

# Is Nonlinearity Important in Non-Breaking Mountain Waves?

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Johnathan J. Metz and Dale R. Durran  
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14 July 2020



# Introduction

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- > Clearest example of nonlinear process is wave breaking, which results in gravity wave drag
- > Are nonlinear effects important without wave breaking?



# A Special Case: Constant N and U Flow

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- > Nonlinear equations reduce (without any small amplitude assumptions!) to the linear equation

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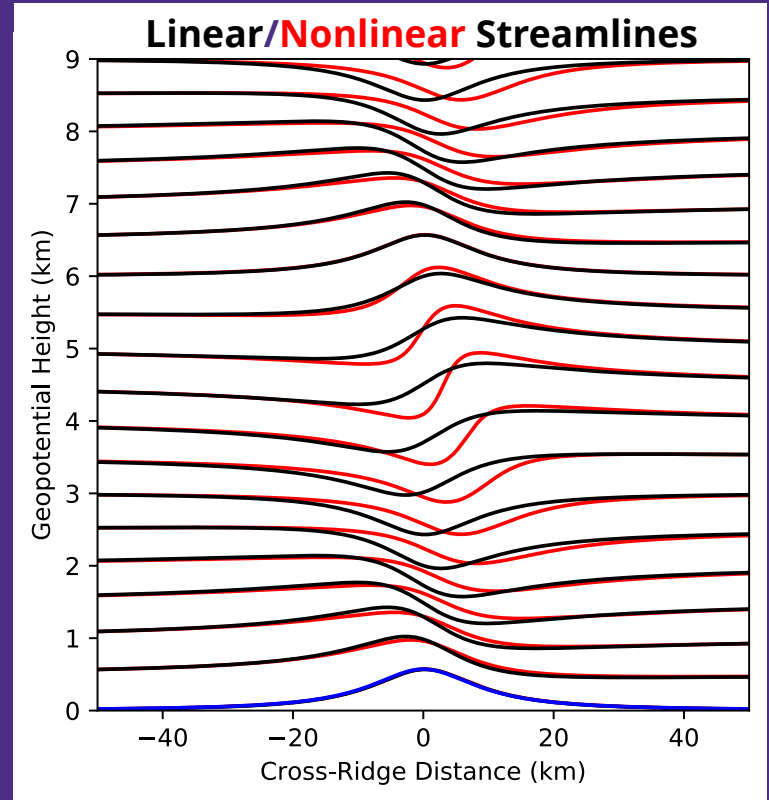


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- > Only difference in solutions is due to the finite-amplitude lower-boundary condition
- > Results in only minor differences between the linear and nonlinear solutions



# Linear Dynamics with a Tropopause

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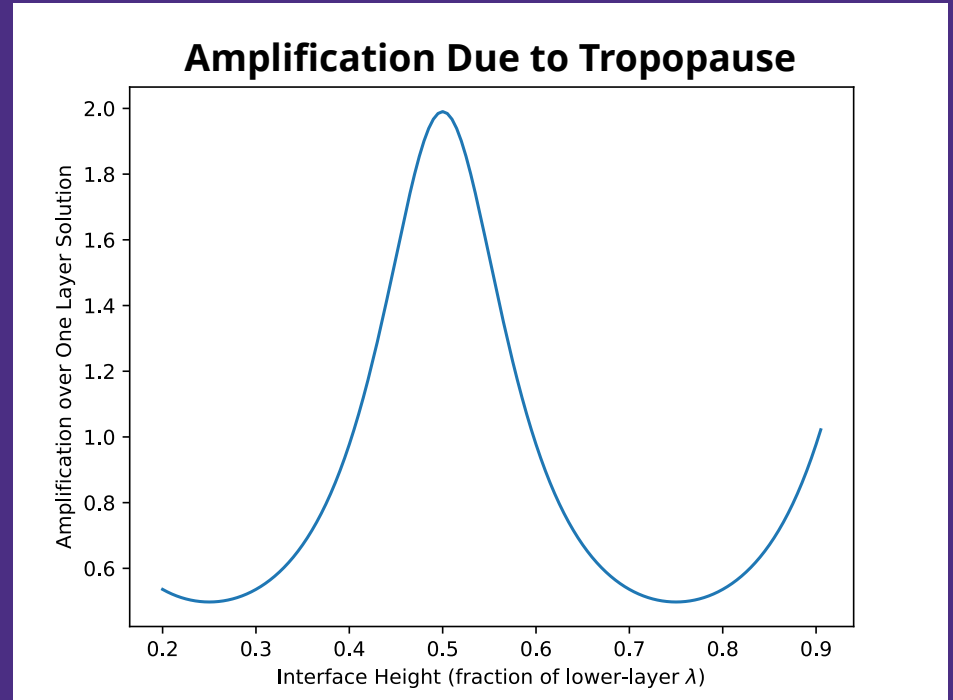
- > For two-layer system with  $N_S/N_T = 2.0$ :



# Linear Dynamics with a Tropopause

- > For two-layer system with  $N_S/N_T = 2.0$ :
- > Up to 2x amplification or deamplification in the surface pressure drag

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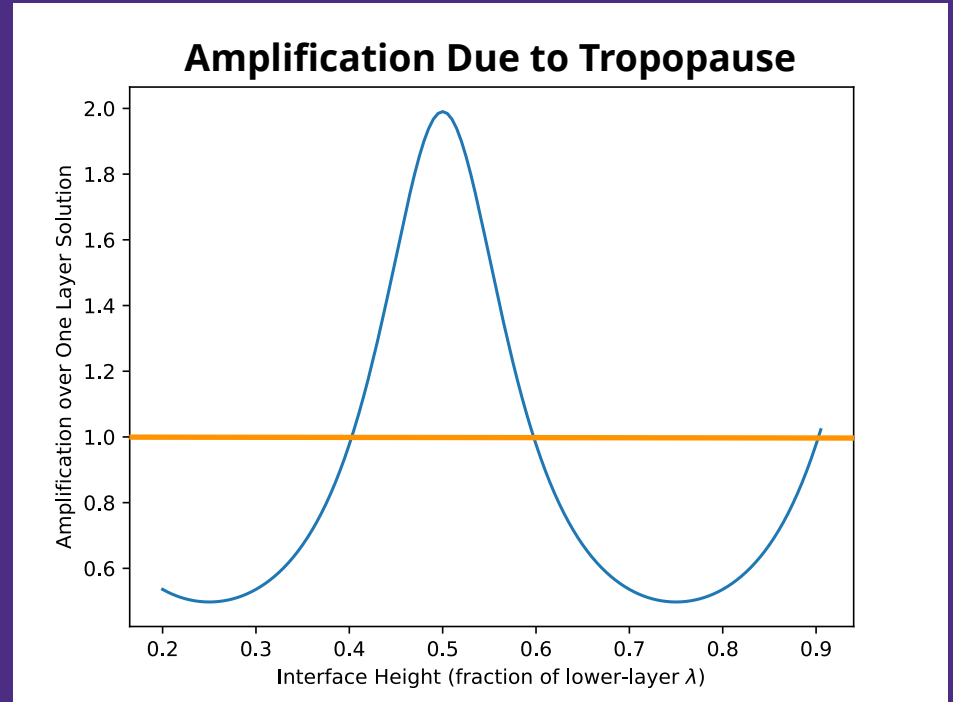


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$$D = \rho_0 \int_{-\infty}^{\infty} p' \frac{dh}{dx} dx$$

- > *In most GWD parameterizations, this curve would be a constant!*



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- > Using a semi-analytic solver, Durran (1992) found significant differences between the linear and nonlinear solutions in the two-layer system
- > However, semi-analytic methods are only available for constant  $N$  and  $U$  layers with infinitesimal transition layers between



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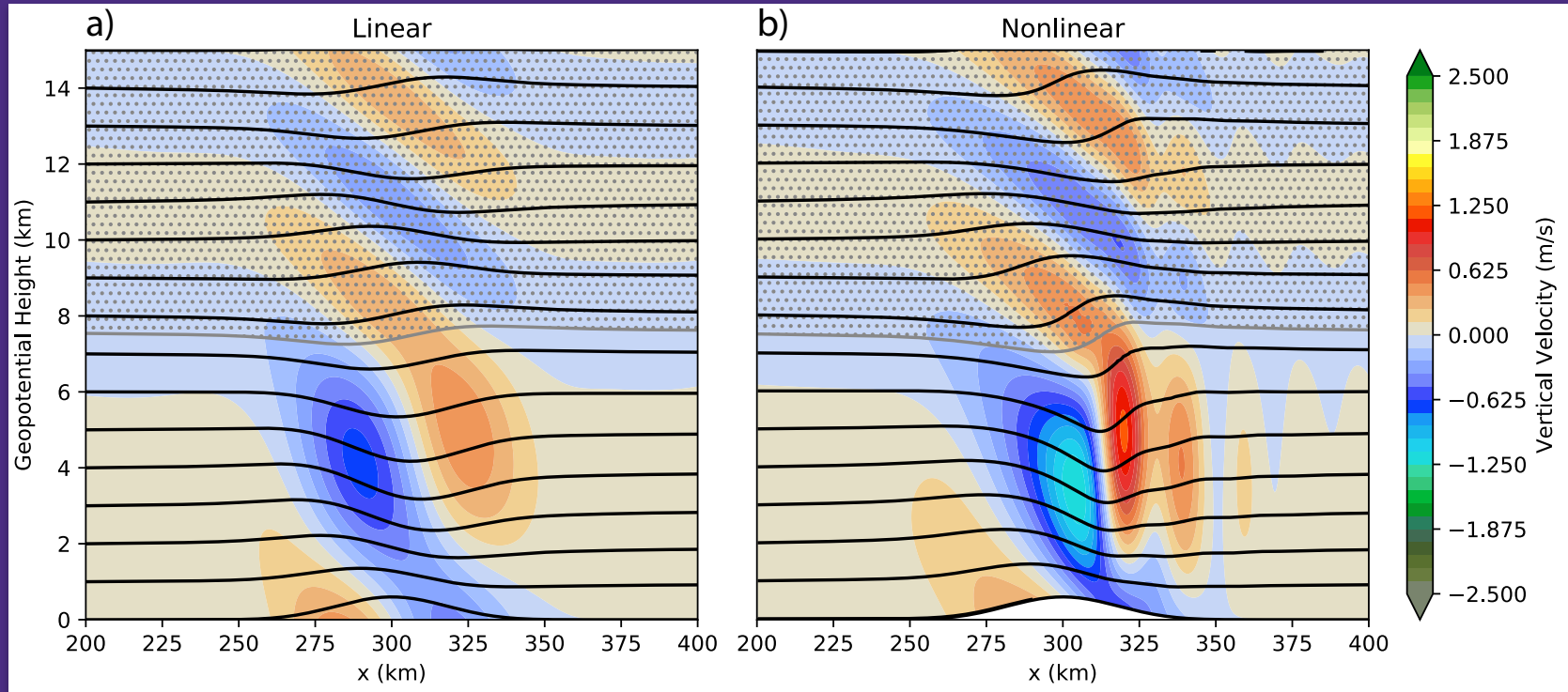
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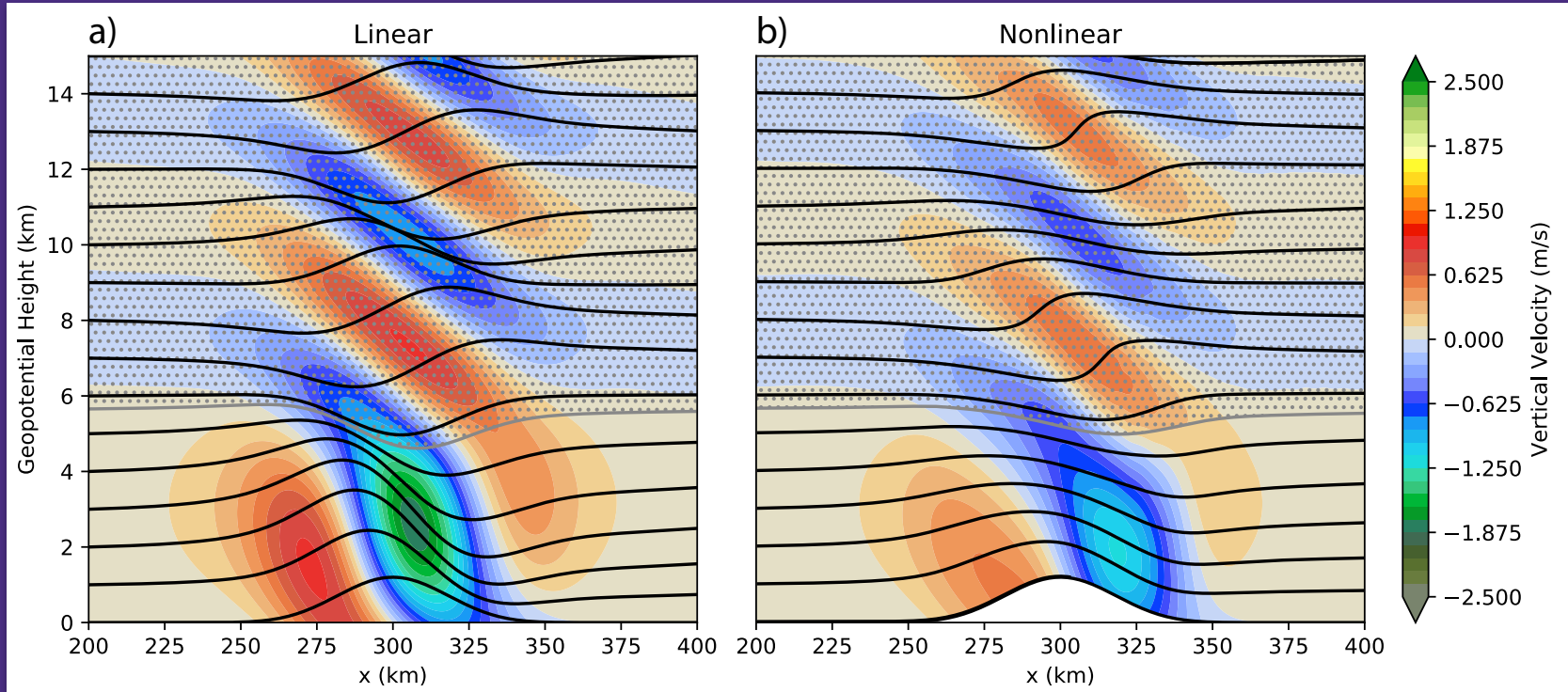
- > We already have a nonlinear time-dependent model (UW meso12)
- > We need a *linear* time-dependent model
- > Take meso12 and linearize advection terms and boundary conditions
- > Run both versions of the model in 2D Boussinesq configuration and compare the differences



# Two-Layer Nonlinear Amplification

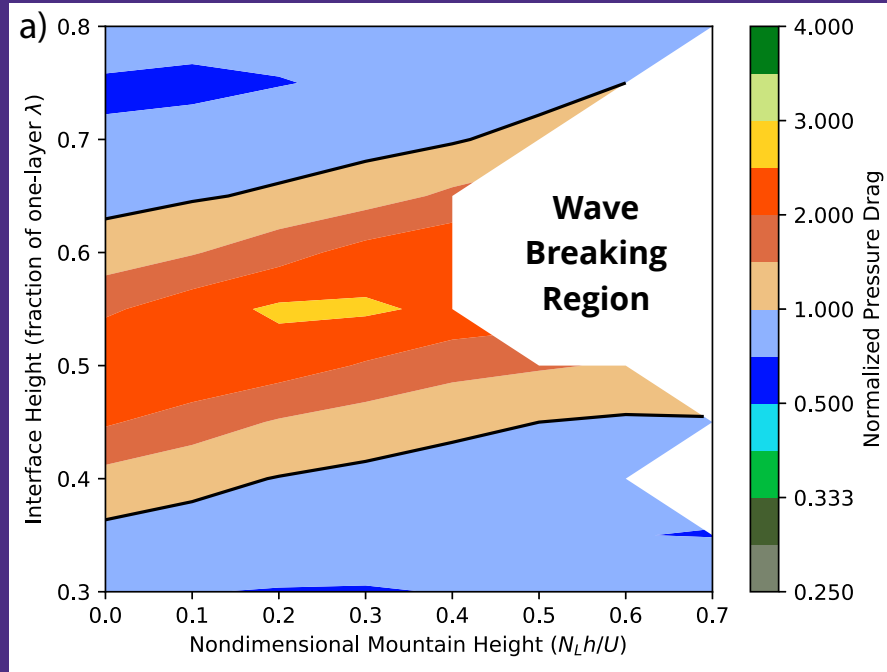


# Two-Layer Nonlinear Deamplification



# Two-Layer Constant Winds

## Normalized Pressure Drag



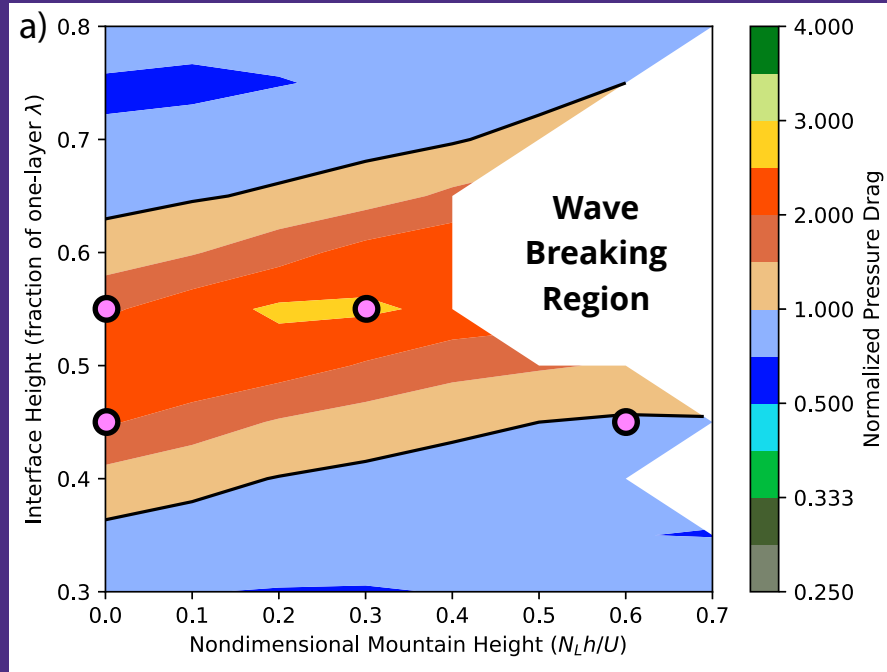
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Minimum: 0.64

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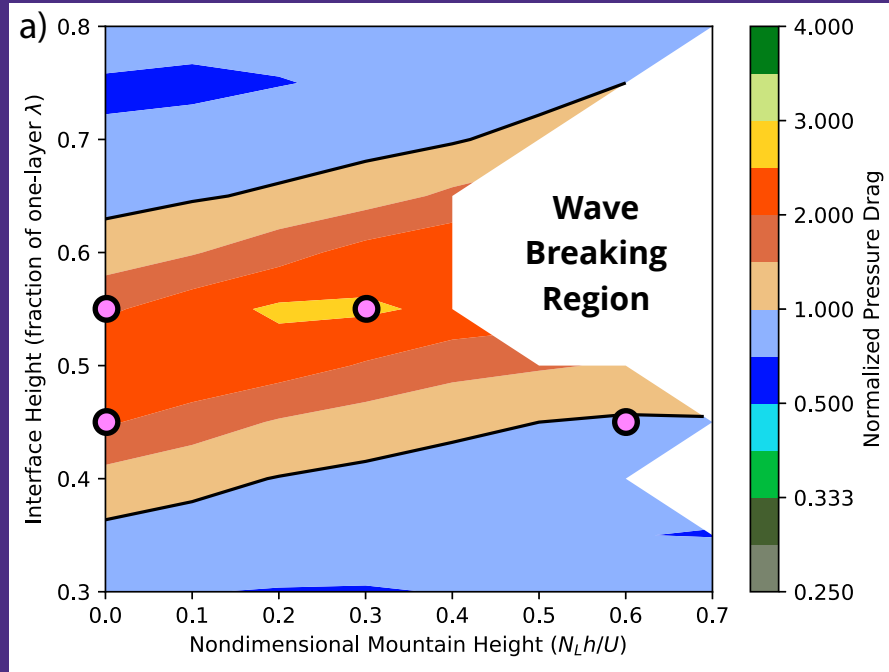
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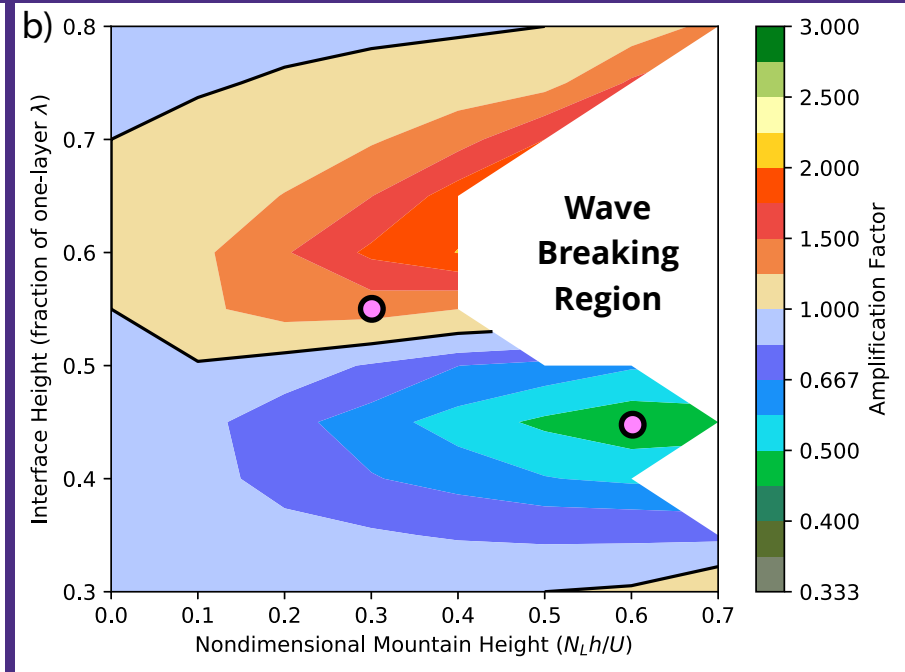
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# Two-Layer Constant Winds

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## Amplification over Linear Solution



# Finite-Depth Tropopause

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$(N = 0.02)$

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- > How sensitive are the previous finite-amplitude results to a smoother transition between layers?

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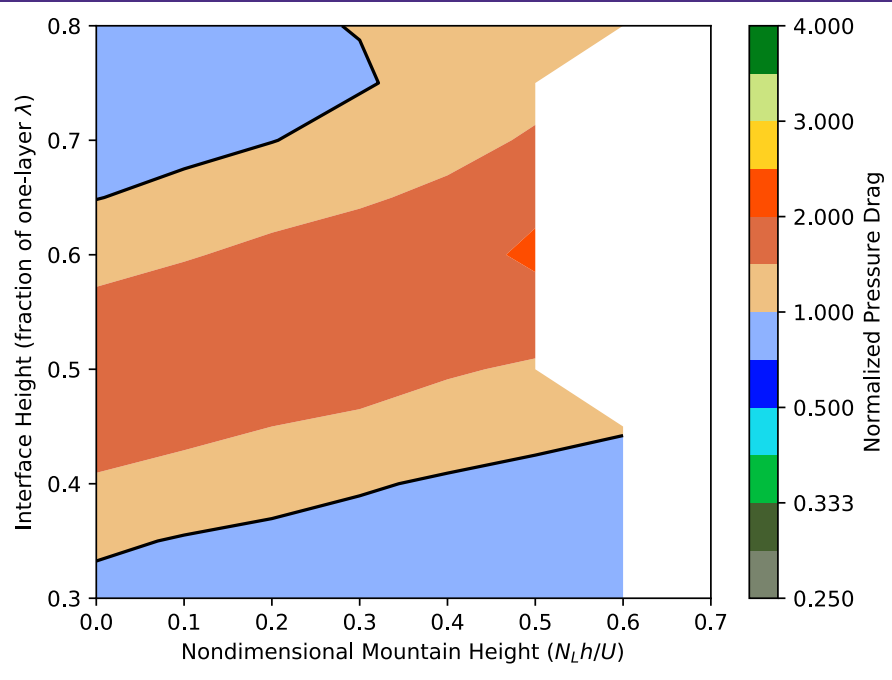
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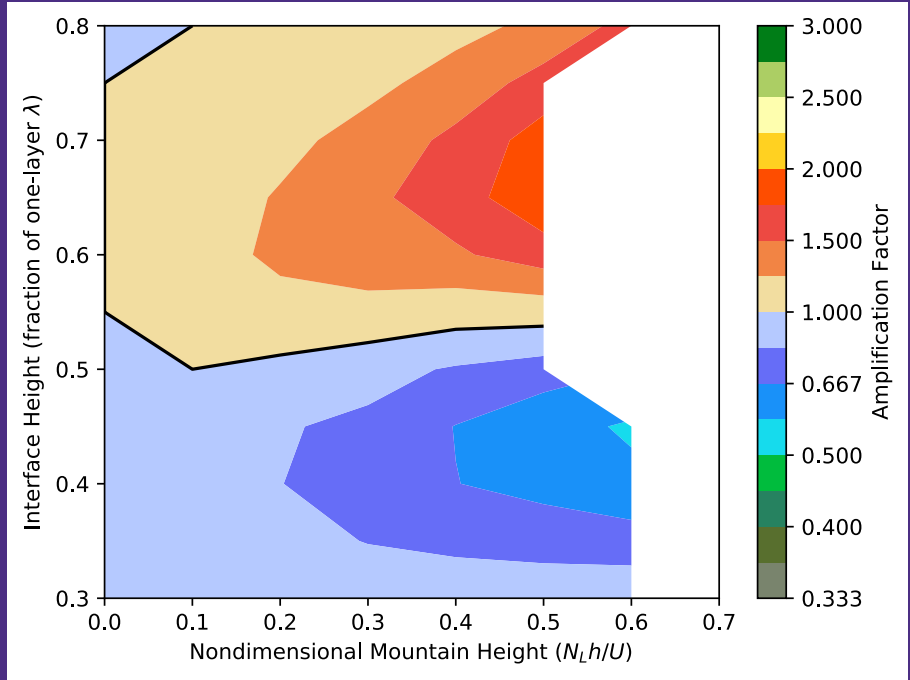
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# 2 km Thick Tropopause Transition

## Normalized Pressure Drag



## Amplification over Linear Solution



# Vertical Wind Shear

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- > The linear steady-state 2D Boussinesq wave equation in the presence of shear is

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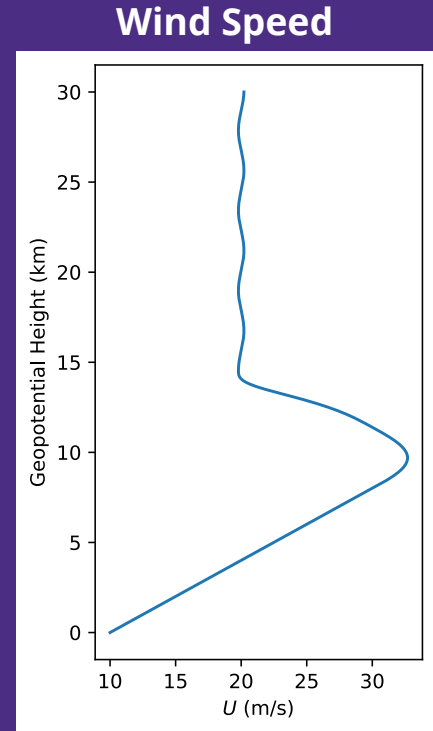
- > Clearly the basic state wind speed  $U$  is an important component of this equation
- > How important are nonlinear processes in a background state with a more realistic profile of  $U$ ?



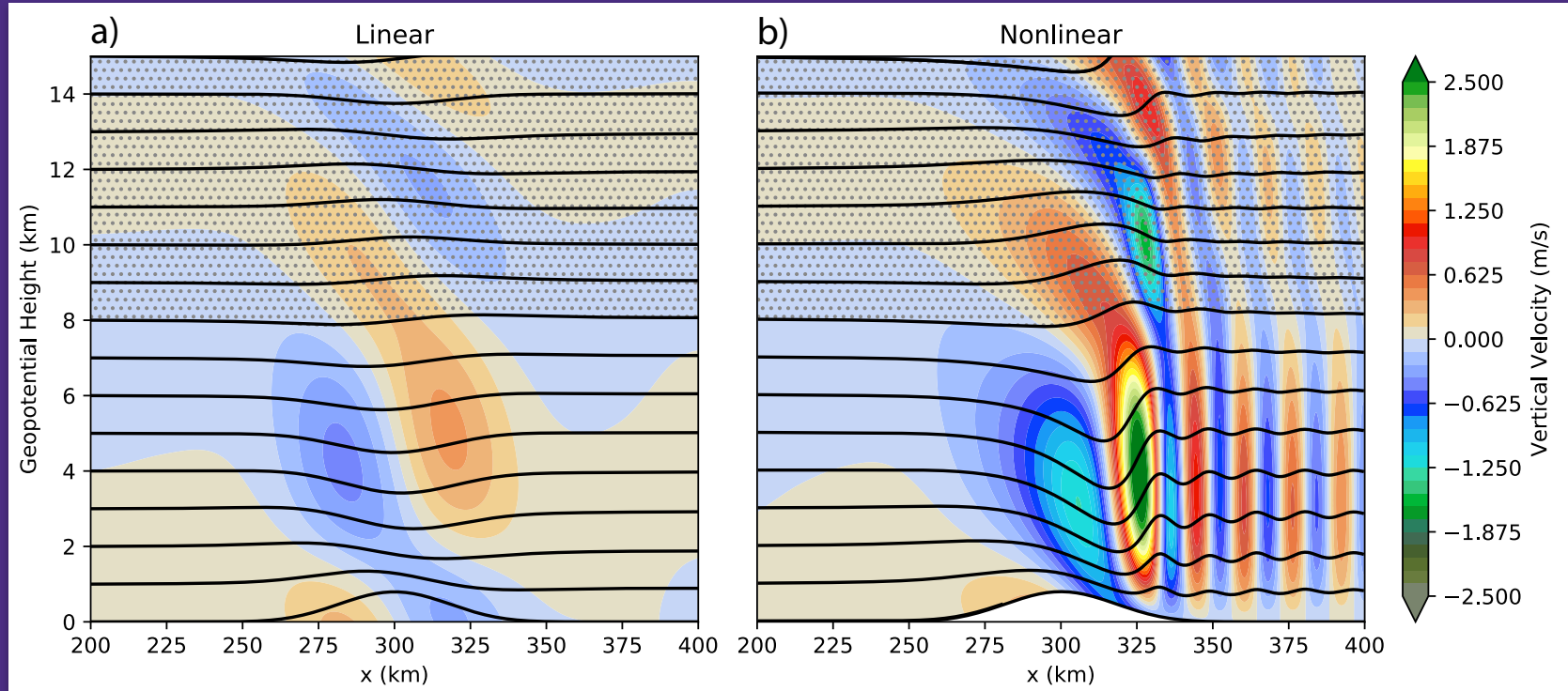


# Realistic Shear Profile

- > Constant shear from  $10 \text{ m s}^{-1}$  to  $30 \text{ m s}^{-1}$  in the troposphere
- > Relaxes back to  $\sim 20 \text{ m s}^{-1}$  in the stratosphere

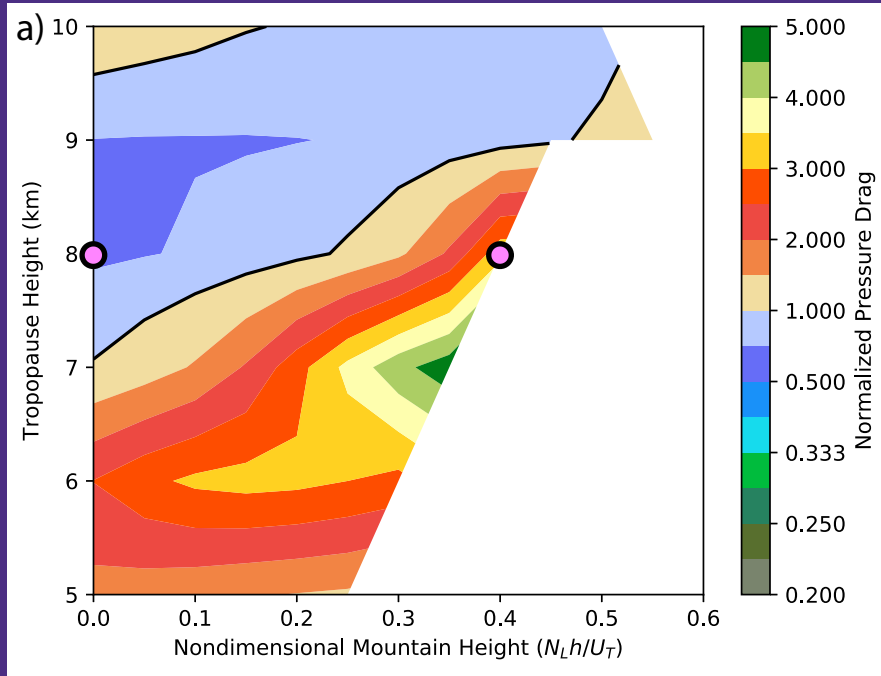


# Realistic Shear Nonlinear Amplification



# Two-Layer N Realistic Shear

## Normalized Pressure Drag



Maximum: 4.80

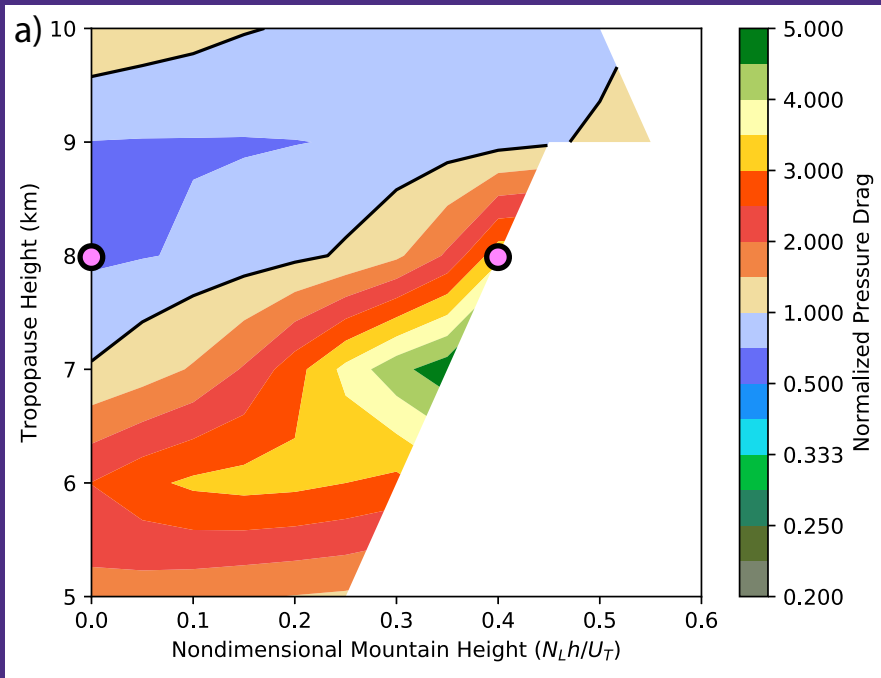
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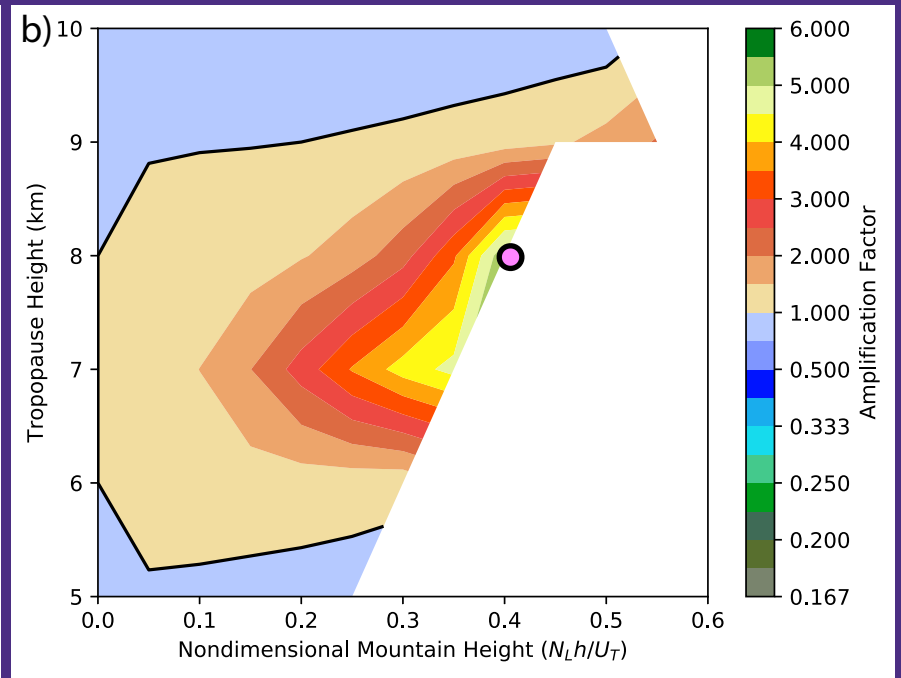
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- > **WKB theory fails at the tropopause**
- > **A parameterization could underestimate the true drag by a factor of 5!**
- > **Even a parameterization that accounts for the tropopause using linear theory would have significant error due to finite-amplitude effects**
- > **Associated difficulty with parameterization is one more reason that increasing resolution to explicitly resolve more of the wave spectrum is an important goal**

