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Displaced Ensemble variational assimilation to incorporate MWI TBs into a CRM for Typhoon Conson

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OUTLINE

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EnVA: min. cost function in the Ensemble forecast error subspace

Minimize the cost function with non-linear Obs. term.

 $J_x = \frac{1}{2}(\bar{X} - \bar{X}_f)P_f^{-1}(\bar{X} - \bar{X}_f) + \frac{1}{2}(Y - H(\bar{X}))R^{-1}(Y - H(\bar{X}))$

- ◆ Assume the analysis error belongs to the Ensemble forecast error subspace (Lorenc, 2003): $\vec{X} \vec{X}^{f} = P_{e}^{f/2} \circ \Omega \qquad \Omega = [\vec{w}_{1}, \vec{w}_{2}, ..., \vec{w}_{N}]$ $P_{e}^{f/2} = [\vec{X}_{1}^{f} \vec{X}^{f}, \vec{X}_{2}^{f} \vec{X}^{f}, ..., \vec{X}_{N}^{f} \vec{X}^{f}]$
- Forecast error covariance is determined by localization $P^{f} \!=\! P_{e}^{f} \circ S$

Cost function in the Ensemble forecast error subspace:

 $J(\Omega) = \frac{1}{2} \operatorname{trace}\{\Omega^{t} S^{-1}\Omega\} + \frac{1}{2}\{H(\bar{X}(\Omega)) - Y\}^{t} R^{-1}\{H(\bar{X}(\Omega)) - Y\}$

Problem in EnVA (1): Displacement error

Large scale
displacement errors of
rainy areas between the
MWI observation and
Ensemble forecasts

 Presupposition of Ensemble assimilation is not satisfied in observed rain areas without forecasted rain.



Mean of Ensemble Forecast (2004/6/9/15UTC FT=7h)









Methodology

- Application results for Typhoon CONSON (T0404)
 - Case
 - Assimilation Results
 - Impact on precipitation forecasts

Displaced Ensemble variational assimilation method

In addition to \overline{X} , we introduced \overline{d} to assimilation. The optimal analysis value maximizes : $\operatorname{arg\,max} P(\overline{X}, \overline{d} \mid Y, \overline{X}^f)$ $P(\overline{X}, \overline{d} | Y, \overline{X}^f) = P(\overline{d} | Y, \overline{X}^f) P(\overline{X} | \overline{d}, Y, \overline{X}^f)$ Assimilation results in the following 2 steps: 1) DEC scheme to derive \bar{d}^a from $P(\bar{d} | Y, \bar{X}^f)$ 2)EnVA scheme using the DEC Ensembles to derive \overline{X}^{a} from $P(\overline{X} | \overline{d}^{a}, Y, \overline{X}^{f})$

Assimilation Procedures



 $J_{x} = \frac{1}{2} \left(\frac{\vec{X}}{\vec{X}} - \bar{X}^{f}(\tilde{d}) \right) P_{f}^{-1} \left(\frac{\vec{X}}{\vec{X}} - \bar{X}^{f}(\tilde{d}) \right) + \frac{1}{2} \left(Y - H(\vec{X}) \right) R^{-1} \left(Y - H(\vec{X}) \right)$





Assimilate TMI TBs (10v, 19v, 21v) at 22UTC









CRM Ensemble Forecasts & RTM

CRM: JMANHM (Saito et al, 2001)

- Resolution: 5 km
- Explicitly predict hydrometeors

Ensemble forecasts

- 100 members started with perturbed initial data at 04/6/9/15 UTC (FG)
- Geostrophically-balanced perturbation (Mitchell et al. 2002) plus Humidity

RTM: Guosheng Liu (2004)

•Plane-parallel model (4stream approx.)

•Mie Scattering (Sphere)

Ensemble mean (FG) Rain mix. ratio

















Summary

Ensemble-based data assimilation can give erroneous analysis, particularly for observed rain areas without forecasted rain.

We developed an ensemble-based assimilation method that uses Ensemble forecast error covariance with displacement error correction.

- The application results (T0404) showed that the assimilation of TMI TBs alleviated the large-scale displacement errors and improved precip forecasts.
- The scheme also avoided misinterpretation of TB increments due to precip displacements as those from other variables.



Thank you for your attention.

End