



89 GHz
166 GHz

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Introduction

Many satellite instruments exploit polarization to observe properties of clouds, precipitation, aerosols and different surface types. For example, observations at millimeter-wave wavelengths have detected frequent polarization signatures generated by large horizontally-oriented non-spherical ice particles in precipitating systems from the tropics to the high latitudes.^{1,2} Using these observations effectively in NWP data assimilation requires an accurate polarized radiative transfer (RT) model. This can be largely achieved by assuming non-spherical particles instead of soft spheres for scattering calculations and matching the particle size distributions (PSDs) assumed in computing the bulk scattering properties with those used in the NWP models.^{3,4}

Our goal is to develop a new RT model for computing polarimetric microwave radiances for use in the JCSDA Community Radiative Transfer Model (CRTM). We will examine the polarization signatures of snow for the GPM Microwave Imager (GMI) 89 and 166 GHz channels for simulated clouds using state-of-the-art single scattering properties of non-spherical particles while ensuring that the PSD used in computing the bulk scattering properties is consistent with the cloud simulation.

Methods

Cloud simulation

10²

10¹

100

0

- Large-scale Weather Research & Forecasting (WRF) model v2.2 simulation for 4 June 2005 (6 km horizontal grid spacing).
- Thompson microphysics (cloud liquid, cloud ice, rain, snow, graupel)⁵; snow PSD is a combination of exponential and gamma distributions (see below); snow habit assumed to be primarily fractal-like aggregates of crystals.



 $T = -4.2^{\circ} C$

Diameter (um)

4000

IWC = 0.528 a/m

6000

8000

10000

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2000

Bulk scattering properties of snow

- Ice particle scattering database for frequencies 1-874 GHz, 12 ice habits and maximum particle dimensions from 2 mm to 10 mm.⁶ Randomly oriented particles are assumed. Calculations are based on Invariant Imbedding T-Matrix method and the Improved Geometric Optics Method; includes temperature dependence of ice refractive index.
- Snow particles were represented as an aggregate of 10 plates (see below).
- Bulk properties (extinction coefficient σ_{ext}, scattering coefficient σ_{sca}, singlescattering albedo ω_0 , phase matrix **P**(Θ)) were integrated over the snow PSD as follows:



Phase Matrix Elements for all Snow Layers



Comparison of current model with a

4.66

4.71

215.43 0.28

2-layer atmosphere test case.9

0.09501 111.89 0.68

0.28160 154 71 2.81

0.61788 200.67 5.44 199.84

0.75540 208.90

0.86563 212.88 3.08 212.35

0.94458 214.70 1.41

0.98940

0.45802 184.41

polarized doubling-adding model for a

112.79

154 54

183.76

208.16

214.38

215.26 0.30

2.14

2 29

4.86

6.21

5.70

3.94

1.79

Polarized microwave model

- Extension of the Successive Order of Interaction (SOI) model⁷, developed at UW and one of the multiple scattering solvers in the CRTM.
- Assumes Fresnel surface reflectance
- . For randomly oriented particles of arbitrary shape; azimuthal mean phase matrix is computed numerically.8
- Only I and Q Stokes parameters are considered since U=V=0 for an azimuthally-averaged radiance field

Results

270

265

260

25



Snow profile

- At 89 GHz, Q almost entirely results from depolarization of the highly polarized ocean surface by atmospheric absorption/emission
- At 166 GHz, a polarization signature (ΔQ) of almost 1 K is generated by the snow particles at the GMI Earth-incidence (zenith) angle of 52.8°. However, the maximum signature of 1.9 K occurs at a zenith angle of about 32°.

Conclusions

- A new microwave RT model has been developed to study the polarization characteristics of randomly-oriented non-spherical snow particles
- Early results show these particles can generate significant polarization signatures.
- · Further tests of the model are needed as well as a reformulation for consistency with the Advanced Matrix Operator Method¹⁰ in CRTM alpha v3.0.

References

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