

NONISOTROPIC AEROSOL SCATTERING EFFECTS ON LONGWAVE IRRADIANCE

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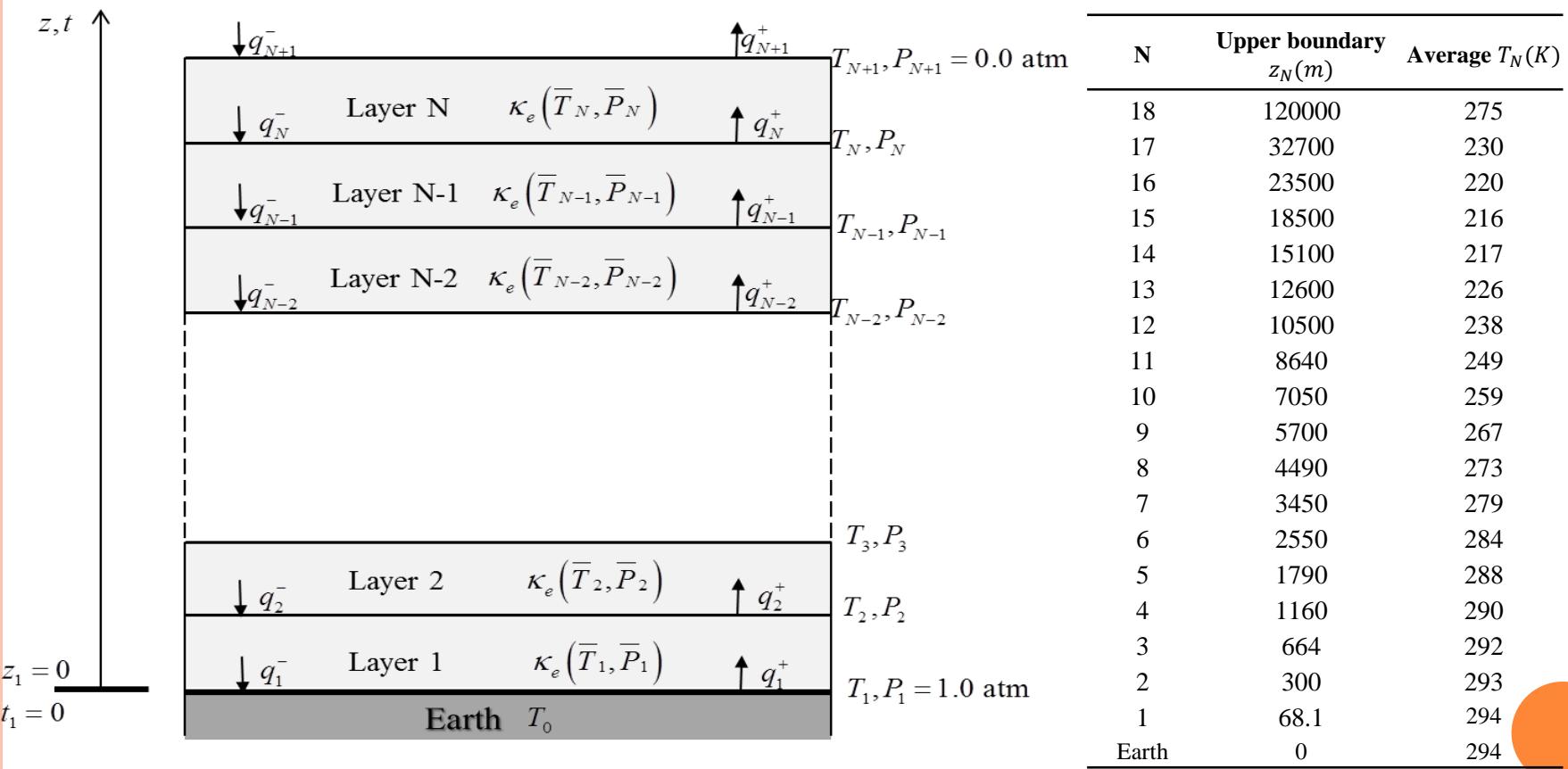
INTRODUCTION

- Solving thermal radiative heat transfer problem in atmosphere in order to
 - Modeling climate change
 - Design solar concentrated power plants
 - Predict solar energy
- Analytical atmospheric model: isotropic assumption
- This study:
 - Monte Carlo approach
 - Mie scattering
 - Scale anisotropic effects



ATMOSPHERIC MODEL

- 18-layer model



ATMOSPHERIC MODEL

- Pressure and temperature in each layer

$$\sigma_n = \frac{2N - 2n + 1}{2N}$$

$$\bar{P}_n = \sigma_n^2 (3 - 2\sigma_n)$$

$$P_n = \frac{1}{N} \sum_{j=n}^N dP_j = \frac{1}{N} \sum_{j=n}^N 6(\sigma_j - \sigma_j^2)$$

$$\bar{T}_n = \frac{T_n(P_n - \bar{P}_n) + T_n + T_{n+1}(\bar{P}_n - P_{1+n})}{P_n - P_{n+1}}$$

- AFGL profiles used for temperature and pressure profile
- HITRAN molecular spectral data for seven gases used to calculate absorption coefficients.



MATHEMATICAL MODEL

- Radiative transfer equation (Boltzmann equation) in a scattering medium:

$$\mu \frac{\partial I(\kappa, \mu)}{\partial \kappa} + I(\kappa, \mu) = (1 - \tilde{\rho})I_b(T) + \frac{\tilde{\rho}}{2} \int_{-1}^1 P(\mu, \mu') I(\kappa, \mu') d\mu'$$

- For plane-parallel, isotropic scattering layers, radiosity, J and irradiance, G can be used:

$$J = (1 - \tilde{\rho})\pi I_b - \tilde{\rho}\pi I$$

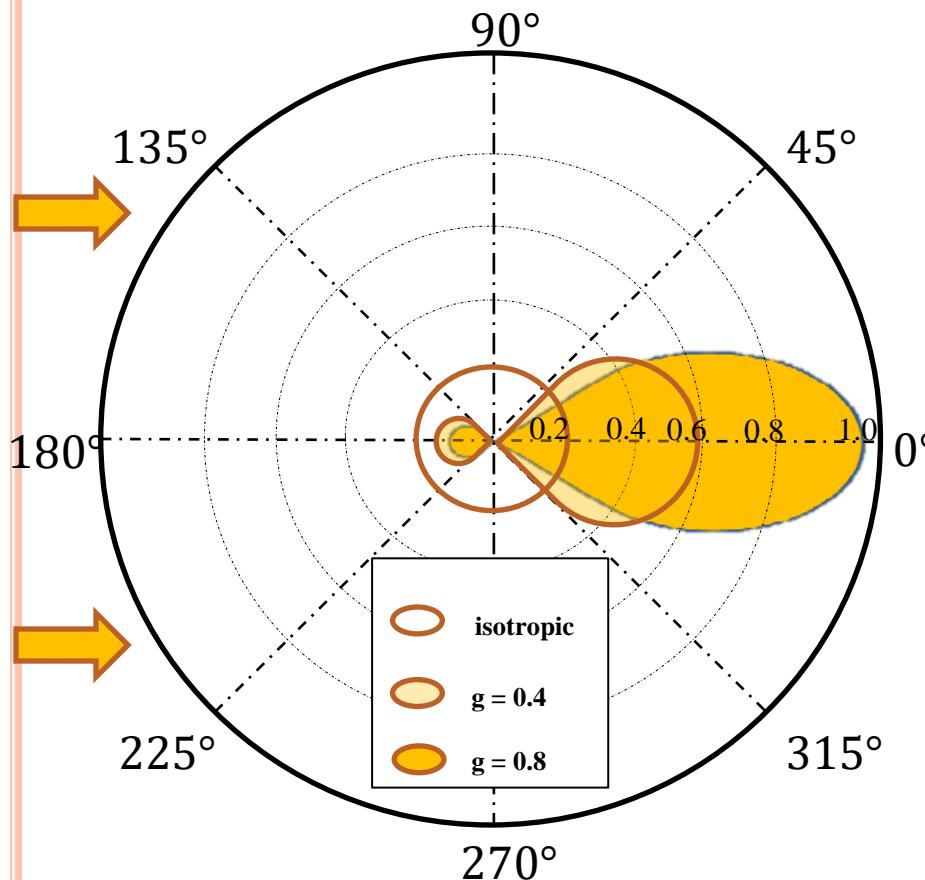
$$G = \int_0^s e^{-t_{s'} J(s')} ds'$$

- Calculate J and G in each layer by solving a matrix constructed by transfer factors and modified transfer factors

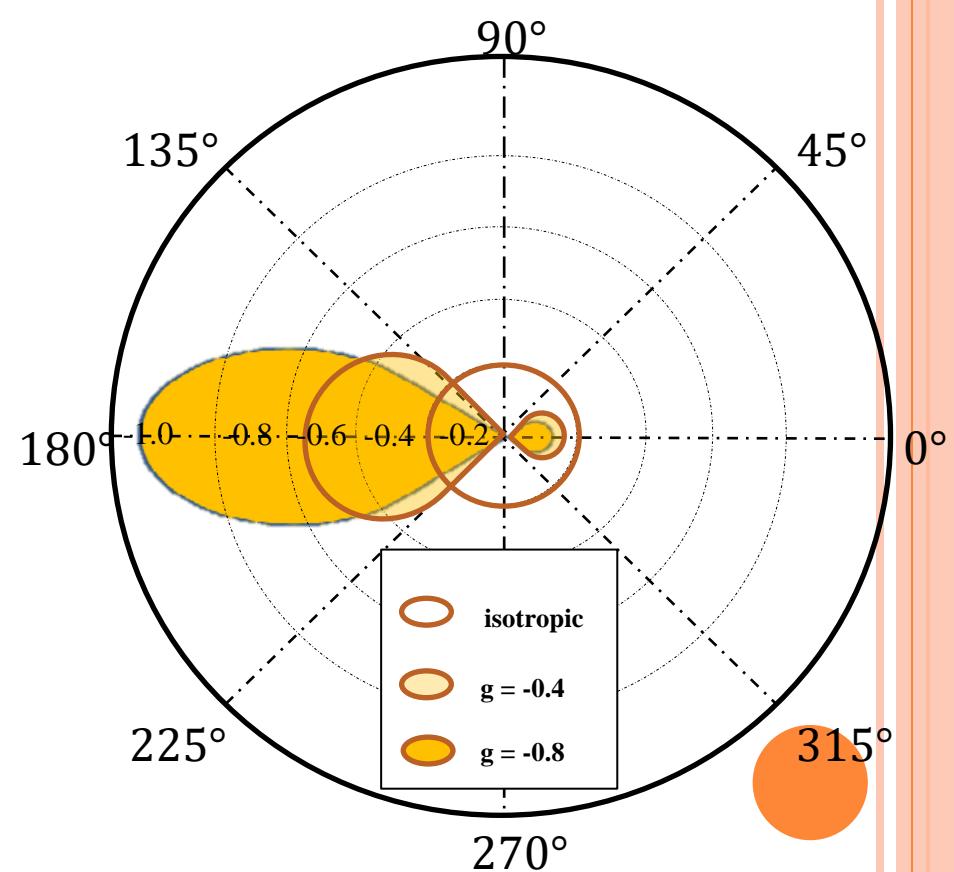


ANISOTROPIC SCATTERING

(a) Forward scattering



(b) Backward scattering



ANISOTROPIC SCATTERING

- Mie theory:

- Used to calculate absorption and scattering coefficients
- Used to determine phase functions
 - Simplified phase function: Henyey-Greenstein

$$P(\cos \theta) = \frac{1 - g^2}{2(1 + g^2 - 2g \cos \theta)^{3/2}} \quad \theta \in [0, \pi]$$

- Scale gross error of final results

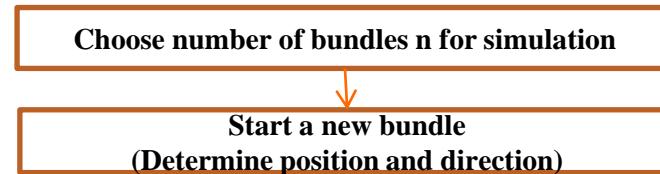
- P-1 approximation scaling

- Scale initial inputs (absorption and scattering coefficients and albedo) to an isotropic scattering medium:

$$\begin{aligned}\hat{\kappa} &= (1 - \tilde{\rho}g)\kappa \\ \hat{\tilde{\rho}} &= \frac{\tilde{\rho}(1 - g)}{1 - \tilde{\rho}g}\end{aligned}$$



METHOD – MONTE CARLO



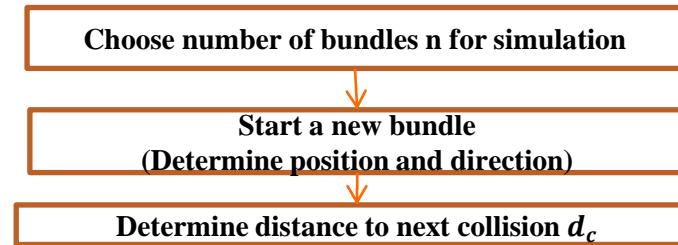
METHOD – MONTE CARLO

- Initialization:
 - Initial angle:
 - Surface: $\theta = \cos^{-1}(\sqrt{\xi_\theta})$ $\varphi = 2\pi\xi_\varphi$
 - Volume: $\theta = \cos^{-1}(2\xi_\theta - 1)$ $\varphi = 2\pi\xi_\varphi$
 - Initial position: random position within an layer
 - Optical depth: $\tau_0 = -\log(1 - \xi)$ since $I \propto e^{-\tau}$
 - Energy contained in each bundle:
 - $E_{bs} = \frac{\varepsilon\sigma T_s^4}{N}$
 - $E_{bg} = \frac{4\kappa_a \Delta z}{N} \int_{\Delta\nu} I_{b\nu}(\nu, T) d\nu$

Where, $I_{b\nu}(\nu, T) = \frac{2hc^2\nu^3}{e^{h\nu c / (\kappa_B T g)} - 1}$



METHOD – MONTE CARLO



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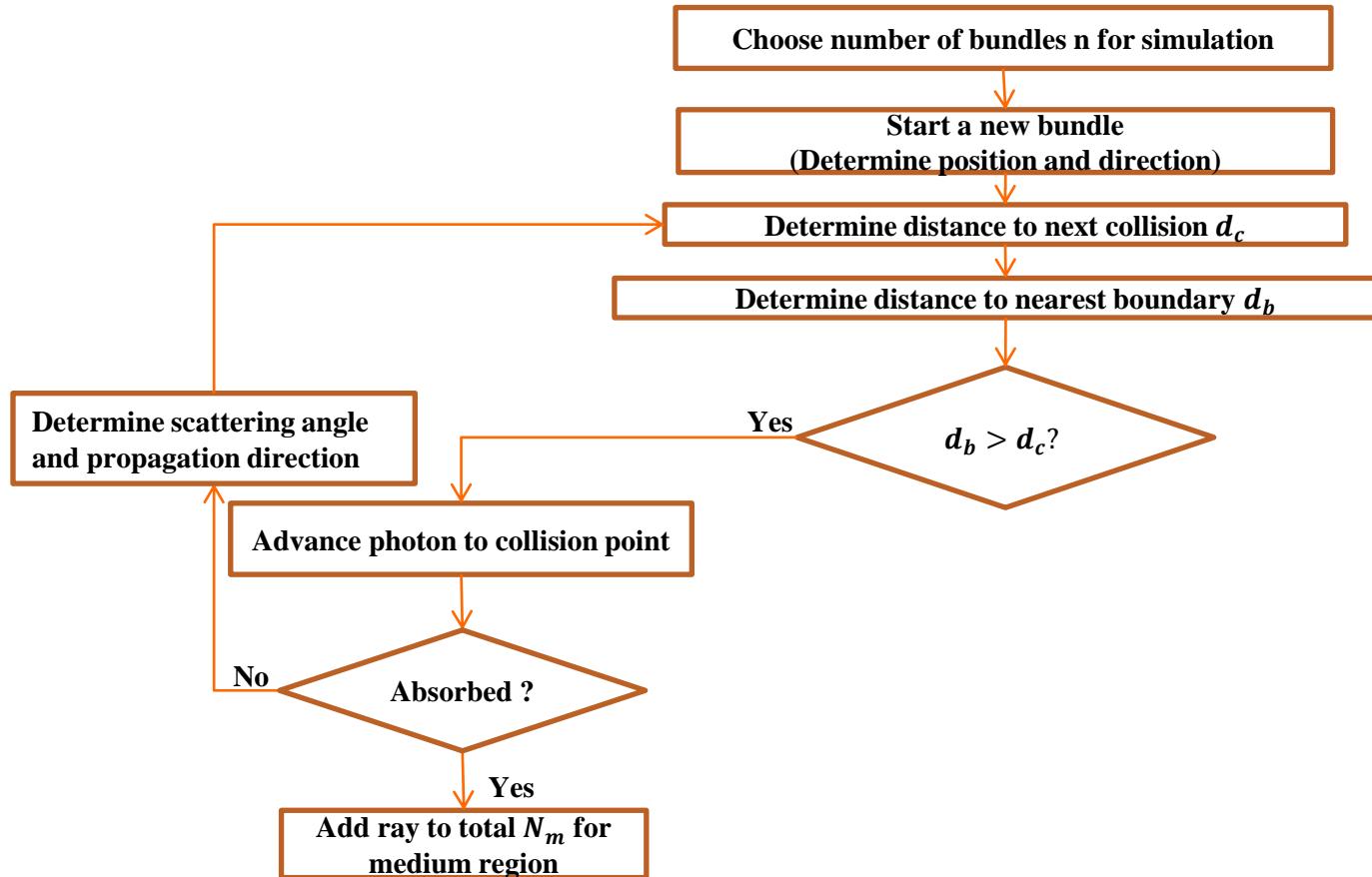
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- Tracing
 - Determine position of interaction
 - $d_c = -\frac{\log(1-\xi)}{\kappa_e}$



METHOD – MONTE CARLO



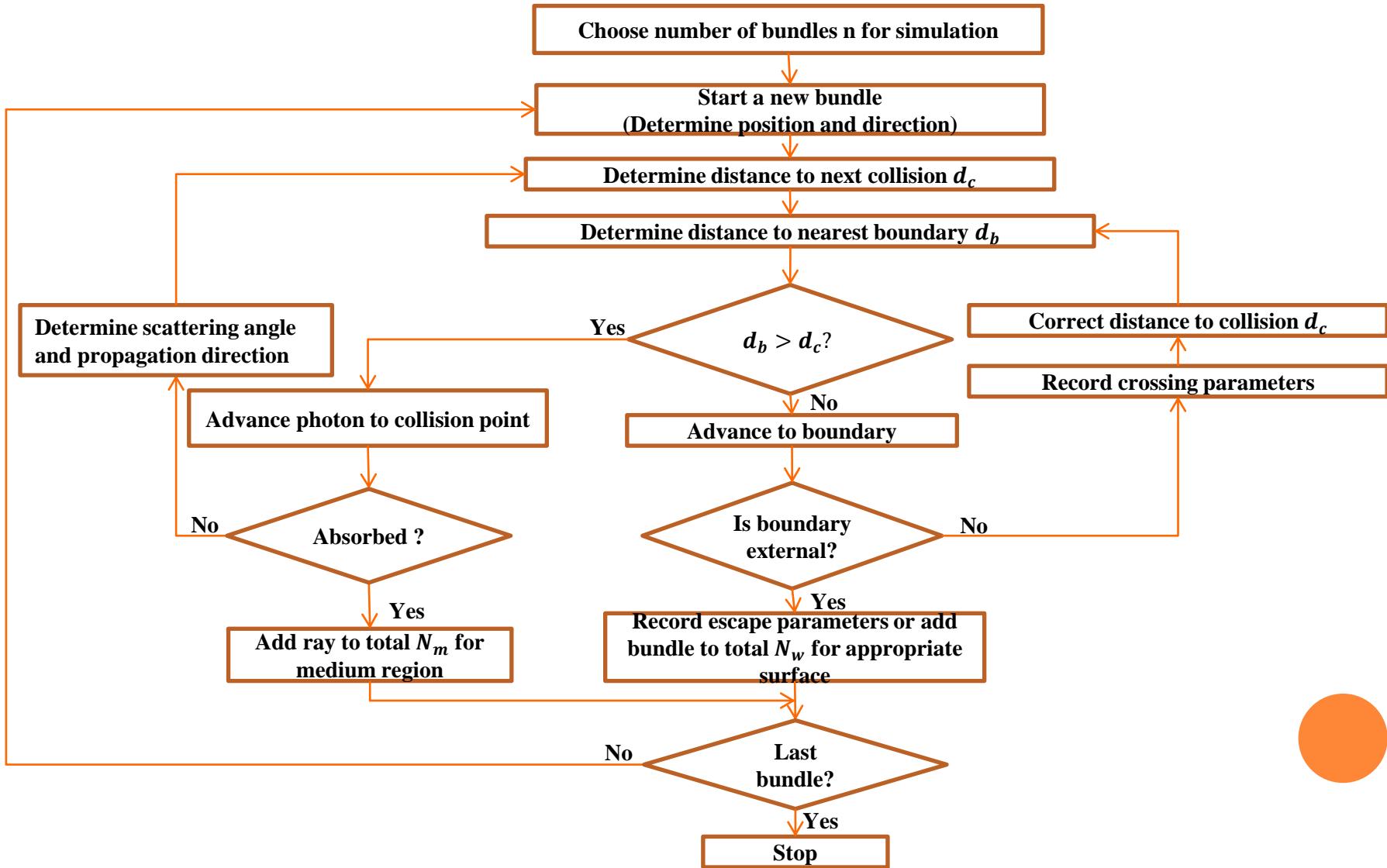
METHOD – MONTE CARLO

- Tracing

- Determine position of interaction
 - $d_c = -\frac{\log(1-\xi)}{\kappa_e}$
- Determine which particle to interact:
 - According to relative ratio of extinction coefficients
- Determine particle destiny:
 - Absorbed: $\xi < (1 - \alpha)$
 - Scattered: $\xi \geq (1 - \alpha)$
 - $\alpha = \frac{\kappa_s}{\kappa_e}$ is single albedo



METHOD – MONTE CARLO



METHOD – MONTE CARLO SCATTERING

Scattering angle:

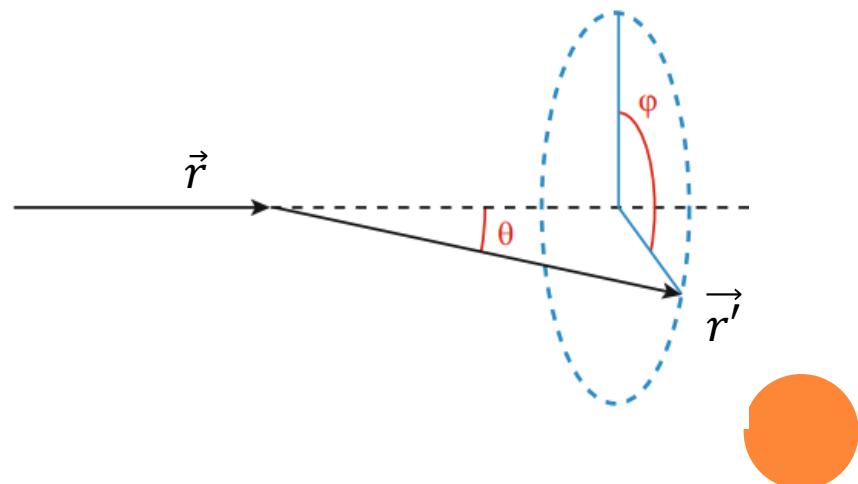
$$\cos \theta = \frac{1}{2g} \left\{ 1 + g^2 - \left[\frac{1 - g^2}{1 - g + 2g\xi} \right]^2 \right\}$$

New Direction:

$$r'_x = r_x \cos \theta - \frac{\sin \theta}{\sqrt{1 - r_z^2}} (r_x r_z \cos \varphi + r_y \sin \varphi)$$

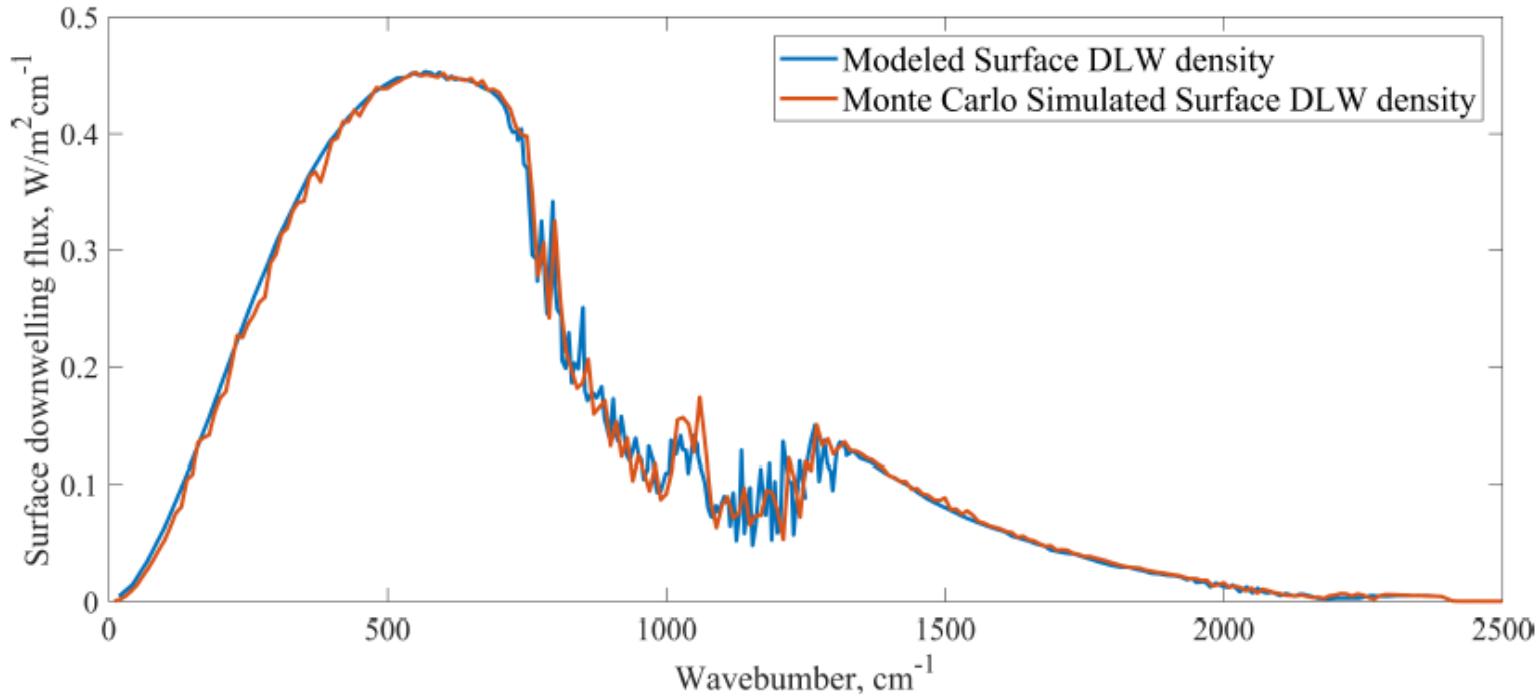
$$r'_y = r_y \cos \theta - \frac{\sin \theta}{\sqrt{1 - r_z^2}} (r_y r_z \cos \varphi - r_x \sin \varphi)$$

$$r'_z = r_z \cos \theta + \sqrt{1 - r_z^2} \sin \theta \cos \varphi$$



RESULTS – MODEL VALIDATION

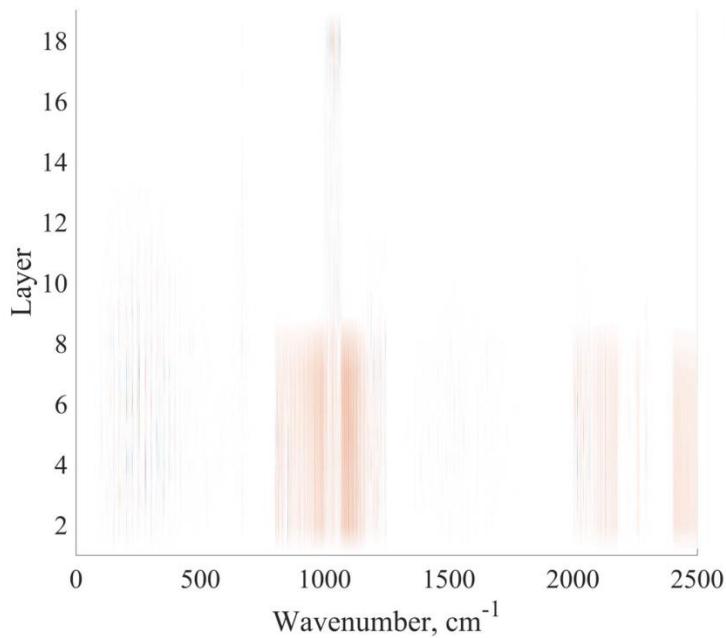
- Compare aerosol-free case
 - Simulation result and ICRCCM model result



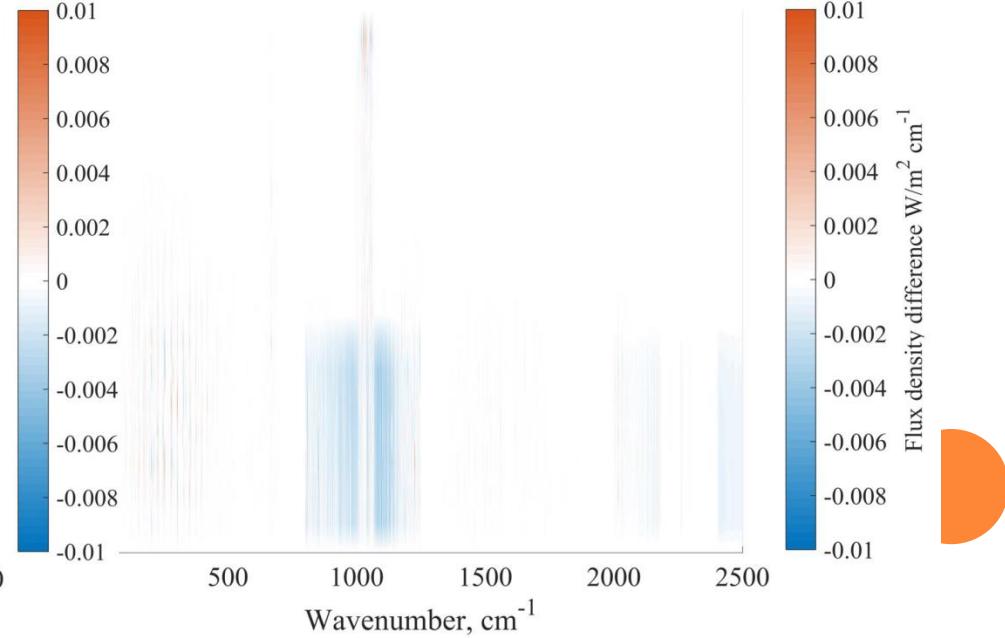
RESULTS – SPECTRAL DIFFERENCE

- The difference of spectral DLW radiative flux density on each layer boundary between a highly forward/backward scattering case, and an isotropic case.

(a) Spectral downwelling longwave radiative flux on each layer boundary in highly backward scattering case ($g = -0.9$)

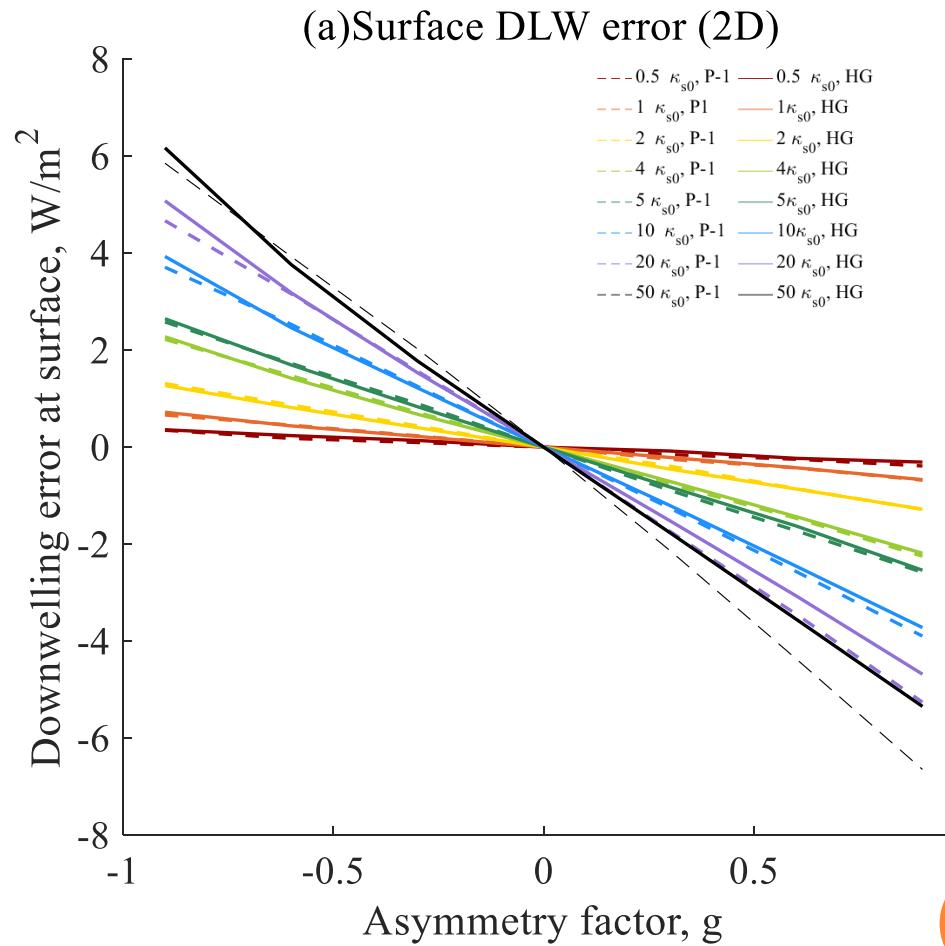


(b) Spectral downwelling longwave radiative flux on each layer boundary in highly forward scattering case ($g = 0.9$)



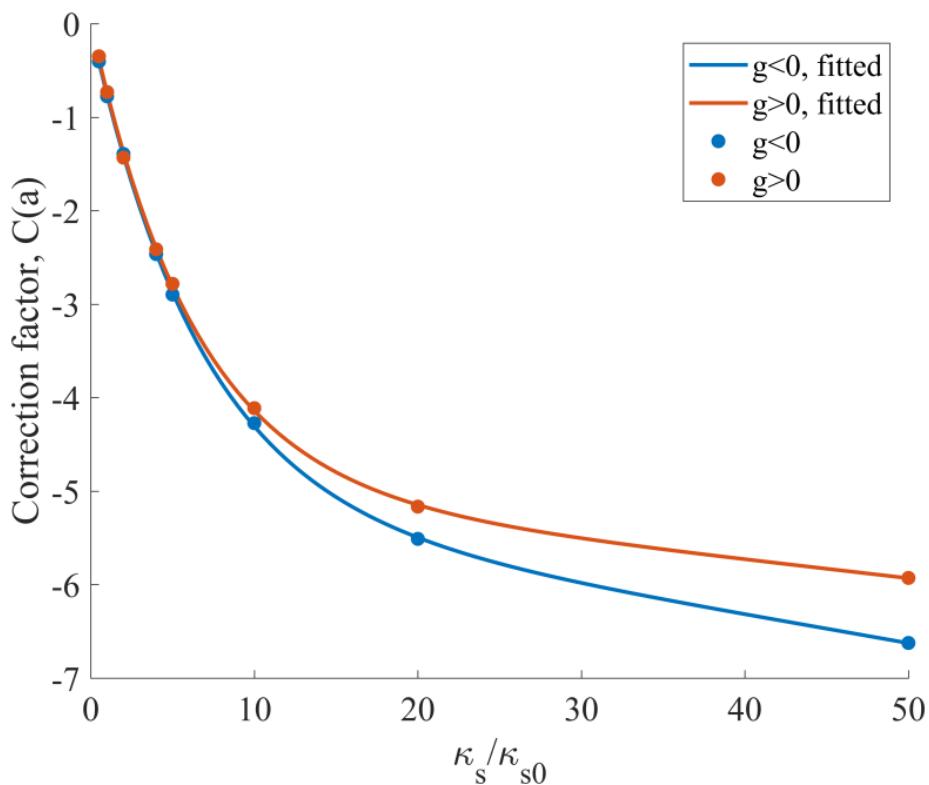
RESULTS

- Downwelling longwave radiation flux errors (W/m²) from isotropic case that vary with asymmetry factor, g and relative aerosol amount



RESULTS

$$C(a) = \frac{\text{DLW}^* - \text{DLW}}{g} = \begin{cases} -5.02e^{0.00335a} + 5.03e^{-0.157a} & \text{for } g \geq 0 \\ -5.27e^{0.00461a} + 5.24e^{-0.147a} & \text{for } g < 0 \end{cases}$$



CONCLUSION

- With specific aerosol concentration and temperature profile, effect can be scaled to isotropic scattering case to assist analytical calculation.

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- P-1 scaling gives pretty good approximation when aerosol concentration is small.

