



I. MOTIVATION AND OBJECTIVES

Motivation:

Wide observational and theoretical evidence exists that ocean spray strongly affects hurricane dynamics thermodynamically and mechanically by modifying the vertical enthalpy and momentum fluxes, respectively. It has been shown that both these influences of ocean spray essentially depend on the size of spray droplets. Hence, accounting for the **droplet size distribution** is important for accurate description and modeling of the spray effect on hurricane dynamics.

Objectives:

- Develop a rigorous and consistent **Eulerian multi-fluid** model of a Marine Atmospheric Boundary Layer (MABL) laden with **polydisperse** spray.
- Carry out an in-depth study of the ocean spray effect on hurricane dynamics using numerical and theoretical analyses of the MABL model.

IV. MATHEMATICAL MODEL OF SPRAY-LADEN MABL

Multi-fluid Eulerian model of MABL laden with polydisperse sea spray

Spray continuity

$$\frac{d}{dz} \left(d_i \frac{ds_i}{dz} + s_i a_i \right) = \frac{1}{\Delta r} (f_{i+1} - f_i) - \dot{q}_i s_i - a_{0i} s_{0i} \delta_{z-z_w} \quad i = 1 \dots N$$

(N=10 for present calculations)

Moisture conservation

$$\frac{d}{dz} \left(k_a \rho_a \frac{dq}{dz} \right) = \rho_w \sum_{i=1}^N \dot{q}_i s_i \quad \leftarrow \text{specific humidity}$$

Momentum conservation

$$\frac{d}{dz} \left(k_a \rho_a \frac{du_a}{dz} \right) = \rho_w \sum_{i=1}^N \frac{s_i (u_a - u_i)}{\tau_{di}} \quad \leftarrow \text{horizontal wind speed}$$

$$\frac{d}{dz} \left(k_i s_i \frac{du_i}{dz} \right) - s_i w_i \frac{du_i}{dz} + \frac{s_i}{\tau_{di}} (u_a - u_i) + (u_i^1 - u_i) a_{0i} s_{0i} \delta_{z-z_w} = 0 \quad \leftarrow \text{horizontal spray motion}$$

$$\frac{d}{dz} \left(\frac{4}{3} k_i s_i \frac{dw_i}{dz} - \frac{2}{3} s_i e_i \right) - s_i w_i \frac{dw_i}{dz} - s_i g + \frac{s_i}{\tau_{di}} (w_i^d - w_i) + (w_i^1 - w_i) a_{0i} s_{0i} \delta_{z-z_w} = 0$$

$$\frac{dP}{dz} = -\rho g \quad \leftarrow \text{vertical spray motion}$$

Thermal energy conservation

$$\frac{d}{dz} \left(k_a \rho_a \frac{d\theta}{dz} \right) = \sum_{i=1}^N \gamma_i (T_a - T_i) \quad \leftarrow \text{potential air temperature}$$

$$\frac{d}{dz} \left(s_i d_i \frac{dT_i}{dz} + H_{z-z_w} s_i a_i T_i \right) = [\gamma_i (T_i - T_a) - l \dot{q}_i] s_i - a_{0i} s_{0i} T_s \delta_{z-z_w} \quad \leftarrow \text{spray temperature}$$

E-epsilon (k-epsilon) 1.5 order model

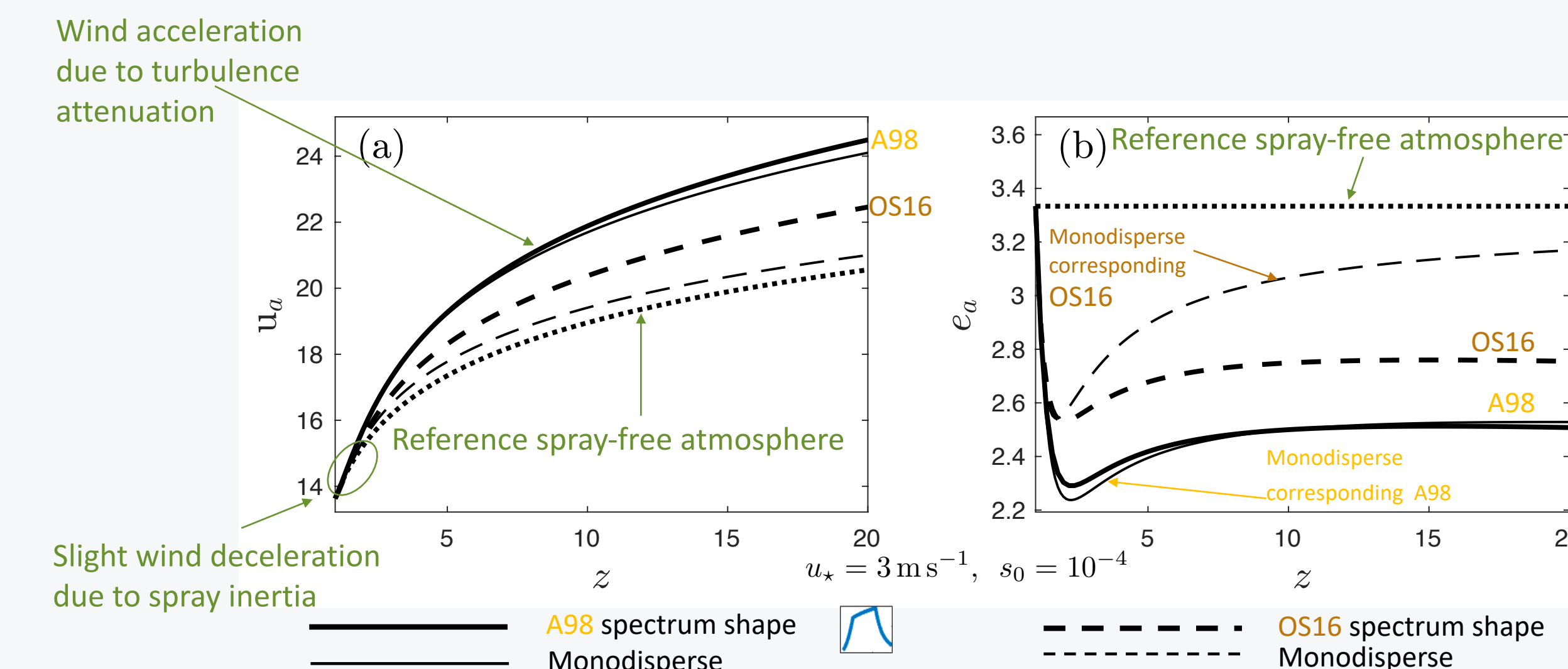
$$\alpha_e \frac{d}{dz} \left(\rho_a k_a \frac{de_a}{dz} \right) = \rho_a (\epsilon_a - p_a) - q_{ae}$$

$$\alpha_\epsilon \frac{d}{dz} \left(\rho_a k_a \frac{d\epsilon_a}{dz} \right) = \rho_a (d_{ae} - p_{ae}) - q_{ae}$$

+ various **constitutive relationships**

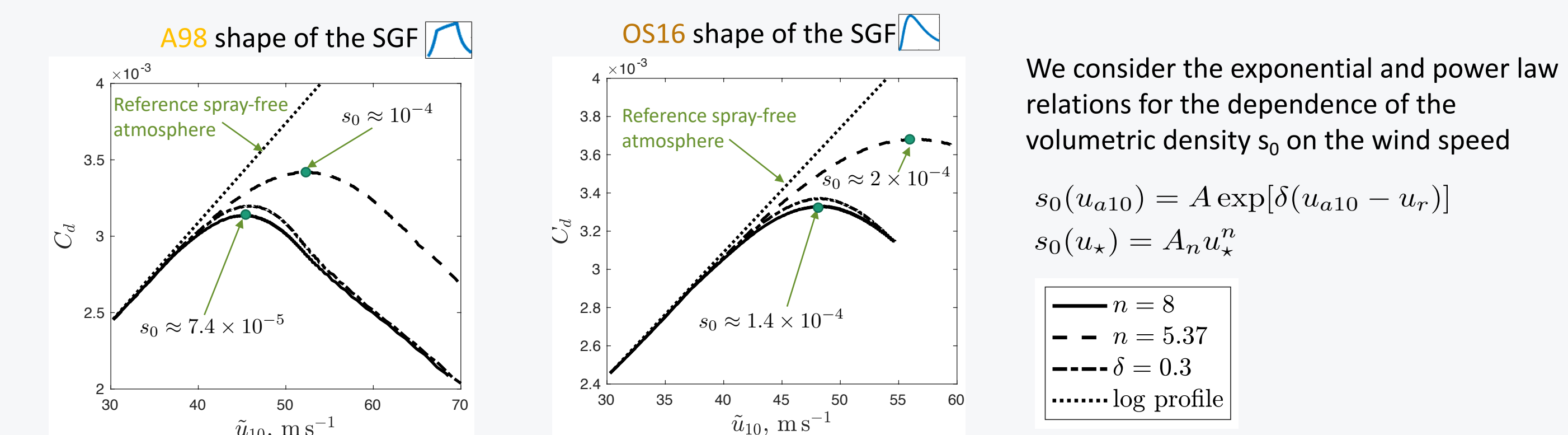
VI. NUMERICAL SIMULATION: MECHANICAL EFFECT OF SPRAY

Horizontal velocity (u_a) and TKE (ϵ_a) of air vs altitude (z)



- Due to the larger fraction of small droplets, sea spray with the **A98** spectrum shape suppresses air turbulence more efficiently than that with the **OS16** counterpart for the same value of the volumetric spray production intensity s_0 causing greater acceleration of the wind above the wave crest level.
- Monodisperse spray assumption leads to a quantitative underprediction of the wind speed.

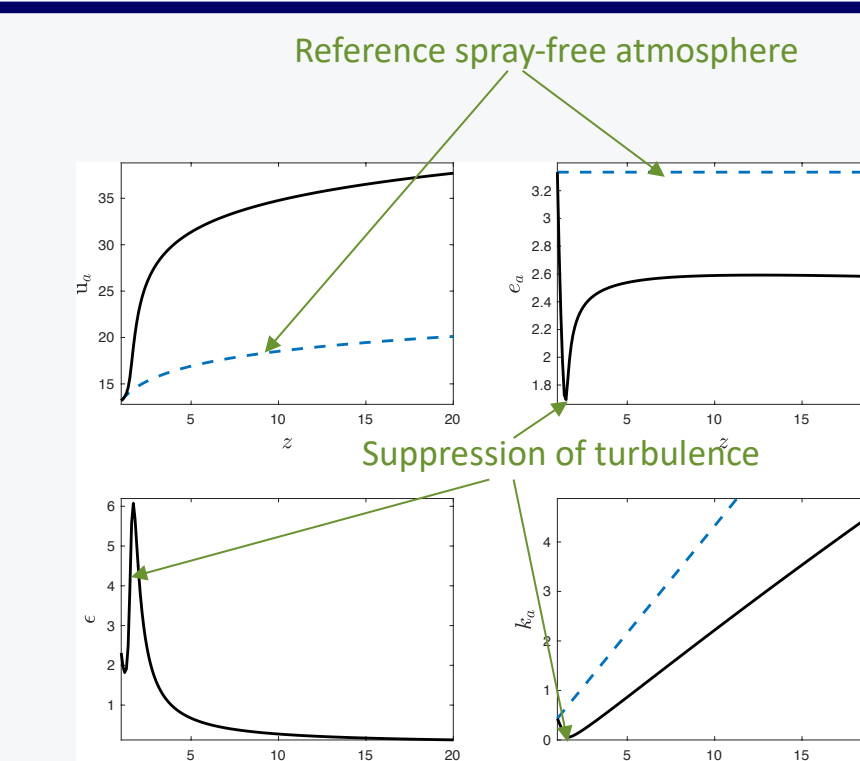
C_d vs u_{10} for different correlation laws between s_0 and u . (u_{a10})



The spray-induced reduction of drag causes the wind to accelerate above the wave crest level. The reduction of the drag coefficient (C_d) occurs even **at relatively low spray concentrations**. C_d reaches its maximum at the volumetric spray concentration $\sim 10^{-4}$, which corresponds to the spray mass fraction of $\sim 10\%$.

Collapse of turbulence in a thin layer (“slip sheets”)

Numerical simulations show that thin layers (~ 1 -2 m), where turbulence is strongly suppressed, develop a few meters above the wave crest level for large values of spray volumetric density s_0 and the friction velocity u_* and the **A98** spectrum. The presence of these structures causes a rapid decrease in the air-sea drag coefficient followed by a strong increase in the wind speed.



II. EFFECTS OF OCEAN SPRAY ON TROPICAL CYCLONE

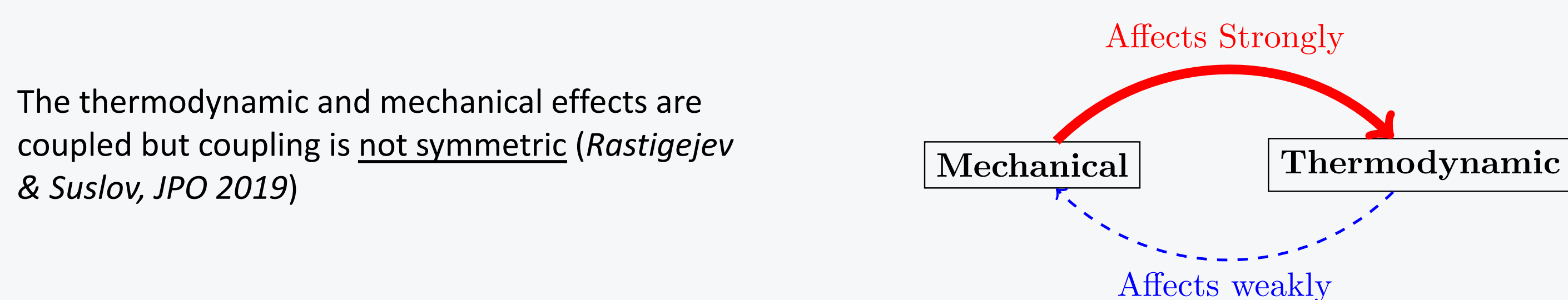
The effect of the ocean spray on the hurricane dynamics is **two-fold**:

I. Thermodynamic (Rastigejev & Suslov 2016, 2019). Evaporating sea spray modifies latent and sensible heat fluxes in a MABL.

II. Mechanical (Rastigejev & Suslov 2022). Sea spray affects the vertical momentum transport.

Accelerating effect due to turbulence attenuation by the spray --- Droplet Slip (DS) and “Gravity Lubrication” (GL)

Decelerating effect due to spray inertia.



The thermodynamic and mechanical effects are coupled but coupling is **not symmetric** (Rastigejev & Suslov, JPO 2019)

Hence, the mechanical effect can be considered **separately**.

III. MECHANICAL EFFECT OF SPRAY

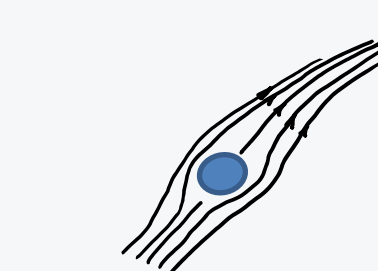
The decelerating effects of spray inertia



Horizontal velocities of injected droplets are much smaller than that of the airflow,
 $u_i^1 \ll u_a$

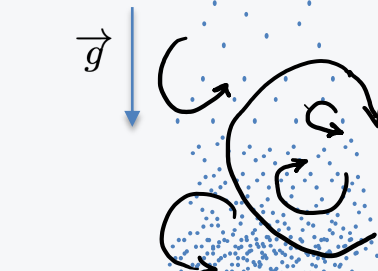
The accelerating effects of spray due to turbulence suppression, two mechanisms

I. Droplet slip (DS) causes air-droplet “friction” that leads to turbulence attenuation



Turbulent flow loses its energy due to air-droplet friction $\Rightarrow$ turbulent transport becomes less intensive $\Rightarrow$ reduction in air-ocean drag $\Rightarrow$ wind acceleration

II. “Gravity Lubrication” (GL) due to upward transport of spray by the turbulent flow



Turbulent flow lifts spray $\Rightarrow$ it works against gravity $\Rightarrow$ loses its energy $\Rightarrow$ turbulent transport becomes less intensive $\Rightarrow$ reduction in air-ocean drag $\Rightarrow$ wind acceleration

DS>GL

V. VOLUMETRIC SPRAY GENERATION FUNCTION (VSGF)

VSGF

$$f_v(u_c, r) = s_0(u_c) p(r)$$

u_c – characteristic speed of the airflow, $u_c = u_*$ or $u_c = u_{a10}$
 r – droplet radius

VSGF Volumetric density of spray in quiescent air shape function

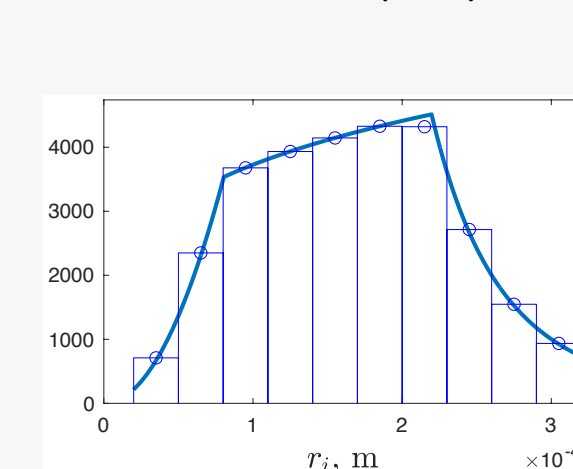
Large uncertainty over spray production rates:

- Six orders of magnitude in field experiments (the wind speeds $\lesssim 20$ m/s)
- Two-three orders of magnitude in laboratory experiments (the wind speeds $\lesssim 60$ m/s)

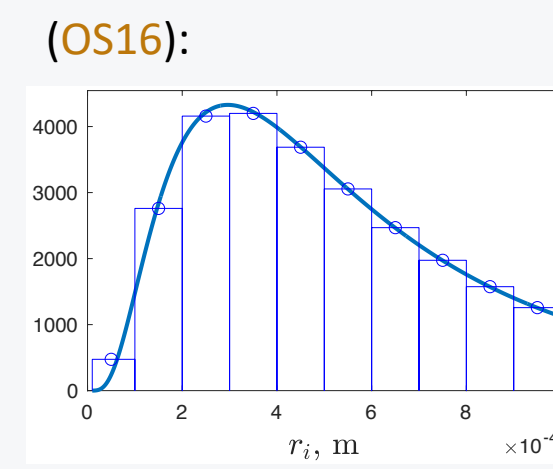
How to relate: Laboratory $\Rightarrow$ Field ??? 2-3 orders of magnitude difference

Shape of VSGF, $p(r)$

I. Andreas 1998 (A98):



II. Ortiz-Suslow et al. 2016 (OS16):



Spray production intensity $s_0(u_c)$

$$s_0(u_*) = A_n u_*^n \quad (\text{Fairall et al. 2009 and others})$$

$n \sim 2 - 16$

$$s_0(u_{a10}) = A \exp[\delta(u_{a10} - u_r)]$$

$$\delta \sim 0.2 - 0.875 \quad (\text{Monahan 1986, Wu 1993})$$

VII. CONCLUSIONS

- ❑ The Eulerian multi-fluid model has been developed to describe a marine atmospheric boundary layer laden with polydisperse ocean spray. The model considers droplets of various sizes and air as separate interacting turbulent continua.
- ❑ This approach enables us to account for the turbulence attenuation due to air-droplet slip (DS) and due to the “gravity lubrication” (GL), and the effect of spray inertia.
- ❑ The acceleration due to DS is much stronger than due to GL and occurs even for low volumetric spray concentrations $s_0 \sim 10^{-4}$ ($\sim 10\%$ spray mass mixing ratio). This can explain a substantial air-ocean drag reduction detected in field observations.
- ❑ Two “shapes” of the volumetric spray generation function, A98 and OS16, have been studied. The fraction of large droplets in OS16 is much higher than that in A98. The shapes yield substantially different values of the air-sea drag coefficient C_d for volumetric spray concentrations $s_0 \gtrsim 10^{-4}$.
- ❑ Polydisperse spray results are generally not quantitatively comparable to those obtained using a simpler monodisperse-spray approach. In particular, the monodisperse spray model may underestimate the spray-induced wind acceleration.
- ❑ Numerical simulations show that thin layers, where the turbulence is strongly suppressed (turbulent viscosity is small), may develop a few meters above the wave crest level for sufficiently large values of volumetric spray density s_0 and friction velocity u_* . Such layers effectively act as slip sheets between the air and the ocean. The presence of these structures causes a rapid decrease in the air-sea drag coefficient and an increase in wind speed.

ACKNOWLEDGEMENTS

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