



# SURFACE COOLING DUE TO PRECIPITATION OVER THE TROPICAL OCEAN

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## 1 - INTRODUCTION:

- Precipitation cools the ocean surface because the temperature of the raindrops is lower than the temperature of the surface. However, this cooling term due to precipitation ( $Q_p$ ) is often not included in the models.
- $Q_p$  can be as high as  $200 \text{ W m}^{-2}$  affecting significantly the skin temperature of the ocean (Gosnell et al. 1995).
- As the skin temperature provides boundary conditions to the atmosphere above it, this might be critical to the processes within the planetary boundary layer (PBL, Chen and Dudhia 2001).

## 2 - OBJECTIVES:

- Provide a documentation of the spatio-temporal variability of  $Q_p$  over the tropical oceans using a variety of observational datasets.
- Discuss the implementation of this process into a simplified 3D ocean model coupled to the WRF model
- Explore the role of  $Q_p$  on large rain events

## 3 - DATA:

**Rain:** TRMM 3B-42,  $0.25^\circ \times 0.25^\circ$ , 3-hrly, from 1998 to 2013

**Pressure:** NCEP Reanalysis 2 (6-hrly, data provided by the NOAA/OAR/ESRL PSD).

**Latent and Sensible Heat, specific humidity and temperature at 2m and surface temperature (skin temperature):** Objectively Analyzed Air-Sea Fluxes (OAFlux,  $1.0^\circ \times 1.0^\circ$ )

**Buoy:** TAO buoy ( $0^\circ\text{N}$ ,  $165^\circ\text{E}$ ) data for December 2006 to compare different components of the surface fluxes with  $Q_p$ .

**Model initial and boundary condition:** ERA-Interim reanalysis. The sea surface temperature data (RTG\_SST) is from NCEP/MMBA.

## 4 - MODEL:

- 2-way nested domains using the **WRF 3.7 model**
- The outer (inner) domain has a grid spacing of 30 km (10 km).
- $Q_p$  and a fresh water input were added to a 3D simplified ocean model (PWP) coupled to the WRF model (Price et al 1994; Price et al 1986).

## 5 - METHODOLOGY:

The  $Q_p$  is given by (Gosnell et al. 1995):

$$Q_p = C_W R (T_0 - T_r) \quad (1)$$

$C_W$  = specific heat of water ( $4186 \text{ J kg}^{-1} \text{ K}^{-1}$ ),

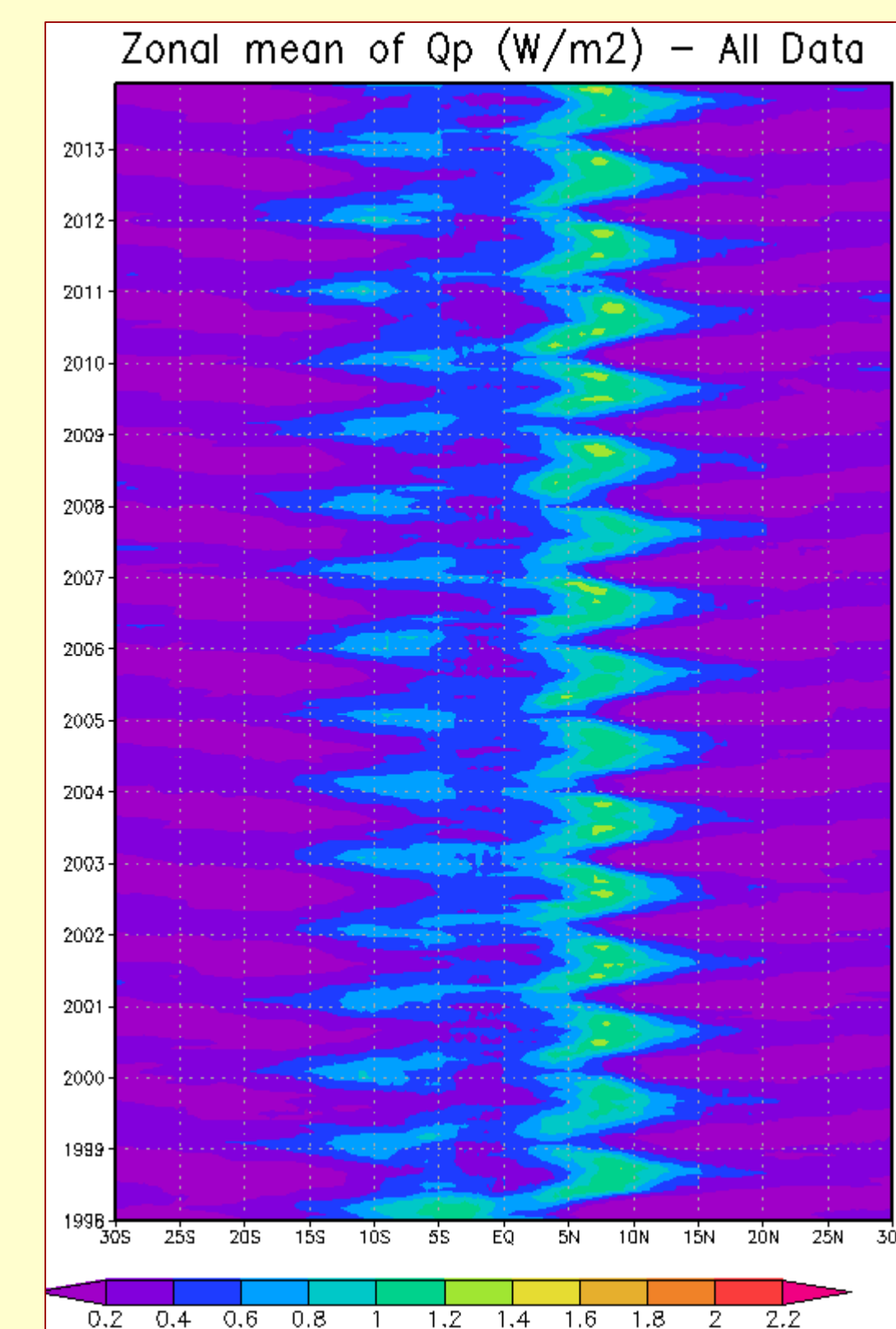
$R$  = rain rate

$T_0$  = bulk SST approximated by the skin temperature

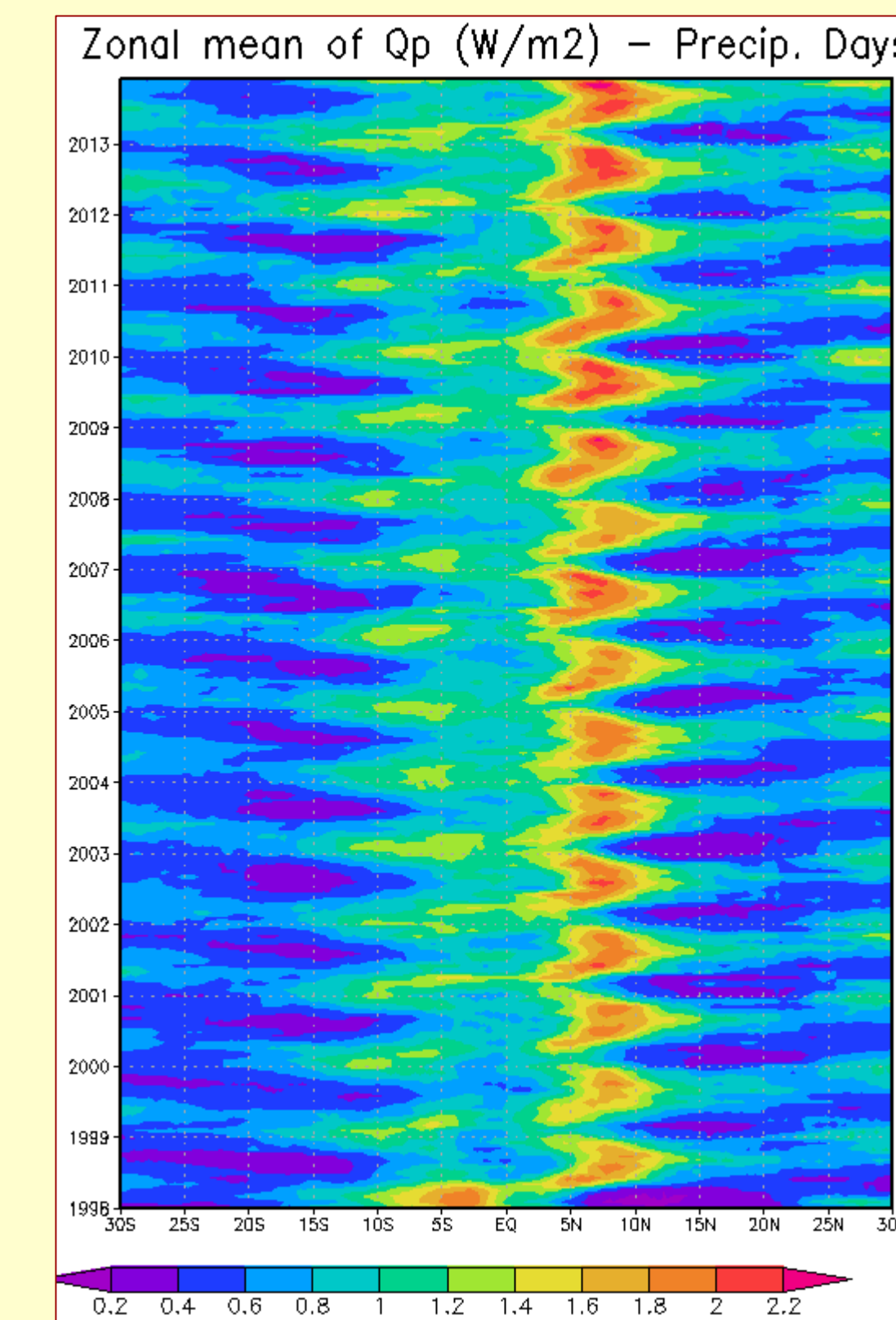
$T_r$  = temperature of raindrops when it reaches the surface.  $T_r$  is approximated by the wet bulb temperature following Stull (2011)

## 6 a – RESULTS AND DISCUSSION

### $Q_p$ Climatology



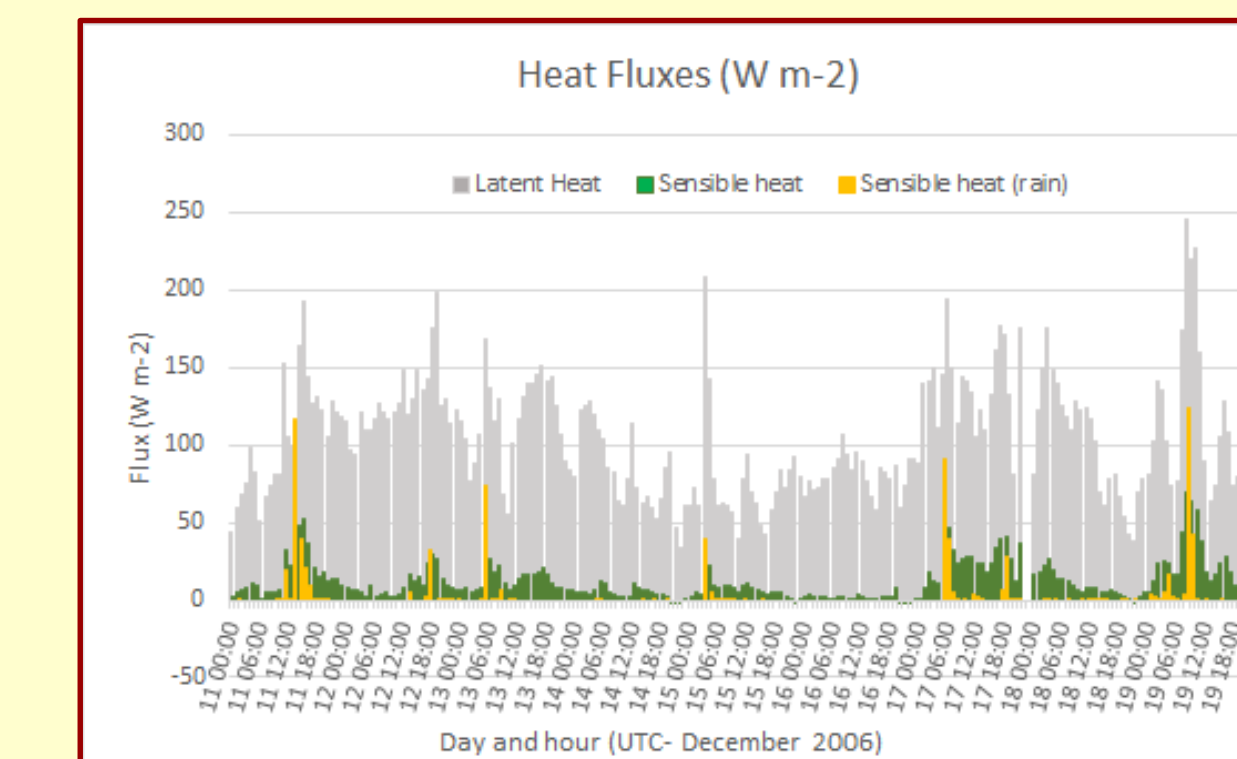
**Figure 1:** Time evolution of monthly averaged  $Q_p$  (zonally averaged; 1998 to 2013 –  $\text{W m}^{-2}$ ) using precipitating and non-precipitating days



**Figure 2:** Time evolution of monthly averaged  $Q_p$  (zonally averaged; 1998 to 2013 –  $\text{W m}^{-2}$ ) using only precipitating days

- Monthly averaged  $Q_p$  was calculated from January 1998 to December 2013 in the tropics ( $30^\circ\text{S}$  to  $30^\circ\text{N}$ ).
- Lowest values of  $Q_p$  correspond to the driest months (February to April)
- Highest  $Q_p$  occurs during the rainiest months (October to December) that coincides with the ITCZ passage from the northern to the southern hemisphere.
- There's also a secondary peak of  $Q_p$  (May to July) that correspond to the Indian Monsoon and the ITCZ.
- Monthly climatology values only go up to  $2 \text{ W m}^{-2}$  but, this average considers precipitating and non-precipitating days.
- When we consider only precipitating days, some areas with values up to  $4 \text{ W m}^{-2}$  appear around  $7^\circ\text{N}$  still following the ITCZ movement.

## 6 b – RESULTS AND DISCUSSION

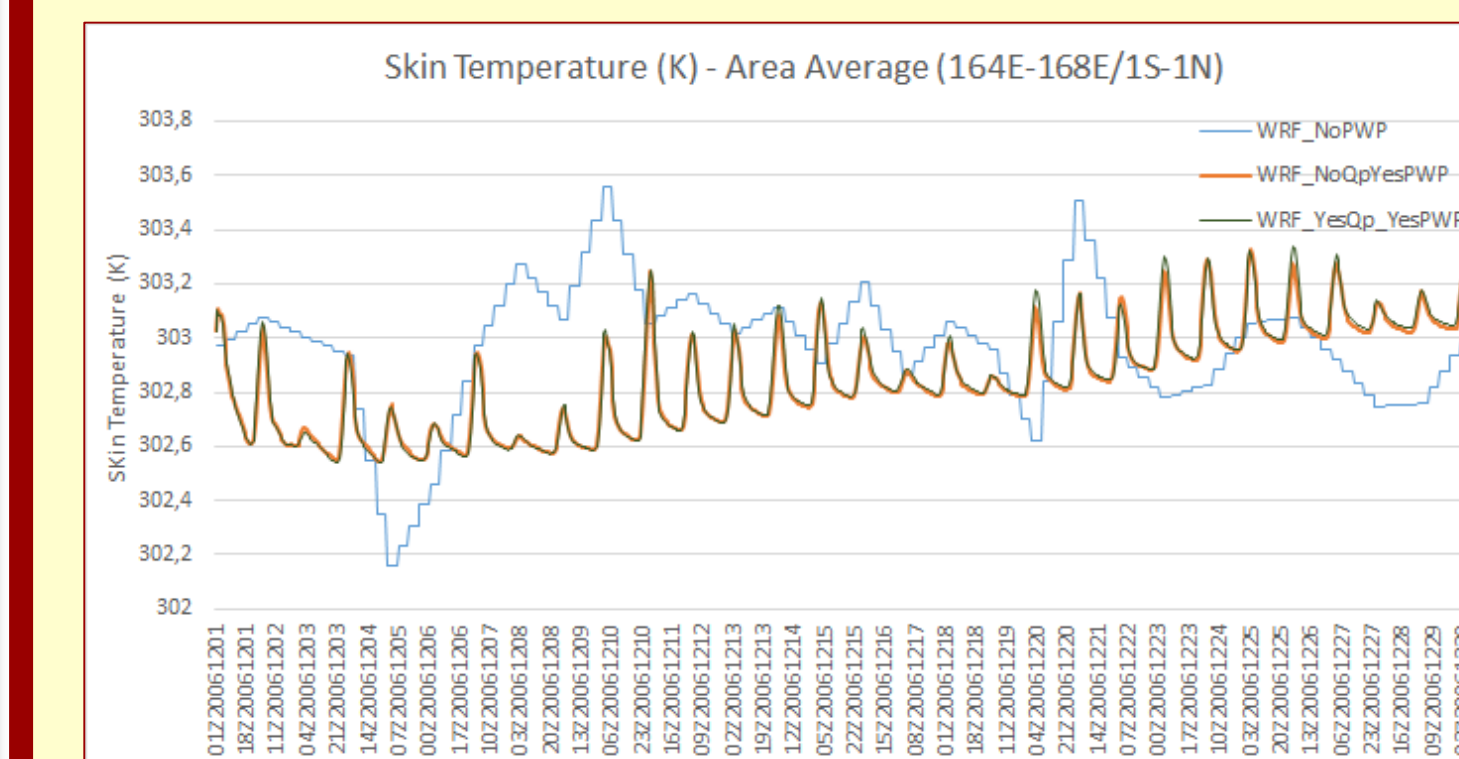


**Figure 3:** TAO buoy - Latent heat flux  $Q_{LH}$  (gray), sensible heat flux  $Q_{SH}$  (green) and  $Q_p$  (yellow) in  $\text{W m}^{-2}$  for 11-19 December 2006 at  $0^\circ\text{N}; 165^\circ\text{E}$

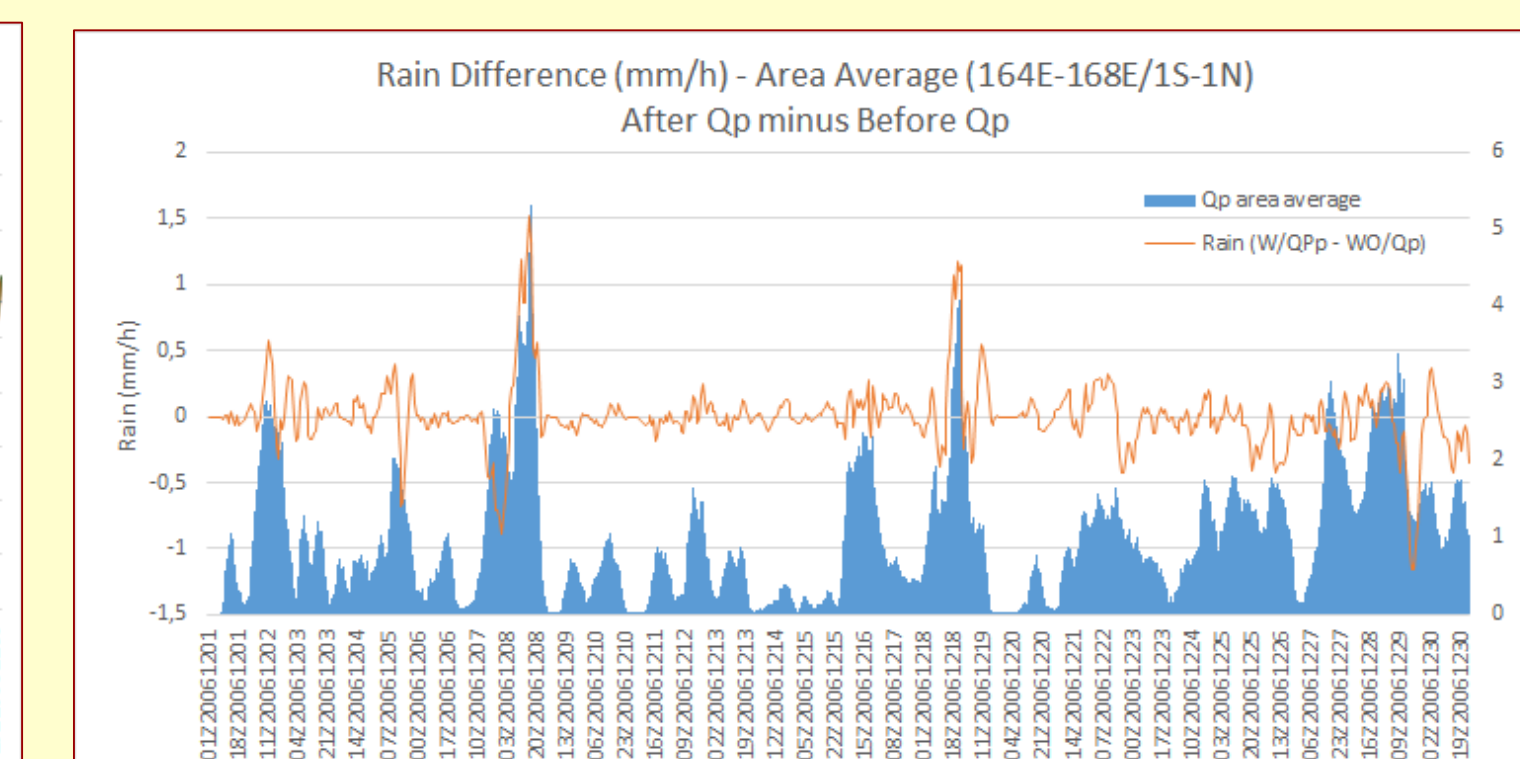
- Hourly observation data with large precipitation shows that  $Q_p$  values can get much larger. (Fig. 3)
- It shows that  $Q_p$  can exceed  $100 \text{ W m}^{-2}$  and be up to 5 times the value of  $Q_{SH}$  and in certain occasions even exceed the value of  $Q_{LH}$

### Simulated $Q_p$

- $Q_p$  term added to the PWP model coupled to the WRF model
- One month simulation over an ocean area (centered on the buoy from Fig. 3), for December 2006. It spans from  $10^\circ\text{S}$  to  $10^\circ\text{N}$  and from  $155^\circ\text{E}$  to  $175^\circ\text{E}$ .



**Figure 4:** Area averaged skin temperature (K) without PWP model (blue), with PWP and with  $Q_p$  (green) and the PWP without  $Q_p$  (orange)



**Figure 5:** Area averaged difference between with and without  $Q_p$ .  $Q_p$  in  $\text{W m}^{-2}$  (blue bars) and skin temperature difference in Kelvin (orange line).

## 7 - CONCLUSIONS:

- $Q_p$  can have the same or larger magnitude  $Q_{SH}$  and even LH.
- Simulations with  $Q_p$  has significant impact on the upper ocean dynamics and thermodynamics and also on the atmosphere. As  $Q_p$  lowers the surface temperature, it tends to delay/reduce the atmospheric convection.

## REFERENCES AND ACKNOWLEDGMENTS:

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