



### **Suitable Background Error Covariances for Radar Reflectivity Direct Assimilation in the Rapid Refresh Forecast System (RRFS)**

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### Outline

- Motivation to improve static background error covariance for radar reflectivity direct assimilation
- Introduction of two methods
  - 1. Ensemble-based Tangent Linear Model (ETLM)
  - 2. Convective-scale Static Background Error Covariance (CSB)
- Cycling tests and discussion



### How to Assimilate Radar Reflectivity

The radar reflectivity observation has a lot of potential to improve precipitation forecasts. However, it is not necessarily used effectively in the operational data assimilation systems.

- Cloud Analysis (e.g., Albers et al. 1996)
  - After data assimilation without reflectivity, hydrometeors and thermodynamical variables are adjusted based on reflectivity.
- 1D+3DVar (e.g., Caumont et al. 2010)
  - Atmospheric variables retrieved from reflectivity with 1DVar are assimilated with 3DVar
- Direct assimilation (e.g., Dowell et al., 2004)
  - Reflectivity is directly assimilated through ensemble covariances of variables estimated with ensemble forecasts



- Nonlinearity of observation operator
- Short correlation length and large bias
- How to create static B (static background error covariance)



https://mrms.nssl.noaa.gov/qvs/product\_viewer/

# $\rightarrow$ How should we defeat these difficulties?



### **Radar Reflectivity Direct Assimilation**

Cost function of Hybrid 3DEnVar:

$$J(\delta \mathbf{x}_{s}, \mathbf{a}_{1}, ..., \mathbf{a}_{K}) = \frac{1}{2}\beta_{s}(\delta \mathbf{x}_{s})^{T} \mathbf{B}^{-1}(\delta \mathbf{x}_{s}) + \frac{1}{2}\sum_{k=1}^{K} [\beta_{e}(\mathbf{a}_{k})^{T} \mathbf{A}^{-1}(\mathbf{a}_{k})] + \frac{1}{2}(\mathbf{H}\delta \mathbf{x} - \mathbf{d})^{T} \mathbf{R}^{-1}(\mathbf{H}\delta \mathbf{x} - \mathbf{d})$$

$$\delta \mathbf{x} = \delta \mathbf{x}_{s} + \sum_{k=1}^{K} (\mathbf{a}_{k} \circ \mathbf{x}_{k}^{e})$$
Linearized observation operator (if original operator is non-linear, the cost is not efficiently minimized)

Here, radar reflectivity is added and analyzed together as:

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servation operator is ar (identity matrix)

• Analysis increment is computed based on the cross-variable covariance to  $\mathbf{x}_{k}^{e(\text{dBZ})}$ 

• However, optimal localization scale of  $\mathbf{x}_{k}^{e(\text{dBZ})}$  is smaller than atmospheric variables



Wang and Wang

(2017. MWR)

## **Variable-Dependent Localization (VDL)**

Wang and Wang (2023, *JAMES*) Yokota et al. (2024, submitted)



- Different localization scales are applied for atmospheric & hydrometeor variables
- However,  $\mathbf{x}_{k}^{e(\text{hydro})}$  is sometimes underestimated (zero in some places)
  - In such places, reflectivity is not assimilated efficiently
- In addition,  $\delta \mathbf{x}_s^{(\text{atmos})}$  is not affected by radar reflectivity assimilation directly
- Static B should be improved for reflectivity assimilation



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### Ensemble-Based Tangent Linear Model (ETLM) B: background error covariance

Cost function of 3DVar:  $J(\delta \mathbf{x}_0) = \frac{1}{2} (\delta \mathbf{x}_0)^T \mathbf{B}^{-1} (\delta \mathbf{x}_0) + \frac{1}{2} (\mathbf{H} \delta \mathbf{x}_0 - \mathbf{d})^T \mathbf{R}^{-1} (\mathbf{H} \delta \mathbf{x}_0 - \mathbf{d})$ 

If ETLM is introduced:  $J(\delta \mathbf{x}_0) = \frac{1}{2} (\delta \mathbf{x}_0)^T \mathbf{B}^{-1} (\delta \mathbf{x}_0) + \frac{1}{2} (\mathbf{H} \mathbf{M} \delta \mathbf{x}_0 - \mathbf{d})^T \mathbf{R}^{-1} (\mathbf{H} \mathbf{M} \delta \mathbf{x}_0 - \mathbf{d})$ 

$$= \frac{1}{2} (\delta \mathbf{x}_t)^T (\mathsf{MBM}^T)^{-1} (\delta \mathbf{x}_t) + \frac{1}{2} (\mathsf{H} \delta \mathbf{x}_t - \mathbf{d})^T \mathsf{R}^{-1} (\mathsf{H} \delta \mathbf{x}_t - \mathbf{d}) \qquad (\delta \mathbf{x}_t = \mathsf{M} \delta \mathbf{x}_0)$$



Cross-variable covariance is included in C<sup>1/2</sup> • X<sub>t</sub>X<sub>0</sub><sup>T</sup> (excluded from I • X<sub>0</sub>X<sub>0</sub><sup>T</sup>)
 Control variables are correlated based on ensemble correlation with each other
 Reflectivity assimilation can change atmospheric variables



R: observation error covariance H: linearized observation operator

 $\delta \mathbf{x}_t$ : analysis increment at time t

d: innovation vector

## **Convective-Scale Static B (CSB)**

Cost function of 3DVar:  $J(\delta \mathbf{x}_0) = \frac{1}{2} (\delta \mathbf{x}_0)^T \mathbf{B}^{-1} (\delta \mathbf{x}_0) + \frac{1}{2} (\mathbf{H} \delta \mathbf{x}_0 - \mathbf{d})^T \mathbf{R}^{-1} (\mathbf{H} \delta \mathbf{x}_0 - \mathbf{d})$ 

#### If CSB is introduced

$$= J(\delta \mathbf{x}_0^{\text{CS}}) = \frac{1}{2} (\delta \mathbf{x}_0^{\text{CS}})^T \mathbf{B}_{\text{CS}}^{-1} (\delta \mathbf{x}_0^{\text{CS}}) + \frac{1}{2} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}} - \mathbf{d})^T \mathbf{R}^{-1} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}} - \mathbf{d})^T \mathbf{R}^{-1} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}}) + \frac{1}{2} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}} - \mathbf{d})^T \mathbf{R}^{-1} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}}) + \frac{1}{2} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}}) + \frac{1}{2} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}} - \mathbf{d})^T \mathbf{R}^{-1} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}}) + \frac{1}{2} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}}) + \frac{1}{2} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}}) + \frac{1}{2} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}} - \mathbf{d})^T \mathbf{R}^{-1} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}}) + \frac{1}{2} (\mathbf{H} \delta \mathbf{x}_0^{\text{CS}}) + \frac{1$$

#### Wang and Wang (2021, MWR)

B: background error covariance R: observation error covariance H: linearized observation operator d: innovation vector  $\delta \mathbf{x}_t$ : analysis increment at time t

$[\Psi_{sf}]$	(stream function)		(zonal wind)	[ <b>f</b> ստ	0	0	0	0		f f <sub>u</sub>	0	0	0	0	0	0	0	0	0	0	0]
$\mathbf{\phi}_{\rm vp}$	(velocity potential)	v	(meridional wind)	r <sub>+</sub> f <sub>+</sub>	f.	0	0	0		$\mathbf{r}_{vu}\mathbf{f}_{u}$	$f_v$	0	0	0	0	0	0	0	0	0	0
t	(temperature)	t	(temperature)	• φ• ψ	φ.	f	Ő	Ő		$\mathbf{r}_{tu}\mathbf{f}_{u}$	$\mathbf{r}_{tv}\mathbf{f}_{v}$	f <sub>t</sub>	0	0	0	0	0	0	0	0	0
D <sub>c</sub>	(surface pressure)	<b>p</b> <sub>s</sub>	(surface pressure)	tψ	U	<sup>t</sup>	0	U		<b>r</b> <sub>pu</sub> f <sub>u</sub>	$r_{pv}f_v$	<b>r</b> <sub>pt</sub> f <sub>t</sub>	<b>f</b> p	0	0	0	0	0	0	0	0
<b>D</b> rh	(relative humidity)	<b>q</b> <sub>rh</sub>	(relative humidity)	r <sub>p</sub> f <sub>ψ</sub>	0	0	fp	0		<b>r</b> <sub>qu</sub> <b>f</b> <sub>u</sub>	$\mathbf{r}_{qv}\mathbf{f}_{v}$	<b>r</b> at <b>f</b> t	<b>r</b> <sub>qp</sub> <b>f</b> <sub>p</sub>	<b>f</b> a	0	0	0	0	0	0	0
	1	w	(vertical wind)	0	0	0	0	∫ <b>f</b> q_		<b>r</b> <sub>wu</sub> <b>f</b> <sub>u</sub>	r <sub>wv</sub> f <sub>v</sub>	<b>r</b> <sub>wt</sub> <b>f</b> <sub>t</sub>	r <sub>wp</sub> f <sub>p</sub>	r <sub>wa</sub> f <sub>a</sub>	f <sub>w</sub>	0	0	0	0	0	0
		$\mathbf{q}_{l}$	(cloud water)	L		γ			J	$\mathbf{r}_{\mathrm{ln}}\mathbf{f}_{\mathrm{n}}$	<b>r</b> <sub>lv</sub> <b>f</b> <sub>v</sub>	r <sub>lt</sub> f <sub>t</sub>	r <sub>ln</sub> f <sub>n</sub>	rlafa	r <sub>lw</sub> f <sub>w</sub>	f	0	0	0	0	0
$\mathbf{o}\mathbf{x}_0$		<b>q</b> r	(rainwater)		E F	$\frac{1}{2}$	2			r <sub>m</sub> f <sub>n</sub>	r <sub>rv</sub> f <sub>v</sub>	r <sub>rt</sub> f <sub>t</sub>	r <sub>m</sub> f <sub>n</sub>	r <sub>ra</sub> f <sub>a</sub>	r <sub>rw</sub> f <sub>w</sub>	r <sub>ei</sub> fi	fr	0	0	0	0
		<b>q</b> <sub>s</sub>	(snow)		- 1					rf.	rf.	r.f.	r <sub>on</sub> f <sub>n</sub>	rafa	rf.	rafi	r."f.	f.	0	0	0
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		<b>q</b> <sub>dbz</sub>	(reflectivity)							gu u	gv v	gt t	gp p	gq q	gw w	gl l	gr r	gs s	gi i	g	
		$\Box$								L <sup>r</sup> zu <sup>T</sup> u	r <sub>zv</sub> r <sub>v</sub>	<b>r</b> <sub>zt</sub> <b>r</b> <sub>t</sub>	г <sub>zp</sub> т <sub>p</sub>	r <sub>zq</sub> r <sub>q</sub>	r <sub>zw</sub> r <sub>w</sub>	r <sub>zl</sub> r <sub>l</sub>	<b>r</b> <sub>zr</sub> <b>r</b> <sub>r</sub>	r <sub>zs</sub> r <sub>s</sub>	<b>r</b> <sub>zi</sub> r <sub>i</sub>	r <sub>zg</sub> r <sub>g</sub>	Tz
		$\delta \mathbf{x}_{o}^{CS}$								L					7						
		00													$B_{cc}^{1/2}$						
															- 65						

Cross-variable covariance is included in B<sub>CS</sub>

Control variables are correlated based on balance operators with each other Reflectivity assimilation can change atmospheric & hydrometeor variables



## **Single-Reflectivity Assimilation Test**

Analysis increment (color) and first guess (contours) of sea-level pressure (hPa) and analysis increment of reflectivity (magenta, 30dBZ) in reflectivity assimilation at 1-km height (innovation: 50dBZ, obs error: 1dBZ)



Flow-dependent analysis within localization

No increment

Flow-dependent and broader, but no change for hydrometeors

Successful hydrometeor analysis, but narrow for surface pressure



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## **Cycling Tests**



VarDA name	Weight of (static B, ensemble B)	ETLM?	CSB?
3DVar	(1.0, 0.0)	-	-
3DVarETLM	(1.0, 0.0)	Yes	-
3DVarCSB	(1.0, 0.0)	-	Yes
EnVar	(0.5, 0.5)	-	-
EnVarETLM	(0.5, 0.5)	Yes	-
EnVarCSB	(0.5, 0.5)	-	Yes

Assimilated observations:

 surface pressure, wind, temperature, relative humidity, precipitable water vapor, radar radial wind, and radar reflectivity

.... : Initial conditions

: FV3LAM-based forecasts

Error covariance / Re-centering

- Localization for ensemble B ( $e^{-20/3}$  scale):
  - 300 km (horizontally) for atmospheric variables
  - 15 km (horizontally) for hydrometeor variables
    - Cross-variable covariances: x0.05 (=15/300)
  - 1.1 Inp (vertically)
- Localization for ETLM ( $e^{-20/3}$  scale)
  - 300 km (only horizontally)

## Weighted RMSE of First Guess



number of analyses

### Hurricane Ian (2022) Analysis

Analysis of composite reflectivity (color, dBZ) and sea-level pressure (contours, every 4 hPa) at **20220930 00z** 

- Both ETLM and CSB improve reflectivity distribution
- ETLM tends to decrease reflectivity

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CSB tends to increase reflectivity

UFS





### **Summary and Future Works**

- Static background error covariance (static B) was improved by ETLM and CSB for radar reflectivity assimilation.
- Ensemble-based tangent linear model (ETLM)
  - Impact of radar reflectivity assimilation is broad and flow-dependent for atmospheric variables, but no impact for hydrometeors
  - ETLM tends to decrease precipitation and makes RMSE smaller mainly for surface pressure, but larger for the other observations

### Convective-scale static B (CSB)

- Hydrometeor is successfully analyzed, but correlation length is too short for atmospheric variables
- CSB tends to increase precipitation and makes RMSE smaller mainly for reflectivity, but larger for the other observations

### • Future works

- Simultaneous application of conventional B, ETLM, and CSB
- Development of multiscale static B



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### **BACK UP**



### **Rapid Refresh Forecast System (RRFS)**

- Next-generation convection-allowing operational forecast system in NCEP
  - One of the UFS applications
  - Based on the FV3 limited area model (LAM) (Black et al. 2021, JAMES)
  - 3-km horizontal grid
  - 65 vertical levels
  - Hourly updated by hybrid 3DEnVar (with 30-member EnKF)
  - Deterministic forecasts to at least 18h every 1h
  - Deterministic & ensemble forecasts to 60h every 6h (6 members x 2 initial times)

# The impacts of ETLM and CSB for radar reflectivity assimilation are clarified in RRFS





### **Single-Reflectivity Assimilation Test with ETLM**





### Single-Reflectivity Assimilation Test with CSB





### **Analysis Increments**





### Analysis increment of surface pressure



(e) EnVarCSB

100934

# covariances

### **Pressure Tendency**

Time series of pressure tendency (domain-wide mean absolute change) in the forecasts from 20220930 00z





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### Hurricane Ian (2022) Analysis

Analysis of composite reflectivity (color, dBZ) and sea-level pressure (contours, every 4 hPa) at **20220930 00z** 



dBZ

dBZ



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dBZ

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dBZ

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