

Observations and Analysis of Atmospheric Waves During the Historic April 27, 2011 Tornado Outbreak

Mobile Integrated Profiling System

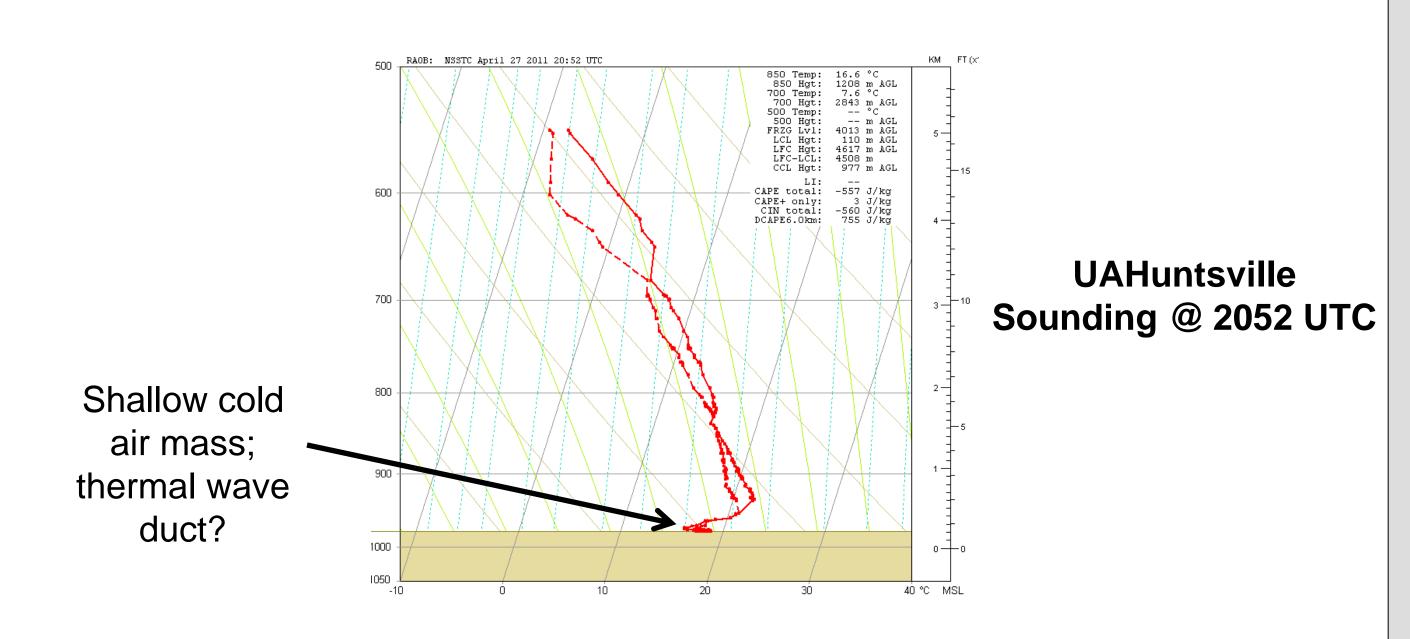
Mobile Alabama X-band
Dual Polarization Radar

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1. Introduction & Background

During the historic deep south tornado outbreak of 27 April 2011, early morning and mid-day convection helped establish and reinforce a thermal boundary just south of the Tennessee Valley. Surface observations around the area indicated temperatures near 60 °F as late as 1800 UTC, while locations ~160 km to the south, in central AL, were approaching 80 °F.

An atmospheric sounding (below) launched from UAHuntsville's campus in Huntsville, AL showed the TN Valley continued to be located on the cool side of the boundary as late as 2100 UTC. A shallow layer of cooler air extended from the surface (212 m MSL) to about 400-500 m MSL. This layer was 200-300 m deeper 2-3 hours preceding the sounding (see MPR time-series in § 3).

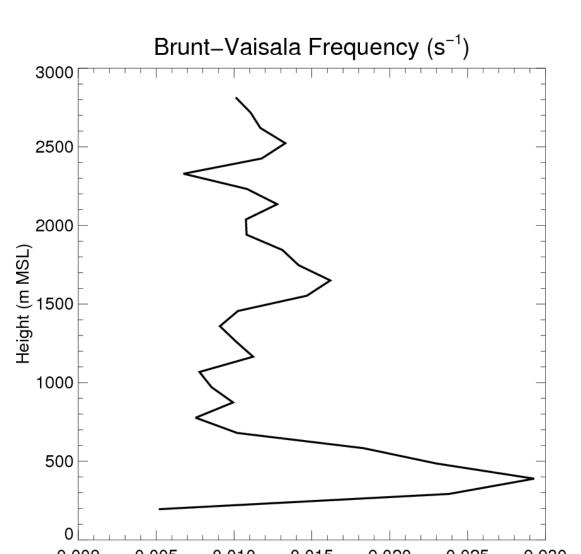


This atmospheric temperature profile provides a potential thermal duct for the generation and propagation of ducted gravity waves (Lindzen and Tung 1976; Koch and O' Handley 1997). Gravity waves can be reflected by the ground and/or by atmospheric layers with vertical gradients of m^2 (vertical wavenumber), thereby trapping the wave within a "duct" (Gill 1982; Nappo 2002). The vertical wavenumber is associated with gradients in the Scorer parameter l^2 (Scorer 1949) through $l^2 = l^2 - k^2$, where l^2 is

$$l^{2} = \frac{N^{2}}{(c - U)^{2}} + \frac{\P^{2}U}{(c - U)^{2}}$$

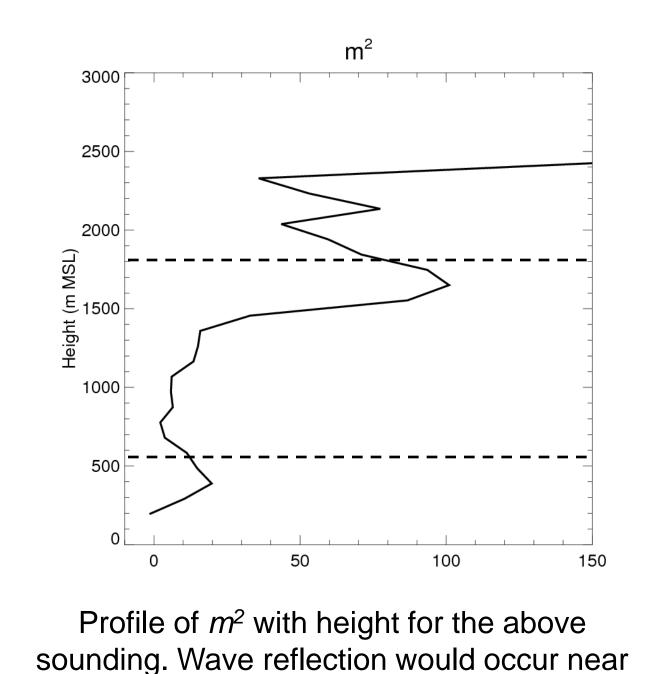
and c is the ground-relative wave phase speed, U is the mean wave-normal background wind, and N is the Brunt-Vaisala frequency.

The (preliminary) thermal duct analysis (below) shows potential reflecting layers near 550 and 1800 m MSL, associated with large vertical gradients of N, indicating an environment supportive of ducted gravity waves. This has important implications, since gravity waves have been shown to initiate or enhance convection (e.g., Uccellini 1975; Stobie et al. 1983; Koch et al. 1988) and also increase the potential for tornadogenesis if they were to interact with a mesocyclone (Coleman and Knupp 2008).



O.000 0.005 0.010 0.015 0.020 0.025 0.030

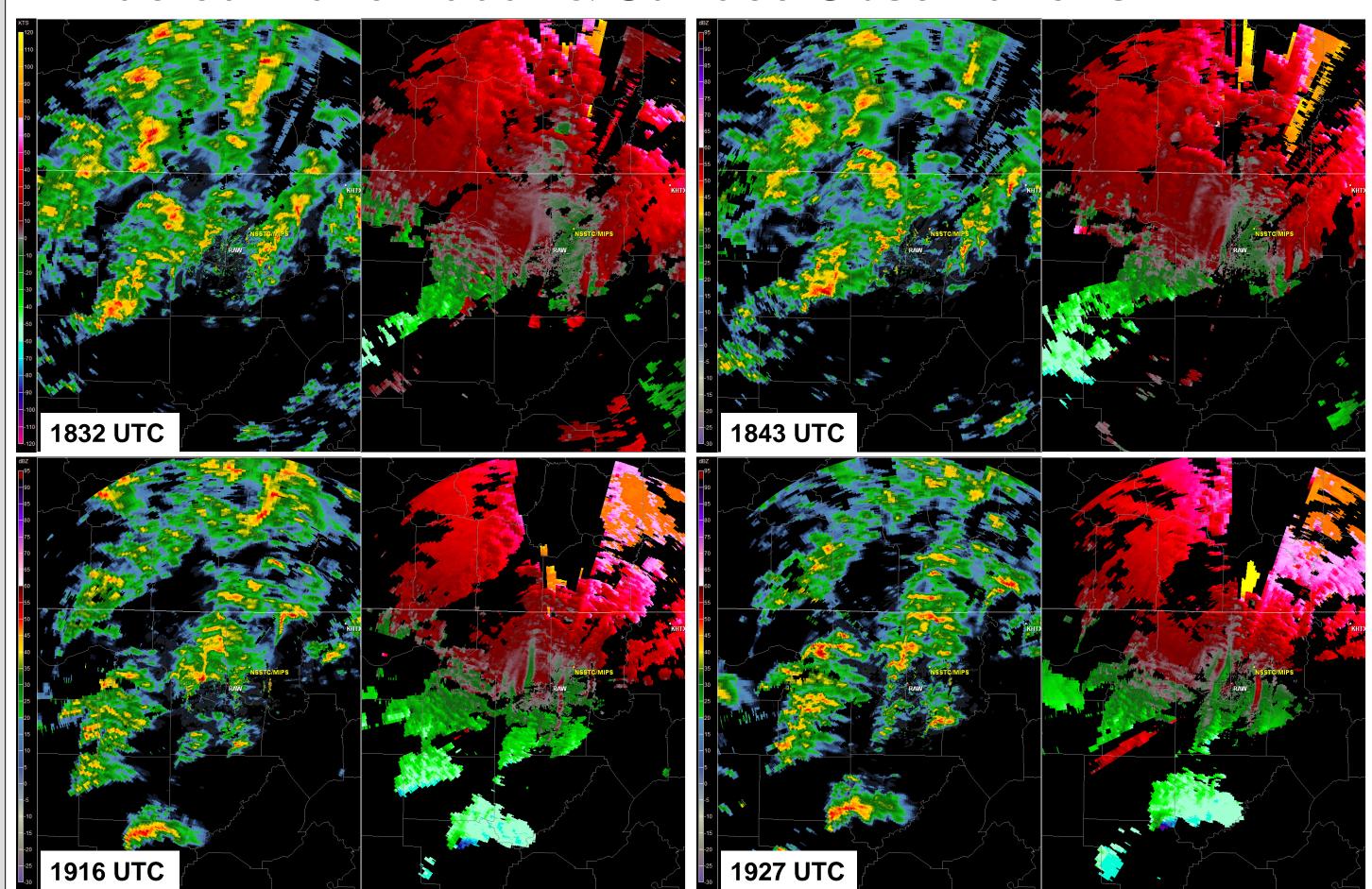
Profile of N with height for the above sounding



the areas indicated by the dashed-lines and

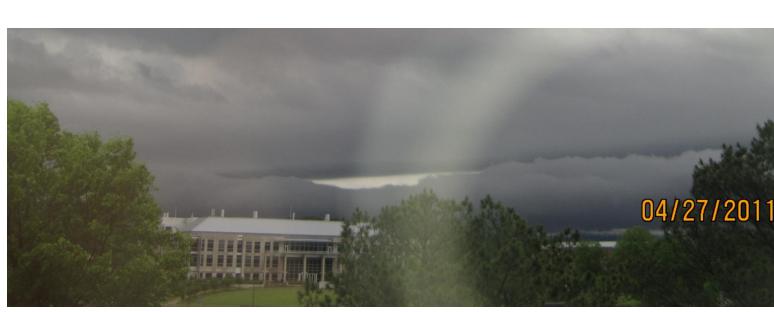
the ground.

2. Ducted Wave Radar & Surface Observations



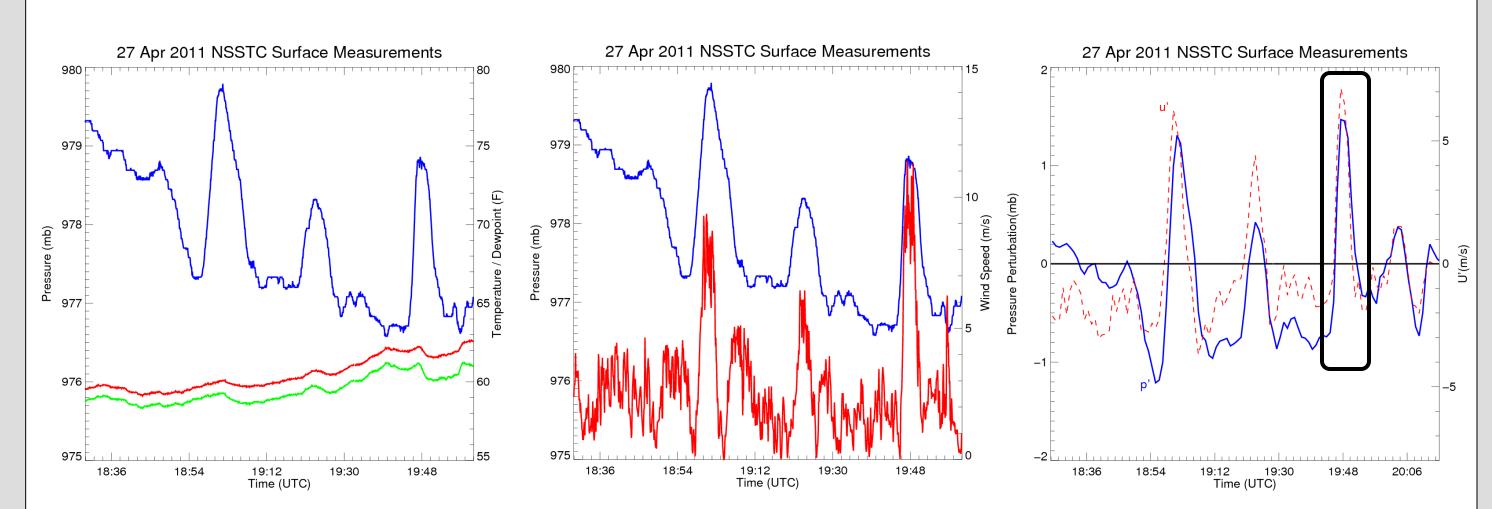
The mid-day QLCS exited the TN Valley near 1800 UTC and was followed by convective showers that formed and propagated along NNE-SSW oriented lines. With a thermal wave duct in place, it is believed these convective showers formed along the ridges (upward moving portion) of ducted gravity waves.

Of particular interest is the observation that the supercell which produced an EF-4 tornado in Cullman, AL appears to link up and propagate with a gravity wave ridge.



Photograph taken near 1910 UTC from the NSSTC on UAHuntsville's campus; view is toward the east. A low-level roll cloud is visible, extending from N to S.

This particular cloud lines up with one of the radar observed convective lines.



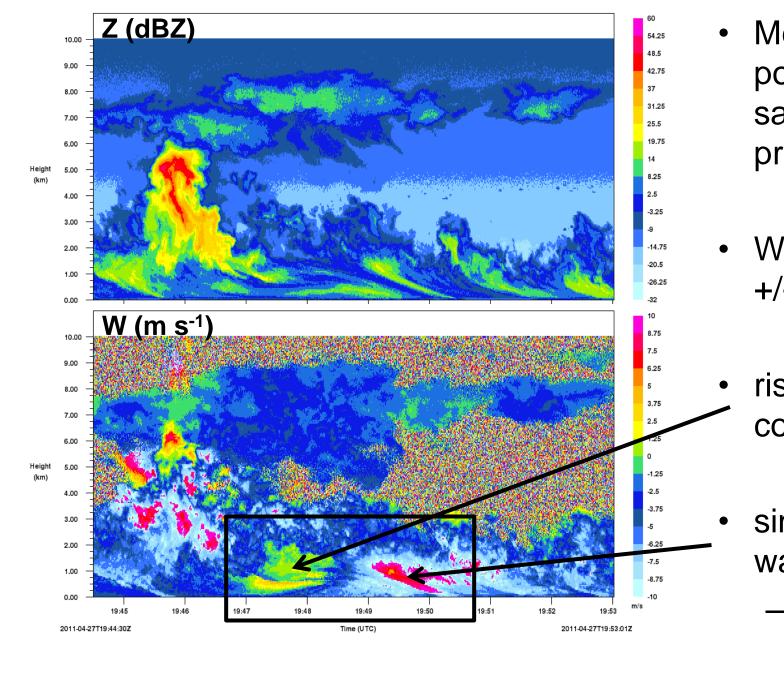
Left: 5-s surface observations of pressure (blue), temperature (red), and dew point (green) at the NSSTC/MIPS location **Middle:** Same as left but for surface pressure (blue) and wind speed (red) **Right:** Same as middle but for surface pressure perturbation (blue) and wind speed perturbation (red dashed). Oscillation near 1948 UTC is examined in greater detail in the next section (XPR measurements).

Significant oscillations in surface pressure and wind speed occurred as these features moved over the NSSTC/MIPS surface instrumentation.

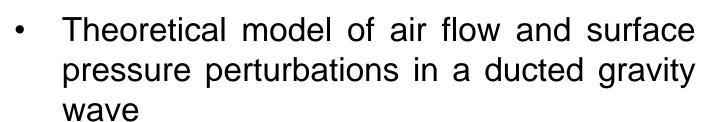
A nonlinear impedance relation between pressure and wind apply in gravity waves, such that the positive/negative pressure and wind perturbations should correlate (Gossard and Hooke 1975; Coleman and Knupp 2010).

This is exactly what is observed as the N-S oriented lines moved over the NSSTC/MIPS site, providing further validation they may be gravity waves.

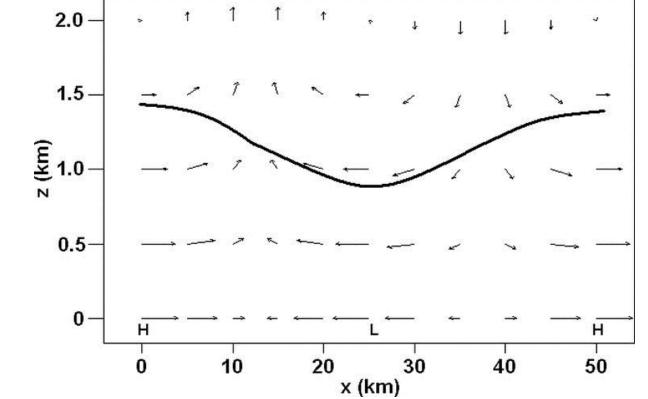
3. Profiling and Dual-Doppler Observations

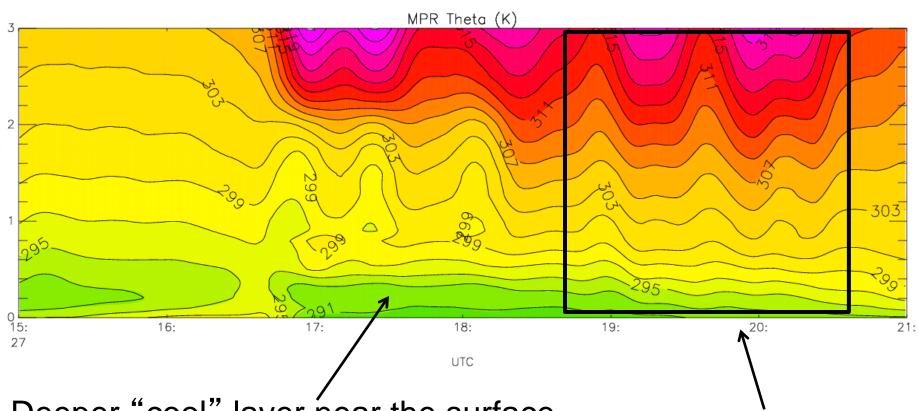


- Measurements from the MIPS vertically pointing X-band radar (XPR; 6 Hz sampling) centered around the pressure/wind oscillation near 1948 UTC
- Wave model (see below) tells us to expect
 +/- vertical velocity couplet before +/- p'
- rising air just prior to +p' (adiabatic cooling; wave ridge)
- sinking air just prior *-p'* (subsidence warming; wave trough)
- Velocity folding (~16 m s⁻¹)

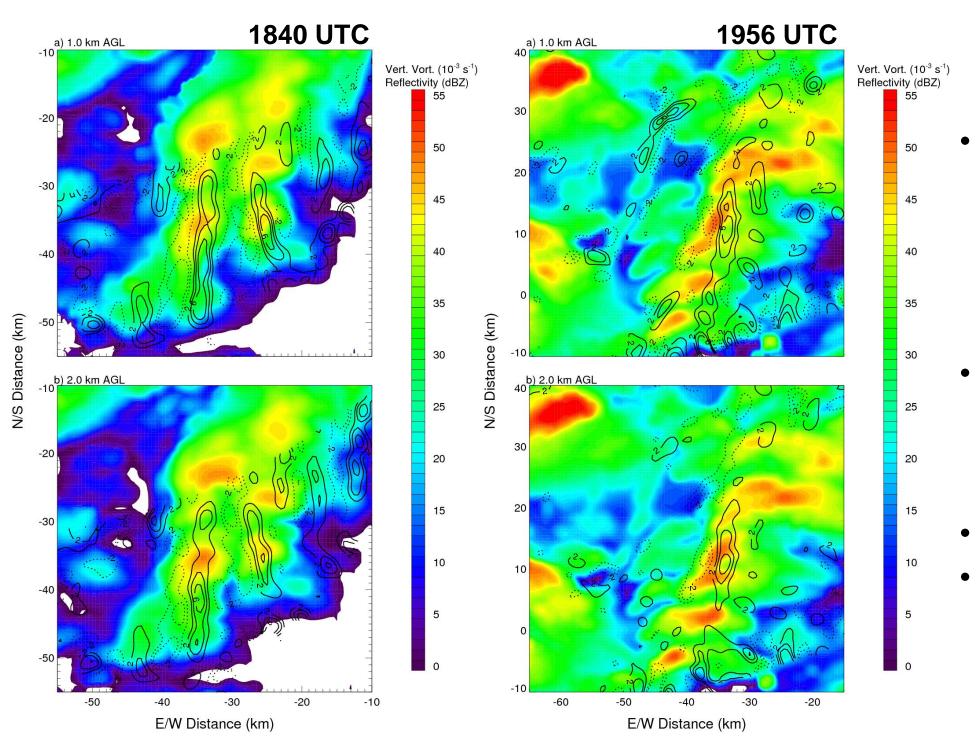


From Coleman and Knupp (2008)





- Potential temperature (θ) derived from the MIPS MPR
- "Jumps" in θ surfaces are often observed during the passage of atmospheric waves (Koch et al. 1991; Knupp 2006; Coleman and Knupp 2011)
- Deeper "cool" layer near the surface
- Time of suspected wave passages
- Clear wave signature in θ



- Vertical vorticity analysis from a (preliminary) ARMOR-KHTX dual-Doppler retrieval of the potential waves
- Indicates +/- ζ couplet associated with the N/S oriented lines
- +ζ along wave ridge
- -ζ along wave trough

This ζ configuration may have important considerations, since the Cullman tornadic supercell was observed propagating along with the potential gravity waves.

It is speculated the ducted gravity waves helped initiate the supercell, while also supplying a source of ambient vorticity. Additional dual-Doppler analyses are being completed, including those which contain the supercell, in an attempt to confirm this hypothesis.