

# Understanding of the roles of global warming and natural variability on monsoon rainfall

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## Extended Abstract

A number of studies have investigated the mechanisms that determine changes in precipitation, including how a wet region gets wetter. However, not all monsoon areas get wetter – there is a need to understand the major factors behind changes in regional monsoon precipitation, in terms of external forcing and internal variabilities in the last six decades by a combination of different observed datasets and model runs. We have found that emerging times of anthropogenic signals are detected in the northern African and Southeast Asian monsoons from the 1990s with the use of CESM Large Ensemble Project. From CMIP5 model runs and three reanalysis datasets, the results are found that the change in monsoon rainfall over African is mainly due to anthropogenic forcing and that over Asian is affected by internal variability. Moreover, the cause of American monsoon rainfall change cannot be known due to a discrepancy among observed datasets.

Monsoon precipitation is the largest component of the global water and energy cycles and a major driver of general atmospheric circulation [C P Chang and World Meteorological Organization., 2011; C P Chang et al., 2018]. The increasing intensity and frequency of heavy rainfall events, heat waves, and severe droughts over the globe have affected the lives of people around the globe and caused huge loss of life, material damages, and associated social costs [Min et al., 2011; Qiu, 2013]. As our climate rapidly changes, we need to improve our understanding of the effect of climate change on monsoons, including on their changes in rainfall and circulation over regionally and geologically different monsoons.

Climate modeling studies [Chou and Neelin, 2004; Chou et al., 2013; Min et al., 2011] predict changes in global precipitation in response to increases in the mean global surface temperature of the Earth. A prevailing theory of precipitation change in a warming world is that wet areas will get wetter [Chou et al., 2009; Held and Soden, 2006]. Under the global warming scenario in phase five of the Coupled Model Intercomparison Project (CMIP5), the Northern Hemisphere (NH) monsoon precipitation will be increased by the increased temperature difference between the NH and Southern Hemisphere (SH) and the enhanced atmospheric moistening [Lee and Wang, 2014]. However, on a basin scale, this is not always the case [Hsu and Li, 2012]. In model simulations, a swing in the South Pacific convergence zone due to global warming (GW) has been predicted that might cause catastrophic droughts and food shortages in the region [Cai et al., 2012], and the increasing global surface temperature induces a decrease in soil moisture due to increased evaporation [Dai, 2013].

As well as GW, natural variability, especially El Niño-Southern Oscillation (ENSO) effect, is also mentioned as the cause of drought [Trenberth et al., 2014]. On the other hand, past variability in rainfall

40 in the Sahel which depends strongly on the West African monsoon is closely linked to the global sea  
41 surface temperature (SST) distributions in both observations and models [Cook, 2008]. These differences  
42 are likely manifestations of the different roles played by natural variability and GW in different regions  
43 [Kim and Ha, 2015 and 2018; Chung et al., 2019]. Multidecadal Pacific SST anomalies have been shown  
44 to contribute significantly to severe droughts over both the western U.S. and the Great Plains [Meehl and  
45 Hu, 2006]. In addition to the influence of the north-south hemispheric thermal contrast in the Atlantic-  
46 Indian Ocean, an east-west thermal contrast in the Pacific Ocean, which is measured by the Interdecadal  
47 Pacific Oscillation (IPO) index, is also predicted to increase the decadal variability of NH land monsoon  
48 rainfall [Wang et al., 2018].

49 However, the key factors of monsoon rainfall change are not clearly defined due to the uncertainty of  
50 climate models and the inconsistency of observed datasets despite many studies on climate change. The  
51 urgency of this question is further complicated by the facts that the observed trends lie outside the range  
52 of natural variability simulated in climate models [Hoerling et al., 2010], but at the same time oppose the  
53 simulated response to GW. This conundrum requires an in-depth re-evaluation of the changes in global  
54 and regional monsoons and causes and drivers of recent global and regional monsoons. The purpose of  
55 this study is, therefore, twofold: 1) to identify long-term climatic change by analyzing robust changes in  
56 precipitation over monsoons from 1958 to 2015 using various reanalysis datasets based on observations,  
57 specifically focusing on different monsoon regions; and 2) impacts of external forcings in driving these  
58 precipitation changes over the individual monsoon region using model simulations from Community  
59 Earth System Model (CESM) and CMIP5.

60 We attempted to draw what drives the linear trend of global monsoon rainfall. According to Figure 1  
61 and 2, which are representing the linear trends of reanalysis rainfall and CMIP5 model rainfall, the results  
62 demonstrate many important points as follows: Major factors of regional monsoon rainfall changes are  
63 investigated by combinations of multi-ensemble, external forcing runs, and observed datasets. The effects  
64 of external forcings on the rainfall are detected from the 1990s over northern Africa and Southeast Asia  
65 and spread to other regions. The African and Asian monsoon rainfall changes, respectively, are caused by  
66 the anthropogenic forcing and internal variability

67 The present study intends to investigate the reliability of available precipitation data and the trends in  
68 rainfall over regional monsoons in order to project the future of monsoons. The CESM does not  
69 accurately simulate the changes seen in global rainfall during the past 60 years (1958-2015) and there is a  
70 significant need to improve our understanding of such changes and the driving factors behind them. By  
71 studying the time of emergence of the effects of human activities on the global rainfall change using the  
72 CESM simulation, we find that changes in precipitation by the anthropogenic forcing is detected over the  
73 northern African and Southeast Asian monsoons from the 1990s. The GHG forcing drives a meridional  
74 dipole in trends over the western Pacific, while natural variabilities such as IPO and NAID induce a  
75 horizontal dipole over the western-central Pacific. This study also suggests that the AMM has a broader  
76 gap in observed land rainfall dataset compared to other monsoons, a distinctive decreasing trend in only

77 land rainfall can be found in the AFM, and an increasing trend is seen in the southern AAM(i.e.  
78 Australian monsoon).

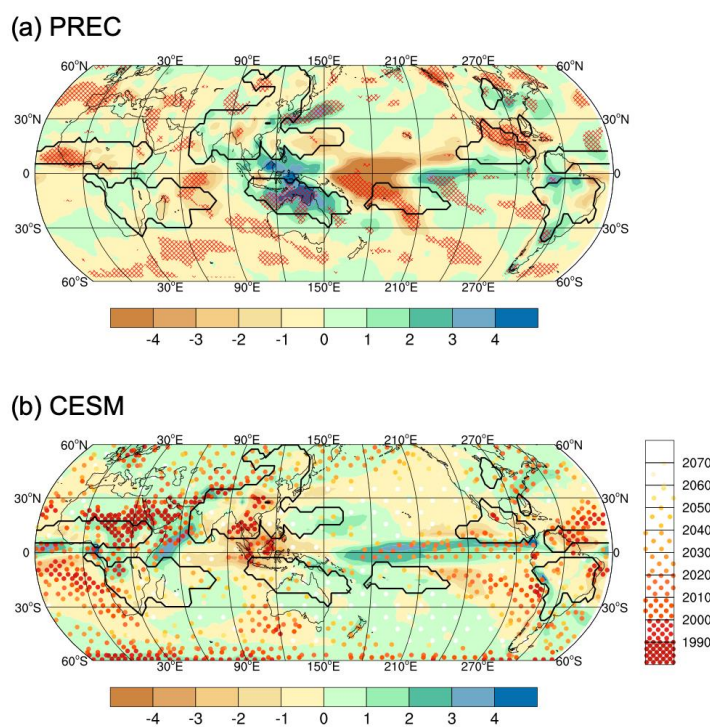
79 In terms of regional monsoon land rainfall changes, the main forcing is different depending on the  
80 region. The decreasing trend in rainfall over the AFM can be seen to come from aerosol and GHG  
81 forcings in northern and southern parts, respectively. The change in Asian monsoon rainfall may be  
82 primarily due to the internal variability rather than external forcings, while the Australian monsoon  
83 rainfall has the potential to be increased by the aerosol forcing. The cause of an increasing trend of the  
84 AMM, especially in the northern part, cannot be known through the response of monsoon rainfall by the  
85 external forcings because the GHG and anthropogenic forcings induce a decreasing rainfall trend. In  
86 addition, it is difficult to understand the observed change in land rainfall over the AMM due to a  
87 discrepancy between reanalysis datasets based on observations.

88 The discussion on quantification of external forcings or internal variability on the individual monsoon  
89 rainfall change still remains unexplored in this study. However, this study shows that a major factor of  
90 rainfall change over the individual monsoon region is distinguishable, which is provided through  
91 combinations of results from different observed datasets and model runs. These results are expected to  
92 provide an improvement in the predictability of individual monsoon rainfall in future climate change.

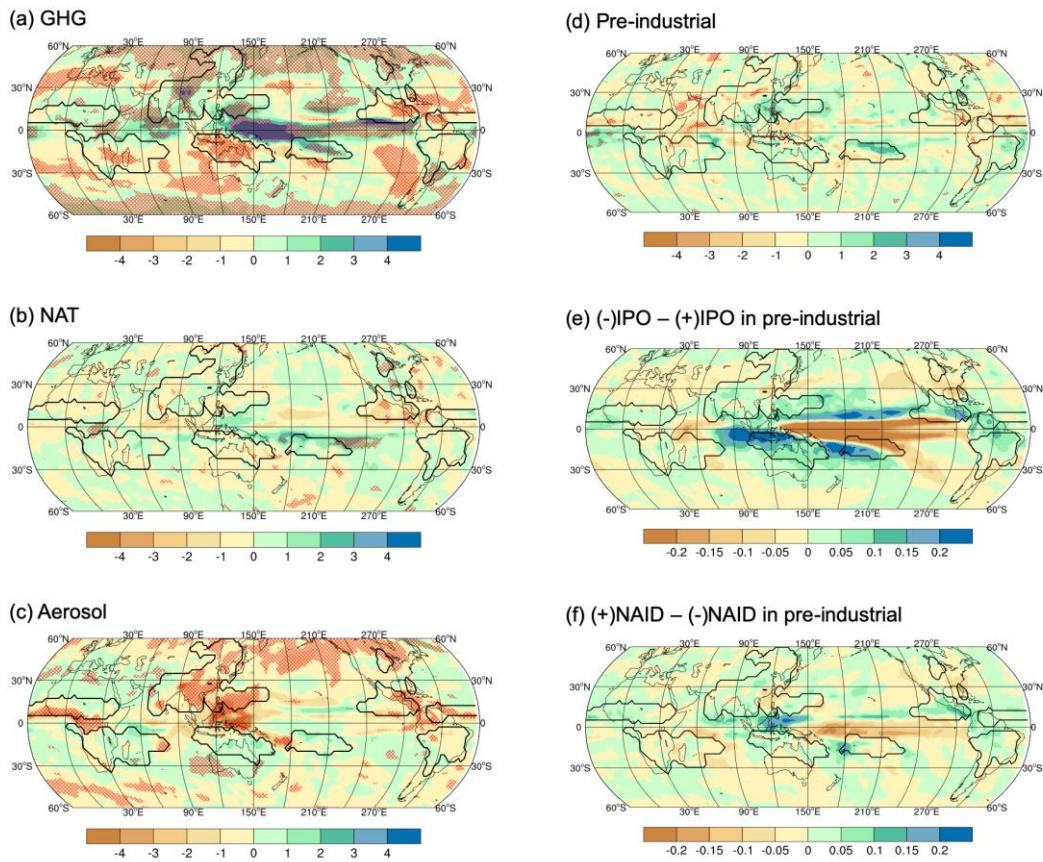
## 93 **References**

- 94  
95 Cai, W. J., et al. (2012), More extreme swings of the South Pacific convergence zone due to greenhouse  
96 warming, *Nature*, 488(7411), 365-+.
- 97 Chang, C. P., and World Meteorological Organization. (2011), The global monsoon system research and  
98 forecast, in *World Scientific series on Asia-Pacific weather and climate v 5*, edited, pp. xii, 594 p., World  
99 Scientific, Singapore ; Hackensack, NJ.
- 100 Chang, C. P., R. H. Johnson, K. J. Ha, D. Kim, G. N. C. Lau, B. Wang, M. M. Bell, and Y. L. Luo  
101 (2018), THE MULTISCALE GLOBAL MONSOON SYSTEM Research and Prediction Challenges in  
102 Weather and Climate, *B Am Meteorol Soc*, 99(9), Es149-Es153.
- 103 Chou, C., and J. D. Neelin (2004), Mechanisms of global warming impacts on regional tropical  
104 precipitation, *J Climate*, 17(13), 2688-2701.
- 105 Chou, C., J. D. Neelin, C. A. Chen, and J. Y. Tu (2009), Evaluating the "Rich-Get-Richer" Mechanism in  
106 Tropical Precipitation Change under Global Warming, *J Climate*, 22(8), 1982-2005.
- 107 Chou, C., J. C. H. Chiang, C. W. Lan, C. H. Chung, Y. C. Liao, and C. J. Lee (2013), Increase in the  
108 range between wet and dry season precipitation, *Nat Geosci*, 6(4), 263-267.
- 109 Chung, E. S., A. Timmermann, B. J. Soden, K. J. Ha, L. Shi, and V. O. John (2019), Reconciling  
110 opposing Walker circulation trends in observations and model projections, *Nat Clim Change*, 9, 405-412.
- 111 Cook, K. H. (2008), CLIMATE SCIENCE The mysteries of Sahel droughts, *Nat Geosci*, 1(10), 647-648.
- 112 Dai, A. G. (2013), Increasing drought under global warming in observations and models, *Nat Clim*  
113 *Change*, 3(1), 52-58.
- 114 Held, I. M., and B. J. Soden (2006), Robust responses of the hydrological cycle to global warming, *J*  
115 *Climate*, 19(21), 5686-5699.
- 116 Hoerling, M., J. Eischeid, and J. Perlwitz (2010), Regional Precipitation Trends: Distinguishing Natural  
117 Variability from Anthropogenic Forcing, *J Climate*, 23(8), 2131-2145.
- 118 Hsu, P. C., and T. Li (2012), Is "rich-get-richer" valid for Indian Ocean and Atlantic ITCZ?, *Geophys Res*  
119 *Lett*, 39.
- 120 Kim, B. H., and K. J. Ha (2015), Observed changes of global and western Pacific precipitation associated  
121 with global warming SST mode and mega-ENSO SST mode, *Clim Dynam*, 45(11-12), 3067-3075.
- 122 Kim, B. H., and K. J. Ha (2018), Changes in equatorial zonal circulations and precipitation in the context  
123 of the global warming and natural modes, *Clim Dynam*, 51(11-12), 3999-4013.

124 Lee, J. Y., and B. Wang (2014), Future change of global monsoon in the CMIP5, *Clim Dynam*, 42(1-2),  
 125 101-119.  
 126 Meehl, G. A., and A. X. Hu (2006), Megadroughts in the Indian monsoon region and southwest North  
 127 America and a mechanism for associated multidecadal Pacific sea surface temperature anomalies, *J*  
 128 *Climate*, 19(9), 1605-1623.  
 129 Min, S. K., X. B. Zhang, F. W. Zwiers, and G. C. Hegerl (2011), Human contribution to more-intense  
 130 precipitation extremes, *Nature*, 470(7334), 378-381.  
 131 Qiu, J. (2013), CLIMATOLOGY Monsoon Melee, *Science*, 340(6139), 1400-1401.  
 132 Schneider, U., A. Becker, P. Finger, A. Meyer-Christoffer, M. Ziese, and B. Rudolf (2014), GPCP's new  
 133 land surface precipitation climatology based on quality-controlled in situ data and its role in quantifying  
 134 the global water cycle, *Theor Appl Climatol*, 115(1-2), 15-40.  
 135 Trenberth, K. E., A. G. Dai, G. van der Schrier, P. D. Jones, J. Barichivich, K. R. Briffa, and J. Sheffield  
 136 (2014), Global warming and changes in drought, *Nat Clim Change*, 4(1), 17-22.  
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 139 **Figure 1.** Linear trend of annual global (60°S-60°N, 0°-360°) rainfall. Plots are made using precipitation  
 140 data from (a) PREC dataset (1958-2015) and (b) the present run which is merged historical run (1958-  
 141 2005) and RCP8.5 run (2006-2015) of CESM Large Ensemble Project (1958-2015). Color shading shows  
 142 linear temporal trends (mm yr<sup>-1</sup>). In (a), the diagonal cross-hatching (red) denotes the statistically  
 143 significant region above the 95% confidence level, as determined using the Mann-Kendall test. The red  
 144 stippled areas in (b) denote the end year of a period starting in 1958 for which the linear temporal trend  
 145 predicts that the signal-to-noise (SNR) ratio exceeds 1 for the first time. The global monsoon domains  
 146 (black solid) are defined according to Wang and Ding (2006) using PREC data.



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**Figure 2.** Linear trend of annual global (60°S-60°N, 0°-360°) rainfall. Plots are made using precipitation data from (a) GHG run (1958-2005), (b) NAT run (1958-2005), (c) Aerosol run (only anthropogenic aerosol forcing, 1958-2005), and (d) pre-industrial run (1-59) from multi-model mean (MMM) in CMIP5. Composite difference of annual precipitation based on (e) -IPO minus +IPO, and (f) +NAID and -NAID in pre-industrial run of CMIP5. Color shading shows linear temporal trends (mm yr<sup>-1</sup>). In (a-d), the diagonal cross-hatching (red) denotes the statistically significant region above the 95% confidence level, as determined using the Mann-Kendall test. The global monsoon domains (black solid lines) are defined according to Wang and Ding (2006) using PREC data.