

# Effects of Midwinter Arctic Leads on Clouds and the Surface Energy Budget

## MOTIVATION

- Extreme air-water temperature difference (20 40°C) and large vertical moisture gradients exist over leads during winter
- Leads may induce extensive plumes and low-level clouds Leads and the ensuing clouds jointly have significant impacts on the wintertime large-scale Arctic surface heat budget
- Accurate determination of the Arctic surface energy budget is particularly important because of the sensitivity of sea ice thickness to the surface radiation fluxes and the ice-albedo climate feedback mechanism
- However, properly representing lead-induced boundary layer clouds and the associated large-scale fluxes remains difficult in climate models due to the unresolved small size of leads and the scarcity of observations in the Arctic etc.



lead and associated clouds. Photo was taken on a flight over the Beaufort Sea, October 12, 1994, by T. Arbetter, University of Colorado

## METHODS

Due to the limited observations, previous studies mostly utilized model simulations only to examine the clouds generated by a lead and the associated effects on surface fluxes (e.g., Glendening et al., 1992; Pinto and Curry, 1995; Zulauf and Krueger, 2003a, 2003b). Our study used both observations and modeling simulations:

- (1) Observational data from ARM were used to derive the statistical associations between lead distributions and the cloud occurrence
- Barrow radiosondes (*T, P, RH*) and surface measurements
- (*T, P, RH, V*), MMCR reflectivity and AMSR-E derived lead fraction • Jan-Apr, Nov-Dec, 2008-2011
- (2) Three dimensional cloud-resolving model (CRM), System for Atmospheric Modeling (SAM), was then used to understand the observed lead-cloud associations

### References

Andreas, E.L. and Cash, B.A., 1999: Convective heat transfer over wintertime leads and polynyas. J. Geophys. Res.: Oceans, 104, 25721-25734.

Glendening, J. W., and S. D. Burk, 1992: Turbulent transport from an Arctic lead: A large-eddy simulation. Boundary Layer Meteorol., 59, 315-339.

Khairoutdinov, M. F., and D. A. Randall, 2003: Cloud resolving modeling of the ARM summer 1997 IOP: Model formulation, results, uncertainties, and sensitivities, J. Atmos. Sci., 60, 607–625. Pinto, J. O., and J. A. Curry, 1995: Atmospheric convective plumes emanating from leads: 2.

Microphysical and radiative processes. J. Geophys. Res., 100, 4633-4642.

Zulauf, M. A., and S. K. Krueger, 2003a: Two-dimensional cloud-resolving modeling of the atmospheric effects of Arctic leads based upon midwinter conditions at the Surface Heat Budget of the Arctic Ocean ice camp. J. Geophys. Res., 108, 4312-4325.

Zulauf, M. A., and S. K. Krueger, 2003b: Two-dimensional numerical simulations of Arctic leads: Plume penetration height. J. Geophys. Res., **108**, 8050-8062.

### Acknowledgement

This work is supported by the NASA grant 80NSSC18K0843. We thank Marat Khairoutdinov for the SAM support

Xia Li, Steven K. Krueger, Courtenay Strong & Gerald G. Mace Department of Atmospheric Sciences, University of Utah



time (hour)



- cross the closed lead surface

## **CONCLUSIONS AND FUTURE WORK**

## Conclusions

- the refrozen lead surface, which is due to
- clouds with ambient dry air
- drier (i.e., relative humidity decreases)
- days may include a large portion of refrozen leads

## **Future Work**

Future work will focus on differentiating open leads with recently refrozen leads. And more case studies using multi-source observations (e.g., satellite, airborne and ground-based observations) will be analyzed and further be simulated to give a better understanding of the interrelationship between leads and low-level cloudiness and their effects on the largescale surface energy budget.

### Contact: xia.li@utah.edu

## SIMULATED CLOUDS

### closed lead refrozen lead t = 6.5 hrs t = 6.5 hrs t = 7.0 hrs t = 7.0 hrs 200 t = 7.5 hrst = 7.5 hrs t = 8.0 hrst = 8.0 hrs 200 100 x (km)

In the open lead case, lead-induced boundary layer clouds are advected downstream over 50 km, which extends the impacts of open lead over a broader region by enhancing the downward SH and infrared radiative flux at the downwind surface

Clouds dissipate within 1.5 hr in refrozen vs. 3 hr in closed lead Clouds cannot advect across the refrozen lead surface, but it can

# Low-level clouds can dissipate very quickly when crossing over

• Large SH over refrozen lead maintains strong convection, mixing

• Large SH increases air temperature, with suppressed LH, air becomes

• Lead-induced boundary layer clouds are quite sensitive to relative humidity of the environmental air, which facilitates the clouds depletion The above findings provide a plausible explanation for the observed counterintuitive results, and suggest that high lead flux