

1. INTRODUCTION

As discussed in the literature, tropical cyclones (TCs) are energized by the released latent heat due to condensation of moist convection within the eyewall and rainbands (Nolan 2007, Sawada and Iwasaki 2010). Fraction of this latent heat energy is transformed into available potential energy (APE) and kinetic energy. According to Emmanuel's MPI theory (Emanuel 1997, Wang 2010), the production rate of the kinetic energy during intensification increases linearly with wind speed, αV^1 ; however, the dissipation rate due to surface friction increases even faster, αV^3 , until the kinetic energy balances and reaches steady state. To examine this energy adjustment mechanism, sensitivity experiments were performed to TC Megi (2010) using the JMA/MRI Non-hydrostatic Model (Saito et al., 2007). This work focuses on the intensification of Megi from the energetics point of view and its sensitivity to planetary boundary layer (PBL) schemes.



Tuleya and Kurihara, 1975: The energy and angular momentum budgets of a three-dimensional tropical cyclone model, JAS, 32(2), 287-301 □ Wang ,and Xu, 2010: Energy Production, Frictional Dissipation, and Maximum Intensity of a Numerically Simulated Tropical Cyclone. JAS, 67, 97–116

A numerical sensitivity study on the energetics of Tropical Cyclone Megi (2010) during intensification

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3. KINETIC ENERGY TENDENCY: Impacts of increasing Cd and changing the DF PBL scheme to MYNN



4. DISCUSSIONS: How did EC and EL increased with Cd1.5 and MYYN scheme? Why did the TC weakened despite the increased in energy gain?

 ∇ MSF Vertical Profile ave(z=0.02, 2km) MYNN impact: 0.010787 Cd impact: 0.224254 4.E+03 3.E+03 3.E+03 2.E+03 MYNN < Deardorff 2.E+03 • MYNN - DF FIG. 5 Cd1.5 >Deardo -0.6 - 0.3 + 0.3 + 0.3 = 0.3Averaged within 300 km radius ; ave(t=0,t=12) □ *The resulting secondary circulation is* stronger with the increased in the *surface momentum exchange* $coefficient \rightarrow a gradient force (Cd1.5).$ □ Near the surface the MYNN scheme produces a negative tendency. Cd 1.5 < Deardorff \Box On the average (z=0.02 to 2km), Cd1.5 7 MSF (1 x 10⁷ kgs⁻¹) and MYNN strengthened the MSF. RADIUS (km FIG. 6 Wind Inflow (v_r) Tangential Winds (v_{ϕ}) — DF —DF -MYNN — MYNN 1.5 ---Cd1.5 **---**Cd1.5 0.5 v_r (m/s) $_{10} v_{\phi}$ (m/s) 12 FIG. 7 \Box Near the surface, the resulting v_{ϕ} increases with MYNN but weakens with Cd1.5 during EI. \Box The resulting inflow v_r , on the other hand, decreases with MYNN but increases with Cd1.5. □ *The reduction of surface cyclonic wind for Cd1.5 case disrupted the gradient wind balance* resulting to a stronger inflow of agradient wind (Montgomery and Smith 2010). Energy Conversion from APE α MSF $EC \approx \left\langle v_r \frac{\partial \Phi}{\partial r} \right\rangle = 2\pi \int \frac{1}{\rho} \frac{\partial \Phi}{\partial r} MSF(r,z) dr$

MSF: Secondary Circulation during Early Intensification



mean axisymmetric secondary circulation MSF \rightarrow EC \rightarrow Intensification? (Gradient wind balance)

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