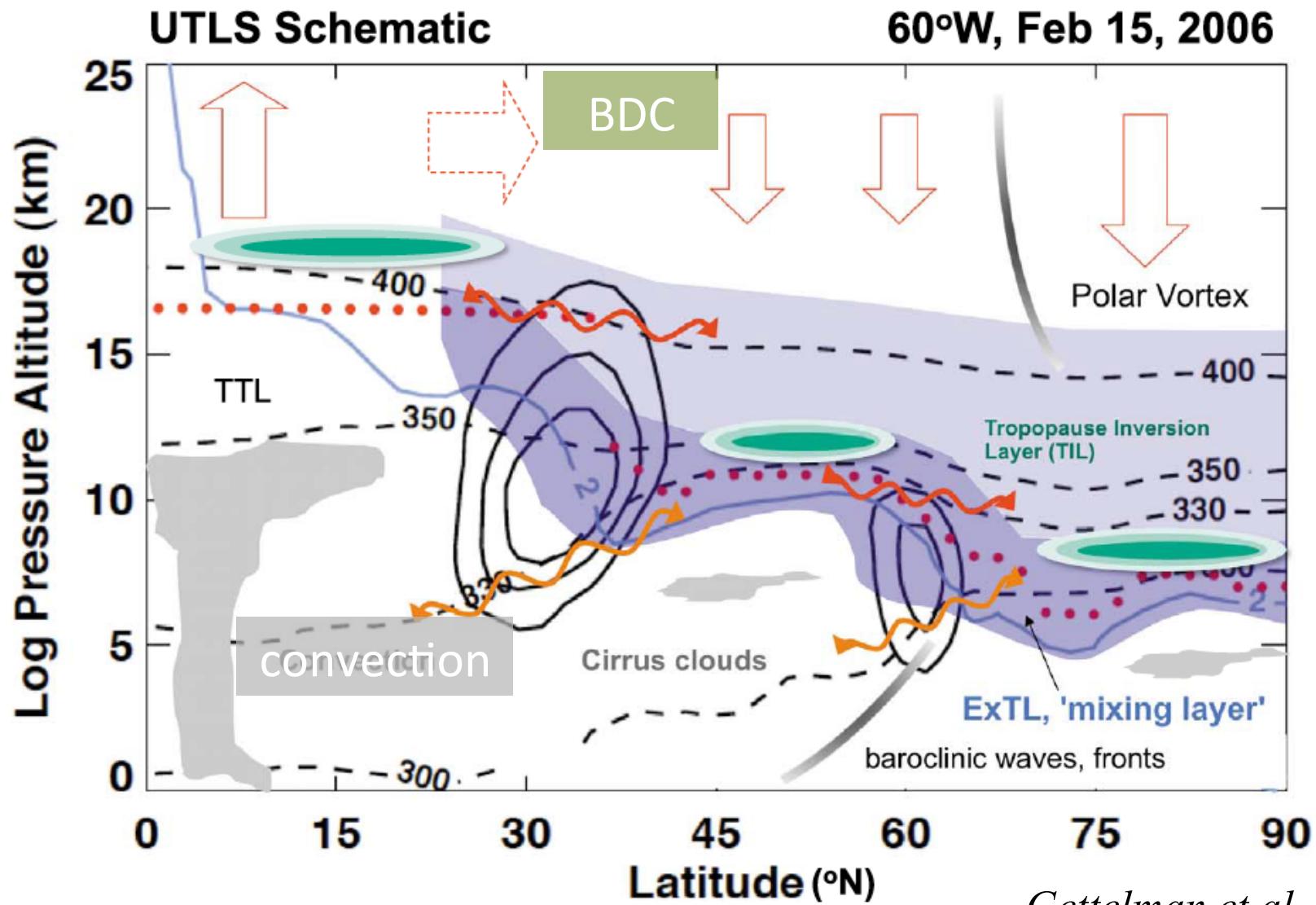
Processes controlling the STE



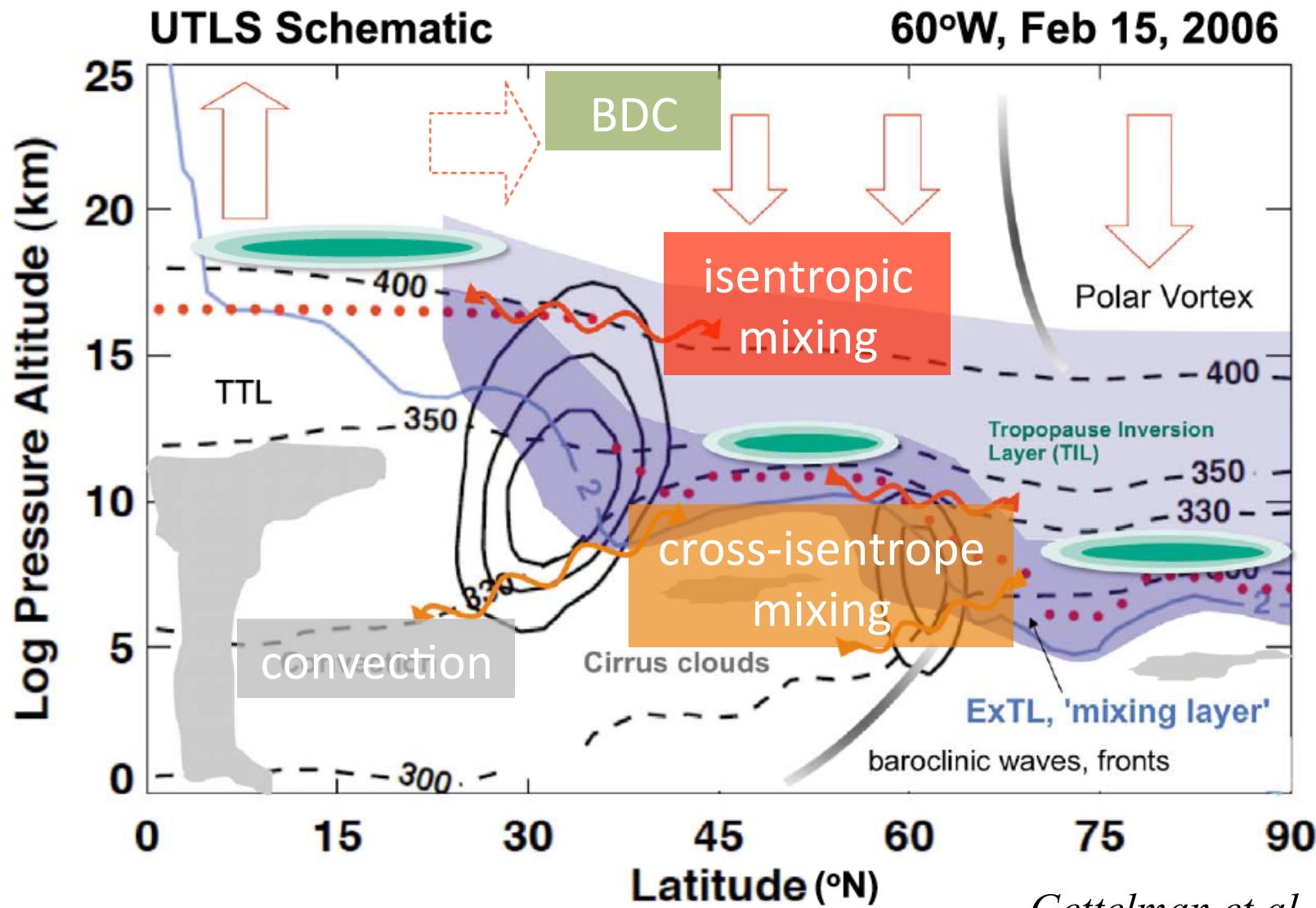
What cause this spatial distribution of STE?



What cause this spatial distribution of STE?



What cause this spatial distribution of STE?



Dynamic Partition – PV Sources

$$F_{STE} = -\frac{\partial M(\dot{q})}{\partial q} \Big|_{q=Q}$$

STE flux (F_{STE}) across a potential vorticity (PV) tropopause Q is affected by the PV tendency dq/dt .

where $M(\cdot) = \int_{STRATO} \sigma(\cdot) dS$ denotes the air mass weighted integration in the stratosphere

$$\dot{q} = \dot{q}_k + \dot{q}_s$$

Isentropic Differential
Mixing Diabatic Heating

$$F_{STE} = F_{mix} + F_{dia}$$

residual component

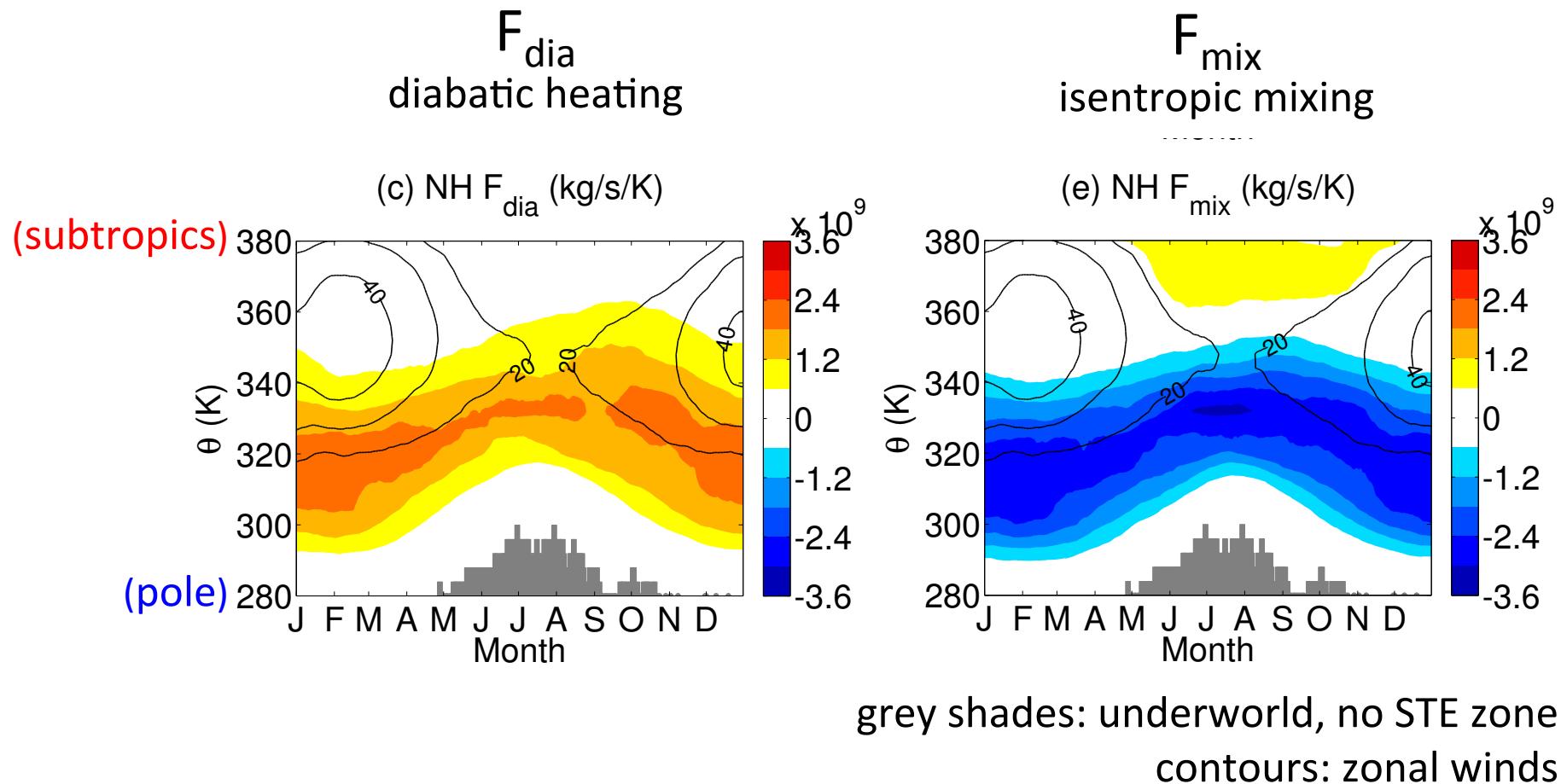
directly calculated

$$\text{where } F_{dia} = -\frac{\partial M(\dot{q}_s)}{\partial q} \Big|_{q=Q}$$

$$\text{and } \dot{q}_s = \frac{q}{\sigma} \frac{\partial}{\partial \theta} (\sigma \dot{\theta})$$

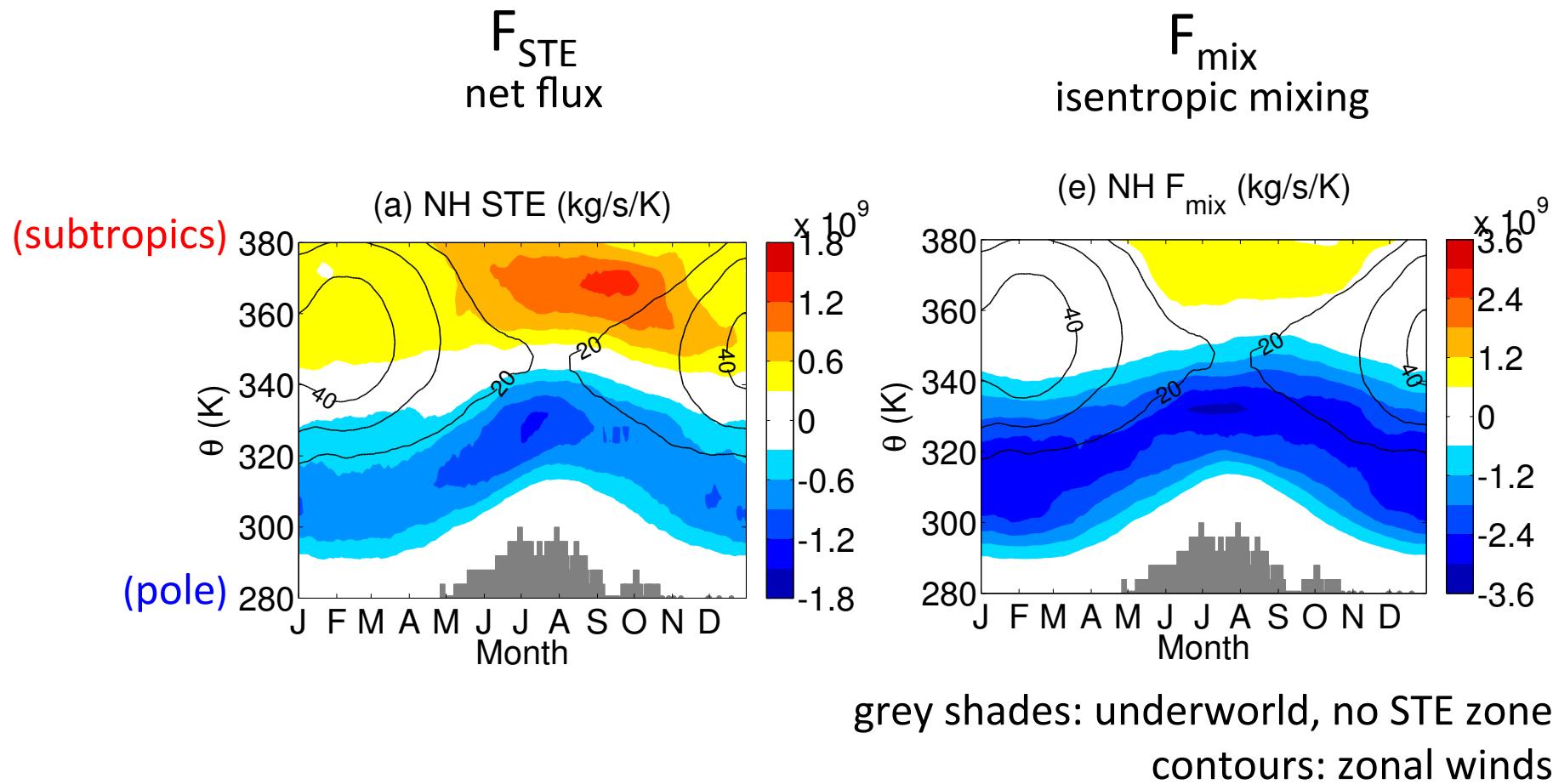
σ isentropic density
 θ isentropic temperature
 $\dot{\theta}$ diabatic heating rate

F_{dia} and F_{mix}



- Upward F_{dia} vs. downward F_{mix}
- Large cancellation

F_{STE} and F_{mix}

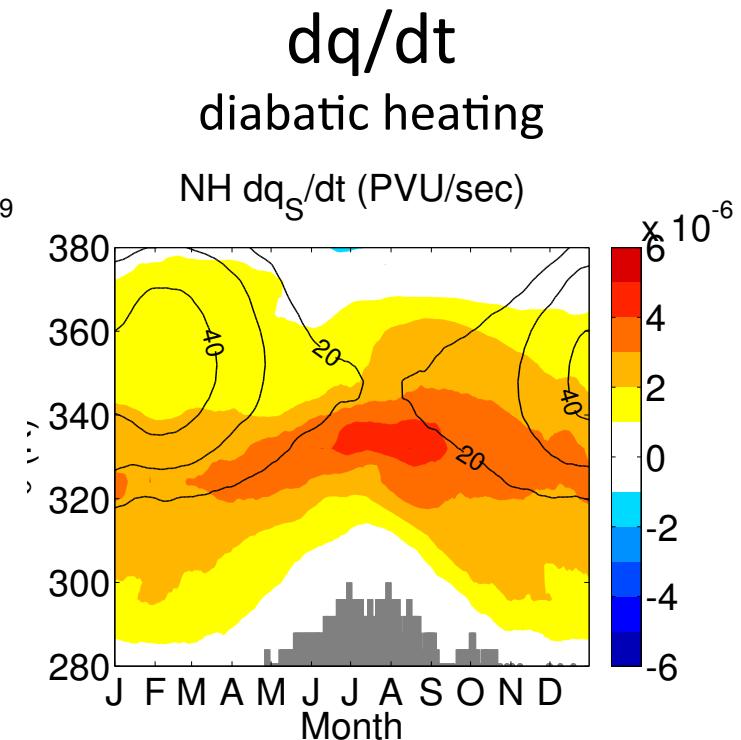
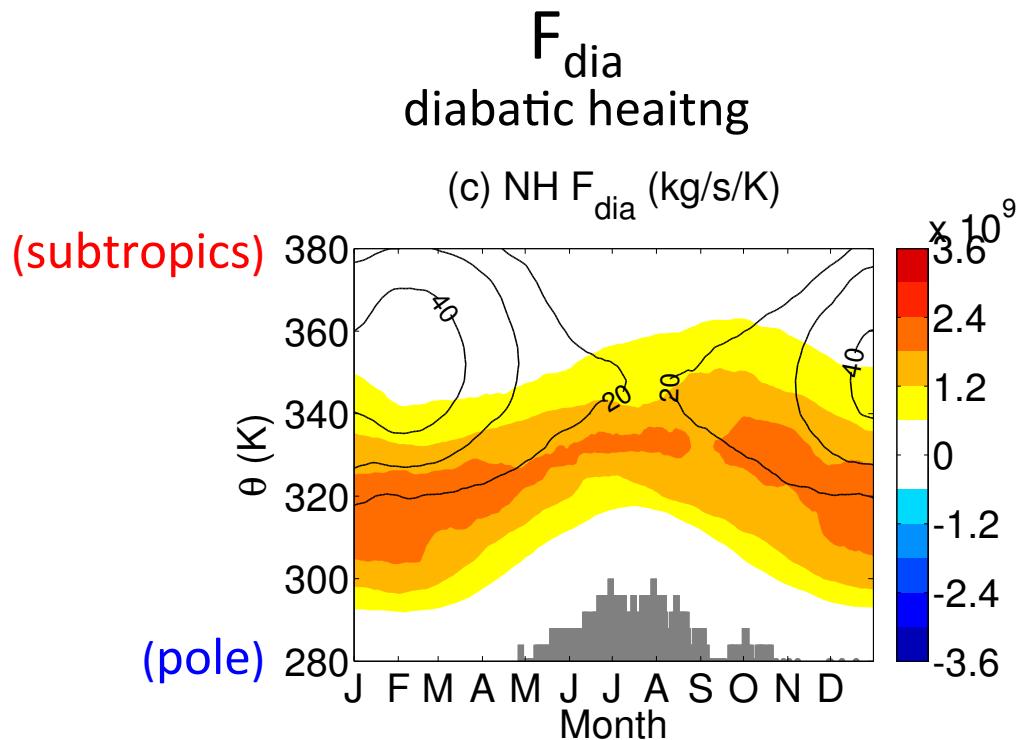


- $F_{\text{dia}} < F_{\text{mix}}$, the net flux F_{STE} displays a similar spatiotemporal pattern as F_{mix}

F_{dia} & diabatic heating

$$F_{dia} = -\frac{\partial M(\dot{q}_S)}{\partial q} \Big|_{q=Q} \propto \dot{q}_S \Big|_{q=Q}$$

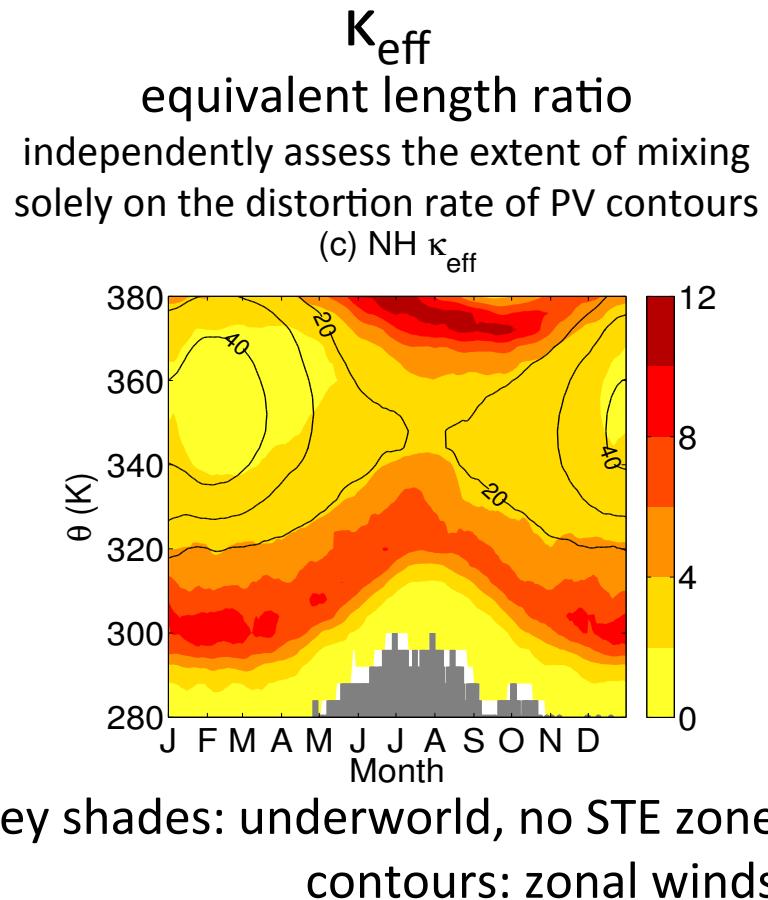
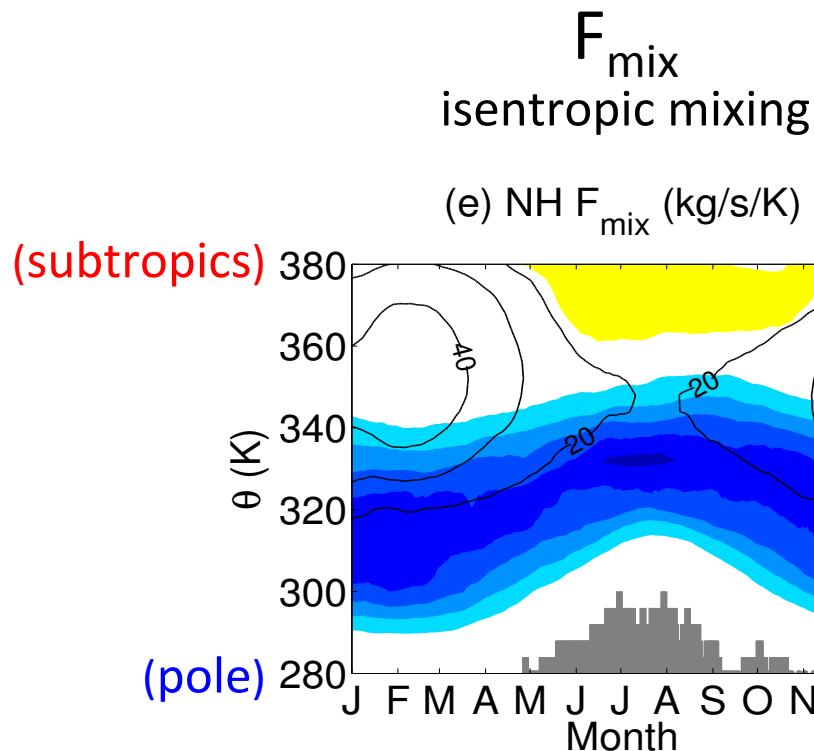
PV tendency due to
diabatic heating



grey shades: underworld, no STE zone
contours: zonal winds

- Strong $dq_S/dt \leftrightarrow$ strong F_{dia}

F_{mix} & isentropic mixing

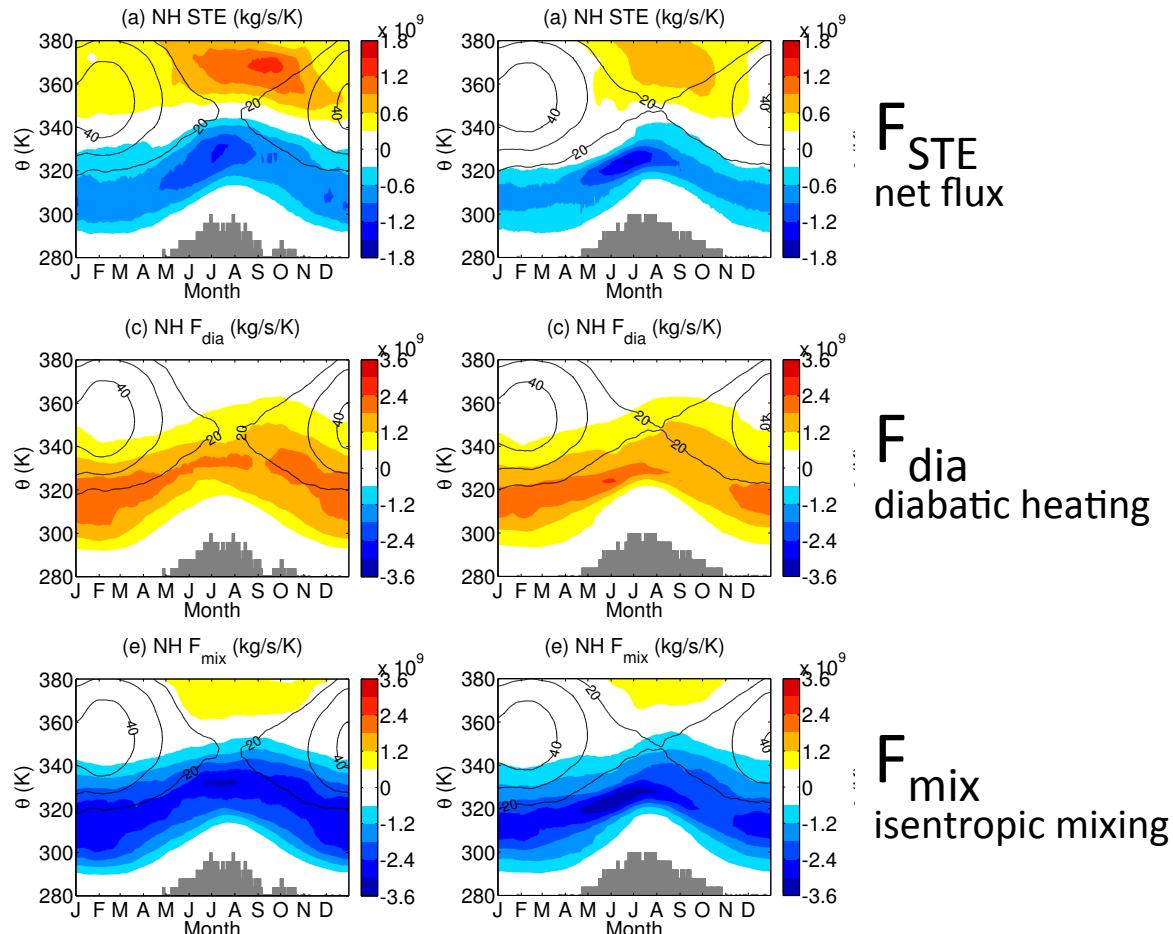


- Strong mixing \leftrightarrow strong F_{mix}
- Poleward flank of jet: weak winds, strong mixing, and strong F_{mix}
- Jet core: strong winds, weak mixing, and weak F_{mix}

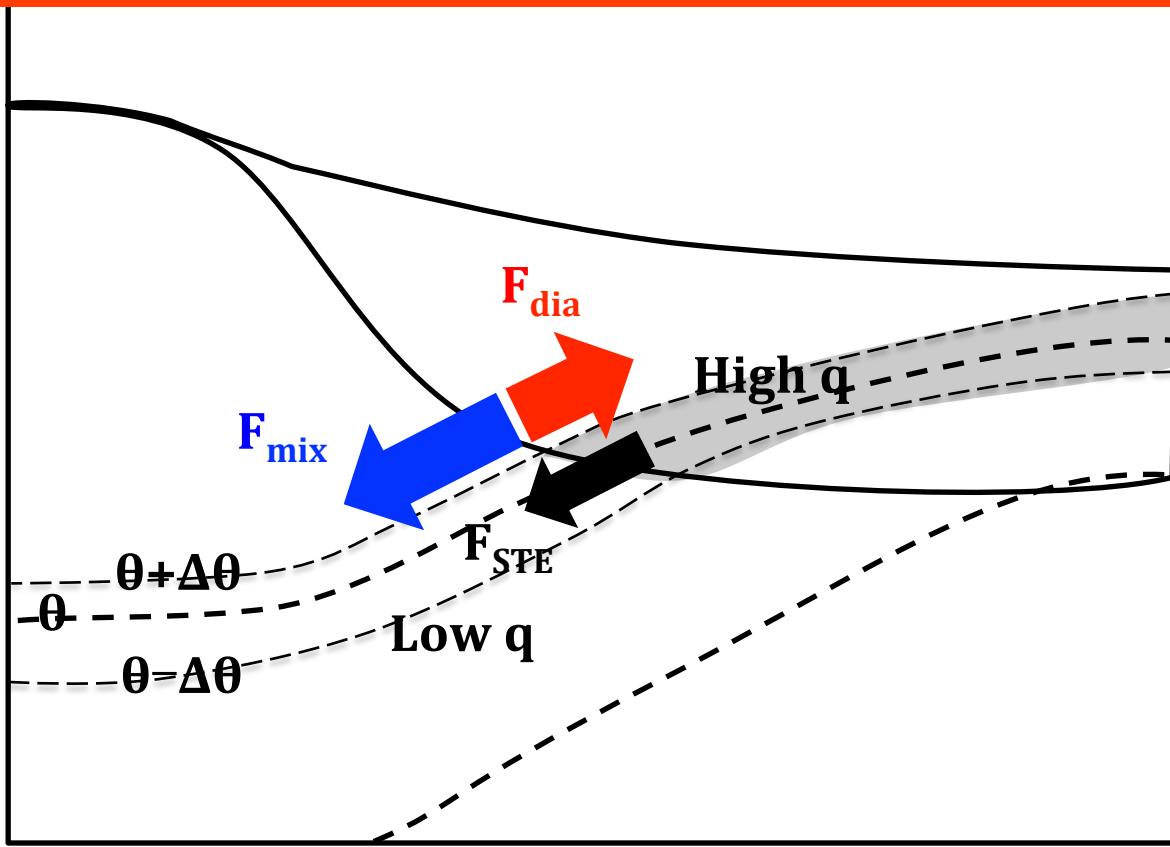
Consistency in the CMAM

WACCM CMAM

Whole Atmosphere Community Climate Model
(1991-2009, 19-yr control run) Canadian Middle Atmosphere Model
(1981-2010, 30-yr control run)



Summary



Equator

Pole

- Upward F_{dia} vs. downward F_{mix} , yielding downward F_{STE}
- F_{dia} – diabatic heating: positive dq/dt , descending tropopause height
- F_{mix} – isentropic mixing: negative dq/dt , ascending tropopause height