Estimating Air-Sea Energy and Momentum Exchanges inside Tropical Cyclones Using the Fetch- and Duration-Limited Wave Growth Properties

> Paul A. Hwang<sup>1</sup> and Yalin Fan<sup>2</sup> <sup>1</sup>Remote Sensing Division <sup>2</sup>Oceanography Division U. S. Naval Research Laboratory

### Fetch- or duration-limited wave growth



# Fetch- or duration-limited growth $\left(\frac{H_s^2 g^2}{16U_{10}^4}\right) = 1.27 \times 10^{-8} \left(\frac{t_d g}{U_{10}}\right)^1$ $\left(\frac{2\pi U_{10}}{T_p g}\right) = 2.94 \left(\frac{t_d g}{U_{10}}\right)^{-0.34}$



$$\left(\frac{H_s^2 g^2}{16U_{10}^4}\right) = 2.94 \times 10^{-3} \left(\frac{2\pi U_{10}}{T_p g}\right)^{-3.42}$$

Hwang and Wang 2004 JPO

Wind-wave triplets 
$$(U_{10}, H_s, T_p)$$

$$H_{s} = 8.10 \times 10^{-4} U_{10}^{1.19} x_{\eta x}^{0.405}$$
$$T_{p} = 9.28 \times 10^{-2} U_{10}^{0.526} x_{\omega x}^{0.237}$$

$$H_{s} = 1.55 \times 10^{-4} U_{10}^{1.47} t_{\eta t}^{0.531}$$
$$T_{p} = 3.53 \times 10^{-2} U_{10}^{0.690} t_{\omega t}^{0.310}$$

### Hwang 2016 JPO

$$U_{10} = 392.95 H_s^{0.681} t_{\eta t}^{-0.362}$$
$$U_{10} = 127.71 T_p^{1.450} t_{\omega t}^{-0.450}$$

$$U_{10} = 397.46H_s^{0.841}x_{\eta x}^{-0.341}$$
$$U_{10} = 91.49T_p^{1.900}x_{\omega x}^{-0.450}$$

$$\frac{d\chi(\omega)}{dt} = Q_{in} + Q_{nl} + Q_{dis} \qquad Q_{in} = \gamma_{in} s \omega \chi(\omega)$$
Energy exchange  

$$E_{t} = \frac{dE}{dt} = \int_{0}^{\infty} \rho_{w} g Q_{in} d\omega = f_{E} (U_{10}, H_{s}, T_{p})$$

$$E_{t} = \alpha_{E} \rho_{a} U_{10}^{3}; \ \alpha_{E} = 0.20 \omega_{\#}^{3.3} \eta_{\#} \qquad \left(\frac{H_{s}^{2}g^{2}}{16U_{10}^{4}}\right) = 2.94 \times 10^{-3} \left(\frac{2\pi U_{10}}{T_{p}g}\right)^{-342}$$
Hwang and Sletten 2008 *JGR*  

$$\tau = C_{d} \rho_{a} U_{\infty}^{2} = C_{\lambda/2} \rho_{a} U_{\lambda/2}^{2} = C_{10} \rho_{a} U_{10}^{2}$$

$$E_{\lambda/2} = 1.22 \times 10^{-2} (u_{s} \omega_{p} / g)^{0.704}$$

$$\int_{C_{\lambda/2}} = 1.289 \times 10^{-3} \omega_{\#}^{0.815} 1$$
Hwang 2004, 2005 *JO*  

$$\chi_{E}(\omega) = \chi_{M}(\omega) c(\omega); \ M_{t} = f_{M}(U_{10}, H_{s}, T_{p})$$

$$M_{t} = \alpha_{M} \rho_{a} U_{10}^{2}; \ \alpha_{M} = 0.40 \omega_{\#}^{4.3} \eta_{\#}$$
Momentum exchange

# Fetch- or duration-limited growth



 $H_{s} = 8.10 \times 10^{-4} U_{10}^{1.19} x_{nx}^{0.405}$ 

 $T_p = 9.28 \times 10^{-2} U_{10}^{0.52} (x_{or}^{0.23})$ 

$$\left(\frac{H_s^2 g^2}{16U_{10}^4}\right) = 2.94 \times 10^{-3} \left(\frac{2\pi U_{10}}{T_n g}\right)^{-3.42}$$

Hwang and Wang 2004 JPO

$$\left(\frac{2\pi O_{10}}{T_p g}\right) = 2.94 \begin{pmatrix} t_d g \\ U_{10} \end{pmatrix}$$
$$H_s = 1.55 \times 10^{-4} U_{10}^{1.47} t_{\eta t}^{0.531}$$

 $\left(\frac{H_s^2 g^2}{16U_{10}^4}\right) = 1.27 \times 10^{-8} \left(\frac{t_d g}{U_{10}}\right)$ 

 $T_p = 3.53 \times 10^{-2} U_{10}^{0.690} t_{\omega t}^{0.310}$ 



Hwang 2016 JPO

### Hurricane fetch and duration

Wind-wave triplets  $(U_{10}, H_s, T_p)$ 

$$x_{\eta x} = 4.24 \times 10^7 U_{10}^{-2.93} H_s^{2.47}$$
$$x_{\omega x} = 2.29 \times 10^4 U_{10}^{-2.22} T_p^{4.22}$$

$$t_{\eta t} = 1.75 \times 10^4 U_{10}^{-2.77} H_s^{3.77}$$
$$t_{\omega t} = 4.81 \times 10^4 U_{10}^{-2.22} T_p^{3.22}$$



$$x_{\eta x}(r,\phi) = s_{\eta x}(\phi)r + I_{\eta x}(\phi)$$
$$x_{\omega x}(r,\phi) = s_{\omega x}(\phi)r + I_{\omega x}(\phi)$$
$$t_{\eta t}(r,\phi) = s_{\eta t}(\phi)r + I_{\eta t}(\phi)$$
$$t_{\omega t}(r,\phi) = s_{\omega t}(\phi)r + I_{\omega t}(\phi)$$
s and I are functions of  $\phi$ ...

Hwang 2016 JPO Hwang and Walsh 2016 JPO (in press)











## Hurricane Wind Waves and Air-Sea Interaction

- Waves inside hurricanes are fetch- and duration-limited
- Full set of (U<sub>10</sub>, H<sub>s</sub>, T<sub>p</sub>) can be derived knowing only one, given fetch or duration
- Wind sea: energy and momentum exchange = f(U<sub>10</sub>, H<sub>s</sub>, T<sub>p</sub>), simple parameterization functions

 $E_{t} = \alpha_{E} \rho_{a} U_{10}^{3}; \ \alpha_{E} = 0.2 \omega_{\#}^{3.3} \eta_{\#}; \ M_{t} = \alpha_{M} \rho_{a} U_{10}^{2}; \ \alpha_{M} = 0.4 \omega_{\#}^{4.3} \eta_{\#}$ 

 Temporal and spatial evolution of hurricane wind wave development from time series of hurricane wind field



#### ESTIMATING AIR-SEA ENERGY AND MOMENTUM EXCHANGES INSIDE TROPICAL CYCLONES USING THE FETCH-AND DURATION-LIMITED WAVE GROWTH PROPERTIES

Paul A. Hwang<sup>1</sup> and Yalin Fan<sup>2</sup>

<sup>1</sup>Remote Sensing Division, Naval Research Laboratory, Washington, DC, USA <sup>2</sup>Oceanography Division, Naval Research Laboratory, Stennis Space Center, MS

#### 1. INTRODUCTION

For wind-generated waves, the wind-wave triplets (reference wind speed  $U_{10}$ , significant wave height  $H_{s}$ , and frequency spectral peak wave period  $T_{p}$ ) are intimately connected through the fetch- or duration-limited wave growth functions (e.g., Hasselmann et al. 1973; Donelan et al. 1985; Kahma and Calkoen 1994; Young 1999; Hwang and Wang 2004). The full set of the wind-wave triplets can be obtained knowing only one of the three variables combined with the fetch  $x_f$  (duration  $t_d$ ) information using the fetch-limited (duration-limited) wave growth functions, which constitute a pair of dimensionless equations describing the growth of wave height and wave period as a function of fetch or duration. The frequency-integrated air-sea energy and momentum exchange rates  $E_t$  and  $M_t$  are functions of the wind-wave triplets and can be quantified with the wind-wave growth functions (Hwang and Sletten 2008; Hwang 2009; Hwang and Walsh 2016).

Previous studies have shown that wave development inside hurricanes follows essentially the same growth functions established for steady wind forcing conditions (Young 1988, 1998, 2006; Hwang 2016; Hwang and Walsh 2016). Here we present the analysis of wind-wave triplets collected inside Hurricanes Bonnie 1998 and Ivan 2004 at category 2 to 4 stages using the airborne scanning radar altimeter (SRA) combined with the NOAA HRD hurricane wind velocity product (Powell et al. 1996; Wright et al. 2001; Moon et al. 2003; Fan et al. 2009). The data yield the windwave triplets inside the hurricanes along several (between 6 and 11) transects radiating from the hurricane center.

Using the wind-wave triplets from the 4 hurricane scenes (Table 1), a parameterization model is formulated for the effective fetch and duration at any location inside a hurricane. The fetch and duration model is applied to the time series of 2D hurricane wind fields from post analysis to investigate the detailed temporal evolution of the wave field and the associated energy and momentum exchanges over the hurricane coverage area. The spatial distributions of energy and momentum exchanges show considerable asymmetry. Referenced to the hurricane heading, the exchanges on the right half plane of the hurricane are much stronger than those on the left half plane, the right-to-left ratio of about 2:1 to 3:1 is common.

### 2. HARMONIC ANALYSIS OF HURRICANE FETCH AND DURATION

The fetch- and duration-limited wave growth functions connect the wind-wave triplets ( $U_{10}$ ,  $H_s$ ,  $T_p$ ). Many different formulas of growth functions have been reported in the literature, in this study we use the results obtained from least squares fitting of five field experiments conducted under quasi-steady and near-neutral stability conditions, collectively called BHDDB (Burling 1959; Hasselmann et al.

1973; Donelan et al. 1985; Dobson et al. 1989; Babanin and Soloviev 1998), as described in Hwang and Wang (2004):

$$\frac{H_s^2 g^2}{16U_{10}^4} = 6.19 \times 10^{-7} \left(\frac{x_f g}{U_{10}^2}\right)^{0.81}, \ \frac{2\pi U_{10}}{T_p g} = 11.86 \left(\frac{x_f g}{U_{10}^2}\right)^{-0.24}, \ (1)$$
$$\frac{H_s^2 g^2}{16U_{10}^4} = 1.27 \times 10^{-8} \left(\frac{t_d g}{U_{10}}\right)^{1.06}, \ \frac{2\pi U_{10}}{T_p g} = 2.94 \left(\frac{t_d g}{U_{10}}\right)^{-0.34}, \ (2)$$

where *g* is the gravitational acceleration (9.8 m/s<sup>2</sup>).  $H_s^2 g^2 / 16U_{10}^4 = \eta_{rms}^2 g^2 / U_{10}^4 = \eta_{\#}$  represents the wave energy  $E = \rho_w g \eta_{rms}^2$ , and  $2\pi U_{10} / T_p g = U_{10} / c_p = \omega_{\#}$  is dimensionless frequency or inverse wave age;  $\rho_w$  is water density. The wind and wave measurements from the hurricane hunter missions provide the necessary data to calculate the effective fetches and durations for the locations where the wind and wave data are acquired (Hwang 2016; Hwang and Walsh 2016). Explicitly, the fetch and duration are derived by rearranging

the variables in (1) and (2)  

$$x_{\eta x} = 4.24 \times 10^7 U_{10}^{-2.93} H_s^{2.47}, x_{ax} = 2.29 \times 10^4 U_{10}^{-2.22} T_p^{4.22}$$
, (3)

$$t_{\eta t} = 1.75 \times 10^4 U_{10}^{-2.77} H_s^{3.77}, t_{\omega t} = 4.81 \times 10^4 U_{10}^{-2.22} T_p^{3.22} .$$
 (4)

The analyses of Hwang (2016) and Hwang and Walsh (2016) show that  $x_f$  and  $t_d$  can be represented by linear functions of the radial distance *r* from the hurricane center:

$$x_{\eta x} = s_{\eta x}(\phi)r + I_{\eta x}(\phi), \ x_{\omega x} = s_{\omega x}(\phi)r + I_{\omega x}(\phi),$$
(5)

$$t_{\eta t} = s_{\eta t} \left( \phi \right) r + I_{\eta t} \left( \phi \right), \quad t_{\omega t} = s_{\omega t} \left( \phi \right) r + I_{\omega t} \left( \phi \right), \tag{6}$$

where  $\phi$  is the azimuth angle referenced to the hurricane heading, positive counterclockwise (CCW). The intercepts *I* and slopes *s* of the linear functions in (5) and (6) from processing the four hurricane scenes are shown in Fig. 1; these fitting parameters can be expressed in Fourier series:

$$q = \sum_{n=0}^{N} \left( a_{n,q} \cos n\phi + b_{n,q} \sin n\phi \right), \tag{7}$$

where *q* can be  $s_{\eta x}$ ,  $I_{\eta x}$ ,  $s_{\omega x}$ ,  $I_{\omega x}$ ,  $s_{\eta \eta}$ ,  $I_{\eta \eta}$ ,  $s_{\omega t}$  or  $I_{\omega t}$ . The curves computed with *N*=1, 2 and 3 are shown with continuous lines in Fig. 1; the data are well represented by (7) with *N*=3.

The harmonics  $a_{nq}$  and  $b_{nq}$  display a systematic quasilinear variation with the radius of maximal wind speed  $r_m$  (Fig. 2)

$$Y = p_{1Y}r_m + p_{2Y} , (8)$$

where Y represents  $a_{n,q}$  and  $b_{n,q}$  in (7). The fitting coefficients  $p_{1Y}$  and  $p_{2Y}$  are listed in Table 2 (the headers s\_ex, l\_ex, s\_ox, l\_ox, s\_et, l\_et, s\_ot and l\_ot represent  $s_{nx}$ ,  $I_{nx}$ ,  $s_{ox}$ ,  $I_{ox}$ ,  $s_m$ ,  $I_m$ ,  $s_{ot}$  and  $I_{ot}$ .

#### 3. DERIVATION OF WAVE PARAMETERS AND AIR-SEA EXCHANGES FROM HURRICANE WIND FIELD

As discussed in Hwang (2016) and Hwang and Walsh (2016), for wind generated waves, ( $U_{10}$ ,  $H_s$ ,  $T_p$ ) are

connected by the growth functions and the full set can be computed knowing only one of the three, coupled with the fetch or duration input. At the present, the hurricane wind field is the most available product so we illustrate the computation of wave parameters and the derived quantities such as  $E_t$  and  $M_t$  from the wind input. Rewriting (1), the equations for  $H_s$  and  $T_p$  with  $U_{10}$  and fetch inputs are

$$H_{sw} = 8.10 \times 10^{-4} U_{10}^{1.19} x_{\eta_x}^{0.405}, T_{pw} = 9.28 \times 10^{-2} U_{10}^{0.526} x_{\omega x}^{0.237}.$$
 (9)

Similarly, for  $U_{10}$  and duration inputs (2),

$$H_{_{SW}} = 1.55 \times 10^{-4} U_{10}^{1.47} t_{\eta t}^{0.531}, T_{_{PW}} = 3.53 \times 10^{-2} U_{10}^{0.690} t_{_{ot}}^{0.310}.$$
 (10)

The subscript *w* is appended to the wave variables in (9) and (10) to emphasize that the wind-sea portion is obtained with the fetch-limited wave growth functions. The fetch and duration in those equations are computed from the harmonic model described in section 2. Once ( $U_{10}$ ,  $H_s$ ,  $T_p$ ) are available, the  $\omega_{\#}$ ,  $\eta_{\#}$ ,  $M_t$  and  $E_t$  can be calculated (Hwang and Sletten 2008; Hwang and Walsh 2016):

$$E_{t} = \alpha_{E} \rho_{a} U_{10}^{3}; \ \alpha_{E} = 0.20 \omega_{\#}^{3.3} \eta_{\#} \\ M_{t} = \alpha_{M} \rho_{a} U_{10}^{2}; \ \alpha_{M} = 0.40 \omega_{\#}^{4.3} \eta_{\#}$$
(11)

where  $\rho_a$  is the air density. Fig. 3 shows an example (approximately the middle point of case I14, 00Z 15 Sep 2004) of using the HWind  $U_{10}$  to compute the wave parameters and air-sea exchange rates.

If a long time series of hurricane wind field is available, the corresponding temporal variation of the wave field and air-sea exchanges can also be computed. Fig. 4 presents an application applied to Hurricane Ivan 2004 (Fan et al. 2009), showing (a) the energy exchange integrated over the hurricane coverage area (here defined as the circle with a 250-km radius), (b) the relative energy exchange contributions of the left- and right half-planes, (c) the momentum exchange integrated over the hurricane coverage area, and (d) the relative momentum contributions of the left- and right half-planes. These results show the temporal evolution and the asymmetric spatial distributions of energy and momentum exchanges, the right-to-left ratio of 2:1 to 3:1 is common.

#### 4. SUMMARY

The wave fields inside hurricanes behave similarly to those generated by steady fetch- or duration-limited wind forcing conditions (Young 1998; 2006; Hwang 2016; Hwang and Walsh 2016). The full set of the wind-wave triplets ( $U_{10}$ ,  $H_s$ ,  $T_p$ ) can be calculated with the fetch- or duration-limited growth functions knowing only one of three variables and accompanied with the fetch or duration information. Using the  $U_{10}$ ,  $H_s$  and  $T_p$  measured inside four different hurricane scenes, we present a fetch and duration model for any location inside a hurricane. The  $E_t$  and  $M_t$  of a wind wave system are functions of the wind-wave triplets, and can be calculated as well. The procedure is very useful for investigating the spatial distribution and temporal evolution of the air-sea interaction processes inside hurricanes.

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Table 1. Some basic information of the four datasets, collected in hurricane Bonnie 1998 and Ivan 2004, used for the analysis in this paper.

File ID	B24	109	l12	I14	
Start time UTC	8/24/1998 20:29	9/9/2004 16:15	9/12/2004 10:39	9/14/2004 20:09	
End time UTC	8/25/1998 1:44	9/9/2004 20:10	9/12/2004 15:41	9/14/2004 2:49	
HRD U10 max (m/s)	44.4	59.4	55.4	61.6	
rm (km)	74.0	13.0	17.0	42.0	
V_a (m/s)	4.5	4.5	3.9	4.5	
theta_a (degN)	13.0	62.0	69.0	25.0	
SRA U10 min	1.4	1.8	0.8	1.2	
SRA U10 max	45.7	74.0	59.9	69.6	
SRA Hs min	4.4	1.6	2.9	3.6	
SRA Hs max	10.9	12.7	12.0	13.1	
SRA Tp min	8.0	5.8	8.2	8.9	
SRA Tp max	13.3	15.2	13.8	14.4	
# SRA spectra	233	376	456	600	

Table 2. Linear fitting coefficients of fetch and duration harmonics as a function of radius of maximum winds  $Y = p_{1Y}r_m + p_{2Y}$  (8), where Y represents  $a_{n,q}$  and  $b_{n,q}$  in (7); see text for further explanation.

	q	s_ex	I_ex	s_ox	I_ox	s_et	I_et	s_ot	I_ot
	a0,q								
p1		-1.40E-02	1.95E+00	-3.04E-02	3.22E+00	-7.88E-04	1.07E-01	-1.47E-03	1.60E-01
p2		1.57E+00	-3.54E+01	3.27E+00	-1.42E+02	9.04E-02	-2.42E+00	1.58E-01	-6.73E+00
	a1,q								
p1		-1.02E-02	1.26E+00	-1.34E-02	1.09E+00	-5.17E-04	5.76E-02	-6.11E-04	4.81E-02
p2		5.10E-01	-4.77E+01	9.85E-01	-5.22E+01	2.39E-02	-2.23E+00	3.99E-02	-2.18E+00
	b1,q								
p1		5.21E-03	-8.23E-01	-1.27E-02	1.16E+00	1.54E-04	-2.81E-02	-5.75E-04	5.30E-02
p2		-6.87E-02	3.02E+01	1.51E+00	-1.02E+02	5.25E-03	8.52E-01	6.90E-02	-4.50E+00
	a2,q								
p1		5.55E-03	-4.11E-01	1.23E-02	-1.10E+00	3.22E-04	-2.31E-02	6.14E-04	-5.28E-02
p2		-4.79E-01	2.25E+01	-1.01E+00	6.93E+01	-2.63E-02	1.27E+00	-4.91E-02	3.24E+00
	b2,q								
p1		-3.32E-03	-1.52E-01	-2.00E-02	1.51E+00	-2.88E-04	5.96E-03	-9.14E-04	6.68E-02
p2		2.11E-01	8.86E+00	1.26E+00	-9.80E+01	1.58E-02	-2.16E-01	5.43E-02	-4.14E+00
	a3,q								
p1		6.44E-03	-5.32E-01	1.45E-02	-9.21E-01	3.19E-04	-2.61E-02	6.23E-04	-3.91E-02
p2		-3.29E-01	2.85E+01	-8.47E-01	5.45E+01	-1.52E-02	1.31E+00	-3.41E-02	2.18E+00
	b3,q								
p1		1.91E-03	-2.28E-01	-5.70E-03	1.01E+00	1.09E-04	-9.53E-03	-1.37E-04	3.62E-02
p2		-2.25E-01	2.11E+01	1.80E-01	-4.71E+01	-1.06E-02	8.72E-01	3.06E-03	-1.68E+00



Fig. 1. Slope  $s_q$  and intercept  $I_q$  of  $q = s_q(\phi)r + I_q(\phi)$ , where q can be  $x_{\eta x}, x_{\omega x}, t_{\eta r}$  or  $t_{\omega r}$ . The results obtained from datasets B24, I09, I12 and I14 (Table 1) are shown in the 4 quadrants, respectively upper left (UL), upper right (UR), lower left (LL) and lower right (LR). The solid, dashed and dotted curves corresponding to the first, second and third order Fourier series are also superimposed in each panel. For each dataset, s and I are plotted in (a) and (b) for  $x_{\eta x}$ , (c) and (d) for  $x_{\omega x}$ , (e) and (f) for  $t_{\eta t}$ , and (g) and (h) for  $t_{\omega t}$ .



Fig. 2. Harmonic representations  $q = \sum_{n=0}^{\infty} (a_{n,q} \cos n\phi + b_{n,q} \sin n\phi)$  of the slopes and intercepts of the linear functions

defining the fetch and duration, where q can be  $s_{\eta x}$ ,  $I_{\eta x}$ ,  $s_{\omega x}$ ,  $I_{\eta x}$ ,  $s_{\eta t}$ ,  $I_{\eta t}$ ,  $s_{\omega t}$  or  $I_{\omega t}$ . The harmonics  $a_{n,q}$  and  $b_{n,q}$  show linear dependence on  $r_m$ , n=0, 1, 2, 3 ( $b_{0,q}=0$ ). (UL): Slopes of fetch for wave height and wave period: (a)  $a_{n,s\eta x}$ , (b)  $b_{n,s\eta x}$ , (c)  $a_{n,s\omega x}$ , and (d)  $b_{n,s\omega x}$ ; (UR): Intercepts of fetch for wave height and wave period: (a)  $a_{n,l\eta x}$ , (b)  $b_{n,l\eta x}$ , (c)  $a_{n,l\omega x}$ , and (d)  $b_{n,l\omega x}$ ; (LL): Slopes of duration for wave height and wave period: (a)  $a_{n,l\eta x}$ , (b)  $b_{n,l\eta x}$ , (c)  $a_{n,l\omega x}$ , and (c)  $b_{n,s\omega x}$ , (LR): Intercepts of duration for wave height and wave period: (a)  $a_{n,l\eta x}$ , (b)  $b_{n,l\eta x}$ , (c)  $a_{n,l\omega x}$ , (LR): Intercepts of duration for wave height and wave period: (a)  $a_{n,l\eta x}$ , (b)  $b_{n,l\eta x}$ , (c)  $a_{n,l\omega x}$ , (c



Fig. 3. An example illustrating the derivation of wave and air-sea interaction parameters using the nurricane wind tield as an input, shown here are (a) input  $U_{10}$ , and outputs of (b)  $H_{sw}$ , (c)  $T_{pw}$ , (d)  $\omega_{\#}$ , (e)  $E_t$ , (f)  $M_t$ ; subscript w emphasizes that the wave parameters derived from the fetch- and duration-growth functions are the wind sea components.



Fig. 4. The temporal evolution of the energy and momentum exchanges of Hurricane Ivan 2004 calculated from the time seried of the HWind time series from 06 to 17 Sep: (a)  $E_t$  integrated over the hurricane coverage area of 250-km radius, and the left and right half contributions; (b) the fractional  $E_t$  contributions of the left and right half planes, (c)  $M_t$  integrated over the hurricane coverage area of 250-km radius, and the left and right half contributions; and (d) the fractional  $M_t$  contributions of the left and right half planes.