

Observations of Aerosol-Cloud Interactions during the North Atlantic Aerosol and Marine Ecosystem Study

Background

- CDNC is a cloud microphysical property that impacts radiative forcing, cloud evolution, precipitation, local and global climate and, through observation, can be used to monitor the cloud albedo effect, or the first indirect effect (Zeng et al., 2013).
- Aerosol-cloud interaction is a mechanism whereby an increase in aerosol concentration results in an increase in cloud droplet number concentration leading to changes in the radiative properties of the cloud (Twomey 1977) as well as the lifetime of the cloud (Albrecht, 1989; Liou and Ou, 1989).
- Number concentration can be related to cloud albedo through the simplified expression:

$$\frac{d\alpha}{dN_d} = \frac{\alpha(1-\alpha)}{3N_d}$$

(Platnick & Twomey, 1994)

• Aerosol-cloud interactions, or the first indirect effect, can be observed through monitoring cloud droplet number concentration.



Research Objectives

- 1) To assess the strength of the connection between cloud properties and aerosol concentration and composition during NAAMES
- 2) Determine which cloud properties are least and most sensitive to aerosol concentrations
- 3) Explore variations in aerosol and cloud properties on daily and seasonal timescales

Instruments

Research Scanning Polarimeter

- Prototype for Aerosol Polarimetry Sensor on the Glory satellite (2011) (Cairns et al., 1999)
- Along track scanning 152 viewing angles per scene $(\pm 60^\circ)$; 14 mrad field of view
- Polarimetric and full intensity measurements in the visible and shortwave infrared over 9 bands: 410, 470, 555, 670, 864, 960, 1593, 1880, 2263 nm
- Measures aerosol: OT, R_{eff}, V_{eff}, R.I., single-scattering albedo, morphology
- Measures cloud: OT, R_{eff}, V_{eff}, CTH, cloud phase, ice asymmetry parameter, CDNC





Condensation Nuclei Counter (CNC)

- CNC measures the concentration of cloud condensation nuclei (CCN) at several superstation levels
- Supersaturation levels range from 0.07% to as high as 2%
- Droplet sizes from $0.75 10 \,\mu$ m across 20 size bins and operates at a 1 Hz sampling rate

Aerosol Mass Spectrometer (AMS)

- Provides organic aerosol analysis of aerosol size and chemical mass loading information
- Particle size range: 0.04 to ~ 1.0 micrometers

Kenneth Sinclair^{1,2}, Bastiaan van Diedenhoven^{2,3}, Brian Cairns², Mikhail Alexandrov^{2,3}, Andrew S. Ackerman², Luke Ziemba⁴, Richard Moore⁴, Ewan Crosbie⁴, Chris Hostetler⁴ AMS 15th Conference on Cloud Physics & 15th Conference on Atmospheric Radiation, Vancouver BC, Canada. July 9-13, 2018.



Measurements

- North Atlantic Aerosols and Marine Ecosystems Study (NAAMES)
- 4 aircraft and ship measurement campaigns over 5 years (Each campaign is aligned to a specific annual event in the plankton cycle)
- Characterize plankton ecosystem properties during primary phases of the annual cycle in the North Atlantic and their dependence on environmental forcings
- Determine how marine aerosols and boundary layer clouds are influenced by plankton ecosystems in the North Atlantic (and their feedbacks)



month Behrenfeld, M., "NAAMES Confirmation Review.' 28 Aug 2015

N_d Retrieval Method



) Cloud Optical Thickness, τ 2) Mean extinction crosssection at cloud top, $\overline{\sigma_{e}}$

3) Cloud Geometrical Thickness, H

Assumptions

- 1) Linearly increasing LWC. L = c h,
- $N_d(h) = c$ 2) Constant N_d profile
- Mean droplet extinction cross section is estimated from polarized reflectances in the cloud-bow near cloud top (figure 1)
- Geometric thickness is estimated by relating incloud water vapor absorption to physical thickness
- This could also be retrieved from polarimetric oxygen A-band
- The method has been validated using data from the NAAMES 1 and 2 campaigns (figure 2)
- It is expected that a polarimeter aboard PACE would be capable of using this method



100 10CDP N_d [Droplets per cm³] *Fig: 12 daily mean N_d values from the RSP and* in situ Cloud Droplet Probe (CDP) collected during the NAAMES 1 and 2 campaigns

Results – Aerosol Mass

Aerosol Type & Mass



Fig: RSP N_d measurements are compared with AMS measured salt, organic and sulfate aerosol masses.

Results N_d and CCN

N_d and CCN CCN at 0.35% Supersaturation Key finding: Boundary layer CCN R=0.83 $[N_d] = -0.3 + [CCN]^{0.8}$ 2015/11/12 * has a strong positive correlation • The strength of all correlations 2016/05/19 < are dependent on supersaturation with weaker 2016/05/29 < 2016/05/30 < correlations being associated with low and high supersaturations CCN [Particles per cm³] Figure: RSP N_d and LARGE in situ CCN geometric means and geometric SD factor (bars) collected during cloud modules Key finding: Liquid water path is **uncorrelated** with N_d **Cloud Property Correlations** 30 R = 0.51 m_{*} = 0.14 R = -0.44 m = -0.07 [•]ੁ ²⁰⁰ R = 0.15 **150** \diamond \diamond \diamond 40 60 20 20 60 40 RSP N₄ [cm⁻³] RSP N₄ [cm⁻³] RSP N₄ [cm⁻³]

Figure: Cloud property variability is compared with CCN concentrations. A linear fit is shown if correlation is over 0.5.

Key finding: Above cloud relative

correlation with liquid water path

• Vertical winds can explain some

variability in the N_d-CCN

• A low number of data points

• Boundary layer relative humidity

did not have a detectable effect

CCN [Particles per cm³]

complicates analyzing

precipitation data

on cloud properties

relationship

humidity has a **robust** inverse

Meteorological Properties



Figure: Horizontal and vertical winds, above and below cloud relative humidity and precipitation data are assessed to determine meteorological contributions to cloud property variability

CCN and Cloud Properties

CCN [Particles per cm³]

20 -



Figure: Cloud property variability is compared with CCN concentrations. A linear fit is shown if correlation is over 0.5.

CCN [Particles per cm³]





Key findings: Through seasons 1) N_d increases 156% 2) R_{eff} decreases 23% 3) τ increases 8%

Conclusions

- Meteorological conditions can explain some variability of aerosol's effect on cloud properties • Strongest influence was found with the above cloud relative humidity (-0.69) and updraft velocity (0.58)
- Precipitation was not found to have a detectable effect
- 2. A strong (0.83) correlation is found between N_d and CCN
 - High correlations also for R_{eff} (-0.39), V_{eff} (-0.40), τ (0.54) and P.T. (-0.53)
- Only weak correlations are found between liquid water path with CCN (0.18) and N_d (0.15) 3. Cloud property and aerosol concentration changes are found to be consistent with the *Twomey* effect
- . Above cloud relative humidity has a *robust* inverse correlation with liquid water path and therefore plays an important role
- A seasonal change in aerosol chemical composition is observed, which correlates with a seasonal change in cloud properties
- A cloud **brightening** is found with τ increasing 8% between campaigns

Future Work

- A. Use a LES model to better determine cloud property dependencies on meteorological properties
- B. Use the in situ UHSAS instrument to assess aerosol size distribution and hygroscopicity on CCN-
- N_d sensitivity

Acknowledgements

Support for this work is provided by NASA grant #NNX15AD44G (ROSES ACCDAM) and an NSERC-PGSD is gratefully acknowledged. Dr. Cairns was primarily supported by the NASA Radiation Sciences Program managed by Dr. Hal Maring.

Authors

- ¹ Ph.D. Candidate, Department of Earth and Environmental Engineering, Columbia University, New York, NY, 10025. Contact: kenneth.sinclair@columbia.edu
- ² NASA/Goddard Institute for Space Studies, 2880 Broadway, New York, NY 10025, USA
- ³ Center for Climate Systems Research, Columbia University, New York, NY 10025, USA
- ⁴ University of Washington, Seattle, Washington, USA
- ⁵ NASA Ames Research Center, Mountain View, California
- ⁶ NASA/Langley Research Center, MS420, Hampton, Virginia 23681, USA

satellites. *Atmospheric Chemistry and Physics Discussions*, 13(11), 29035-29058.

References

- Cairns, B., L. D. Travis, and E. E. Russell (1999), The Research Scanning Polarimeter: Calibration and ground-based measurements, Proc. SPIE Int. Soc. Opt. Eng., 3754, 186–196.
- Collins, W. D., W. C. Connant, and V. Ramanathan (1994), Earth radiation budget, clouds, and climate sensitivity, in: The Chemistry of the Atmo- sphere: its Impact on Global Change, edited by J. G. Calvert, pp. 207–215, Blackwell Scientific Publishers, Oxford, UK.
- Liou, K. N., and S. C. Ou, 1989: The role of cloud microphysical processes in climate– An assessment from a one-dimensional perspective. J. Geophys. Res., 94, 8599-8607.
- Sanchez, K. J., Chen, C. L., Russell, L. M., Betha, R., Liu, J., Price, D. J., ... & Müller, M. (2018). Substantial Seasonal Contribution of Observed Biogenic Sulfate Particles to Cloud Condensation Nuclei. Scientific reports, 8(1), 3235. Twomey, S. (1977). The influence of pollution on the shortwave albedo of clouds. *Journal of the atmospheric sciences*, 34(7),
- 1149-1152 Zeng, S., Riedi, J., Trepte, C. R., Winker, D. M., & Hu, Y. X. (2013). Study of cloud droplet number concentration using the A-Train