IMPACT OF DOPPLER WEATHER RADAR RADIAL WINDS REFLECTIVITY AND CUMULUS PHYSICS ON THE PREDICTIONS OF SQUALLS OVER BANGLADESH

Mohan K. Das^{1,3,*}, Sujit K. Debsarma¹ and Someshwar Das²

¹SAARC Meteorological Research Centre (SMRC), Dhaka, Bangladesh ²National Centre For Medium Range Weather Forecasting (NCMRWF), NOIDA, India ³Jahangirnagar University, Savar, Bangladesh

1. Introduction

Modern radars are important equipments for nowcasting Local Severe storms which appear in the form of Nor'westers or 'Kal Baisakhis' in Bangladesh and eastern region of India during pre-monsoon season (March to May). These systems develop mainly due to merging of cold dry northwesterly winds aloft and southerly low level warm moist winds from the Bay of Bengal. The Three Dimensional Variational (3DVAR) data assimilation approach is one of the most promising techniques available to directly assimilate heterogeneous mesoscale observations in order to improve the estimate of the model's initial state. In this study, Doppler Weather Radar (DWR) observations (radial winds and reflectivity) of Bangladesh Meteorological Department (BMD) are used during 11 May, 19 May, 21 May and 23 May 2011 squall events in order to update the initial and boundary conditions through 3DVAR technique within the Weather Research Forecasting (WRF) modeling system. It indicates that NWP models are very important for obtaining guidelines for the prediction of local severe thunderstorms [1].

2. Data and Methodology

The WRF Model has been used over the study domain (Fig. 1). The model is run at 3 km resolution with 27 vertical levels by using 6 hourly NCEP-FNL Data (1°x1°) as IC/BC. GTS and non-GTS data (AWS), BMD Khepupara and Cox'sbazar DWR derived radial wind and reflectivity data are utilized in the Experiments through 3DVAR techniques. Kain-Fritsch (KF) and Grell-Devenyi (GD) ensemble cumulus scheme have been used in the model run. Realized weather phenomena over Bangladesh during May 2011 shown in Table 1.

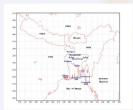


Fig. 1: Model domain under study.

Table 1. Squalls and gusty wind events during May 2011 in Bangladesh

Date	Reporting	Events	Wind Speed	Wind	Occurrence
	Stations		(kph)	Direction	Time (UTC)
	Rangpur	Squall	66	NW	1530
	Mymensingh	Squall	56	NW	1900-2000
01/05/11	Sylhet	Squall	44	W	2045-2056
	Dhaka PBO	Gusty wind	37	NW	2300
	Dhaka MMO	Gusty wind	48	N	2307
02/05/11	Rangpur	Squall	70	NW	2135-2137
04/05/11	Chittagong	Gusty wind	41	W	0907
	Chittagong	Gusty wind	52	NW	1000-1030
05/05/11	Feni	Squall	56	S	1300-1330
09/05/11	Bogra	Squall	111	NW	1030-1100
	Dhaka	Gusty wind	41	W	1330
	Dinajpur	Squall	56	NW	1100- 1130
10/05/11	Dhaka	Gusty wind	59	W	1535
	Sylhet	Gusty wind	56	W	1722-1723
	Rajshahi	Squall	41	W	0300-0330
11/05/11	Dhaka	Gusty wind	56	W	0542-0544
	Chittagong	Squall	70	WNW	0932-0934
12/05/11	Rangpur	Squall	63	NW	0130-0200
13/05/11	Dhaka	Gusty wind	56	W	1640
	Chittagong MMO	Squall	52	NW	2000-2030
14/05/11	Bogra	Squall	52	NE	2030-2100
17/05/11	Dhaka	Gusty wind	37	N	0802
18/05/11	Ishurdi	Squall	59	NW	2044-2045
	Dhaka	Gusty wind	56	W	2225
	Patuakhali	Squall	78	NW	2235-2240
19/05/11	Chittagong MMO	Gusty wind	56	NW	100
	Chittagong PBO	Squall	44	NW	0530-0600
20/05/11	Chittagong PBO	Squall	44	NW	2200-2230
21/05/11	Dhaka MMO	Gusty wind	74	WNW	1209
	Dhaka PBO	Gusty wind	67	W	1215
	Faridpur	Gusty wind	56	NW	1230
	Patuakhali	Squall	59	NW	1400
	Chittagong MMO	Squall	80	N	1500-1530
22/05/11	Chittagong MMO	Squall	56	NW	0950-0953
23/05/11	Bogra	Squall	45	NW	0130-0200
24/05/11	Dhaka	Gusty wind			0800
25/05/11	Dhaka	Gusty wind	37	NW	0710
	Chittagong	Gusty wind	44	N	1230
31/05/11	Chittagong	Gusty wind	48	W	0950-1005

2.1 Model configuration

Configuration of the WRF-ARW Model used in the study is given at Table 2.

Model Feature	Configuration				
Horizontal spatial resolution	3 km				
Grid points	629X461				
Vertical Levels	27				
Topography	USGS				
	Dynamics				
Time Integration Semi Implicit					
Time Steps	15 seconds				
Vertical Differencing	Arakawa's Energy Conserving Scheme				
Horizontal Diffusion	2 nd order over Quasi-pressure, surface, scale selective				
	Physics				
Cumulus parameterization	Kain-Fritsch scheme				
	Grell-Devenyi ensemble scheme				
PBL	Mellor-Yamada-Janjic scheme				
Surface Layer	Monin-Obukhov (Janjic Eta) sheme				
Cloud Microphysics	Eta microphysics				
Radiation	RRTM (LW), Dudhia (SW)				
Land Surface Process	Unified Noah Land Surface Model				

3. DWR derived reflectivity and radial wind analysis

BMD DWR recorded reflectivity of 45 dBZ (Fig. 2) and radial velocity between -45 to 50 m/s (Fig. 3). Study of Radar data showed that the Nor'westers propagate in the form of parallel bow shaped squall lines having horizontal length of more than 250 km at the time of squall.

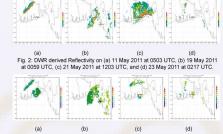


Fig. 3: DWR derived Radial wind on (a) 11 May 2011 at 0503 UTC, (b) 19 May 2011 at 0059 UTC, (c) 21 May 2011 at 1203 UTC, and (d) 23 May 2011 at 0217 UTC.

4. RESULTS AND DISCUSSION

4.1 Convective parameterization schemes (CPSs)

The first simulation used the Kain-Fritsch scheme (KF) and the second one used Greli-Devenyi (GD) ensemble parameterization. To compare the differences between the CPSs, simulations are performed for a particular time period utilizing the same IC/BC and other physical parameterizations for each CPSs and then model outputs are compared with observation. An attempt is made to examine 3 hourly accumulated rain with the comparison of TRMM 3B42RT (Fig. 4). Here KF scheme is comparatively better simulated than GD scheme and 3DVAR shows some improvement.

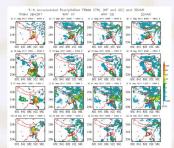


Fig. 4: 3 hrs accumulated precipitation (mm) patterns from TRMM, KF, GD and 3DVAR. Row: row-1: 11 May 2011(a-d), row-2: 19 May 2011(e-h), row-3: 21 May 2011(i-l) and row-4: 23 May 2011(m-p).

4.2 T- phi gram analysis.

T- phi grams of Dhaka are analyzed and instability indices are tabulated below (Table 3). 3DVAR results indicate minor improvement.

Table 3. Simulated stability indices over Dhaka at 0000 UTC of 11, 19, 21 and 23 May 2011

Stability Indices	Thunderstorm	11 May	19 May	21 May	23 May
	Events	2011	2011	2011	2011
CAPE	Obs.	688.01	10.65	1837.68	1475.31
	CTRL	1498	54	1064	1899
	3DVAR	1451	154	1066	2050
Lifted index (LI)	Obs.	-5.53	-0.34	-5.74	-3.75
	CTRL	-8	0	-8	-5
	3DVAR	-8	-2	-5	-5
K index (KI)	Obs.	44	42.7	43.4	34.4
	CTRL	35	32	39	34
	3DVAR	38	35	40	34
Totals Totals Index (TTI)	Obs.	57.7	47.8	48.9	42
	CTRL	52	42	50	43
	3DVAR	53	47	50	43
Precipitable water (mm)	Obs.	58.81	71.87	73.74	61.74
	CTRL	33.4	46.7	40.7	37.7
	3DVAR	33.8	46.8	43.1	37.9

4.3 Sea Level pressure (SLP) Time Series

SLP seems to be less than 1002 hPa (Fig. 5) over Dhaka due to existence of heat low during the pre-monsoon and 3DVAR showed the pressure drop and increasing tendency due to convection which is helpful for forecasting.

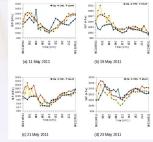


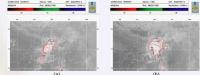
Fig. 5. Comparison of Sea Level Pressure (SLP) among obs, CTRL and 3DVAR (a) 11May 2011, (b) 19 May 2011 (c) 21 May 2011 and (d) 23 May 2011.

3DVAR shows improvements of predictions in three out of four cases (Table 4)

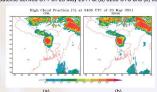


4.4 Satellite Derived Cloud Top Temperature (CTT) (°C) and model simulated high cloud fraction (%)

On 23 May 2011, moderate to strong convection was seen over northwest of Bangladesh by Kalpana-1 satellite imagery (CTT -60 °C) which was moving eastwards and expanded over northeast of Bangladesh (Fig. 6). Model simulated results showed some spatial shiftment of high cloud fraction (Fig.7).



(a) (b)
Fig. 6: Satellite derived CTT on 23 May 2011 at (a) 0230 UTC and (b) 0500 UTC.



(a) (b)
Fig. 7: Simulated High cloud fraction on 23 May 2011 at 0400
UTC (a) CTRL and (b) 3DVAR run.

4.5 850 hPa Wind Vector and 10m Wind Speed analysis

850 hPa horizontal winds in the experimental forecasts are having higher gustiness over the region of the observed system compared to the CTRL forecasts. There is a strong trough at 850 hPa simulated by the 3DVAR; this is absent in the CTRL run (Fig. 8).

The 3DVAR run has simulated high 10m wind speed (12 -15 m/s) along the coastal belt of West Bengal.

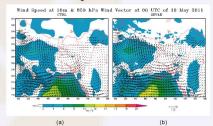


Fig. 8: Vector Wind at 850 hPa and 10 m wind speed (m/s) forecasts valid at 0000

5. Conclusion

The impact of data assimilation is clearly highlighted as the experimental simulations from the WRF-3DVAR are able to capture the thunderstorm closer to the observations compared to the CTRL run. The position and intensity of the simulated thunderstorms in the experimental runs is close to the observed values, as compared to the DWR derived products.

The thunderstorm event from the model simulation is best highlighted from the spatial distribution of 10 m wind speed and SLP. It is observed that the signature of the thunderstorm activity is well represented in the model generated CAPE. The 850 wind fields also reflects the system captured by the model simulations [2].

The model is still not able to accurately produce the exact location and time of occurrences of the thunderstorms. The genesis and growth of cells, their horizontal distribution, vertical structure, the direction and speed of their movement and time of dissipation are still some of the challenges that remained unresolved.

Acknowledgements

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References

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[2] Someshwar Das, R. Ashrit, M. W. Moncrieff (2006): Simulation of a Himalayan Cloudburst Event. *Journal of Earth System Science* 115 (3): 299-313.

SAARC Meteorological Research Centre (SMRC)
Plot: E-4/C, Agargaon
Sher-E-Bangla Nagar, Chaka-1207
Bangladech
"E-mail: mkdas@saarc-smrc.org
Website: www.saarc-smrc.org