

Land-surface response to shallow cumulus



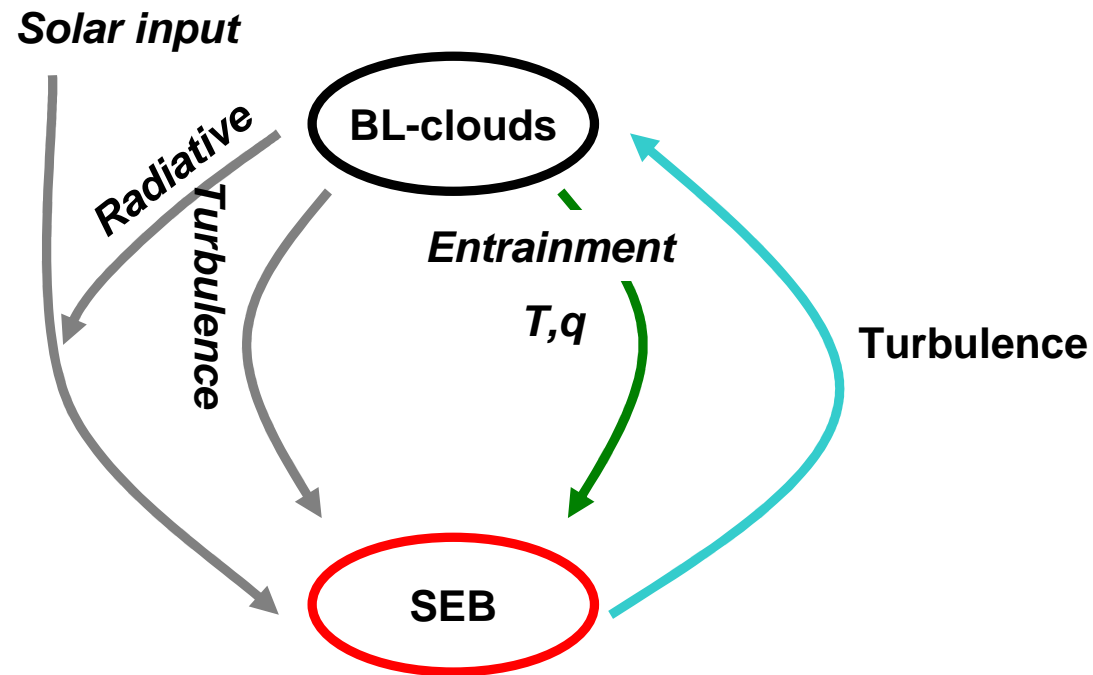
Fabienne Lohou and Edward Patton



Outline

1/ SEB response on average and locally

(Radiative and turbulent effects (not shown here))



2/ Boundary layer response to cloud-induced surface flux heterogeneity

3/ Entrainment associated to BL-clouds

NCAR LES code

Moeng, 1984, 1986 / Sullivan et al., 1996 / Patton et al., 2005

Cloud transmission coefficient
(Joseph et al., 1976)

Optical depth (Stephens, 1984)

Atmos. Forcings
Ch, Cm, θ , q, u, SWdn

NOAH Land surface model

1D set of equations for thermodynamic and hydrologic variables: Mahrt and Pan, 1984, 1987

Surface energy balance: Chang et al, 1999

Surface Forcings: H, LE, G, Tskin, SMCskin

Simulated case

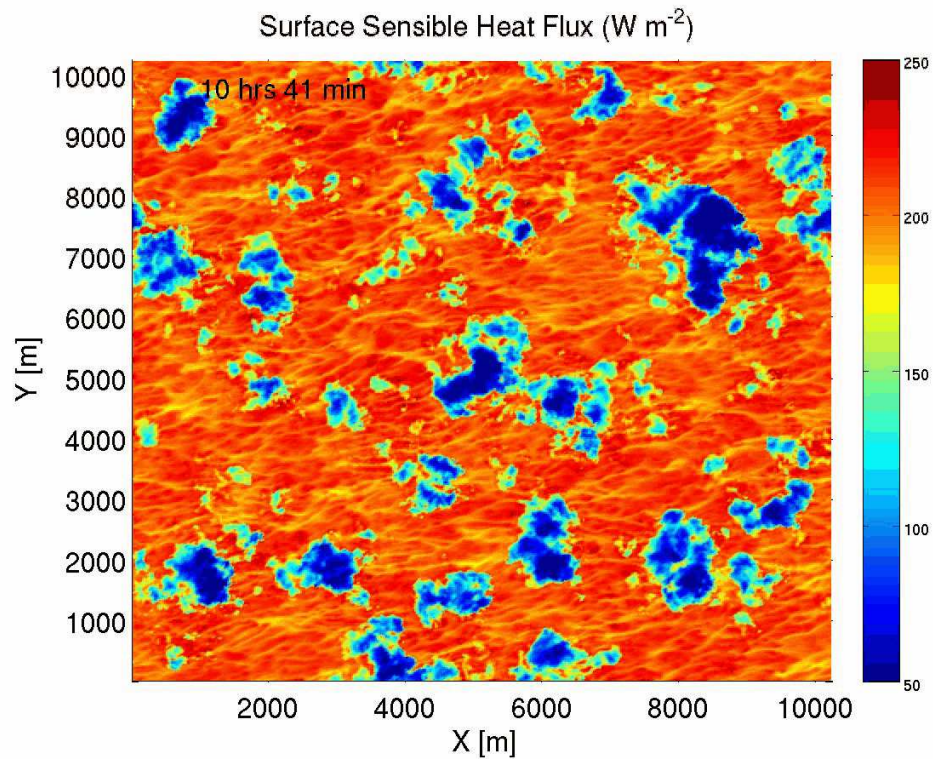
Case: 21 June 1997 ARM SGP: Golaz et al. 2001, Brown et al. 2002, Vilà-Guereau et al., 2005

Domain: 10 km² x 4 km

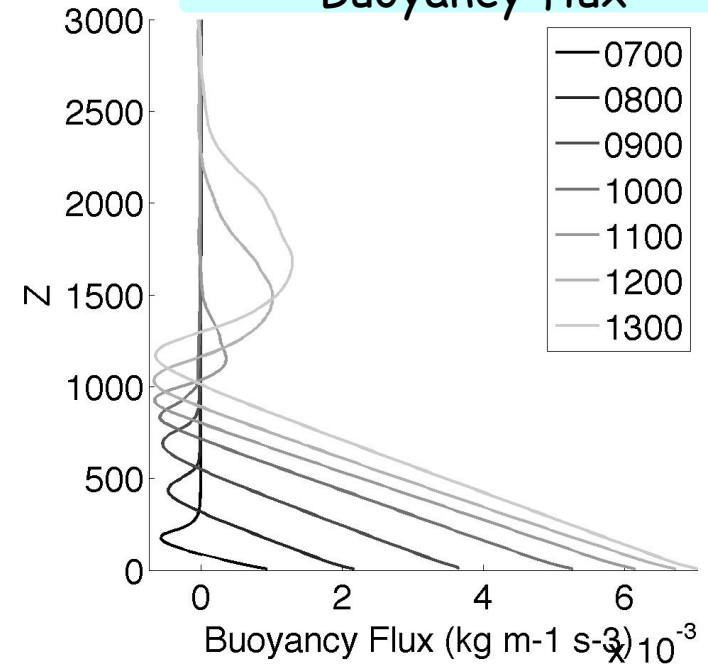
Resolution: $\Delta x = \Delta y = 20$ m and $\Delta z = 8$ m **4 levels in the ground:** 5, 20, 60, 100 cm

Soil type: bare / clay-loam (*wilting point* = 0.103 m³/m³ // *field capacity* = 0.465 m³/m³) soil moisture content: 0.44 m³/m³// soil moisture availability: 0.93

Surface sensible heat flux



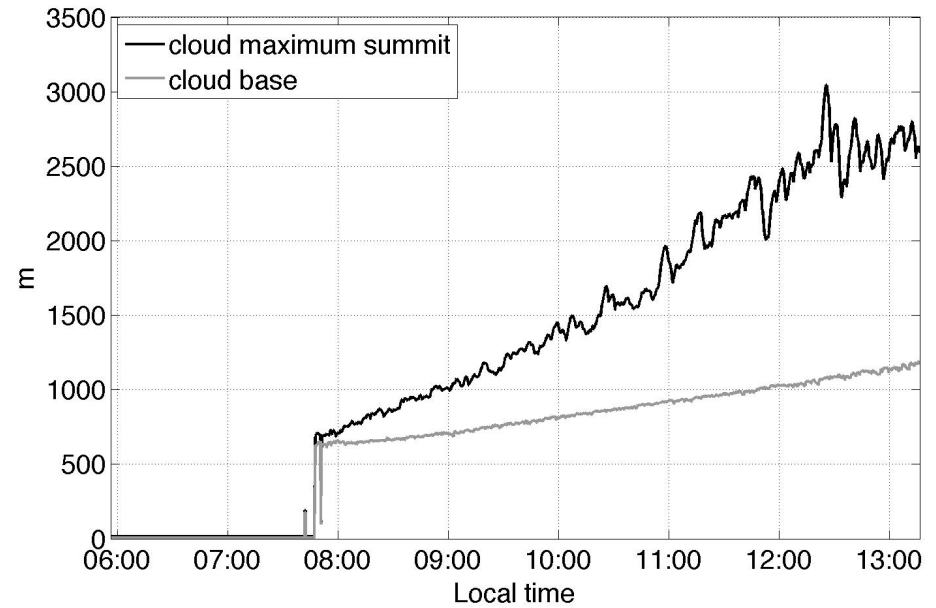
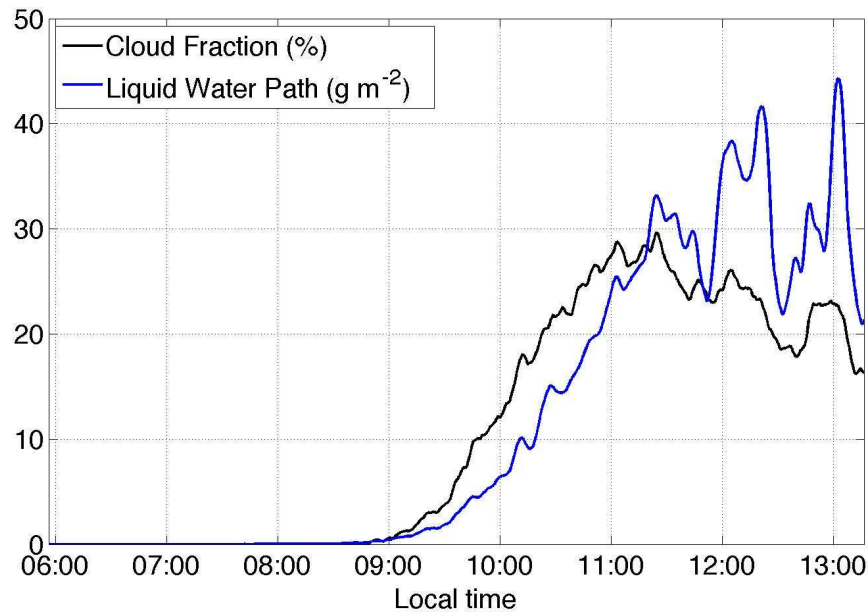
Buoyancy flux



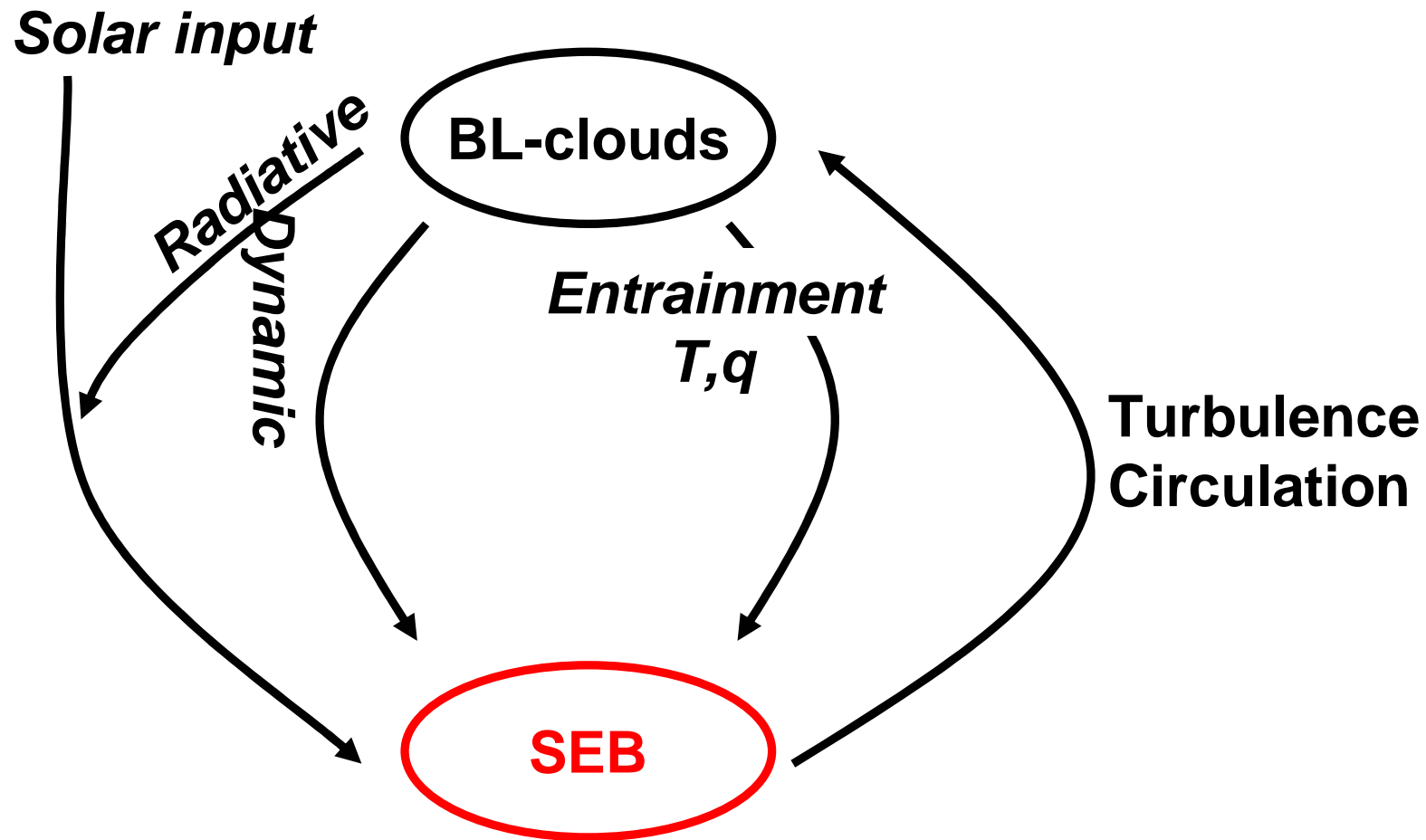
un-shaded area

shaded area

Cloud layer characteristics

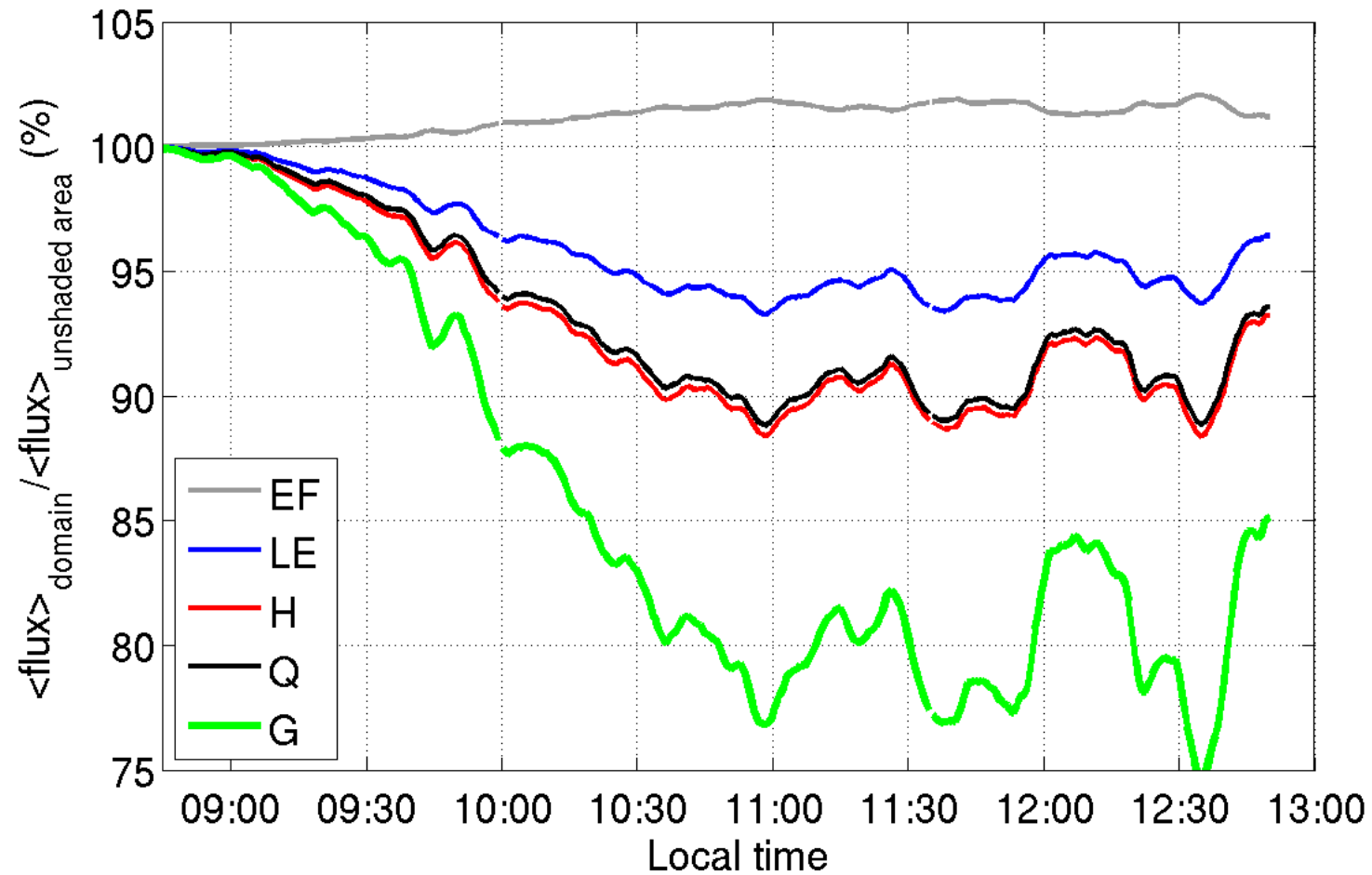


- ✓ Forced Cumuli: cloud top < level of free convection
(Stull, 1988 ; *Otles and Young, 1996*)
- ✓ Depth up to 1.4 km
- ✓ Diameter up to 1 km
- ✓ Cloud Fraction up to 30%

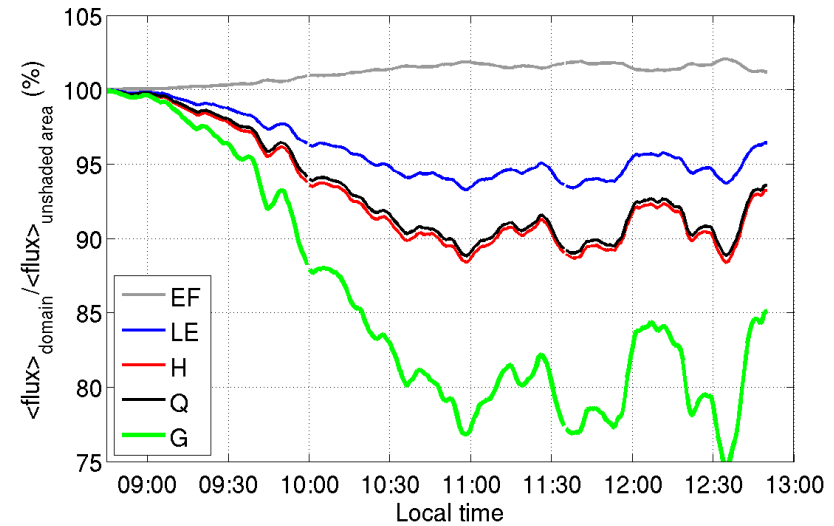


What is the SEB response to BL-Clouds on average and locally?

SEB response to BL-clouds on average



SEB response on average



SEB response locally



Shaded area

$$Q = H + LE + G \quad \text{EF} = 0.8$$

$$Q = 0.3 Q + \underline{1.2 Q} + (-0.5 Q)$$

Un-shaded area

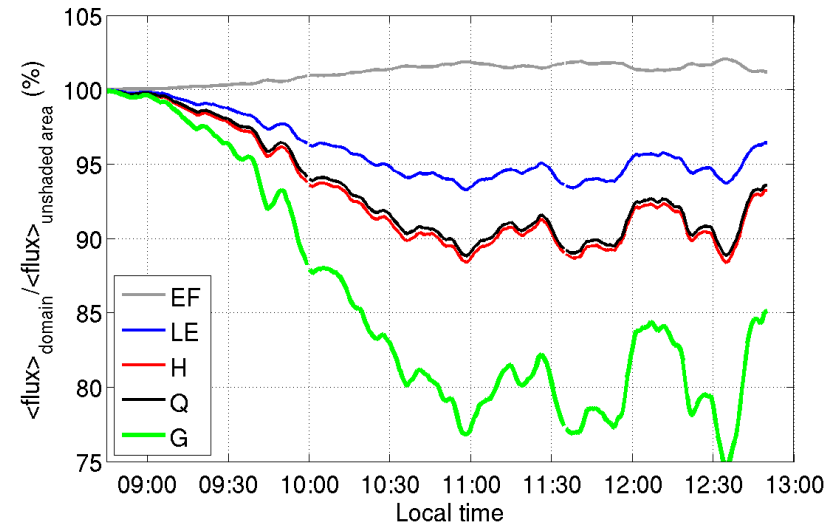
$$Q = H + LE + G \quad \text{EF} = 0.6$$

$$Q = 0.3 Q + \underline{0.5 Q} + 0.2 Q$$

SEB response locally



SEB response on average



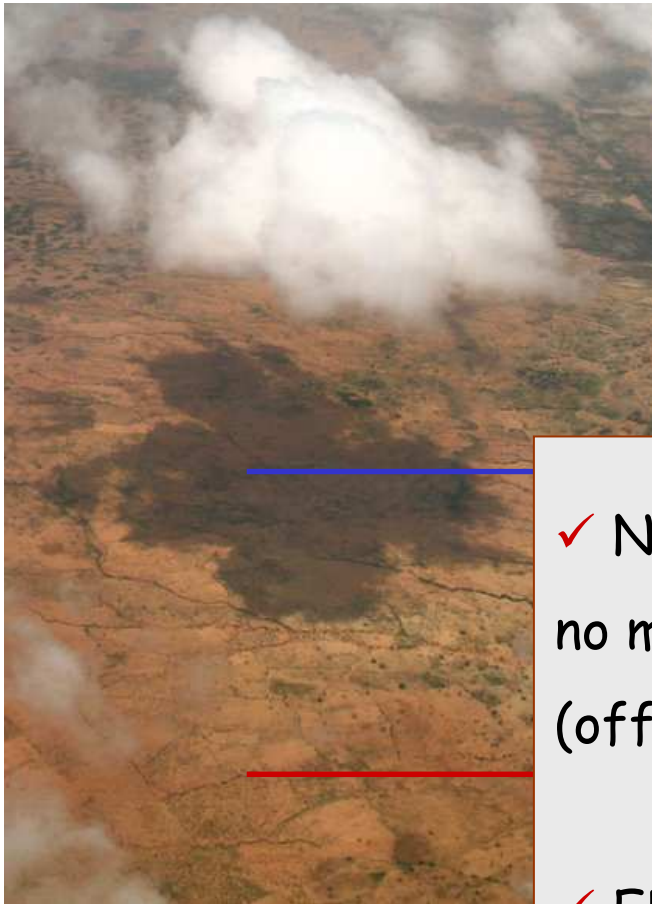
Penman equation

$$\lambda_v E = \frac{\frac{dq_{sat}}{dT_a} Q + \rho_a c_p \delta_e r_a^{-1}}{\frac{dq_{sat}}{dT_a} + \gamma}$$

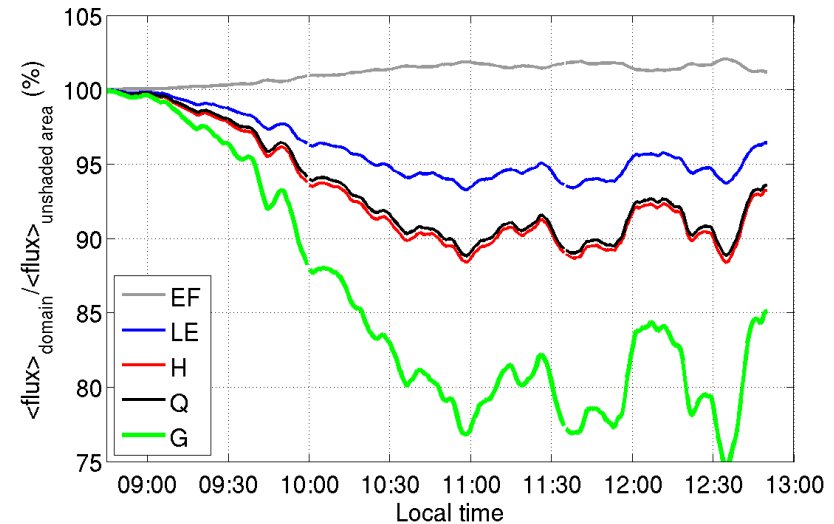
$$LE = SMI \times \lambda_v E$$

$$H = \frac{\rho_a c_p}{r_a} (\theta_s - \theta)$$

SEB response locally



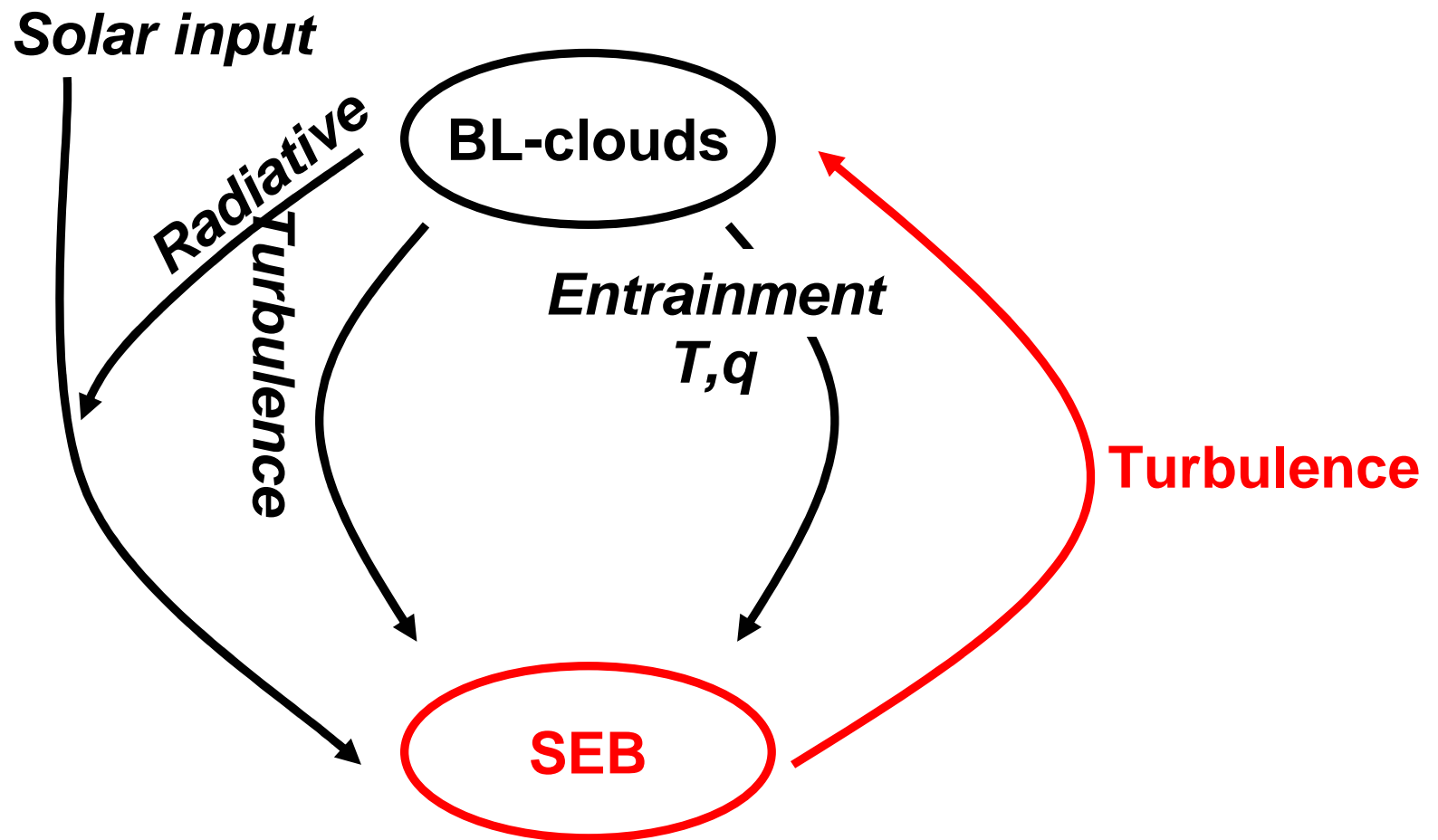
SEB response on average



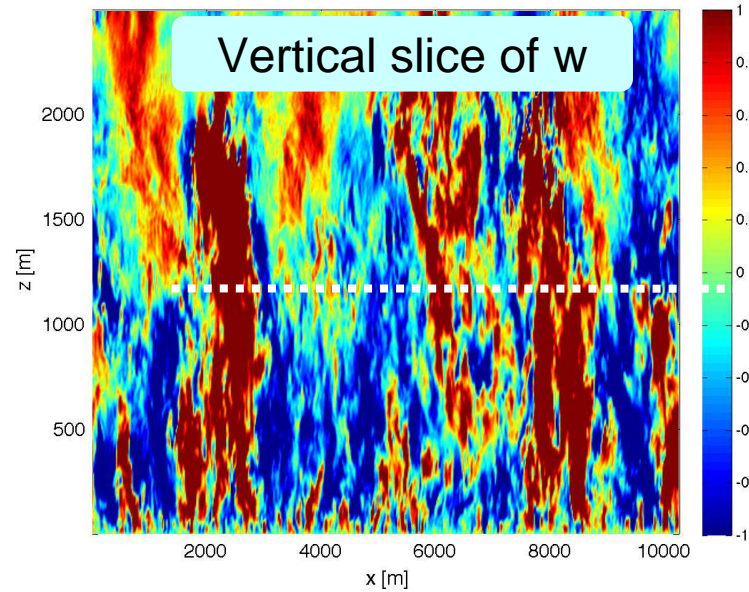
✓ Non linear response of the SEB to BL-clouds no matter the soil type and soil moisture.
(offline tests with 1D NOAH LSM (not shown))

✓ EF increases of 3%

Boundary layer response to cloud-induced surface flux heterogeneity

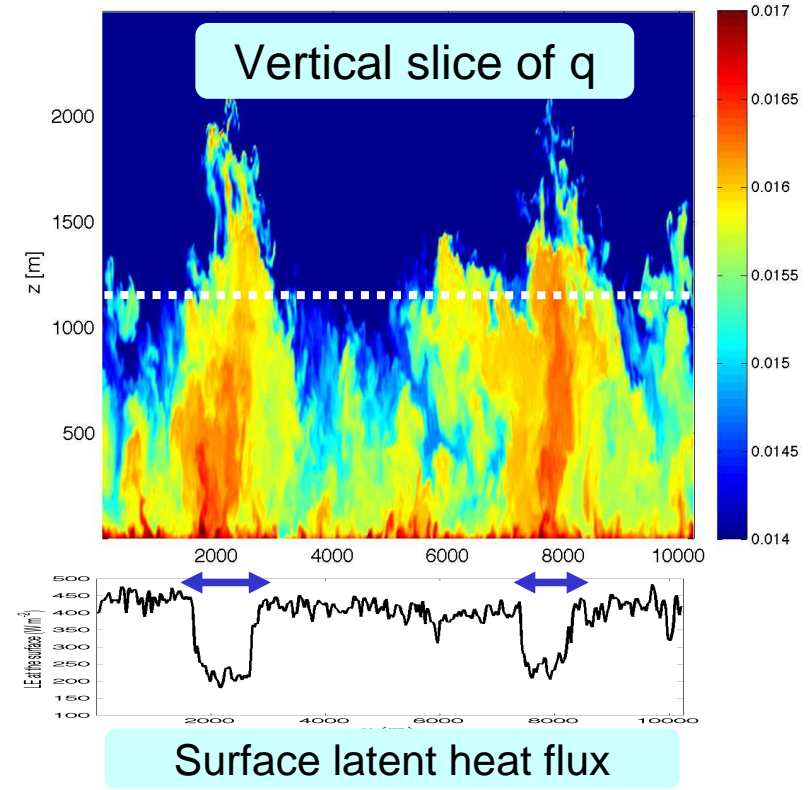
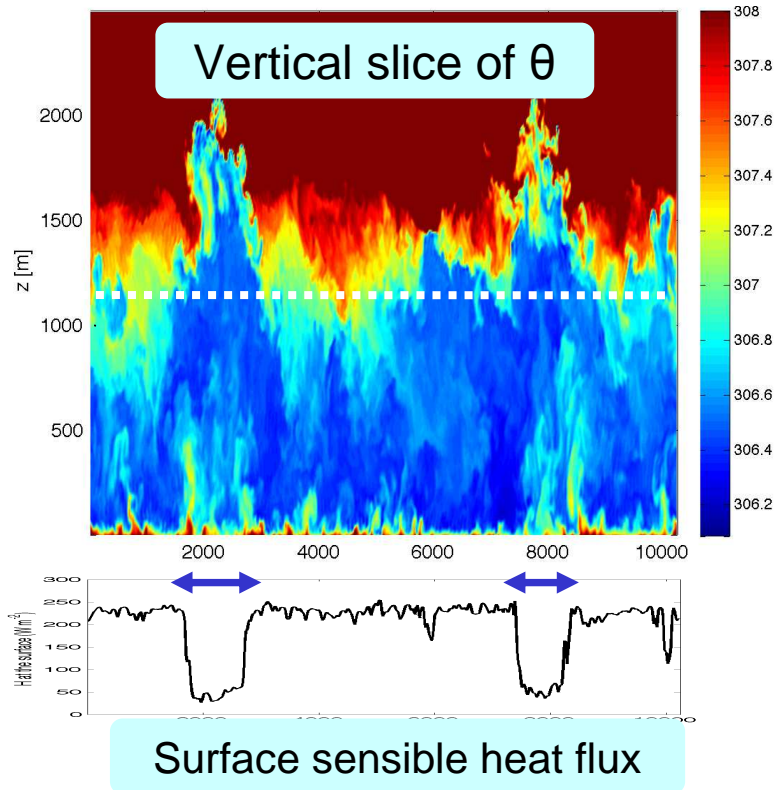


Cloud roots characteristics: mean parameters



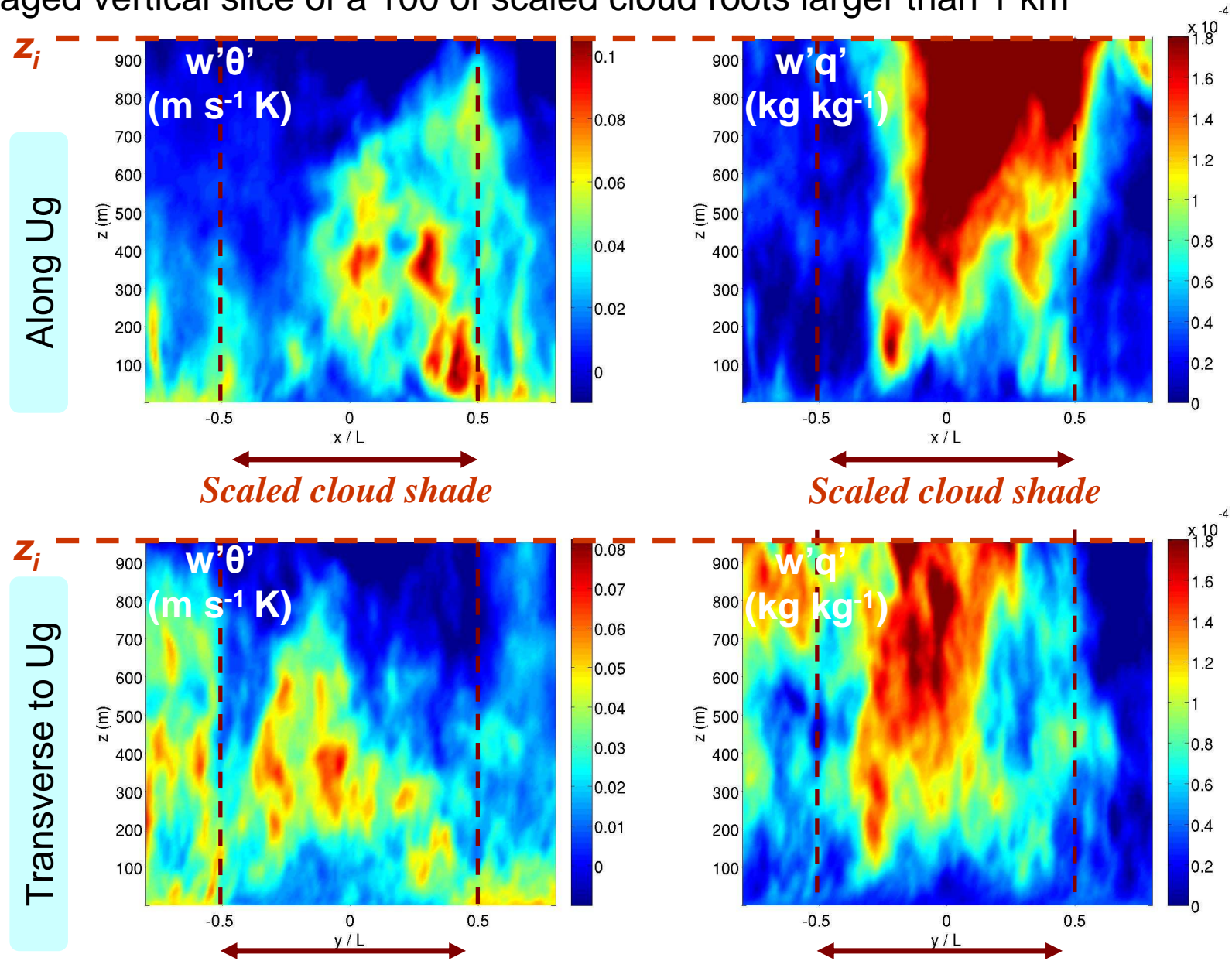
Example at 1300 LST

Cloud root \equiv column above shaded area



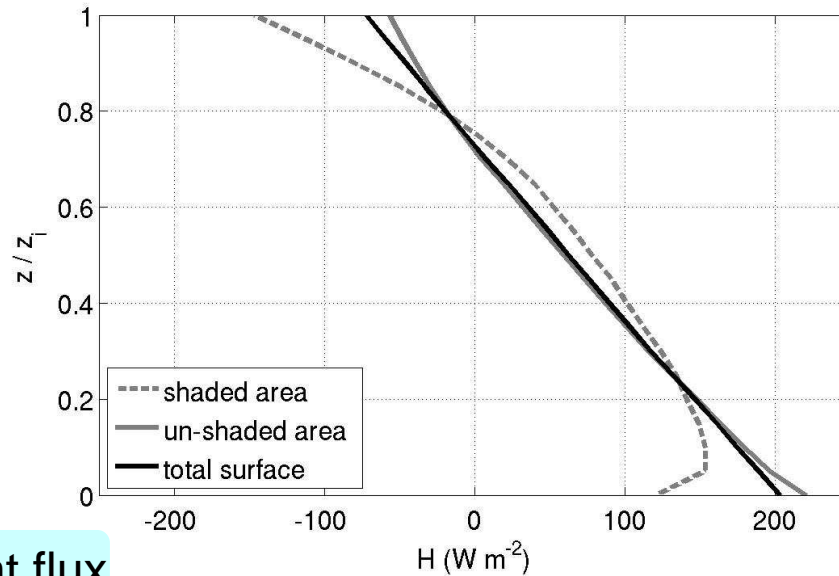
Cloud roots characteristics: vertical fluxes

Averaged vertical slice of a 100 of scaled cloud roots larger than 1 km

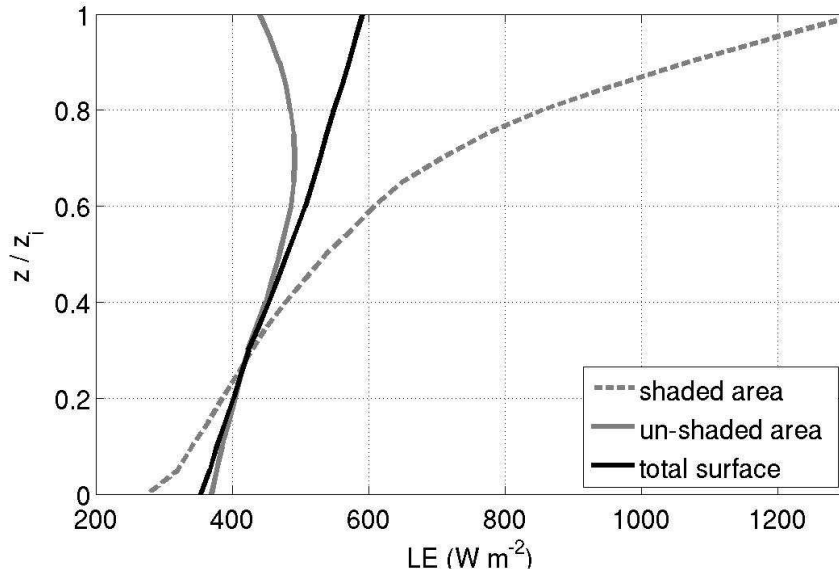


Vertical flux profiles in the shaded and un-shaded areas

Sensible heat flux

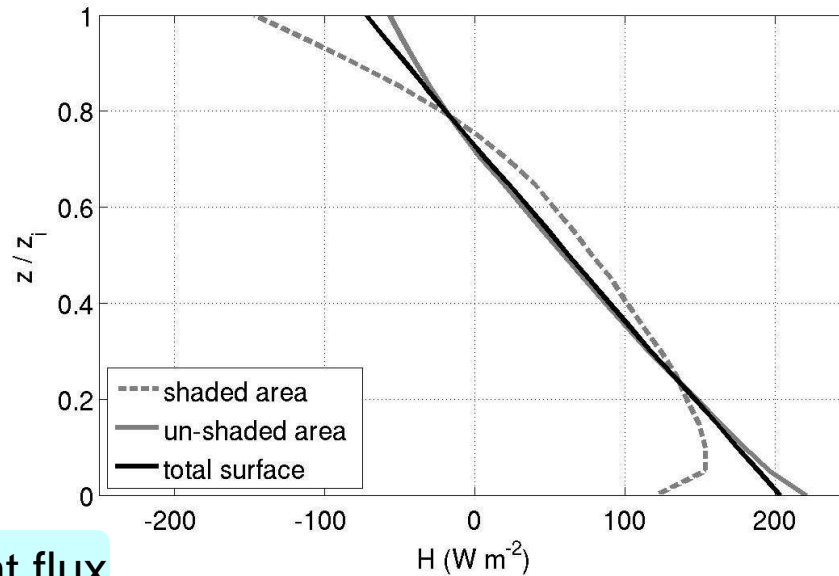


Latent heat flux

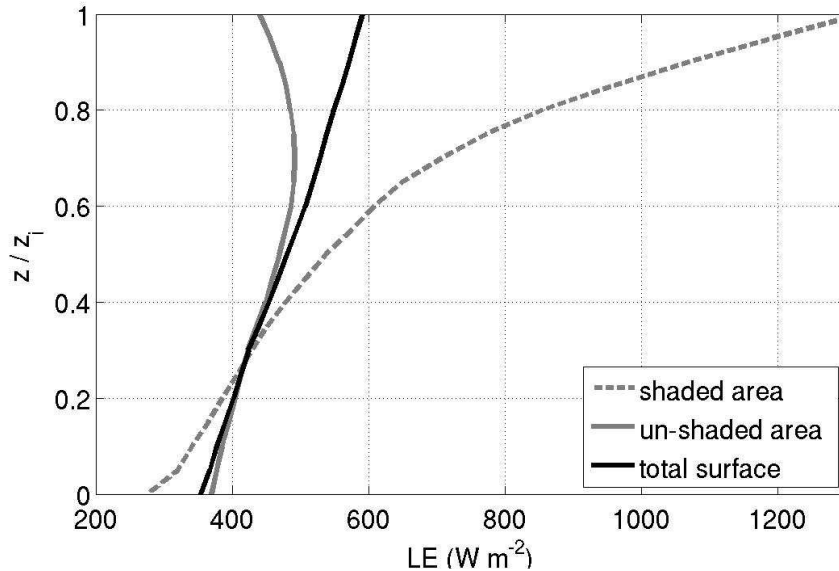


Vertical flux profiles in the shaded and unshaded areas

Sensible heat flux

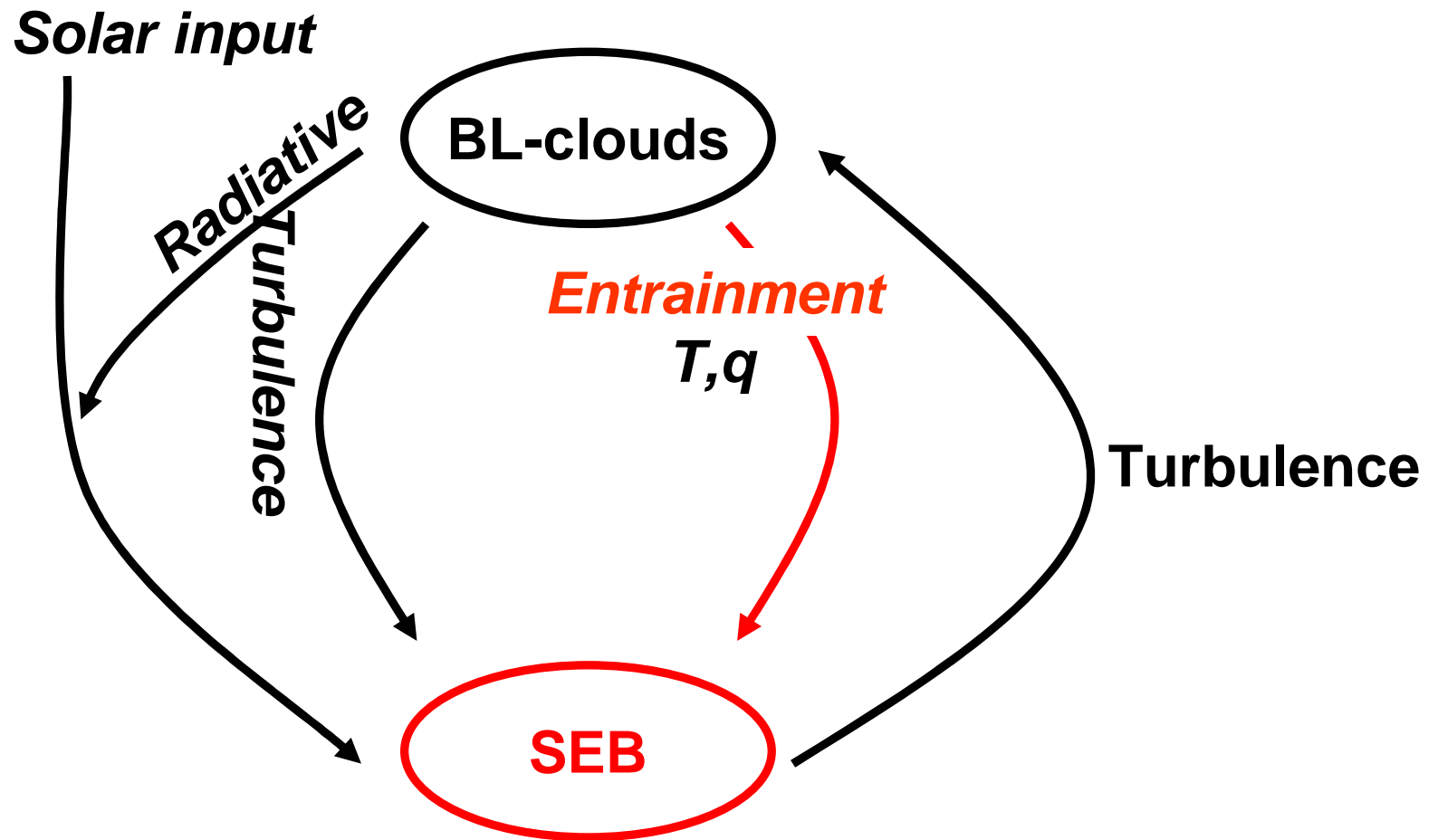


Latent heat flux



Small and moving surface heterogeneities due to cloud shading and secondary circulation are able to affect the atmospheric flux above the surface layer

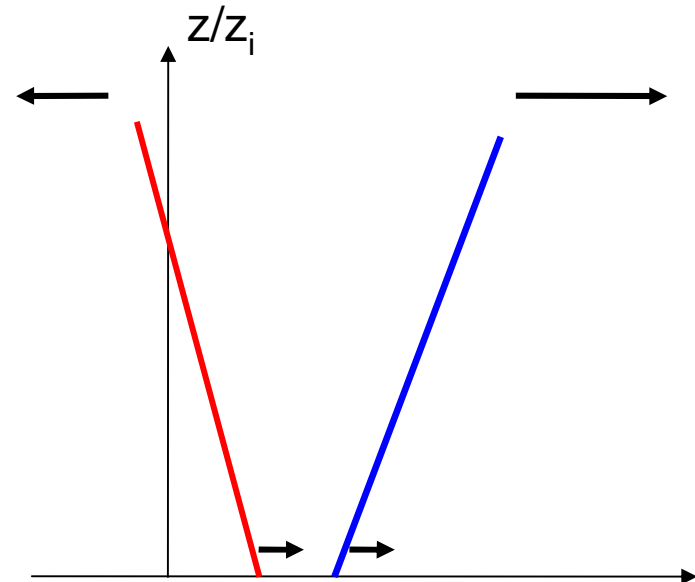
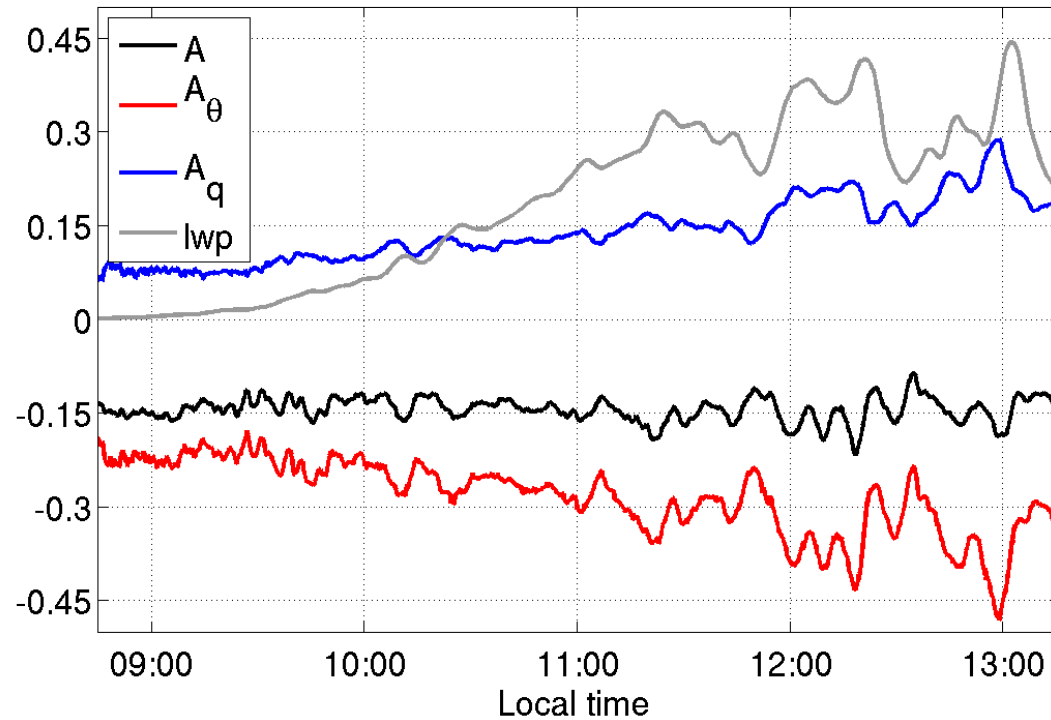
Entrainment associated to cloud BL-clouds



$$A = \frac{\overline{w' \theta'_v}_{zi}}{\overline{w' \theta'_v}_0} \approx \frac{\overline{w' \theta'_v}_{zi}}{\overline{w' \theta'_v}_0} (1 + 0.61 \overline{q}) + 0.61 \overline{\theta} \frac{\overline{w' q'_{zi}}}{\overline{w' \theta'_v}_0}$$

Sensible contribution A_θ

Latent contribution A_q



$-0.2 \leq A \leq -0.1$

Otles and Young, 1996

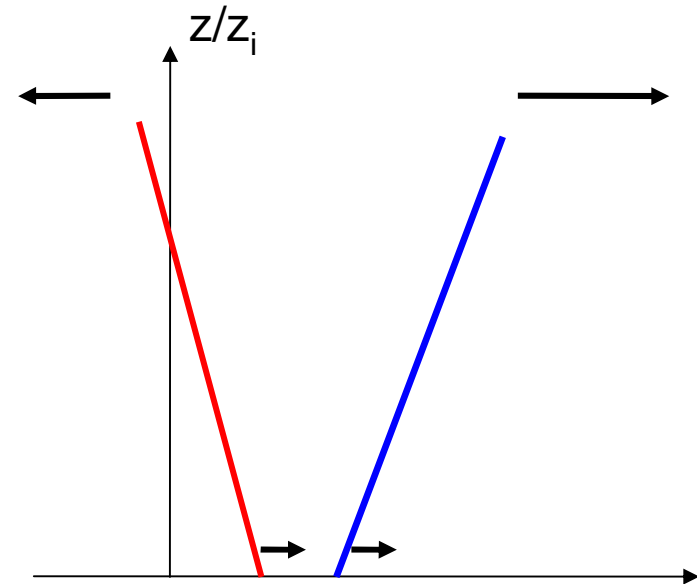
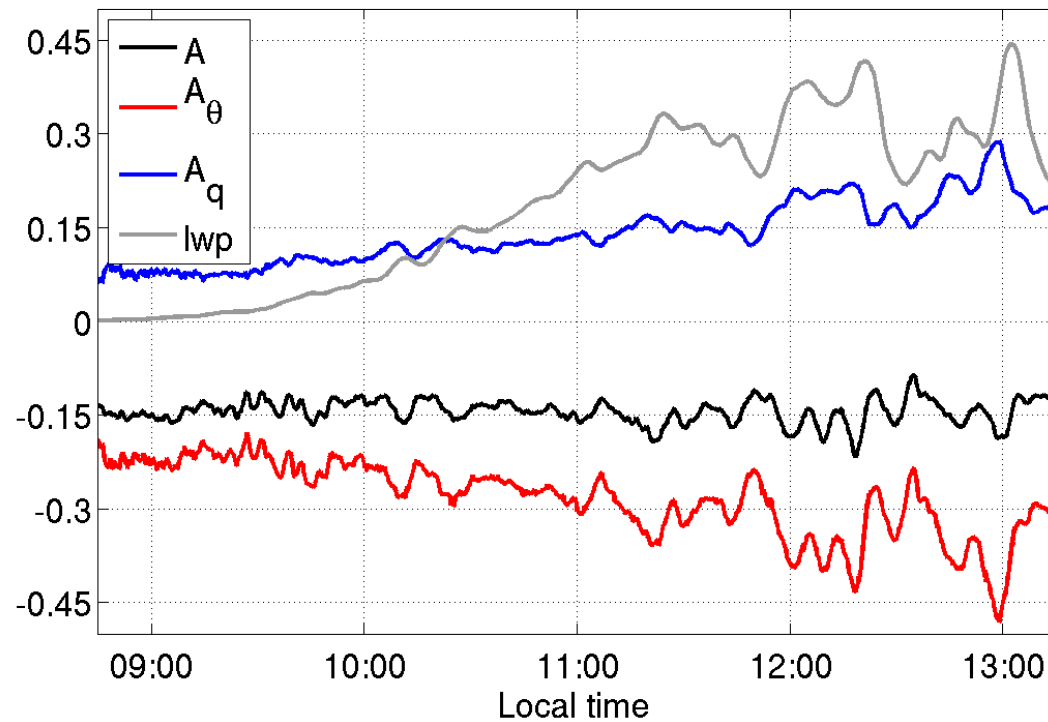
LeMone and Pennel, 1976

Nicholls and LeMone, 1980

$$A = \frac{\overline{w' \theta'_v}_{zi}}{\overline{w' \theta'_v}_0} \approx \frac{\overline{w' \theta'_v}_{zi}}{\overline{w' \theta'_v}_0} (1 + 0.61 \overline{q}) + 0.61 \overline{\theta} \frac{\overline{w' q'_{zi}}}{\overline{w' \theta'_v}_0}$$

Sensible contribution A_θ

Latent contribution A_q



Entrainment rate A is not changed on average.

But at short time scale, cloud activity tends to slightly increase A .



Thank you for your attention



Additional slides

Land-surface / Atmosphere coupling

flux at the surface

$$H = \frac{\rho_a c_p}{r_a} (\theta_s - \theta)$$

$$LE = \frac{\rho_a L}{r_a + r_s} (q_{sat}(T_s) - q)$$

Need of the skin temperature



Pronostic equation for skin temperature

$$C_{skin} \frac{dT_s}{dt} = Q + G - H - LE$$

C_{skin} : skin layer heat capacity

... hard to define

Penman-Monteith equation

$$Q + G = H + LE$$

No skin heat capacity

Immediate response of the skin layer

NCAR LES code

Moeng, 1984, 1986 / Sullivan et al., 1996 / Patton et al., 2005

Set of Navier-Stokes equation
(u,v,w)

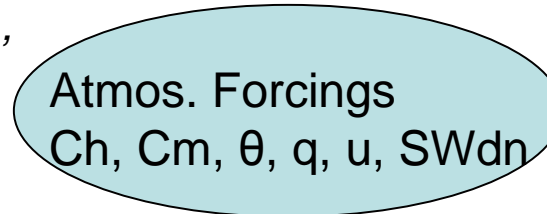
- ✓ Subgrid-scale terms are parameterized with the subgrid TKE and a stability-dependent length scale.
- ✓ Condensation scheme is based on all-or-nothing assumption / Non-precipitating clouds.
- ✓ 1D radiation

Cloud transmission coefficient
(Joseph et al., 1976)

$$t_c = \frac{5 - e^{-\tau}}{4 + 3\tau(1-f)}$$

Optical depth (Stephens, 1984)

$$\tau = \frac{3lwp}{2 r_e} \text{ with } r_e = 10\mu\text{m}$$



NOAH Land surface model

1D set of equations for thermodynamic and hydrologic variables: Mahrt and Pan, 1984, 1987 // Surface energy balance: Chang et al, 1999



NOAH Land surface model

1D set of equations for thermodynamic and hydrologic variables: *Mahrt and Pan, 1984, 1987* // *Surface energy balance: Chang et al, 1999*

Atmos. Forcings
Ch, Cm, θ , q, u, SWdn

$$r_a = \frac{1}{u C_h}$$

$$C_h = f(z_0, L, \psi,$$

$$SMA = \frac{SMC - WL TSMC}{SMC_{MAX} - WL TSMC}$$

Surface Forcings
H, LE, G, T_{skin} , SMC_{skin}

1/ Penman equation

$$\lambda_v E = \frac{\frac{dq_{sat}}{dT_a} Q + \rho_a c_p \delta_e r_a^{-1}}{\frac{dq_{sat}}{dT_a} + \gamma}$$

$$LE = SMA \times \lambda_v E$$

2/ Soil moisture update

3/ Soil heat flux

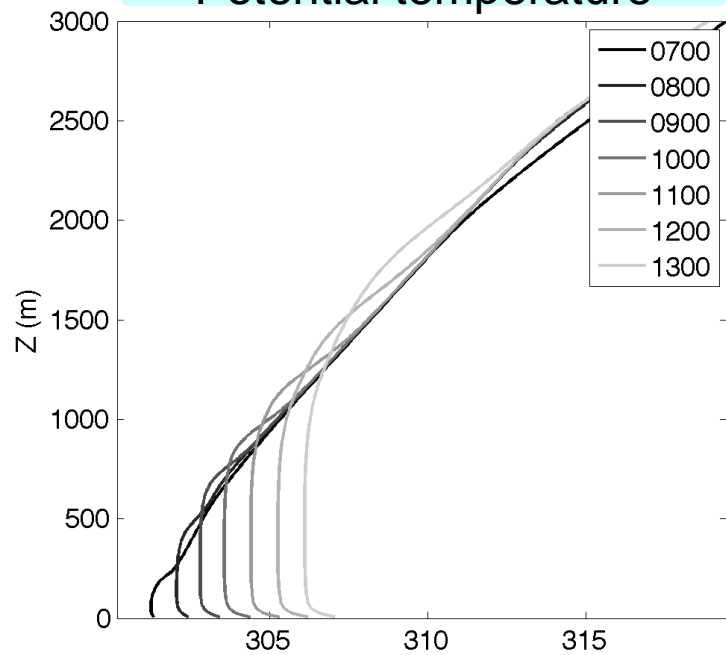
$$G = \gamma \frac{\partial T_g}{\partial z}$$

4/ Soil Temperature update

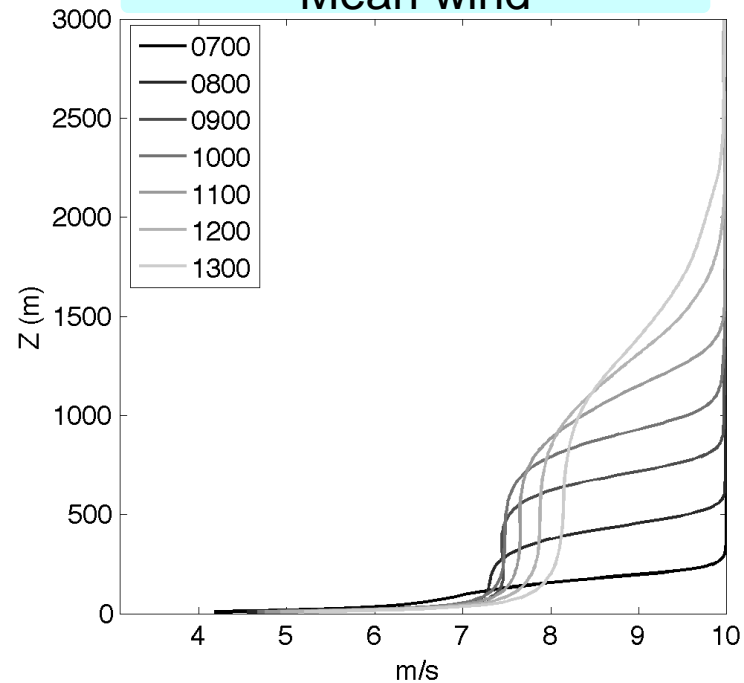
5/ Sensible heat flux

$$H = \frac{\rho_a c_p}{r_a} (\theta_s - \theta)$$

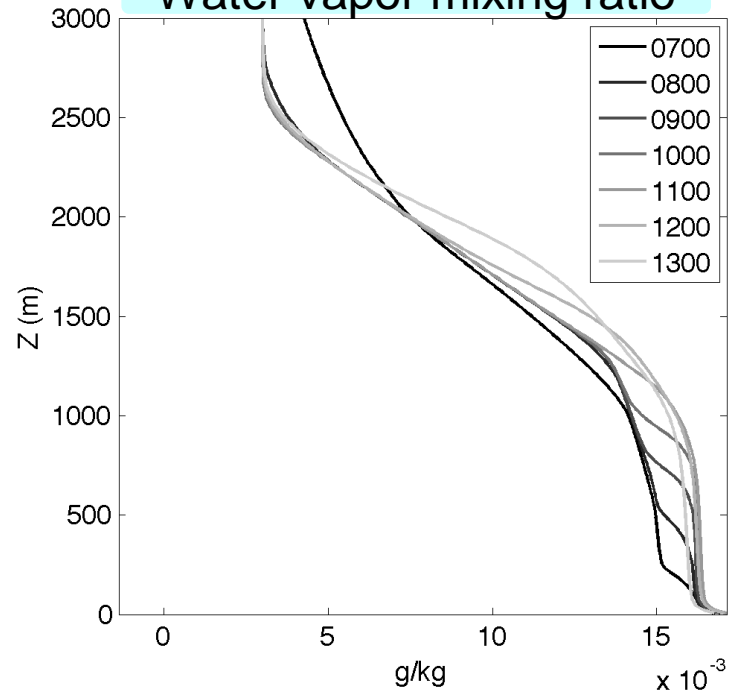
Potential temperature



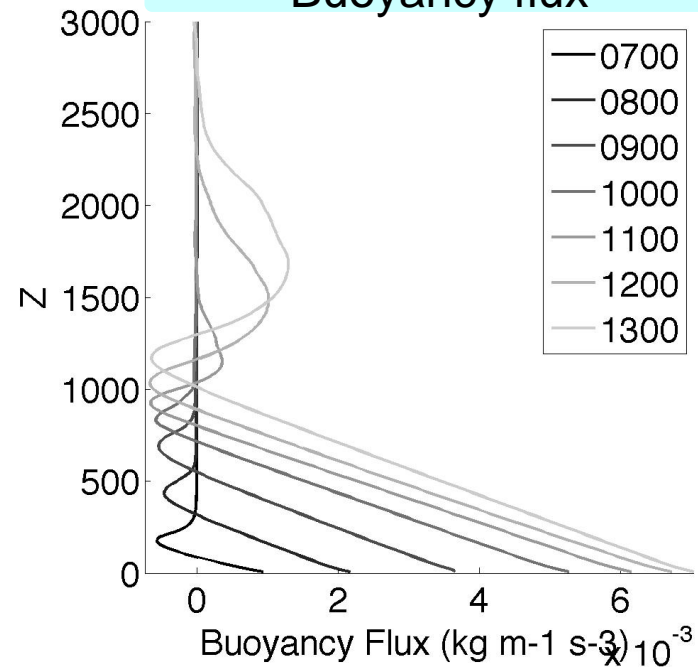
Mean wind



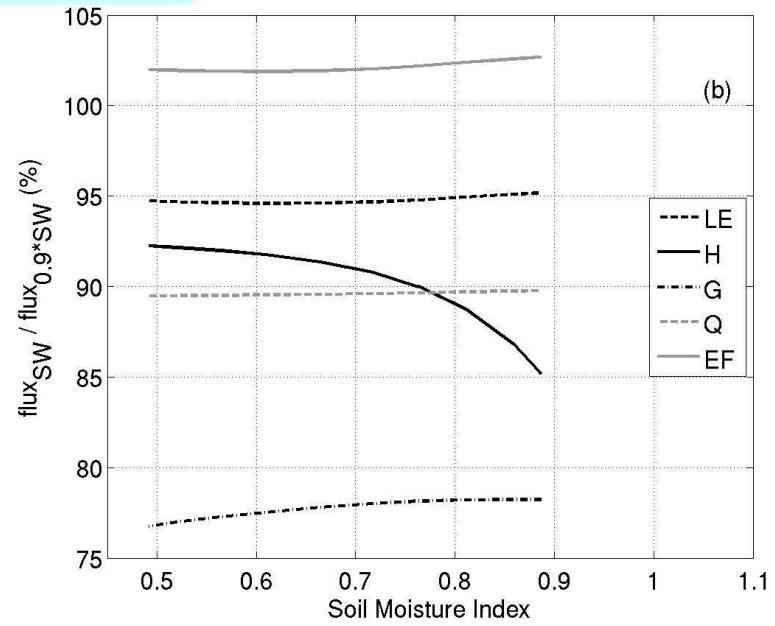
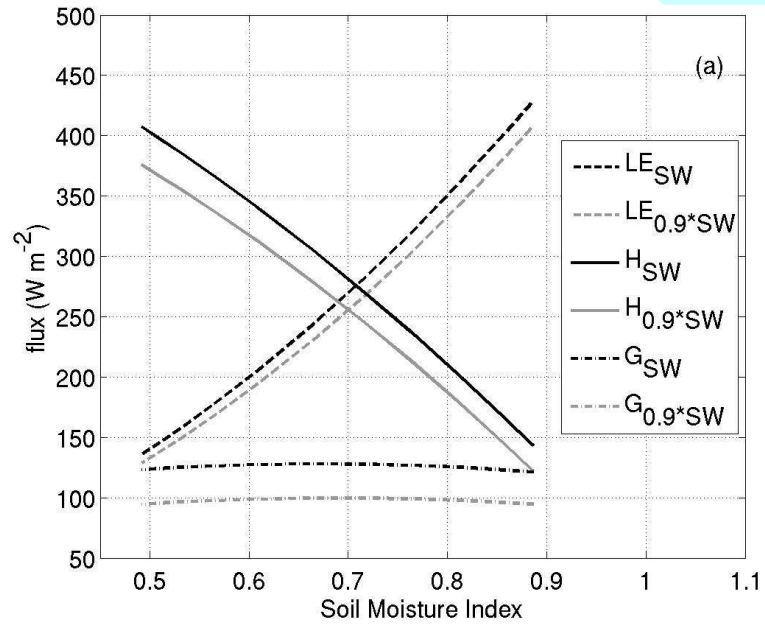
Water vapor mixing ratio



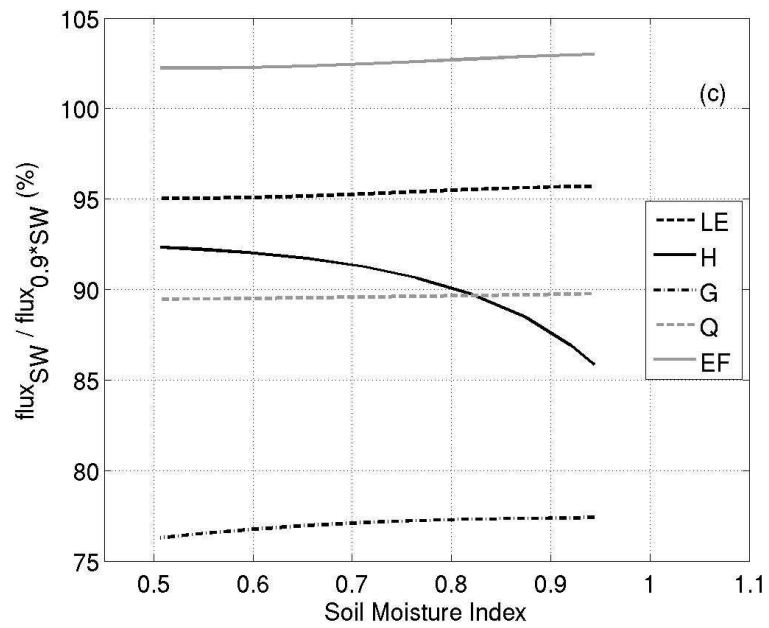
Buoyancy flux



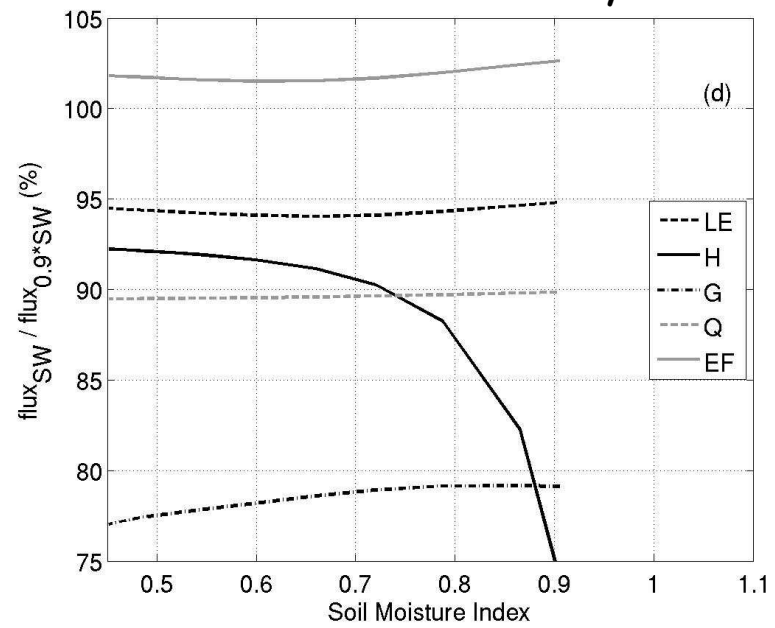
Clay -loam



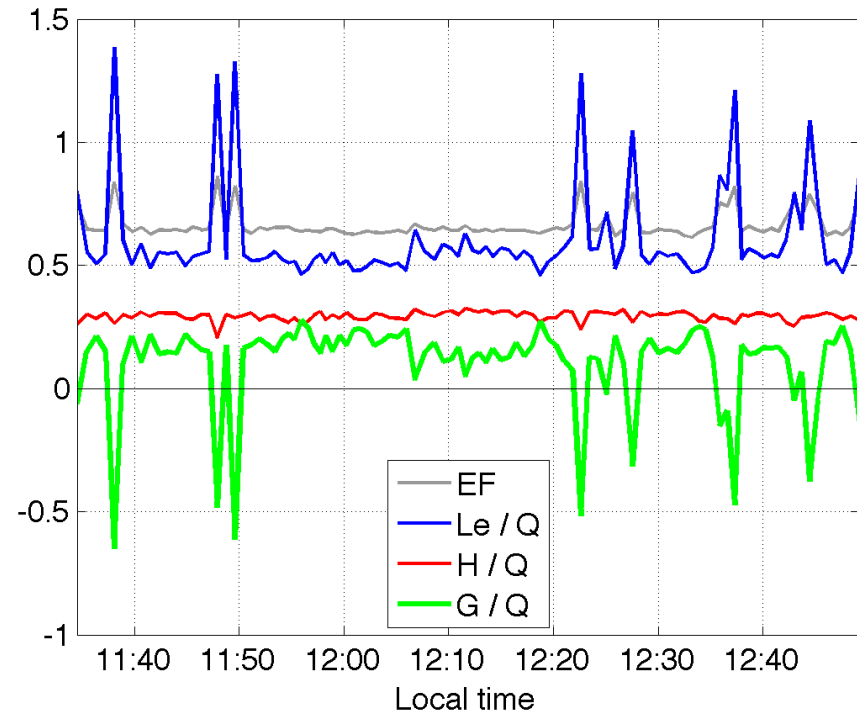
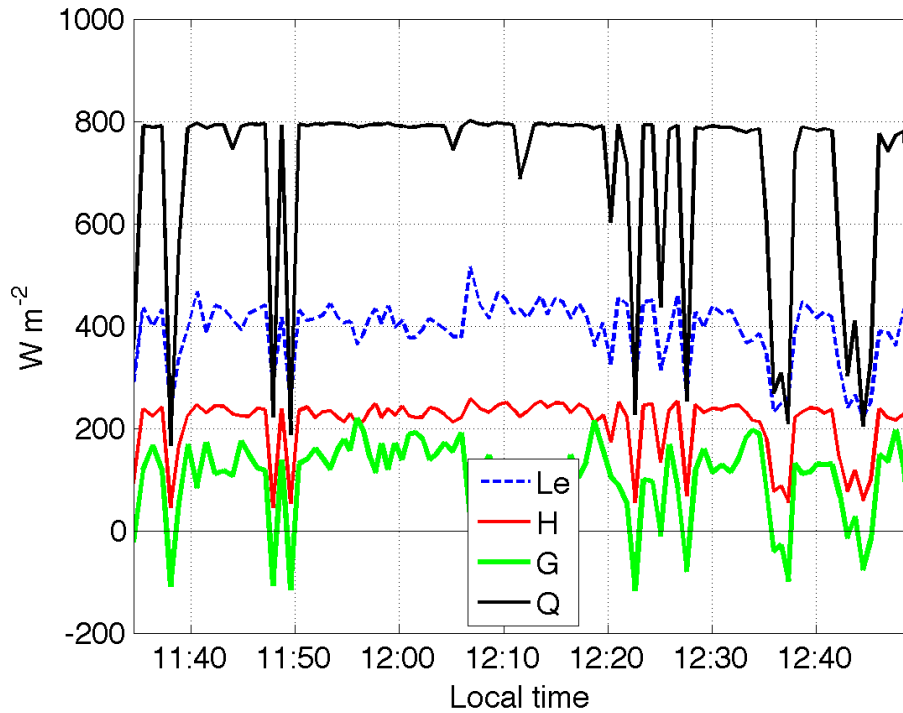
Loam



Clay



Local SEB response to sudden net radiation changes



un_shaded area

shaded area

H ***0.3 Q***

0.3 Q

LE ***0.5 Q***

1.2 Q

G ***0.2 Q***

-0.5 Q

Radiative and turbulence effects on surface flux

In NOAH LSM (Penman, 1948; Monteith, 1981):

$$\lambda_v E = \frac{\frac{dq_{sat}}{dT_a} Q + \rho_a c_p \delta_e r_a^{-1}}{\frac{dq_{sat}}{dT_a} + \gamma}$$

$$r_a = \frac{1}{u C_h}$$

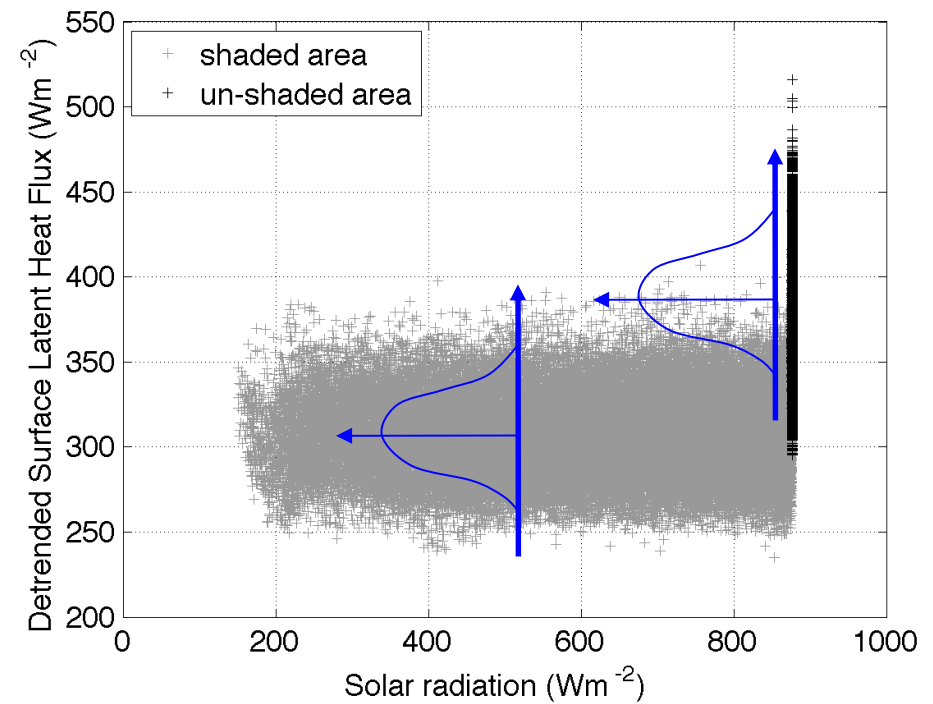
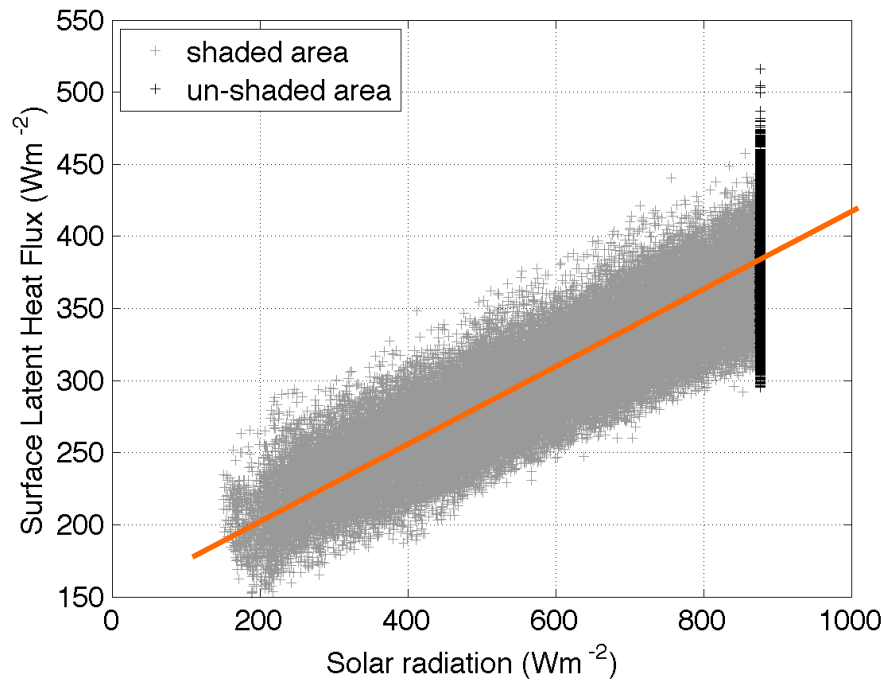
$$C_h = f(z_0, L, \psi)$$

Mean variation of potential evaporation strongly driven by radiative effect.

Cloud-induced turbulence and secondary circulation tend to increase the surface flux variability.

➡ Does this variability participate to convection triggering.

At one time step, over the 2D surface



Radiative effect via the
correlation coefficient

SWdn-LE ~ 0.9

SWdn-H ~ 0.98

Turbulent effect via the
temporal evolution of H and LE distribution

For un-shaded area

$$\frac{\sigma_{F_{US}}}{\langle F \rangle}$$

For shaded area

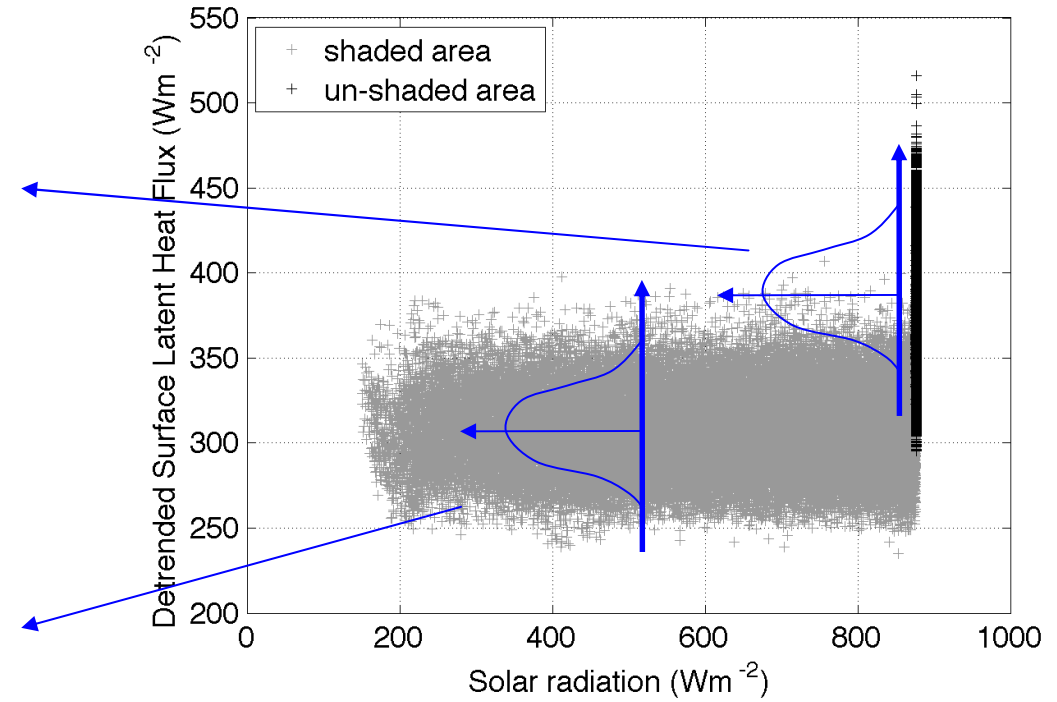
$$\frac{\sigma_{Fd_S}}{\langle F \rangle}$$

F: *H* or *LE*

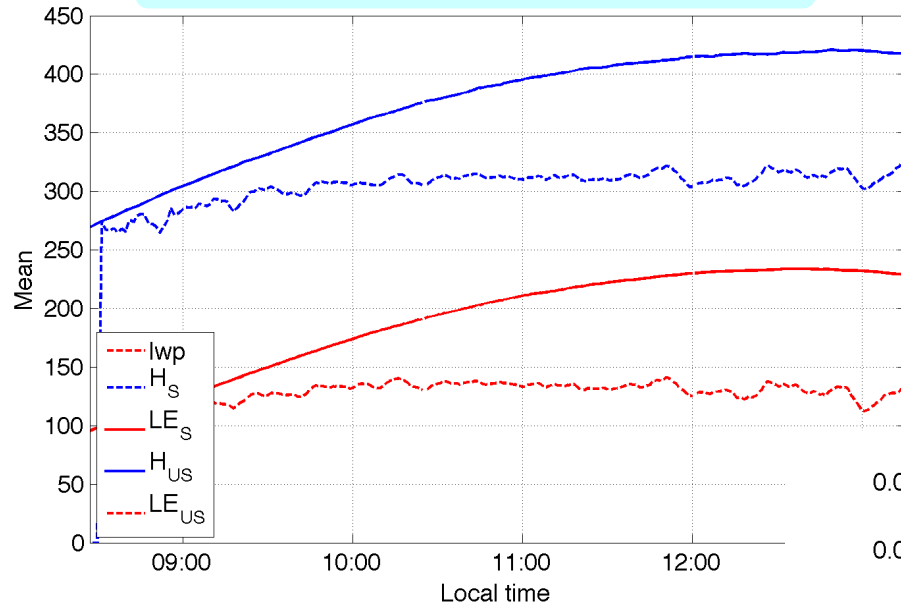
d: *detrended flux*

S: *shaded area*

US: *un-shaded area*

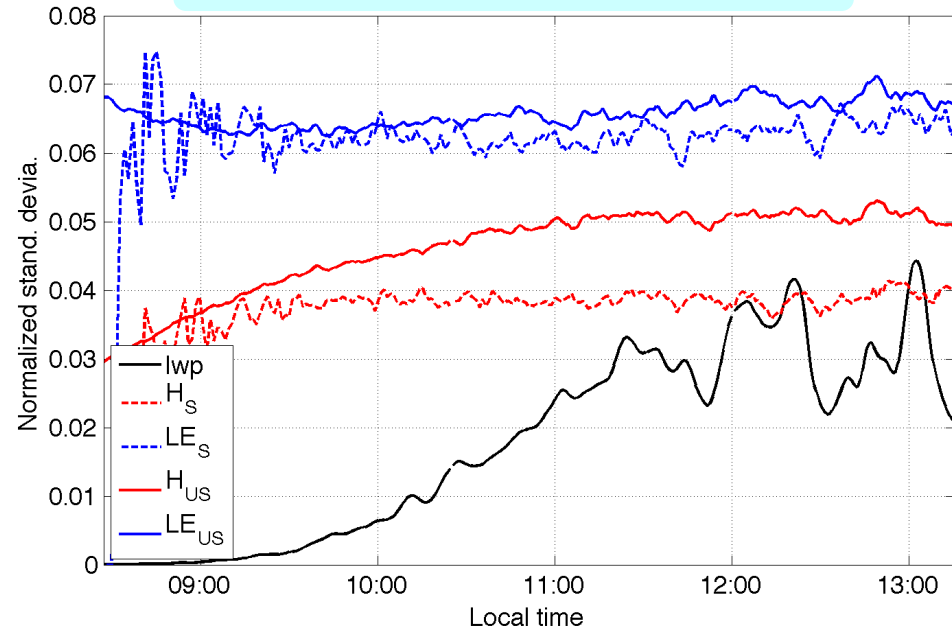


Mean flux



— un-shaded area
- - - shaded area

Normalized standard deviation



The secondary circulations and turbulence associated to cloud activity slightly change the surface flux distribution.