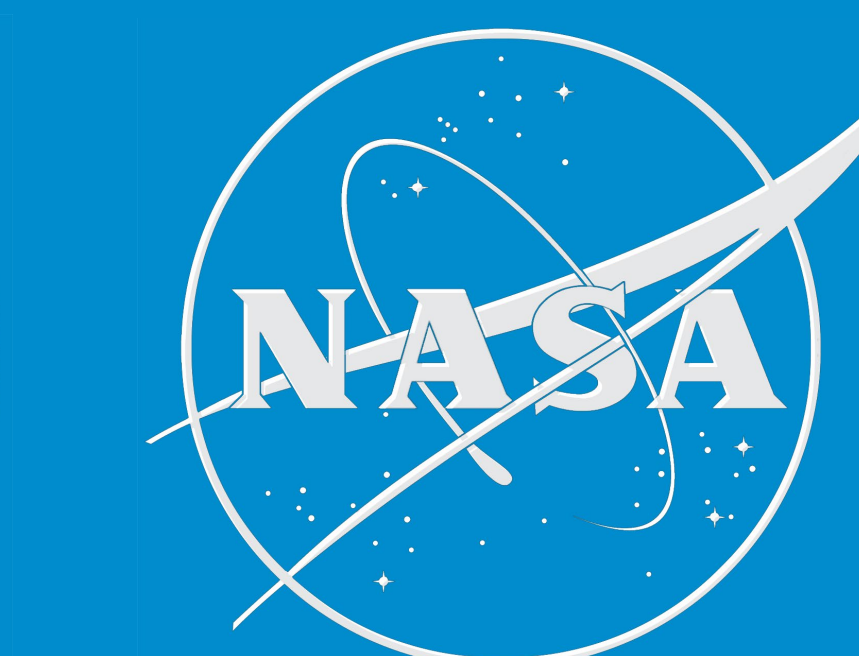
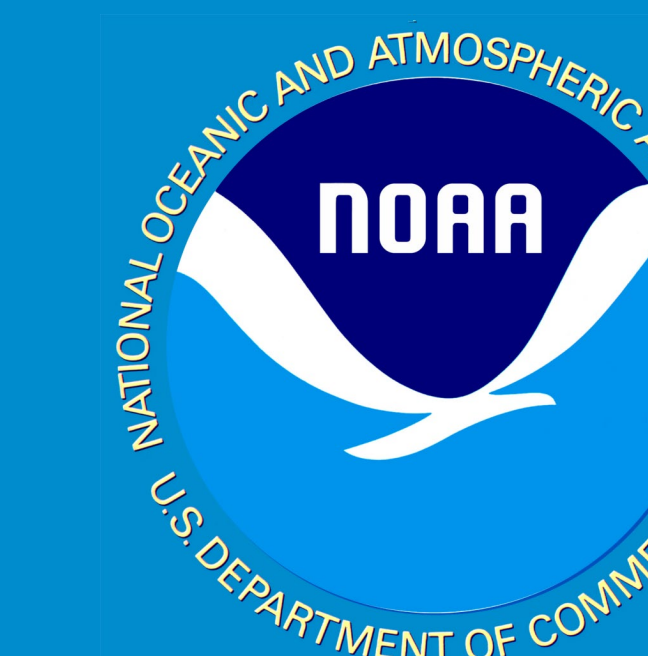
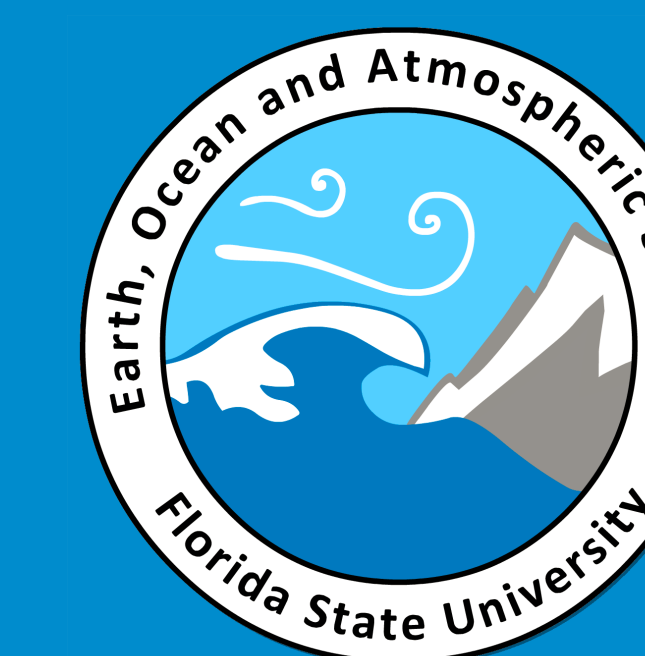


SIMULATED AZIMUTHAL STRUCTURE OF THE HURRICANE BOUNDARY LAYER IN HURRICANES IRMA ('17) AND EARL ('10) DURING INTENSITY CHANGE

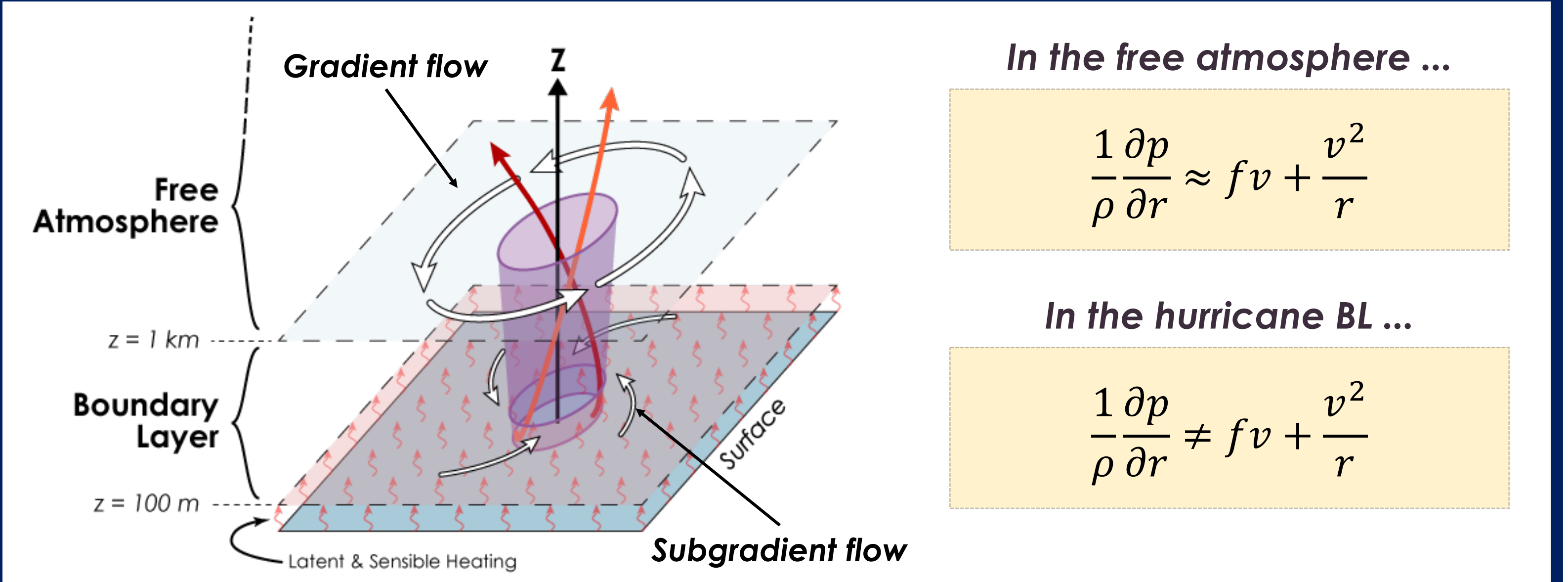
KYLE AHERN (KYLE.K.AHERN@GMAIL.COM), MARK A. BOURASSA, AND ROBERT E. HART
FLORIDA STATE UNIVERSITY; DEPARTMENT OF EARTH, OCEAN, AND ATMOSPHERIC SCIENCE



Introduction

The hurricane boundary layer (BL), constituting the lowest 1–2 km of the atmosphere, represents a key component of a hurricane's "heat engine." Friction promotes inflow in the hurricane BL^[1], which draws heat and moisture sourced from the warm ocean toward the center of a hurricane^[2].

Entropy-rich air in the BL can **converge and ascend** in inner-core convection.



At least to this effect, **the BL plays a key role in storm intensity.**

Observed **azimuthal-mean BL inflow** has been shown to **differ** between hurricanes that intensify and hurricanes that do not intensify^[3], implying differences in **BL convergence** and possibly **convection**.

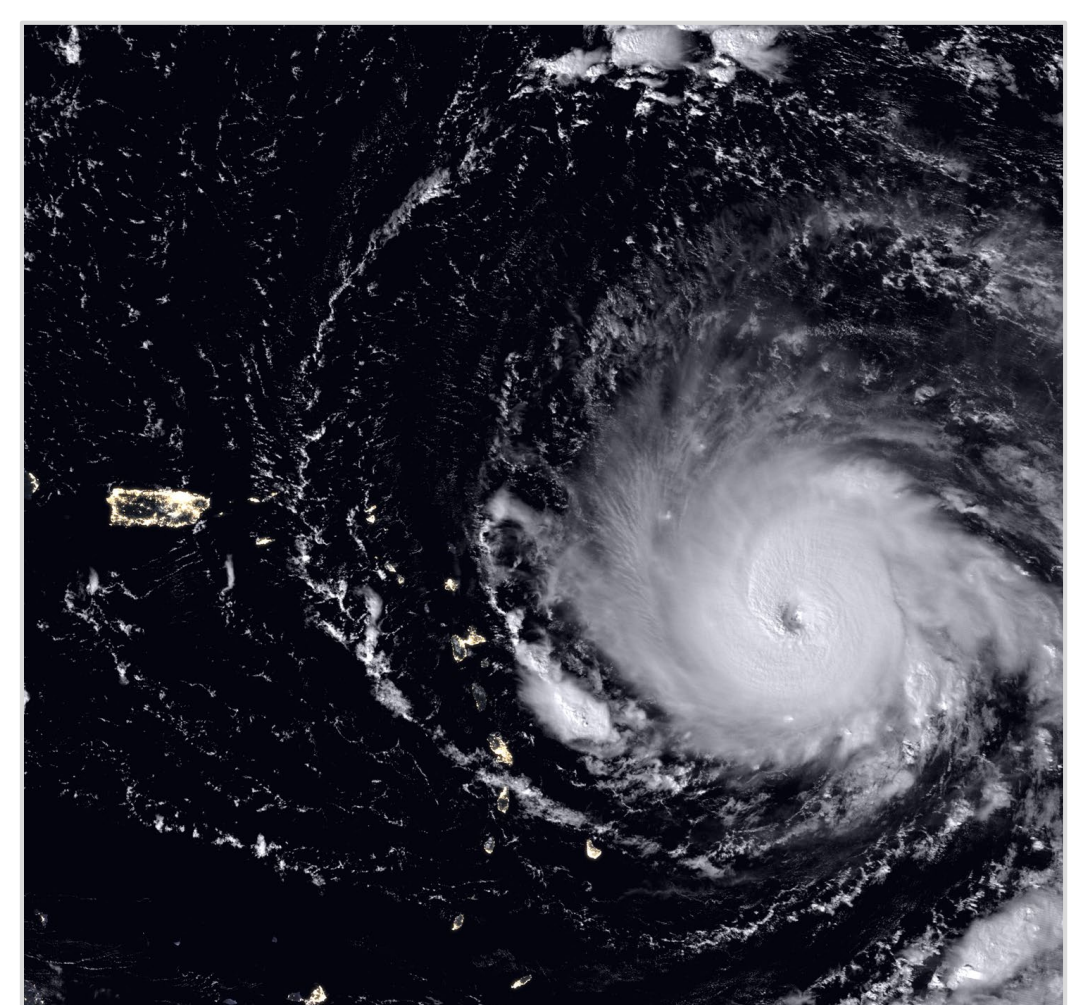
Radial inflow can also be **asymmetric**, with inflow depth and/or magnitude **enhanced downstream** of storm motion and **vertical wind shear**^[4,5].

How does **azimuthal** BL structure differ between hurricanes that intensify and those that do not? How do these differences compare from **case to case**?

Case Subjects

High-resolution WRF-ARW^[6] simulations were run for **Hurricanes Irma (2017) and Earl (2010)**. These cases were chosen for their differing environmental shear and comparable intensities.

Hurricane Irma
Weak/moderate shear



Hurricane Earl
Moderate/strong shear

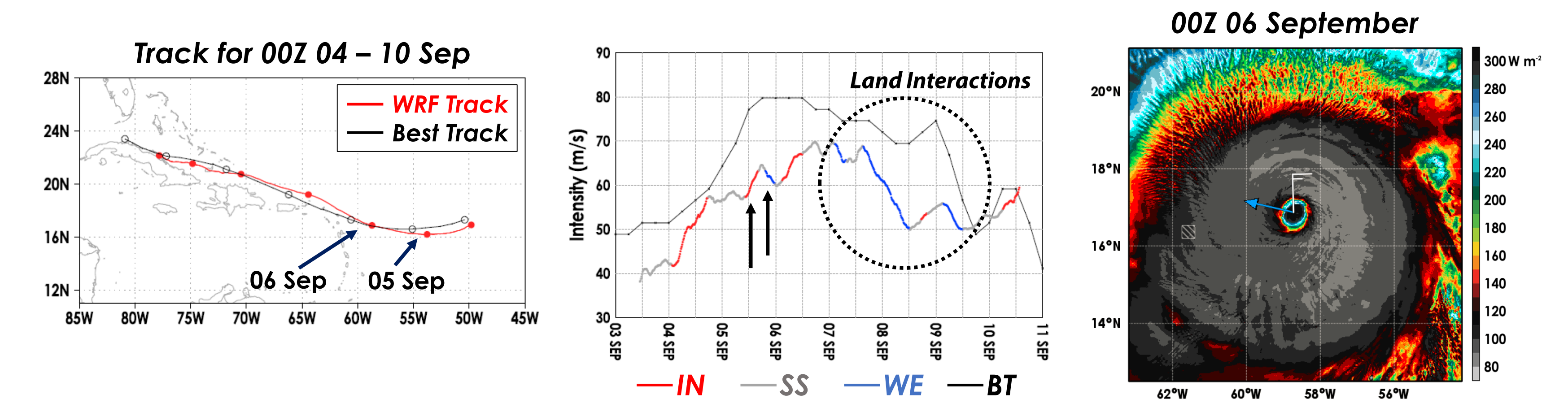


Satellite imagery taken from <http://earthobservatory.nasa.gov>

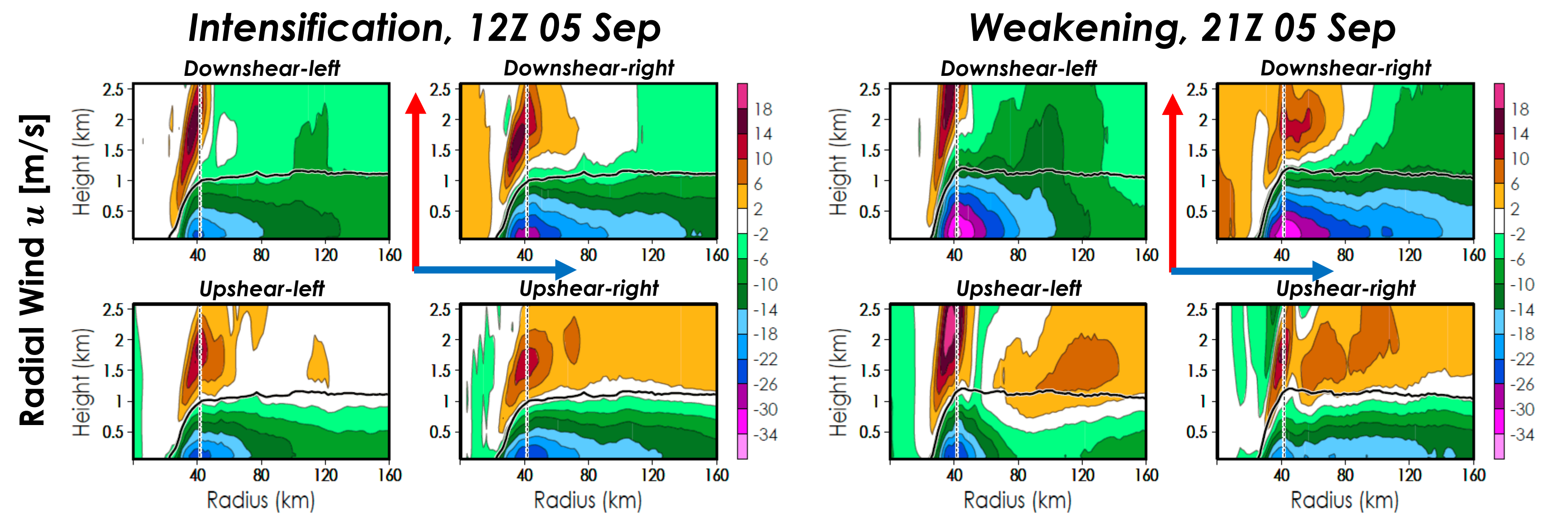
Hurricane Irma

For this case, we focus on two periods on 05 September: one intensifying, and one weakening. Irma tracked WNW at ~11 kt, while northerly shear increased from ~15 kt to ~20 kt just prior to weakening. Shear decreased to ~15 kt near the end of the weakening period.

Irma's track and intensity are depicted below, with relevant points of time highlighted by arrows. On the right, we show a snapshot of Irma's simulated IR, with motion and shear drawn as a vector and a wind barb, respectively.

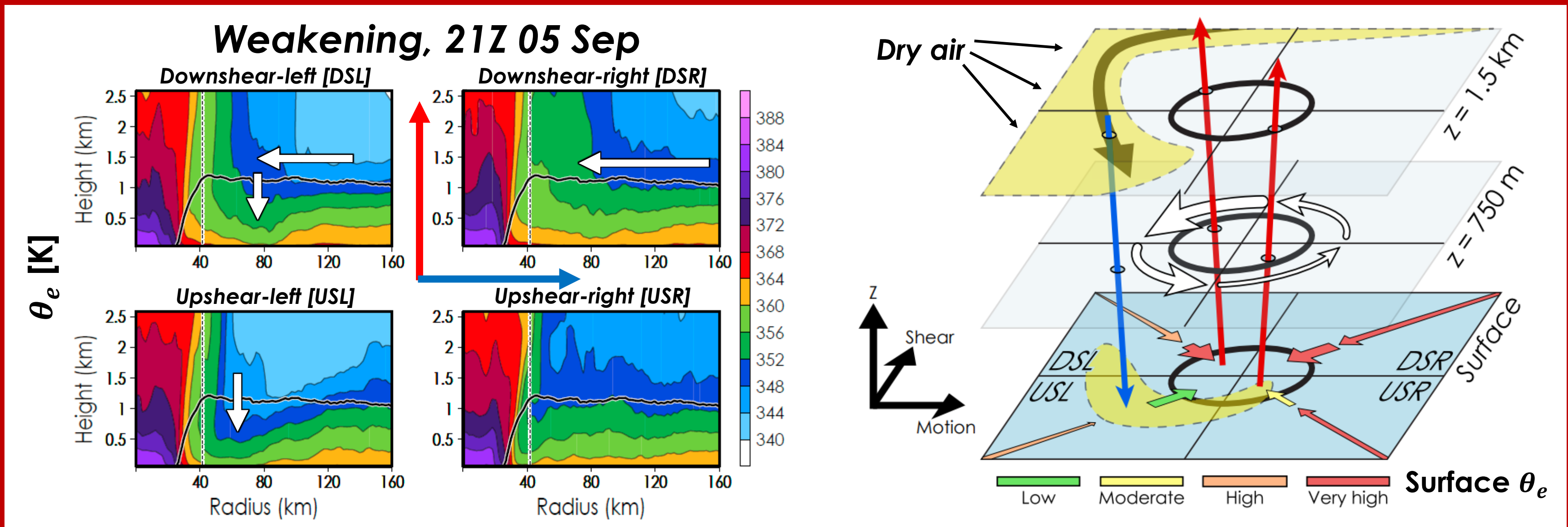


Below are radius-height cross sections of radial wind from each shear-relative quadrant during intensification (left) and weakening (right) on 05 September. The solid black line is the azimuth-mean inflow layer top, and the dotted line is the RMW. Red and blue vectors respectively denote the shear and storm motion directions.



Asymmetry is amplified during weakening, with **ample BL divergence upshear-left**.

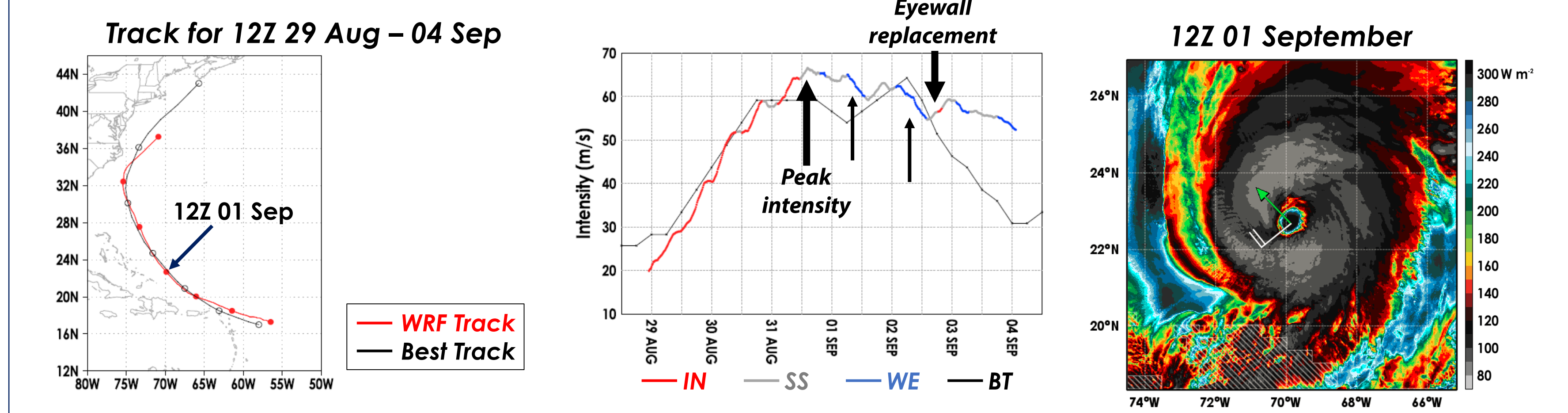
During weakening, deep inflow downshear transports drier air above the BL inward. The air then descends into the BL left of shear, in coincidence with BL divergence there.



As shown in the schematic on the right, the dry air that descends into the BL left of shear is then ingested by convection downwind, likely reducing Irma's intensity.

Hurricane Earl

In this case, we highlight a period of prolonged weakening following peak intensity, which culminated in an eyewall replacement.



Following peak intensity, shear increased from ~15 kt to ~20 kt, and then became nearly orthogonal to the storm motion. Earl's BL exhibited strong asymmetry after 01 September.

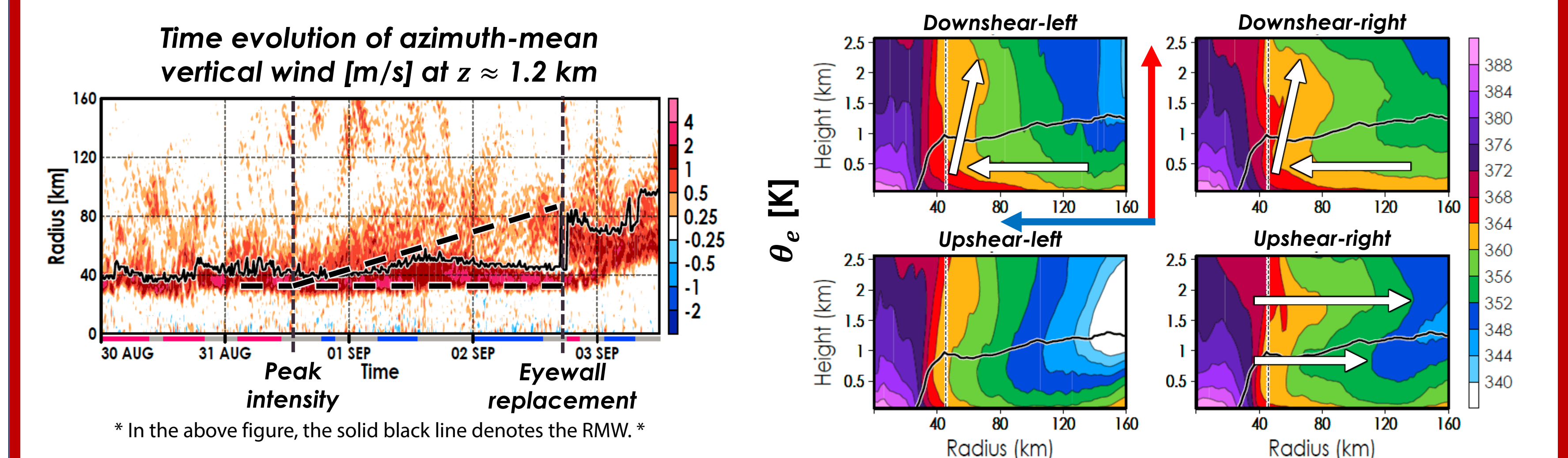
Unlike Irma, **Earl develops strong outflow in the USR quadrant.**

The outflow is linked to the intense inflow left of shear, which positively advects angular momentum inward ...

$$M_a = rv + \frac{1}{2} f_0 r^2$$

... and results in a local spin-up of v . This leads to strongly supergradient flow in the USL quadrant.

The supergradient flow is associated with an outward-directed agradient force. The USR outflow broadens the v -field, thereby increasing the inertial stability I^2 of the vortex outside of the RMW. The USR outflow also advects high-entropy air outward, which may recirculate via replenished BL inflow downshear.



Earl's weakening is coincident with the emergence of an outer region of BL ascent. The outwardly exhausted high- θ_e air may recirculate into the outer region of ascent, possibly helping to form the roots of convection outside the RMW.

KA thanks his advisors, Mark Bourassa and Bob Hart, for their guidance and support. KA also thanks the scientists at HRD/AOIML for their recommendations and advice on this modeling work. This research is supported financially by NASA / JPL (#1419699) and NOAA (#NA11OAR4320199).

References:
 [1] Smith, R. K., M. T. Montgomery, and N. Van Sang, 2009: Tropical cyclone spin-up revisited. *Quart. J. Roy. Meteor. Soc.*, **135** (642), 1321–1335.
 [2] Emanuel, K. A., 1986: An air-sea interaction theory for tropical cyclones. Part I: Steady-state maintenance. *J. Atmos. Sci.*, **43** (6), 585–605.
 [3] Ahern, K. M., A. Bourassa, R. E. Hart, J. A. Zhang, and R. F. Rogers, 2019: Observed kinematic and thermodynamic structure in the hurricane boundary layer during intensity change. *Mon. Wea. Rev.*, **147** (8), 2765–2785.
 [4] Zhang, J. A., R. F. Rogers, P. D. Reasor, E. W. Uhlhorn, and F. D. J. Marks, 2013: Asymmetric hurricane boundary layer structure from drizzle composites in relation to the environmental vertical wind shear. *Mon. Wea. Rev.*, **141** (11), 3968–3984.
 [5] Barnes, G. M., and K. P. Dolling, 2013: The inflow to Tropical Cyclone Humberto (2001) as viewed with azimuth-height surfaces over three days. *Mon. Wea. Rev.*, **141** (4), 1324–1336.
 [6] Skamarock, W. C., and Coauthors, 2008: A description of the Advanced Research WRF version 3. NCAR Tech. Note NCAR/TN-475+STR, 113 pp.