

# Stable and symmetric barotropic flows over topography

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## Quasi-geostrophic framework

The nondimensional equation

$$J(\psi, Q) \equiv \hat{\mathbf{k}} \cdot \nabla \psi \times \nabla [\nabla^2 \psi + h(x, y)] = 0 \quad (2.1)$$

governs the steady advection of potential vorticity  $Q = \nabla^2 \psi + h(x, y)$  by the geostrophic current  $\mathbf{u} = \hat{\mathbf{k}} \times \nabla \psi$  in the presence of the  $O(1)$  bathymetry  $z = h(x, y)$ . The assumed fluid domain

$$D = [0 \leq x \leq \pi] \times [0 \leq y \leq \pi] \quad (2.2)$$

is included in a certain  $f$ -plane, and the usual no mass-flux boundary condition

$$\psi(x, y) = 0 \quad \forall (x, y) \in \partial D \quad (2.3)$$

will be applied to single out the model solutions. If

$$\frac{dQ}{d\psi} = -2 \quad (2.5)$$

then the solutions of problem (2.1), (2.3) are stable in the

norm  $\|\phi\| = \left( \int_D \phi^2 dx dy \right)^{1/2}$  where  $\phi = \phi(x, y, t)$  is any

perturbation superimposed to the basic state  $\psi$  and satisfying the time-dependent version of (2.1), (2.3).

## A class of stable solutions

A special solution is Taylor's vortex

$$\Psi = \sin(x)\sin(y) \quad (2.7)$$

over a flat bottom, i.e., for  $h(x, y) = 0$ ; in fact, stream function (2.7) identically verifies (2.3) and (2.6). This solution

suggests to derive further solutions in the presence of a modulated bathymetry, that is for  $h(x, y) \neq 0$ , setting

$$\psi(x, y) = \Psi(x, y) [1 + \eta(x, y)] \quad (2.8)$$

where  $\eta(x, y)$  is determined by the request

$$\nabla^2 [\Psi(1 + \eta)] + h = -2\Psi(1 + \eta) \quad (2.9)$$

Note that, owing to (2.7), boundary condition (2.3) is satisfied by (2.8) provided that  $\eta(x, y)$  takes finite values everywhere in the fluid domain (2.2). Criterion (2.5) is still satisfied by (2.9), so solution (2.8) is stable. Equation (2.9) yields the following link between  $\eta(x, y)$  and  $h(x, y)$

$$h(x, y) = -2 \left[ \cos(x)\sin(y) \frac{\partial \eta}{\partial x} + \sin(x)\cos(y) \frac{\partial \eta}{\partial y} \right] - \sin(x)\sin(y) \nabla^2 \eta \quad (2.10)$$

In the present investigation only the class of functions

$$\eta(x, y) = -\frac{h_0}{2} (px + qy + rxy) \quad (2.11)$$

is taken into account, where  $h_0, p, q, r$  are  $O(1)$  parameters. Then, substitution of (2.11) into (2.10) yields the bathymetric profiles

$$h(x, y) = h_0 [\cos(x)\sin(y) (p + ry) + \sin(x)\cos(y) (q + rx)] \quad (2.12)$$

together with the stream functions

$$\psi(x, y) = \Psi(x, y) \left[ 1 - \frac{h_0}{2} (px + qy + rxy) \right] \quad (2.13)$$

which are produced according to (2.8).

Equation  $1 - \frac{h_0}{2} (px + qy + rxy) = 0$ . This line separates two current systems; once  $p, q, r$  are fixed, the shape and the extension of these currents depend on  $h_0$  in a way that will be clarified by means of some examples in what follows. Integration of (2.6) on  $D$ , results in the equation

$$\int_{\partial D} \mathbf{u} \cdot \hat{\mathbf{t}} ds = - \int_D (h + 2\Psi) dx dy \quad (2.14)$$

Then, substitution of (2.12) and (2.13) into (2.14) gives

$$\int_{\partial D} \mathbf{u} \cdot \hat{\mathbf{t}} ds = -8 + 2\pi h_0 (p + q) + \pi^2 h_0 r \quad (2.15)$$

If  $h_0 = 8[2\pi(p + q) + \pi^2 r]^{-1}$ , then the so obtained solution is constituted of two counter-rotating vortices, with opposite relative vorticity.

The fluid domain (2.2) is transformed into itself by the mirror reflections

$$R_1 : (x, y) \rightarrow (x', y') = (y, x) \quad (3.1a)$$

$$R_2 : (x, y) \rightarrow (x', y') = (\pi - y, \pi - x) \quad (3.1b)$$

$$R_3 : (x, y) \rightarrow (x', y') = (\pi - x, y) \quad (3.1c)$$

$$R_4 : (x, y) \rightarrow (x', y') = (x, \pi - y) \quad (3.1d)$$

If the bathymetry is invariant

$$h(x, y) = h(x', y') \quad (3.2)$$

under a certain reflection among (3.1a)-(3.1d) then the stream function is invariant under the same reflection for which (3.2) holds true, that is

$$\psi(x, y) = \psi(x', y') \quad (3.4)$$

In the following examples some bathymetric reliefs satisfying (3.2) will be considered.

## Current fields over some special bathymetric reliefs

a) If  $p=1, q=r=0$ . Then

$$h = h_0 \cos(x)\sin(y) \quad (4.1)$$

and

$$\psi = \sin(x)\sin(y)(1 - h_0 x/2) \quad (4.2)$$

according to (2.15) the two vortices contribute equally (and oppositely) to relative vorticity for  $h_0 = 4/\pi$ . These results are illustrated in [left panel](#).

b) If  $p=q=1, r=0$ . Then

$$h = h_0 \sin(x + y) \quad (4.4)$$

while (2.13) becomes

$$\psi = \sin(x)\sin(y) \left[ 1 - \frac{h_0}{2} (x + y) \right] \quad (4.5)$$

Equation (2.15) shows that the two vortices have opposite vorticities for  $h_0 = 2/\pi$ . [Middle panel](#) shows the isobaths and some snapshots of the model solution

c) If  $p=q=0, r=1$  then

$$h = h_0 [y \cos(x)\sin(y) + x \sin(x)\cos(y)] \quad (4.6)$$

and

$$\psi(x, y) = \sin(x)\sin(y) \left( 1 - \frac{h_0}{2} xy \right) \quad (4.7)$$

If  $h_0 > 2/\pi^2$ , the streamlines of (4.7) exhibit the formation of a couple of counter-rotating vortices. [Right panel](#) shows the isobaths and some snapshots of the model solution.

