



Ground perspective:

-Stratocumulus clouds span thousands of square kilometers over the Earth's surface. Due to drizzle presence at the cloud dominan rada the cloud backscatter. microphysical structure asserts itself as a blind spot for ground based profiling instruments.

Satellite perspective:

Meteosa[,] board on Generation provides Second Brightness day/night Temperatures. From retrievals point of view, using the infrared part of the spectra is practica (channels 3.9 and 10.8µm) due to their sensitivity to droplet size and cloud phase.

D. 5.6 x scene [km]







1. Drizzling stage Sc during ASTEX, early phase of depleting

2. Evolved Sc, via phase of Pocket of Open Cells (POCs), into scattered fractions

Large Eddy Simulation output \rightarrow ECSIM synthetic observations

Liquid Water Content \rightarrow [Fixed number concentration] \rightarrow Droplet Effective Radius [Assumed γ -distribution] \rightarrow Drop Size Distribution \rightarrow 3D cloud scene model



Figure above: Extracted Liquid Water Path (LWP) variability during a nonintensive precipitation process (hour14 of the ASTEX simulation)

Remote Sensing of the Environment Corresponding author: I.Stepanov@tudelft.nl

The satellite observation of drizzle in Stratocumulus clouds **I. Stepanov**, S.B. Velasco, H.W.J. Russchenberg Delft University of Technology, Delft, The Netherlands

n ²]	Cloud scenes	utilization
60		
	 Quantity 	extraction

ion: effective radius, liquid water content, liquid water path, optical depth, droplet number concentration.

• Simulated retrievals using 3D Monte Carlo or Independent Column Approximation long wave Radiative Transfer Model: Radiance Brightness $(Wm^{-2}sr^{-1}cm^{-1})$ and *Temperature* (K)

Pockets of Open Cells (POCs) - Location where drizzle forced change of structure in Stratocumulus cloud can be found^[1]. Compared to surrounding cloud field, POCs are flagged as areas with significant precipitation.

 ΔT_{h} - a proxy for the droplet size (see Figure below)^[2], which is an indicator of the size of the droplets at the top of the cloud (location of POCs), related to precipitation triggering mechanism.



Non-scaled LWP



15 20 W-E direction [km] Effective radius hour 24



20 15 W-E direction [km] BTD 11,4µm



W-E direction [km]



due to evaporation or precipitation. For the fine scale microphysical processes and dynamical analysis of a drizzling Stratocumulus, synthetic remote sensing observations were used. In this study, a large eddy simulation cloud scene is simulated based on the ASTEX (The Atlantic Stratocumulus Transition Experiment) campaign, on domain size 25.6x25.6km, and ingested into the EarthCARE Simulator (ECSIM) Radiative Transfer Model.

Meteosat Second Generation Imager data

- Data available in **Pixel Counts** (*binary pixel value*)
- Conversion necessary to obtain infrared radiance (physical value)
- Radiance conversion to **Brightness Temperature**

References

^[1]VanZanten, Margreet C., and Bjorn Stevens. "Observations of the structure of heavily precipitating marine stratocumulus." Journal of the atmospheric sciences 62, no. 12 (2005): 4327-4342.

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25

10

BTD (11-4)µm [K]



- → Running mean smoothing of raw BTD signals shows a qualitative match to the droplet size when compared to
- → Range of BTD errors provides restrictions to categorization
- → Sensitivity of BTD to the adiabatic profile of effective radius to be studied \rightarrow proxy \rightarrow flag for drizzle at cloud

Radiance = CAL offset + CAL slope * Count

$$T_b = \left[C_2 v_c / \log \left(\frac{C_1 v_c^3}{R} + 1 \right) - B \right] / A$$

 $C_1 = 1.19104 \ 10^{-5} \text{ mW m}^{-2} \text{ sr}^{-1} (\text{cm}^{-1})^{-4}$

 $C_2 = 1.43877 \text{ K}(\text{cm}^{-1})^{-1}$

 v_c = central wavenumber of the channel

A, B coefficients

- ^[2]Pérez, Juan C., Félix Herrera, Fernando Rosa, Albano Gonzalez, Melanie A. Wetzel, Randolph D. Borys, and Douglas H. Lowenthal. "Retrieval of Marine Stratus Cloud Droplet Size from NOAA–AVHRR Nighttime Imagery."
- ^[3]Lensky, Itamar M., Daniel Rosenfeld, 2003: Satellite-based insights into precipitation formation processes in

Delft University of Technology