

The 26 June 2015 Nocturnal Mesoscale Convective System during PECAN

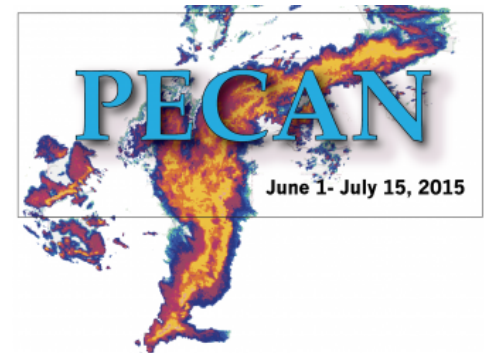
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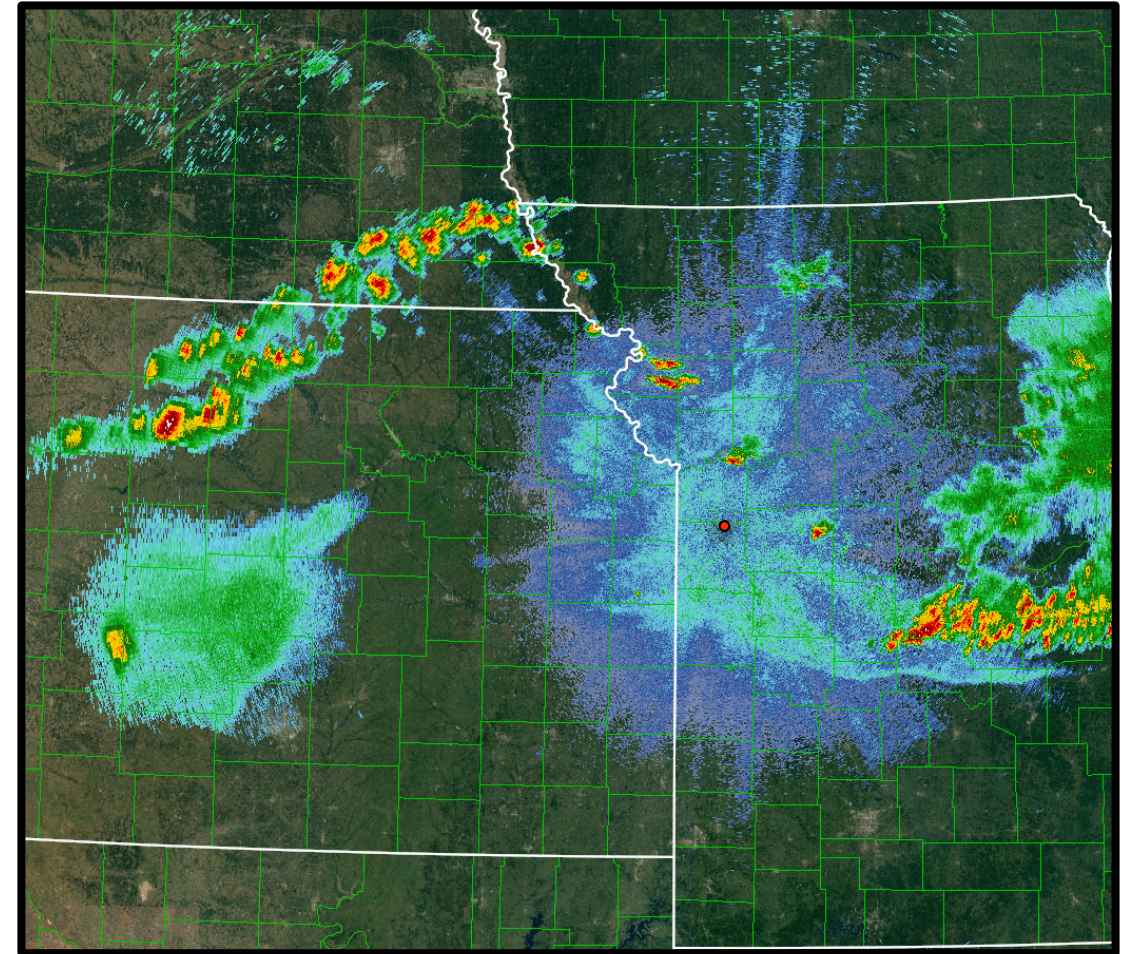
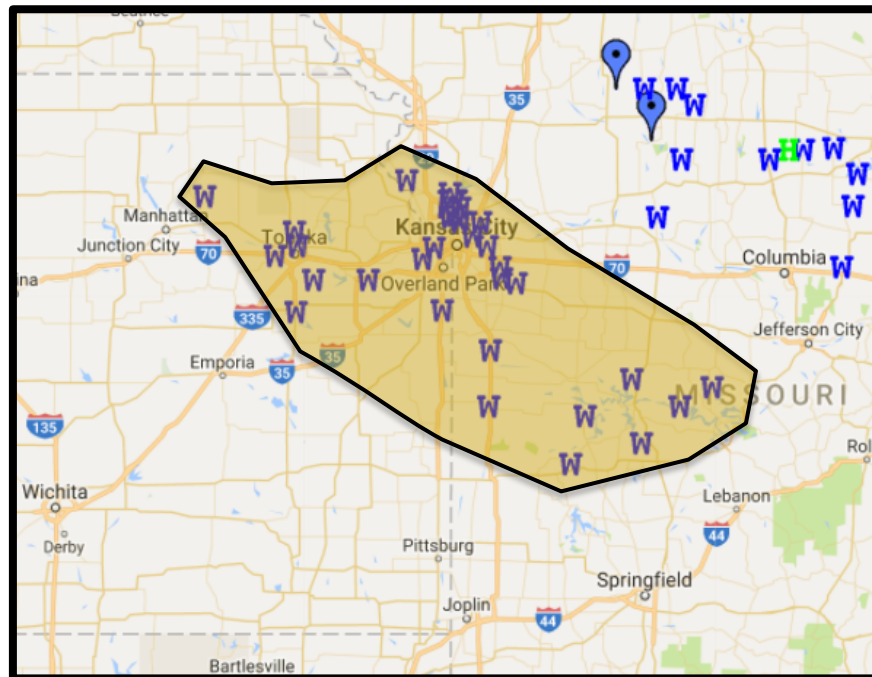
Acknowledgements:

NSF Grant AGS-1259709, Conrad Ziegler, Michael Coniglio, Michael Biggerstaff,
Russ Schumacher, Ted Mansell, Terry Schuur, George Bryan

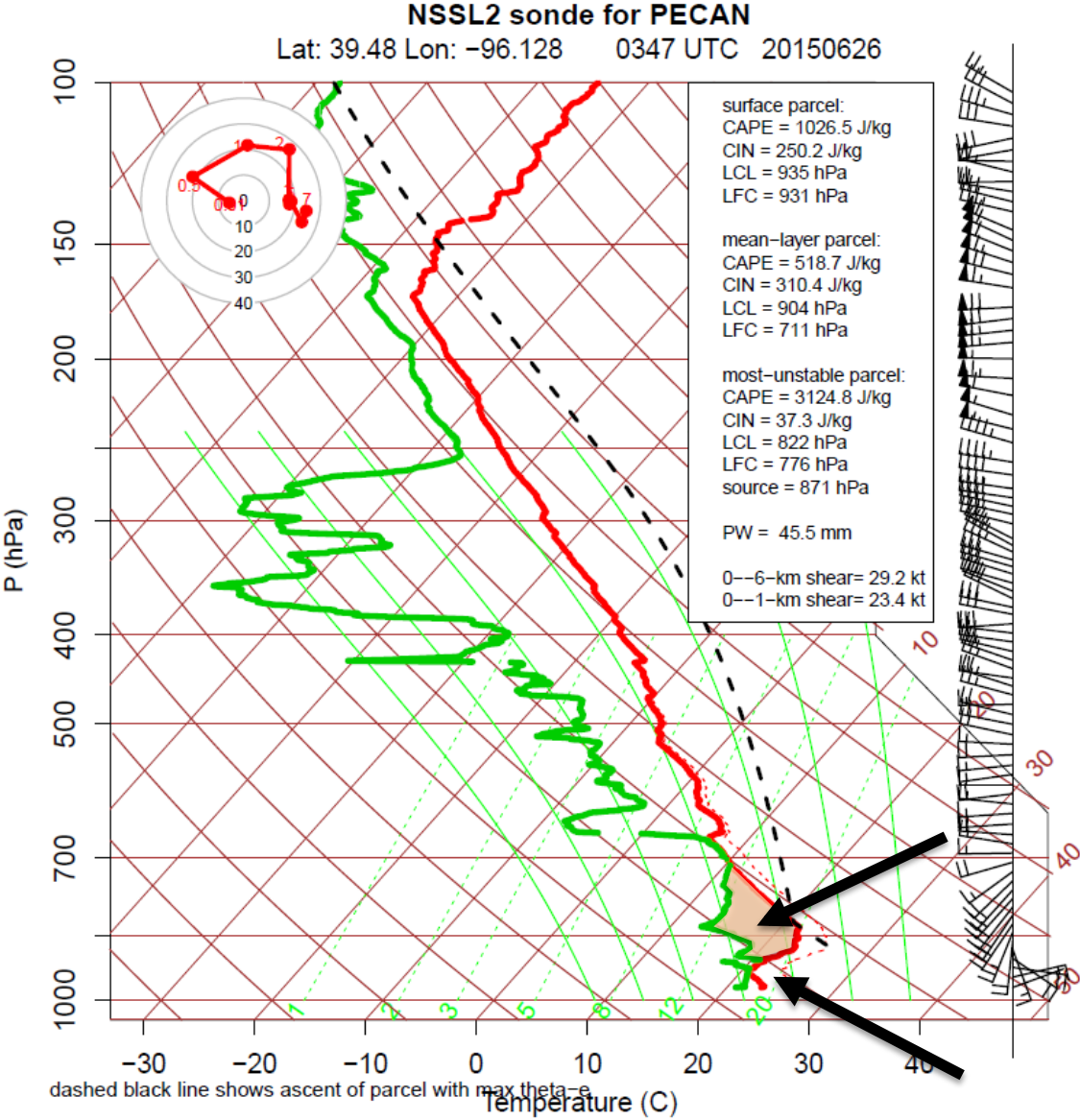
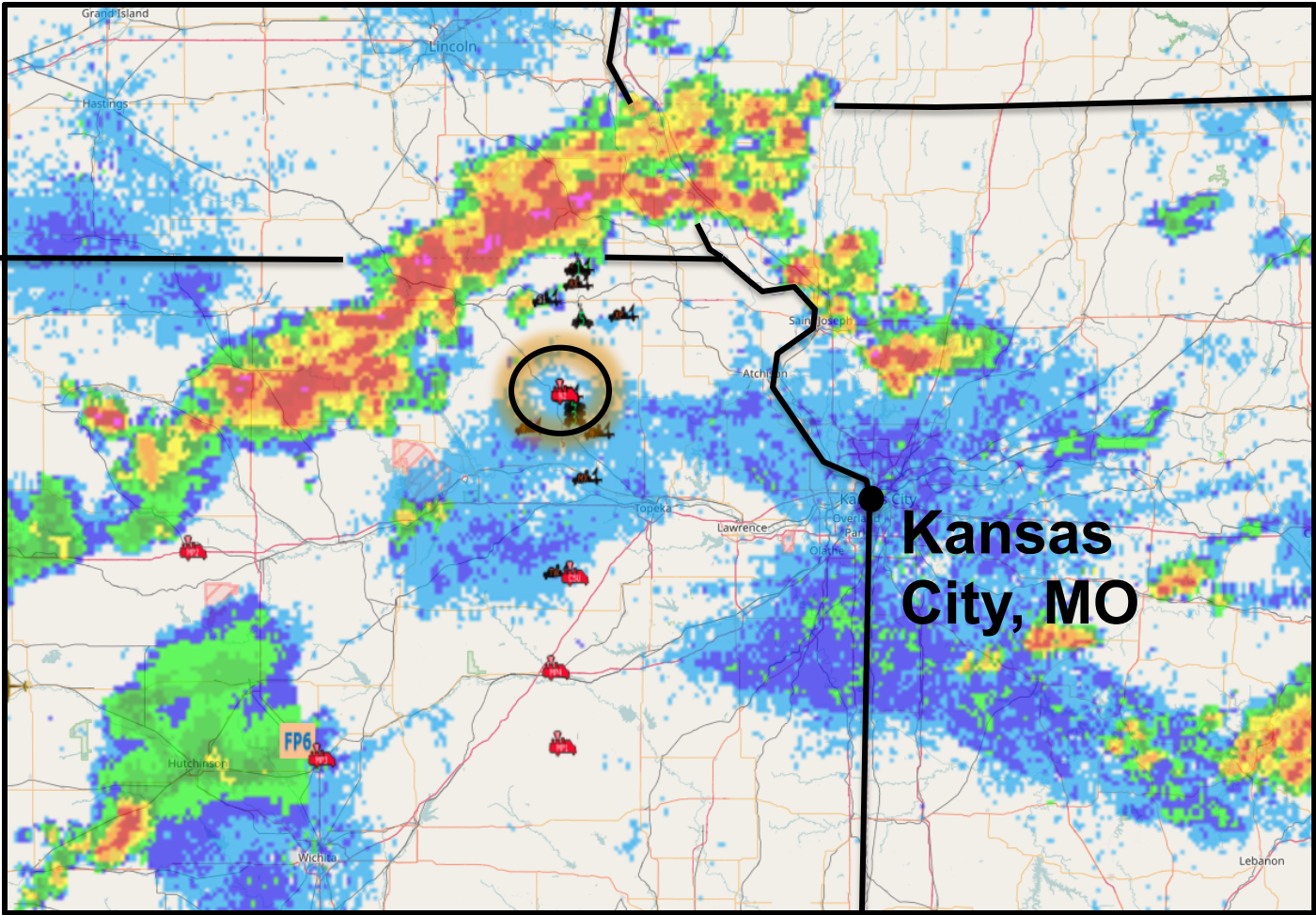


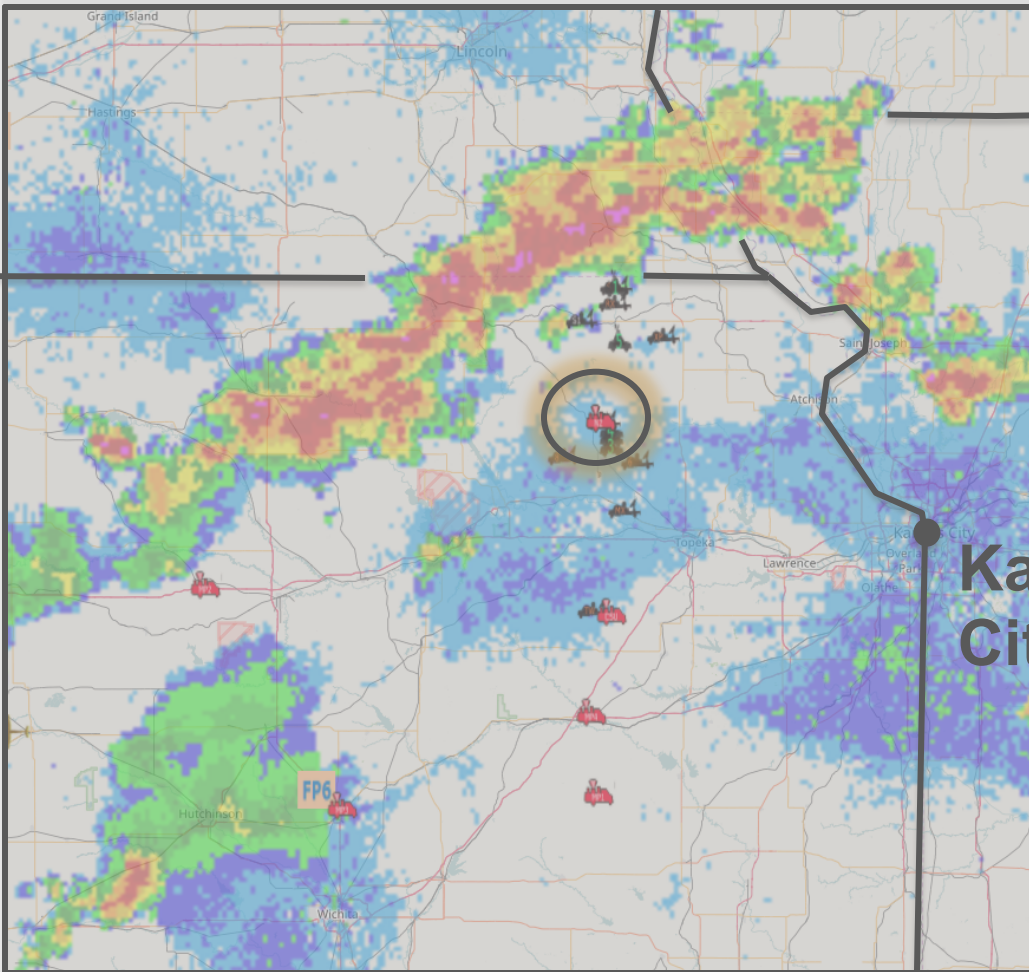
June 25/26 2015: Severe Nocturnal MCS Impacts Kansas City, Missouri and Surrounding Areas

- Thunderstorms developed along cold front, progressed southeast overnight
- Kansas City, MO area hit hard after midnight



Base-scan reflectivity from KEAX between
0300 - 0900 UTC





surface parcel:

CAPE = 1026.5 J/kg

CIN = 250.2 J/kg

LCL = 935 hPa

LFC = 931 hPa

mean-layer parcel:

CAPE = 518.7 J/kg

CIN = 310.4 J/kg

LCL = 904 hPa

LFC = 711 hPa

most-unstable parcel:

CAPE = 3124.8 J/kg

CIN = 37.3 J/kg

LCL = 822 hPa

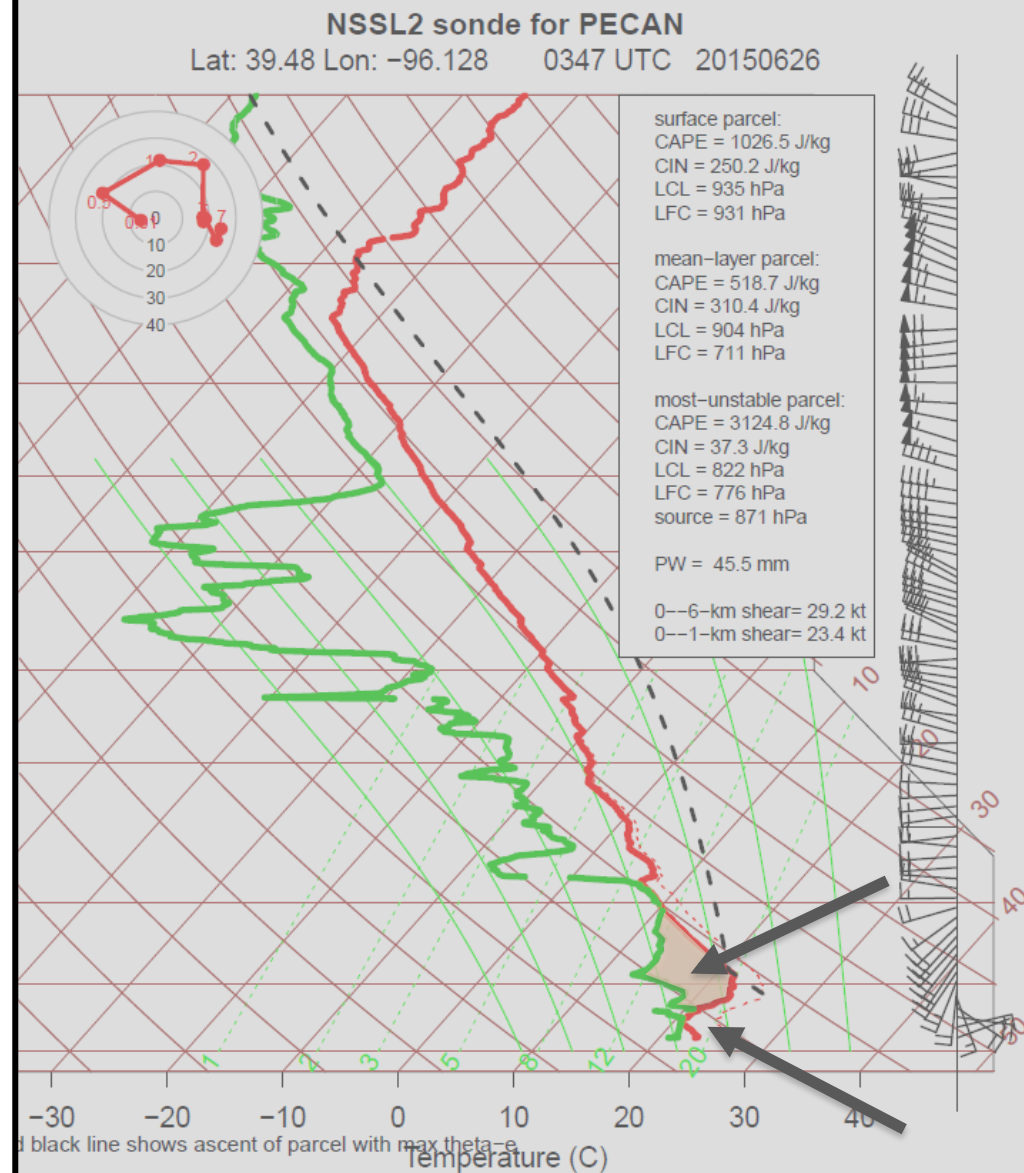
LFC = 776 hPa

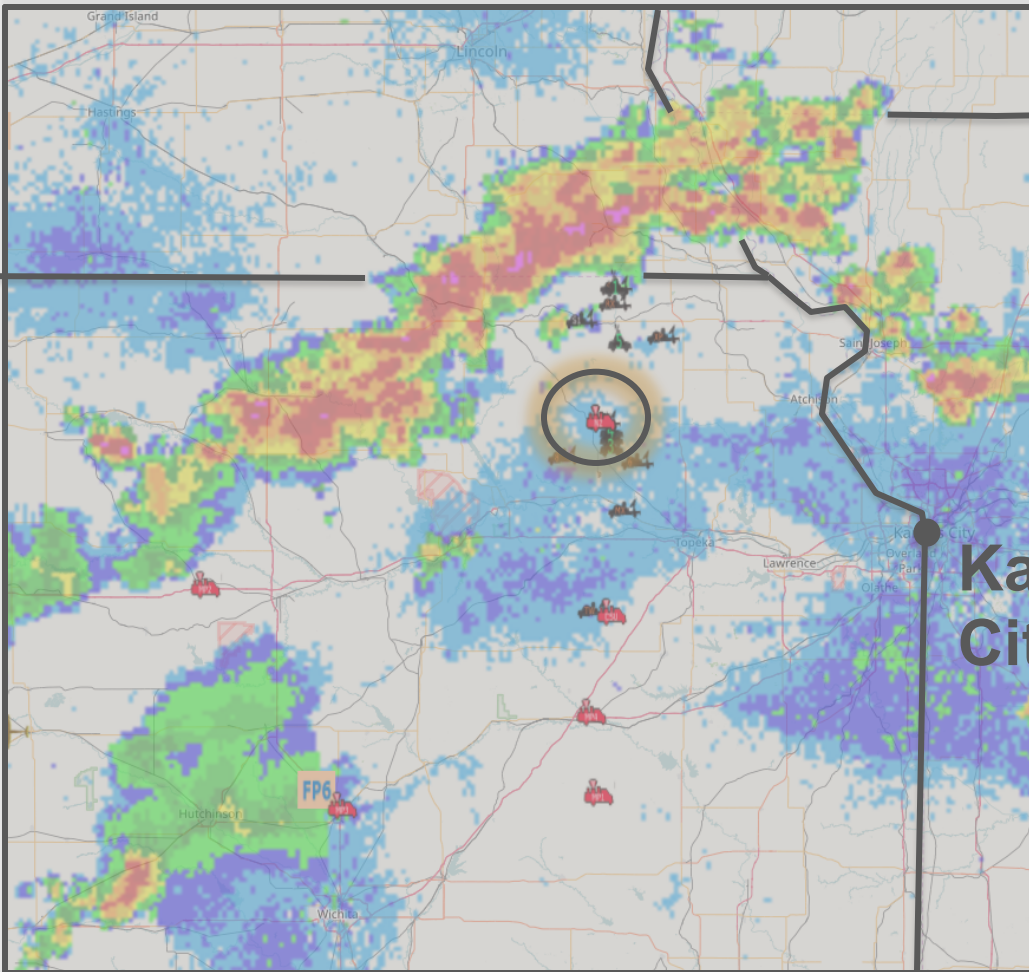
Source = 871 hPa

PW = 45.5 mm

0--6-km shear= 29.2 kt

0--1-km shear= 25.4 kt





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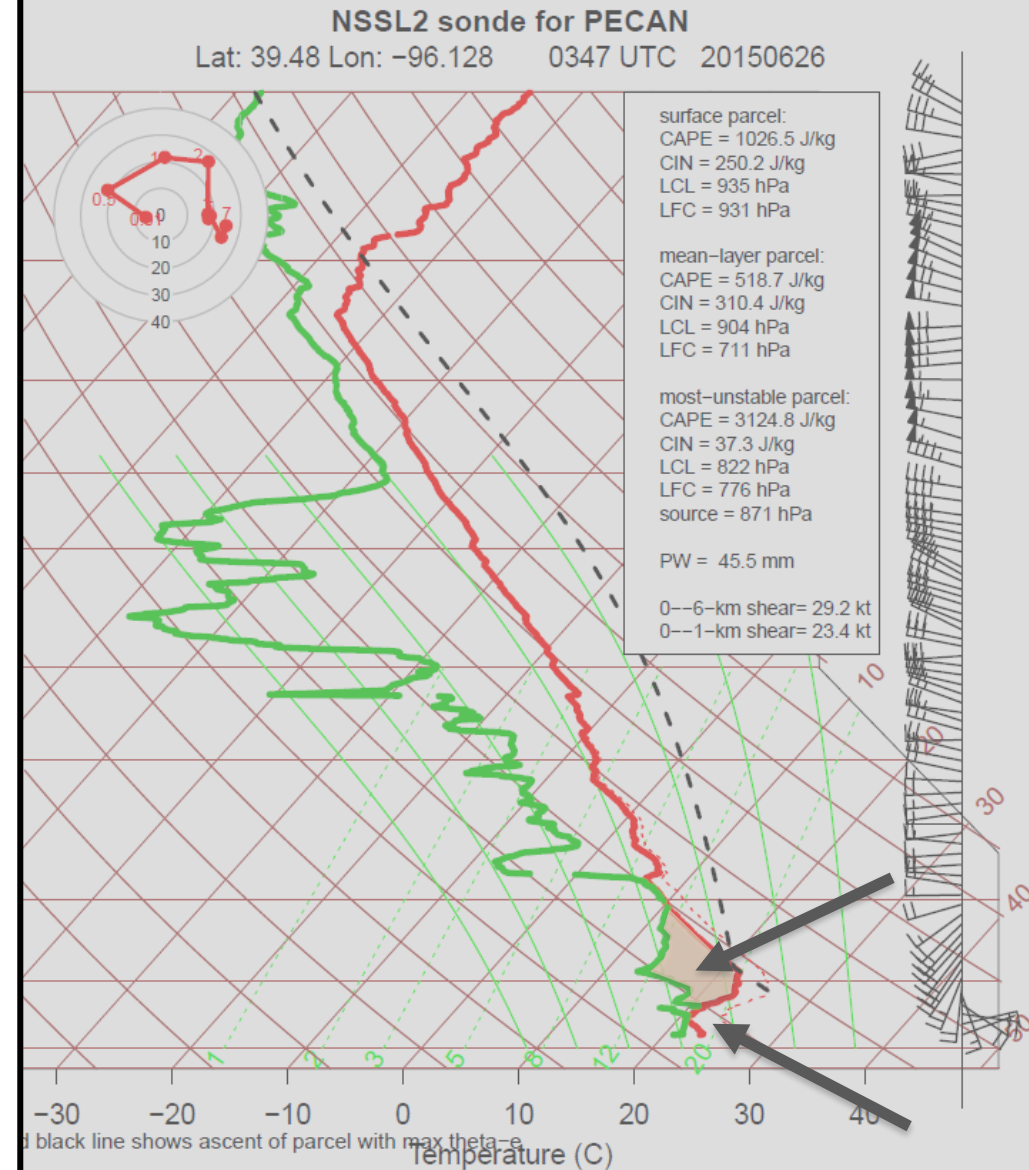
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Two Main Research Questions

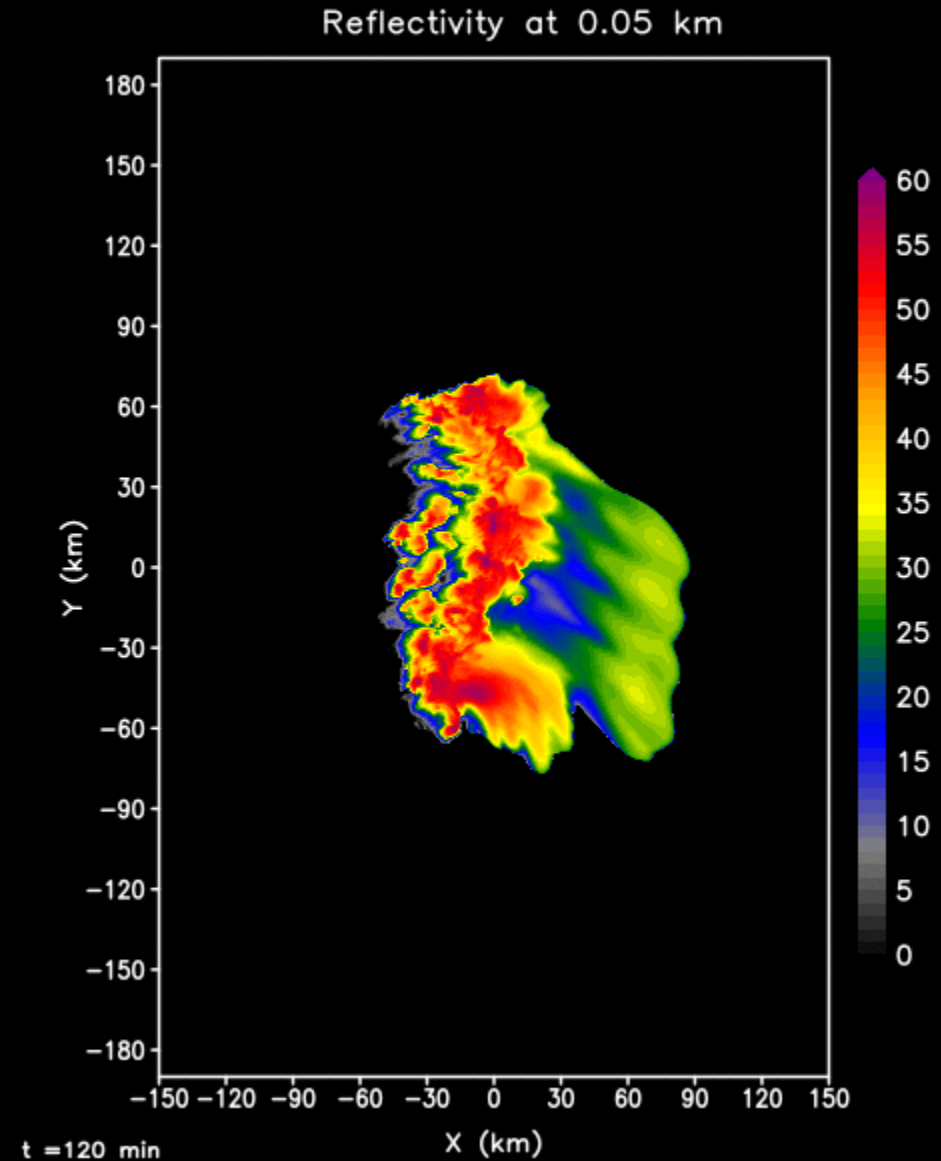
1. What physically drives the transition of an elevated to surface-based MCS?
2. What are the primary drivers of the severe surface winds?

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Model Specifications

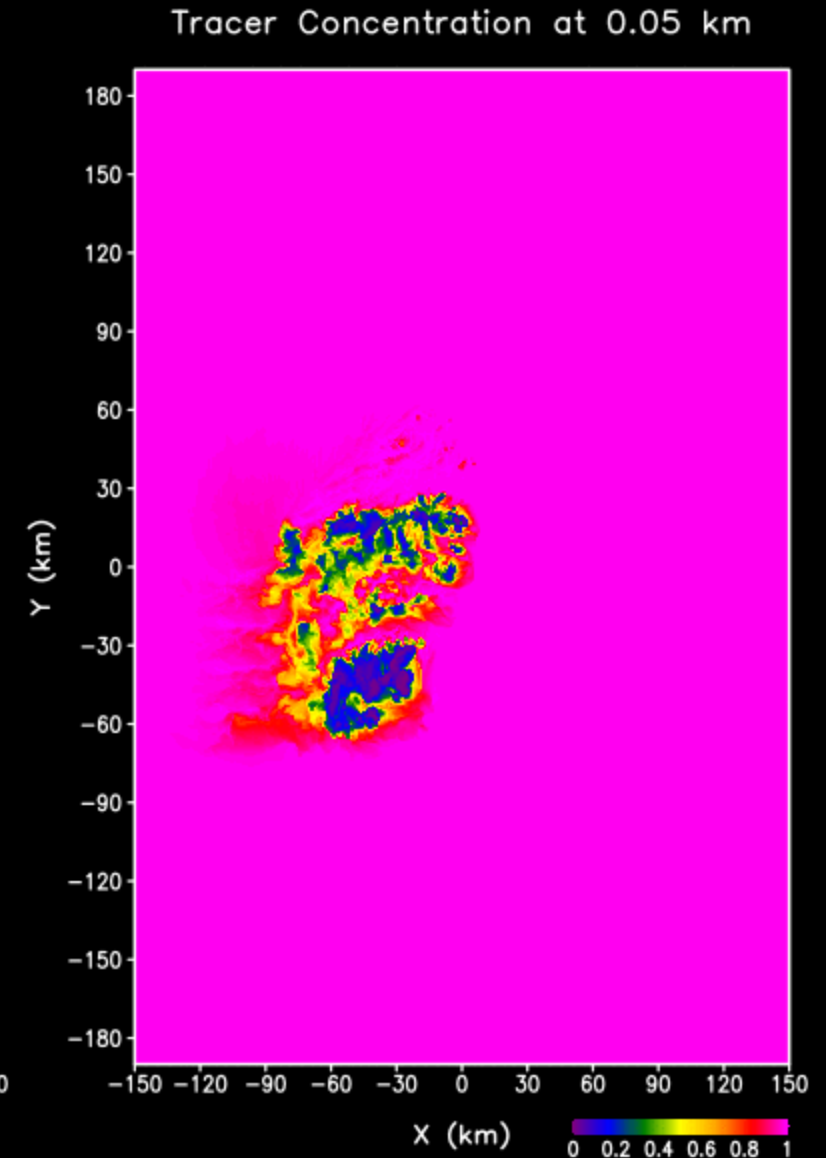
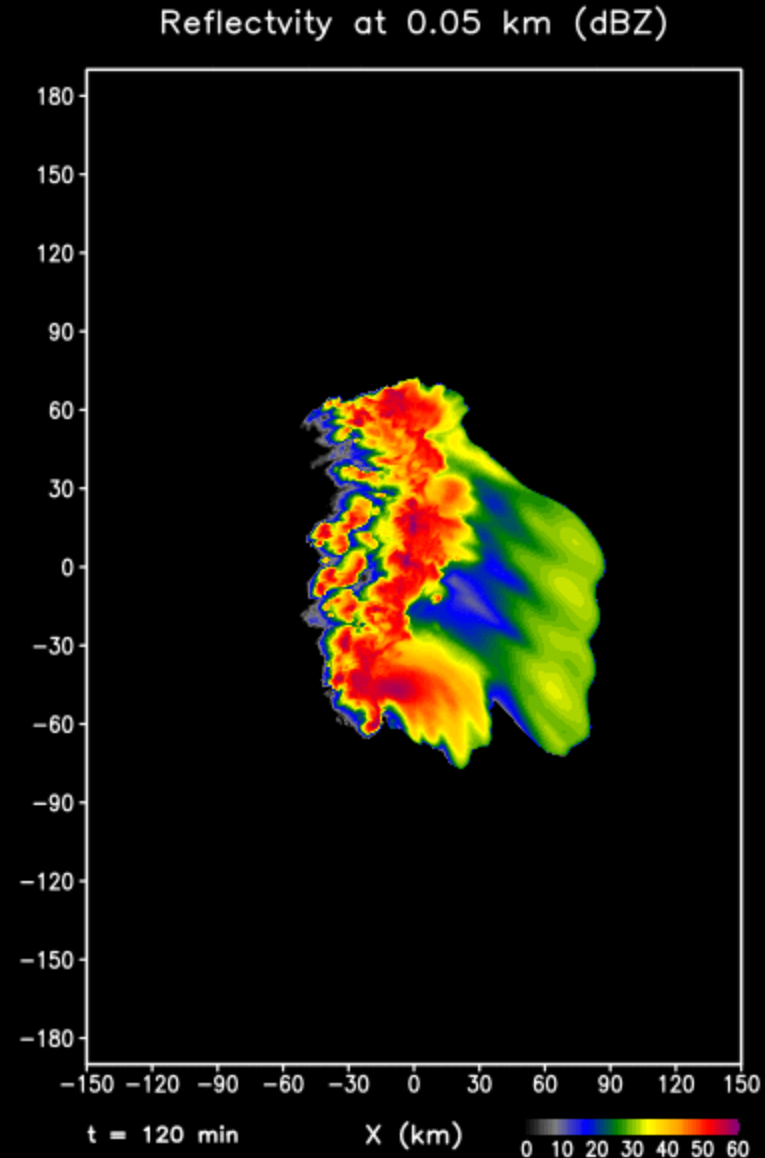
- CM1, Version 17
- 250 m horizontal grid spacing
- 100 to 250 m (stretched) vertical grid spacing
- NAM point sounding close to frontal boundary
- Morrison double-moment moisture scheme
- Initiation via line of warm bubbles
- Free slip upper/lower boundaries
- No Coriolis



Main Points:

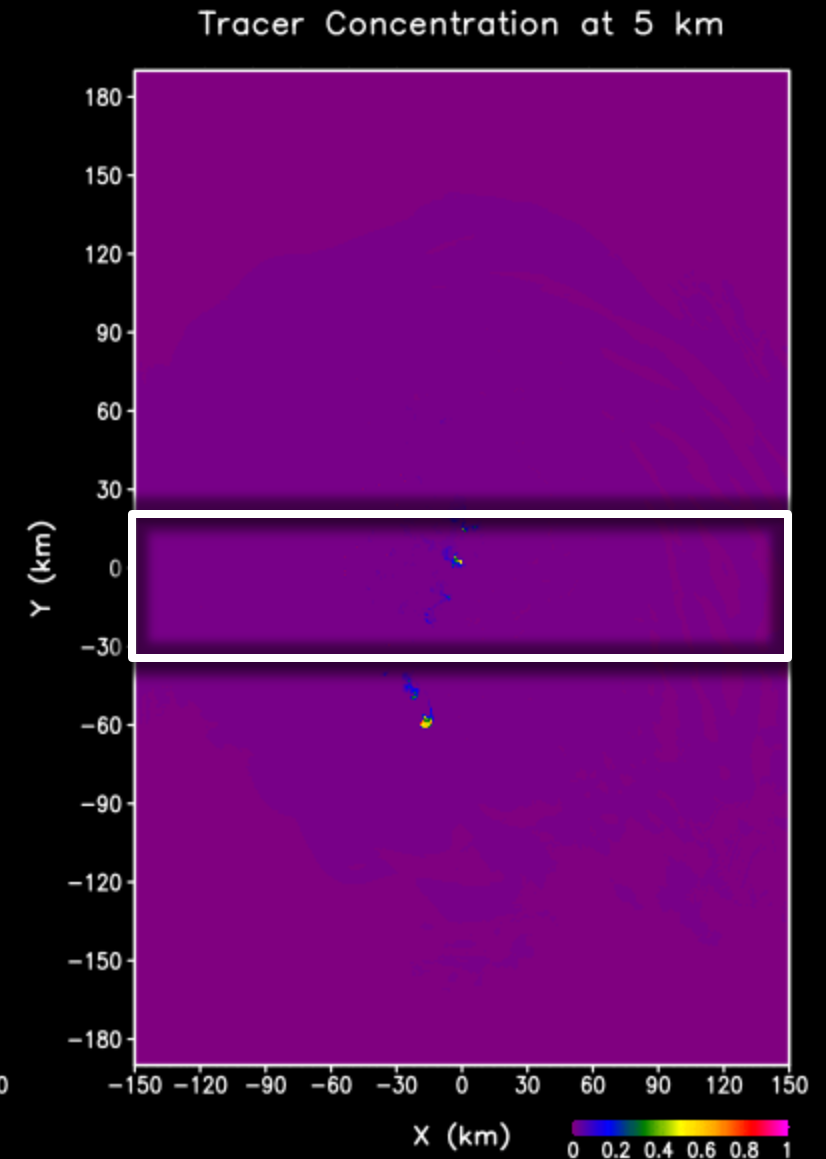
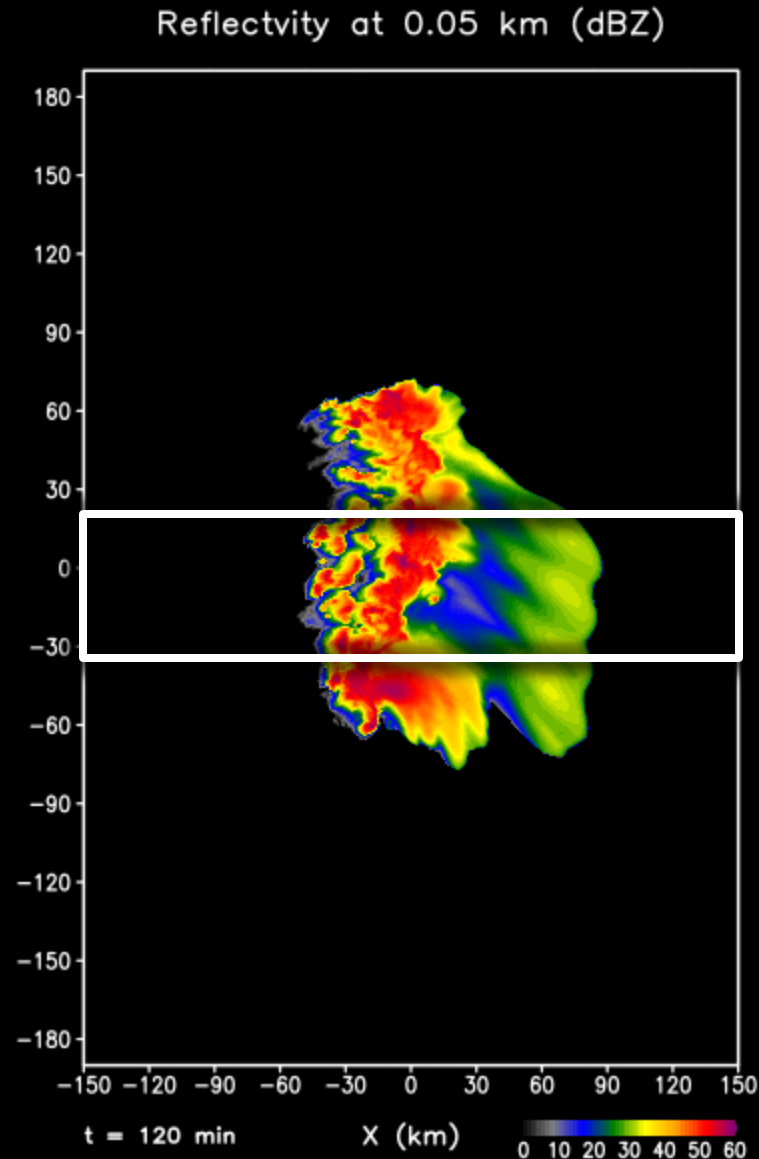
- Passive tracer originates between 0 – 500 m
- Gust front of simulated MCS lifts low-level tracers along majority of leading edge

...But not necessarily to parcel LFC

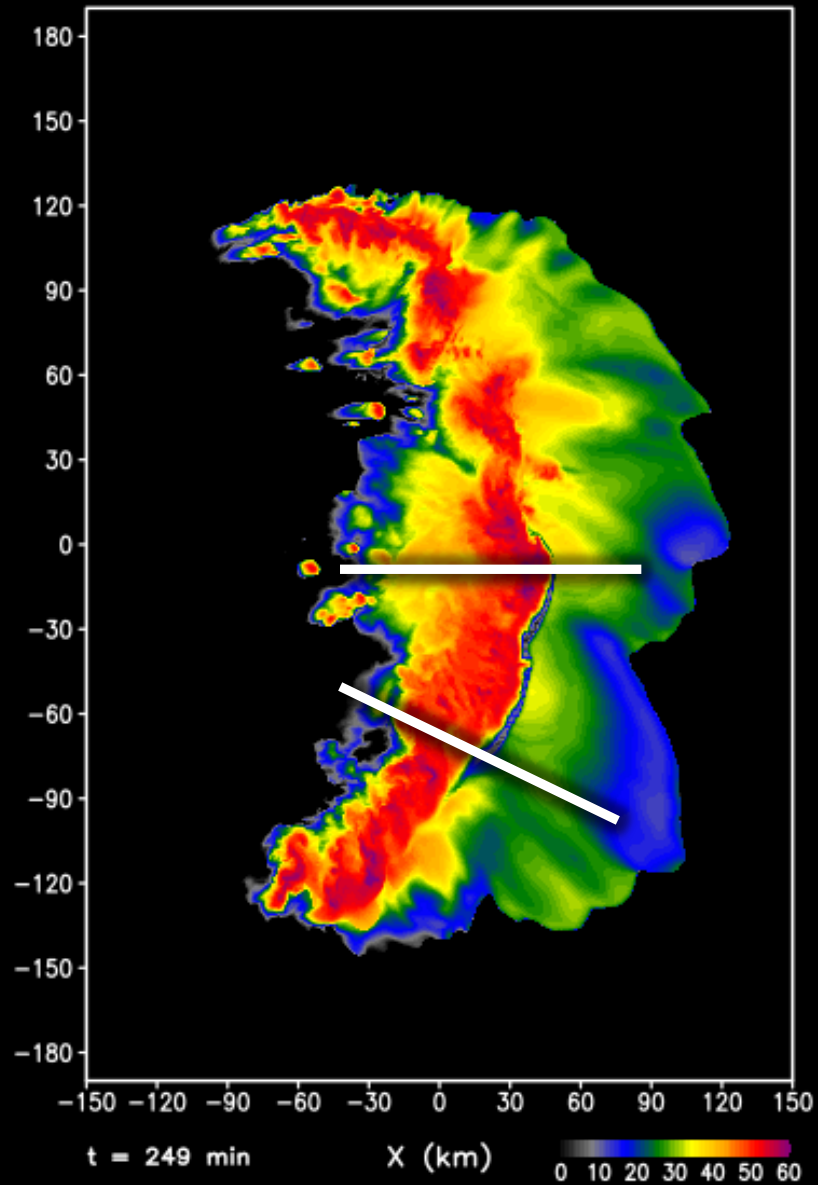


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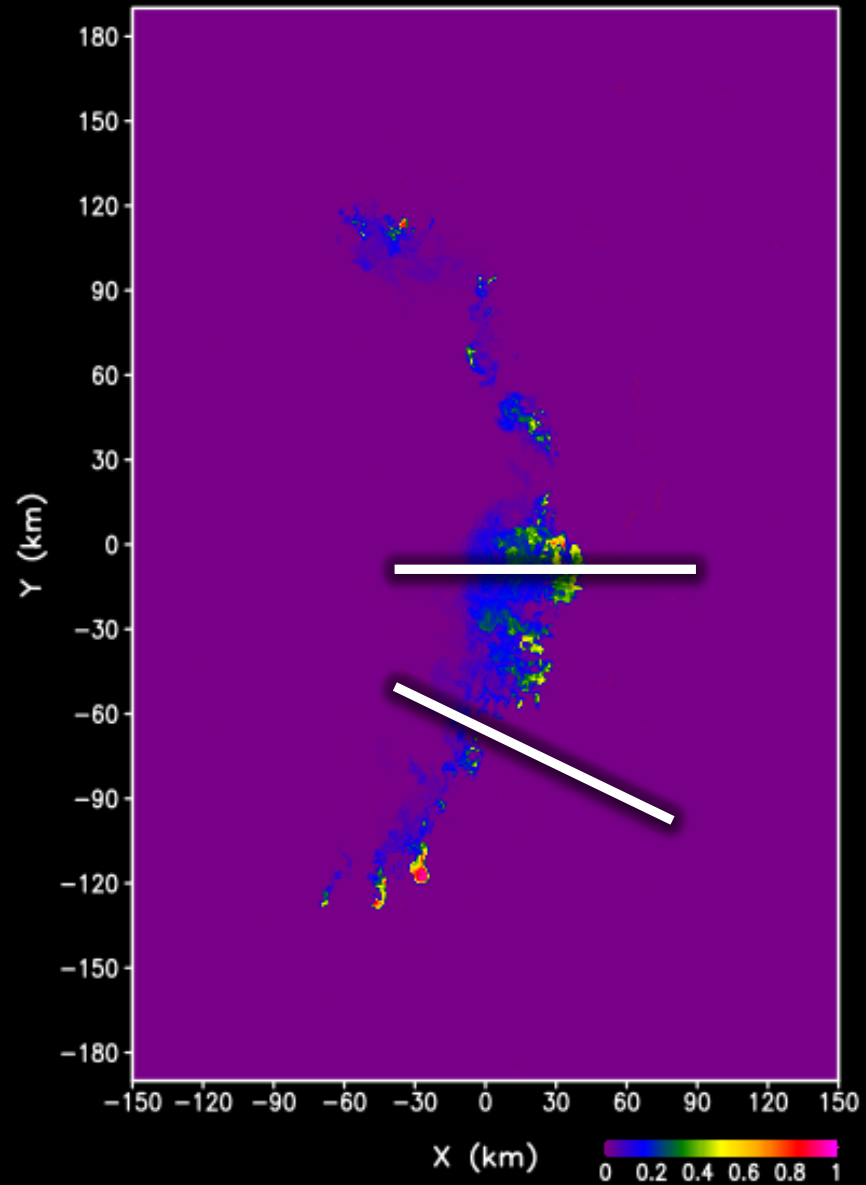
- Tracer originating below 0.5 km at 5 km is an indication of a surface-based updraft
- Bow echo is surface-based while southern segment is elevated
- Trajectory analyses (not shown) confirms the dichotomy

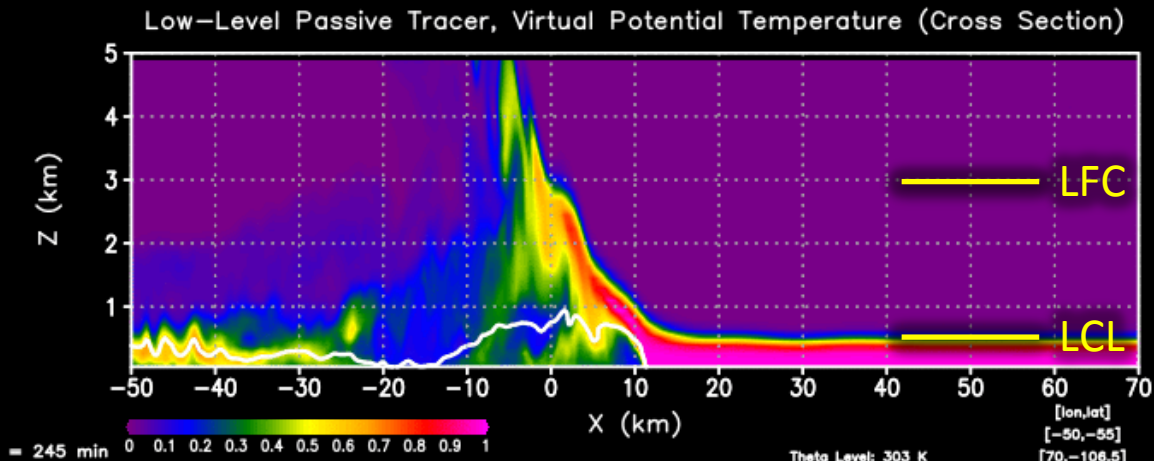
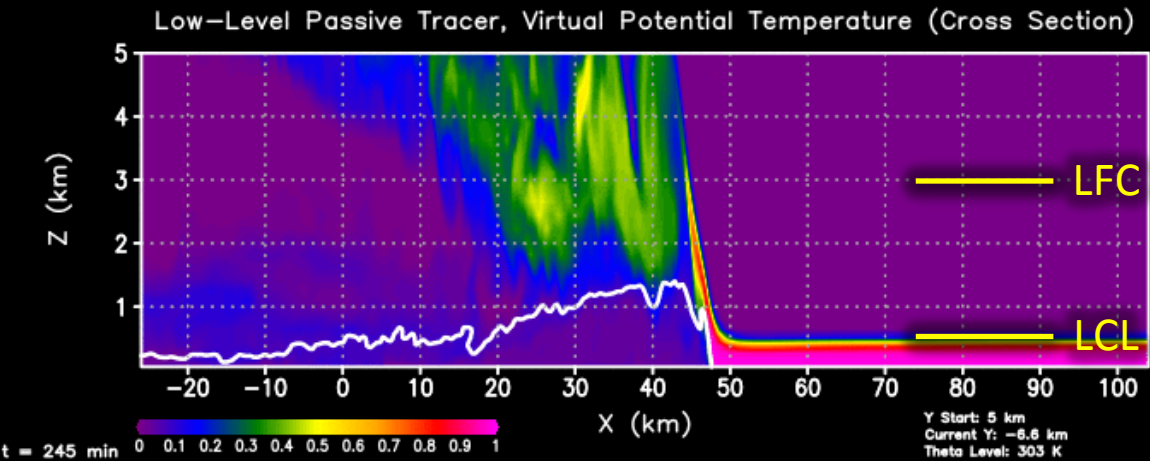
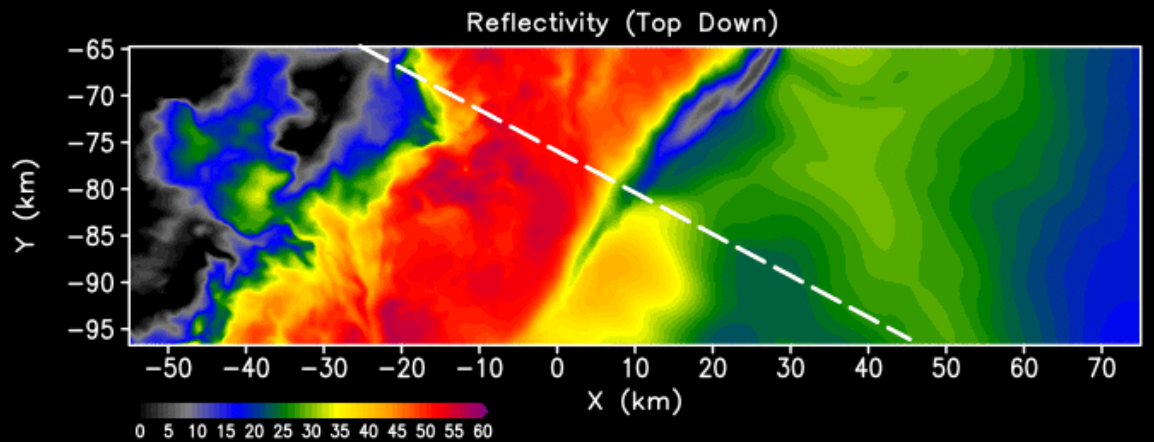
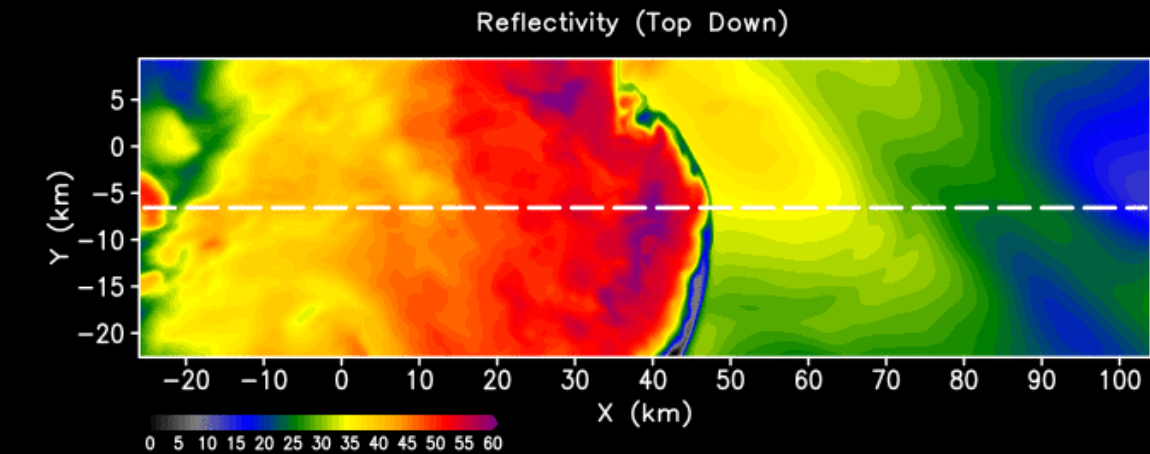


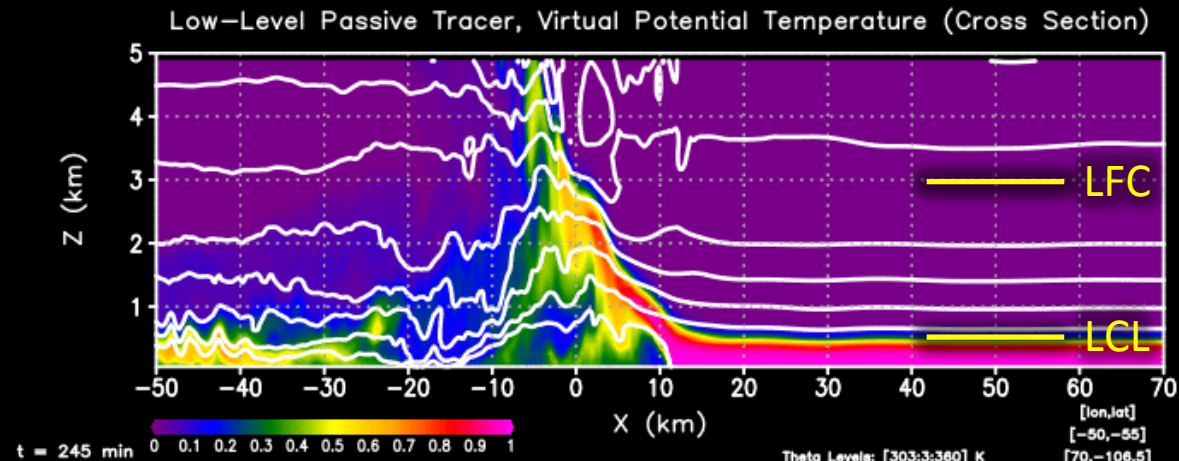
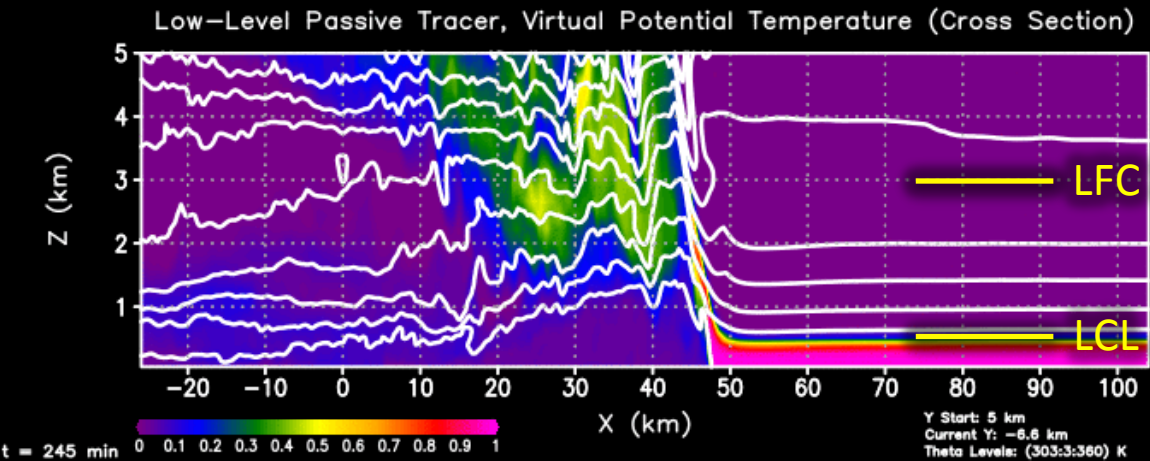
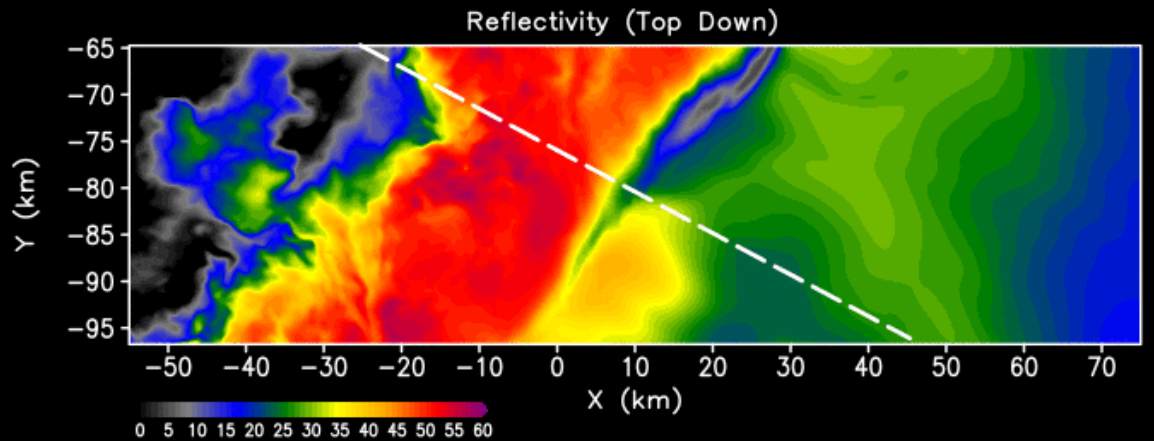
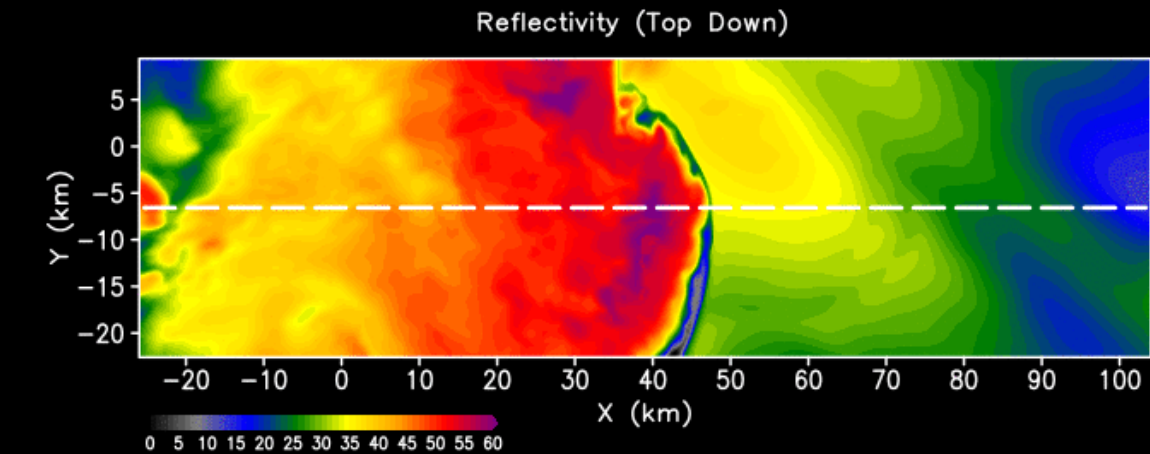
Reflectivity at 0.05 km (dBZ)



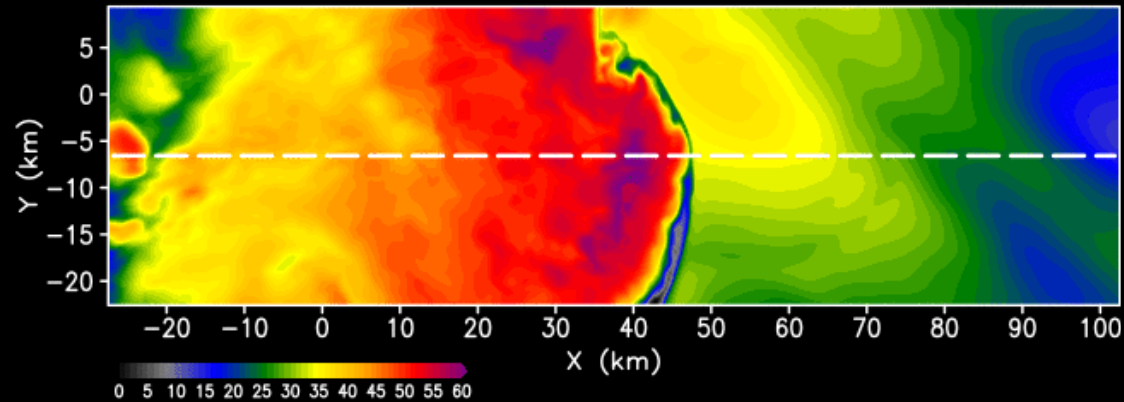
Tracer Concentration at 5 km



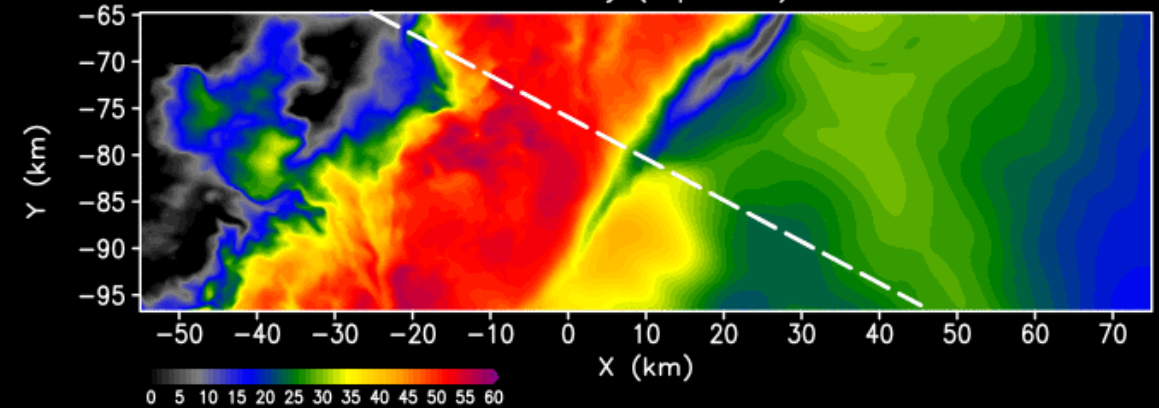




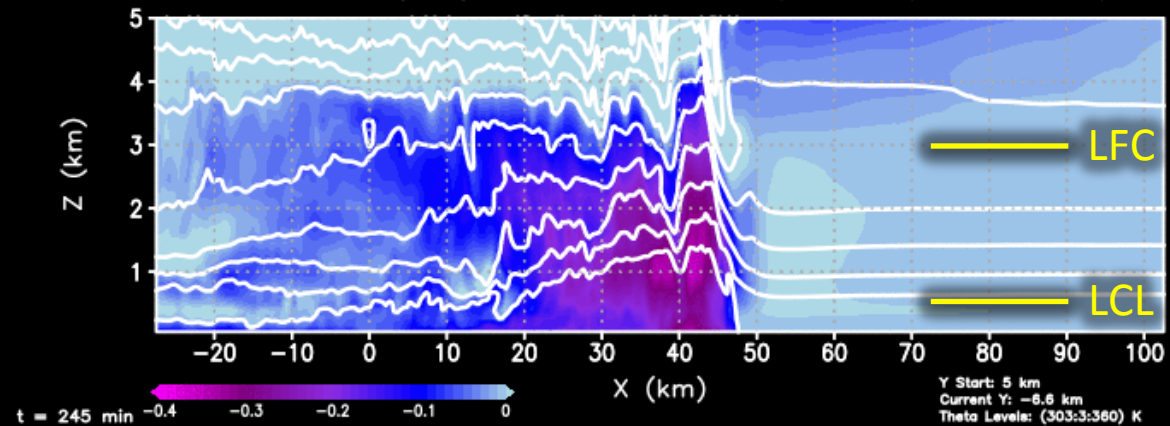
Reflectivity (Top Down)



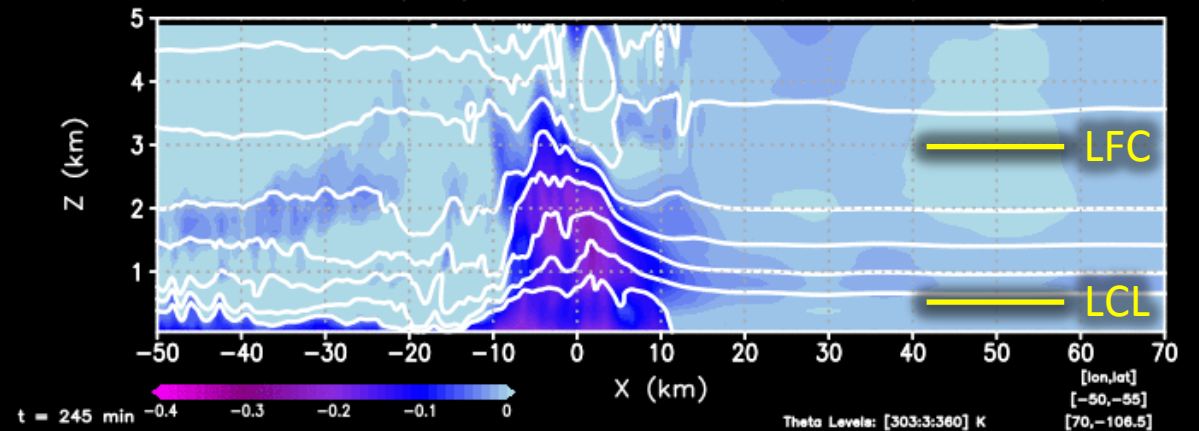
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Cold Pool Buoyancy, Virtual Potential Temperature (Cross Section)

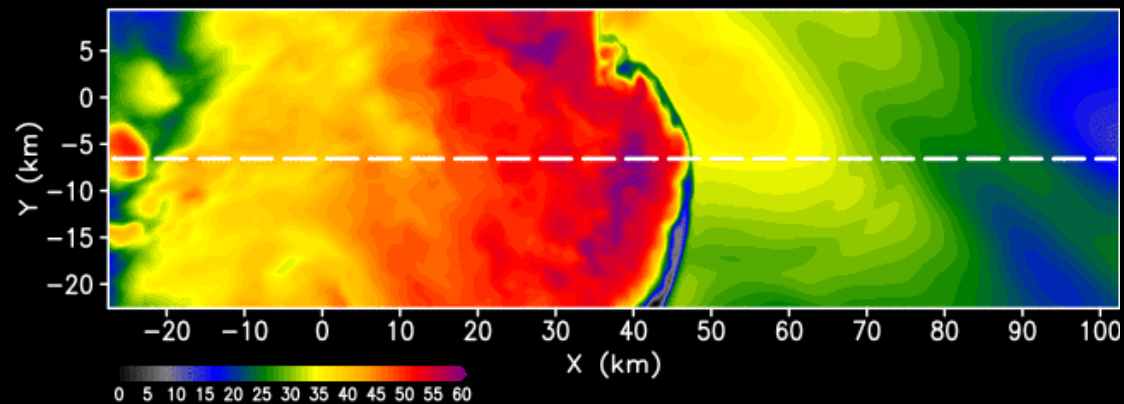


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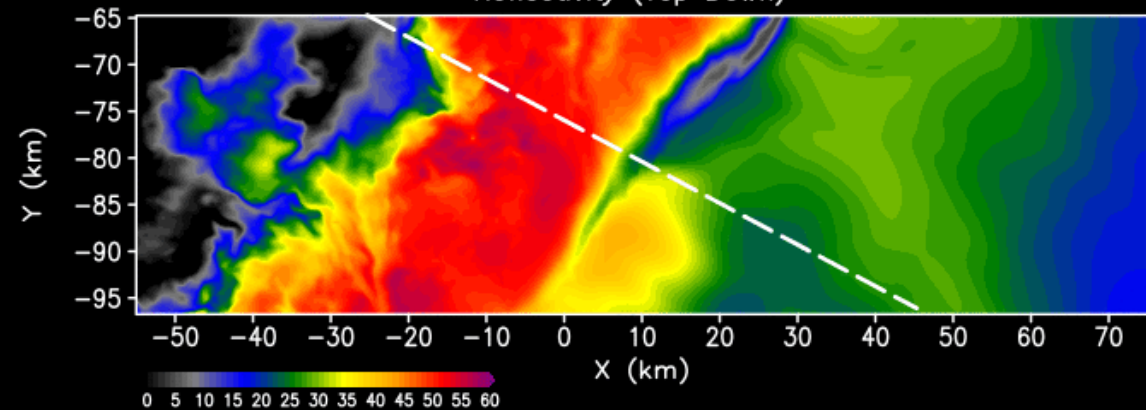


$$B = g \left(\frac{\theta'_v}{\theta_v} - r_c \right)$$

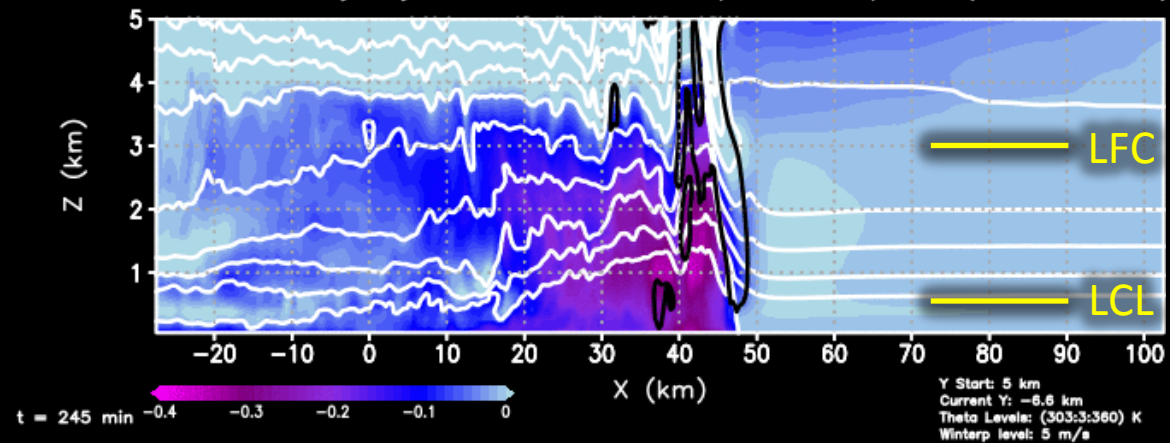
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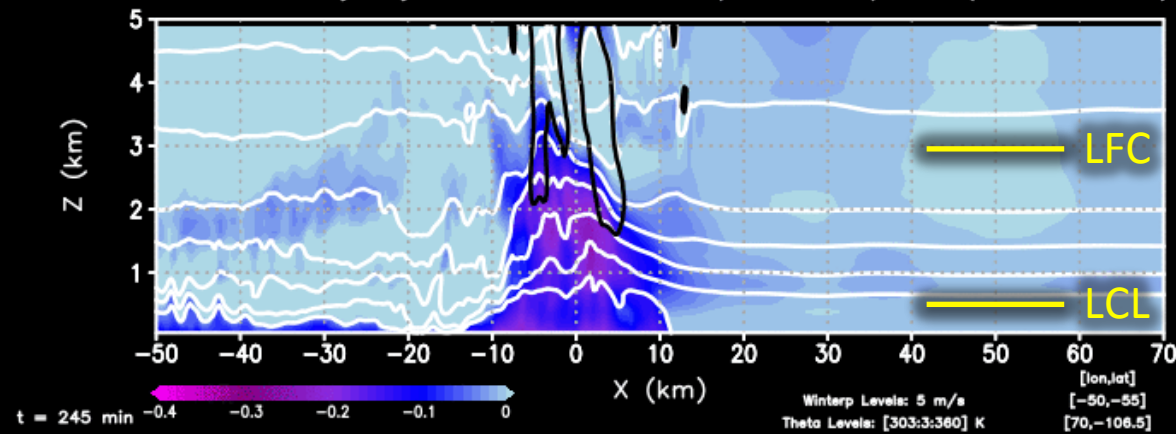
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Cold Pool Buoyancy, Virtual Potential Temperature, Updraft (Cross Section)



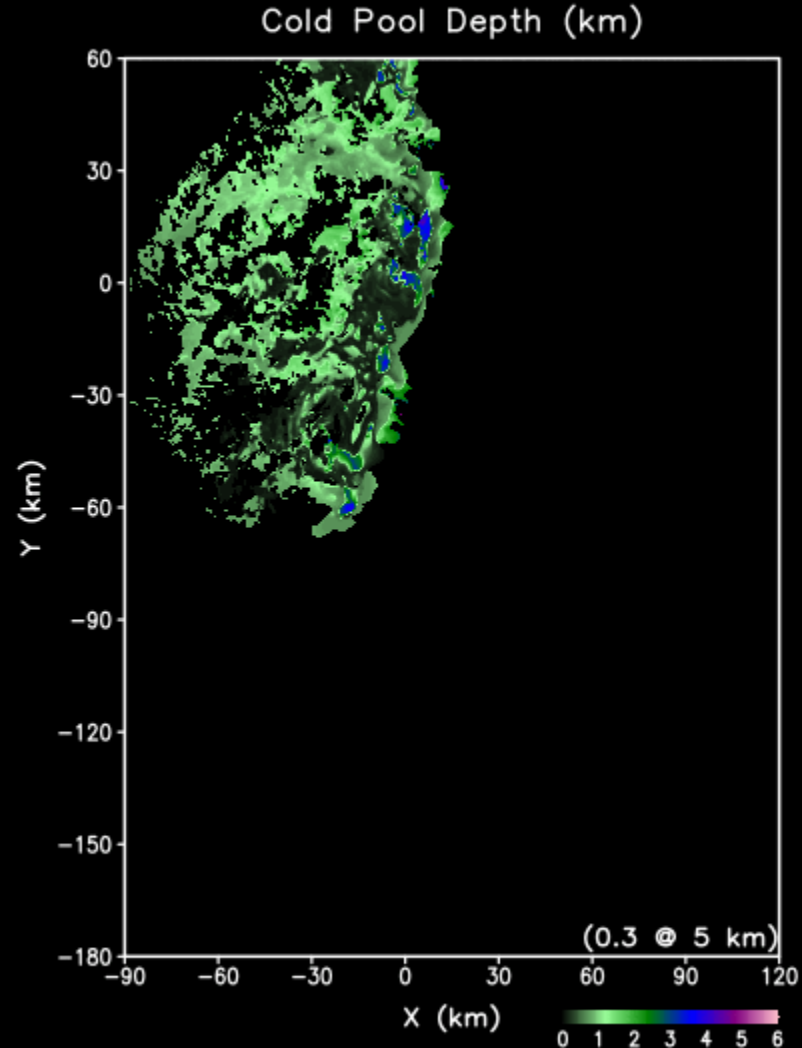
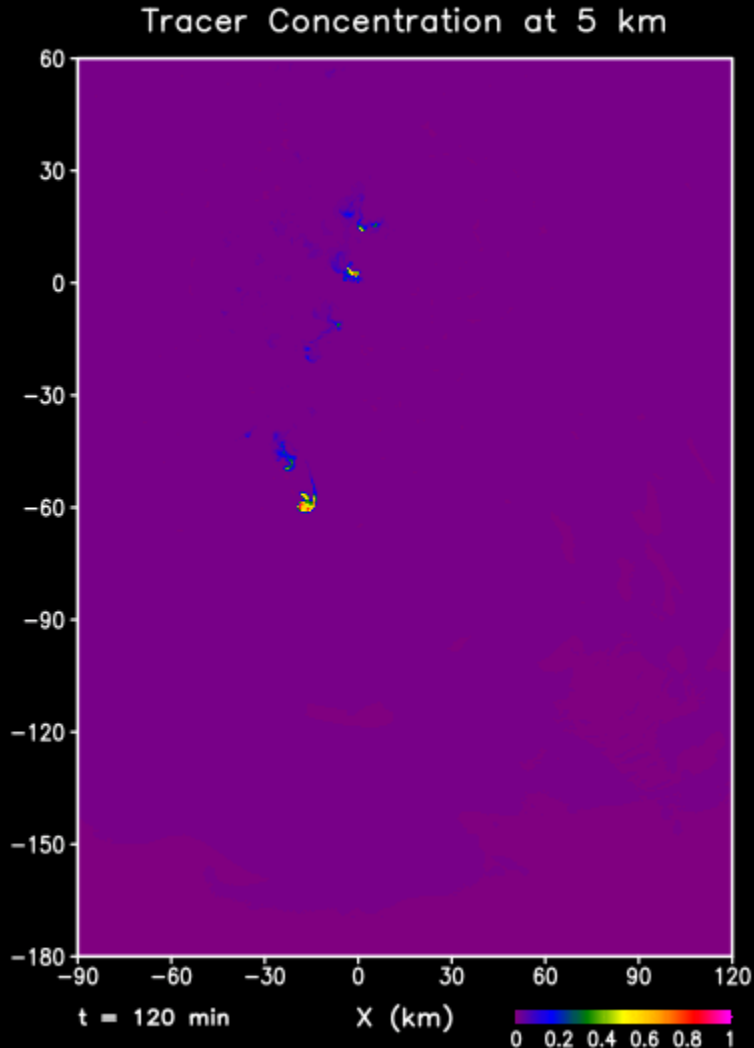
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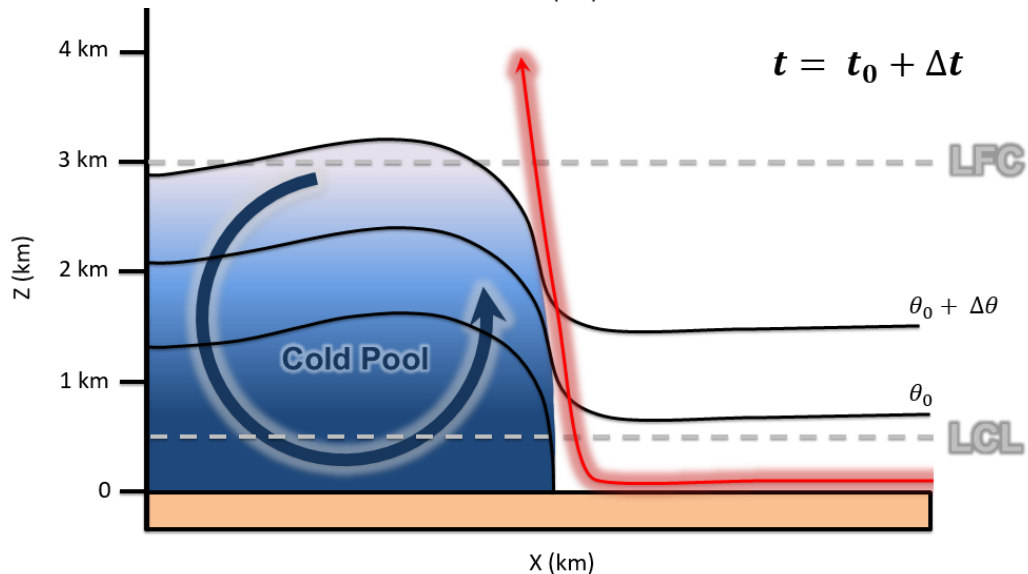
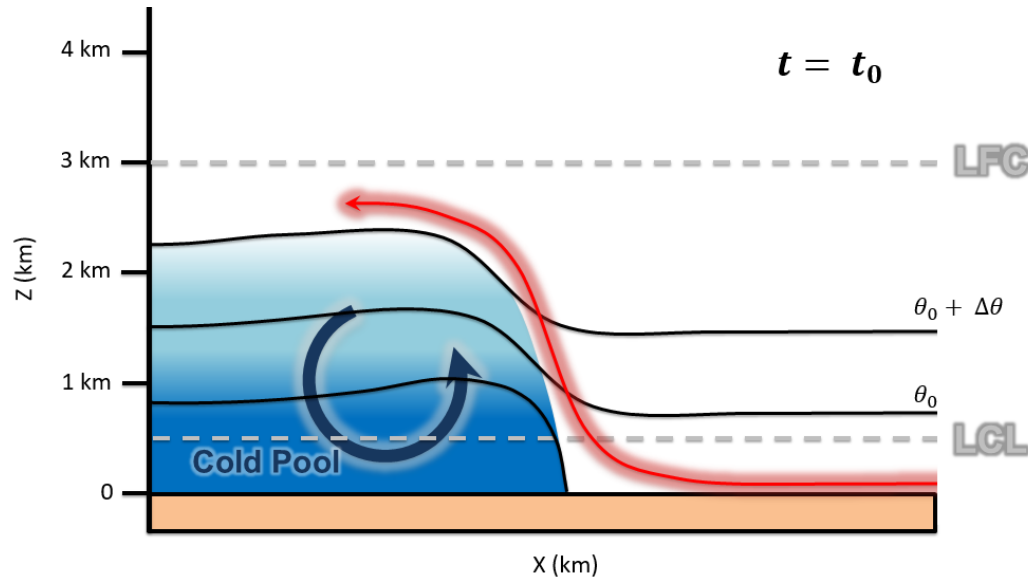
When the depth of the cold pool reaches the LFC of surface-based parcels, *the system becomes surface-based.*

**This
process
*takes time!***




**0-3 km
shear vectors**

Concluding Remarks: Time is the answer!



General Idea:

1. As elevated convection matures, cold pool strengthens & deepens
2. After some time, cold pool depth reaches LFC of surface-based parcels
3. System subsequently strengthens, enabling self-feedback loop

Other Factors:

- Strong line-normal shear may help build cold pool faster (e.g. RKW theory)
- Sensitivities?

