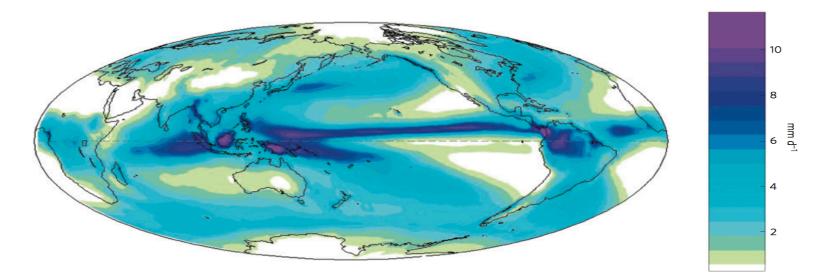
Dynamical constraints on ITCZ migration in planetary atmospheres using a moist GCM

Sean Faulk, Jonathan Mitchell (UCLA) Simona Bordoni (Caltech) AMS Conference on Hurricanes and Tropical Meteorology April 20th, 2016

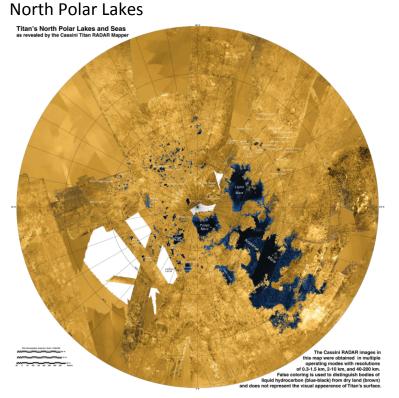
ITCZ on Earth

Observed annual-mean precipitation in 2013



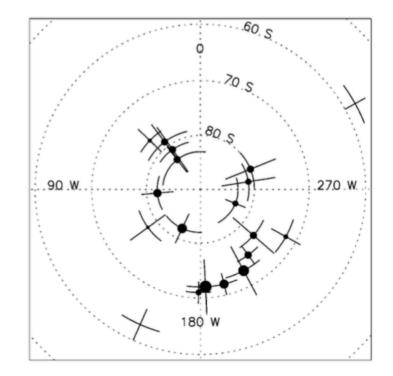
⁽Frierson et al., 2013)

"ITCZ" on Titan reaches pole



(Image source: NASA/ JPL/ ASI/ USGS)

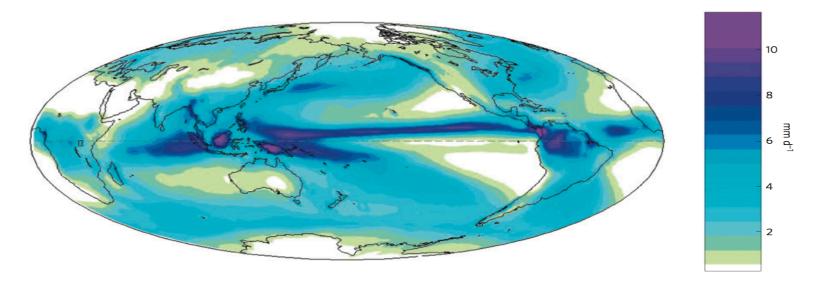
South Polar Clouds



(Bouchez and Brown, 2005)

ITCZ on Earth

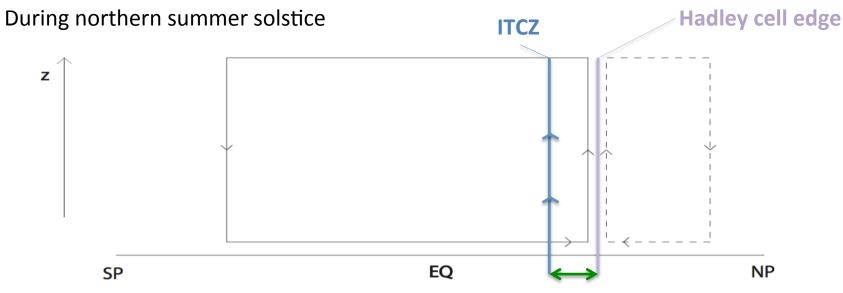
Observed annual-mean precipitation in 2013



⁽Frierson et al., 2013)

Our question: Why doesn't Earth's ITCZ reach summer pole?

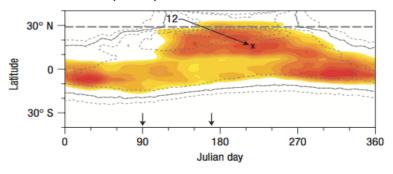
ITCZ is associated with ascending branch of Hadley cell



Hadley cell edge = latitude where meridional flow vanishes; well correlated with energy flux equator (Kang et al., 2008) and maximum zonal MSE in most of our aquaplanet simulations

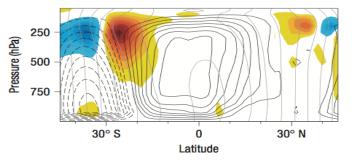
ITCZ = latitude of maximum zonal precipitation

Observations of Asian monsoon sector



Zonal mean precipitation

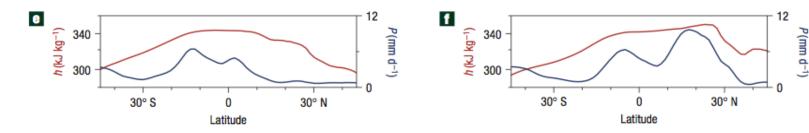
Zonal and time mean streamfunction



(Bordoni and Schneider, 2008)

Moist static energy (MSE) vs. precipitation Equinox

Monsoon onset



Thermodynamic arguments for ITCZ position on Earth

- 1) Convection and uplift generally occur over warmest surface waters, just equatorward of latitude of **maximum MSE** (Privé and Plumb, 2007) or at energy flux equator (Kang et al., 2008, 2009)
 - Continental distribution (Philander et al., 1996) and ocean heat transport (Frierson et al., 2013) help explain ITCZ's northern annualmean position as a result of warmer waters north of equator

- 2) Surface **heat capacity** affects movement of ITCZ on Earth (Bordoni and Schneider, 2008; Donohoe et al., 2013)
 - Often invoked as part of explanation for far-reaching Hadley cells on Titan

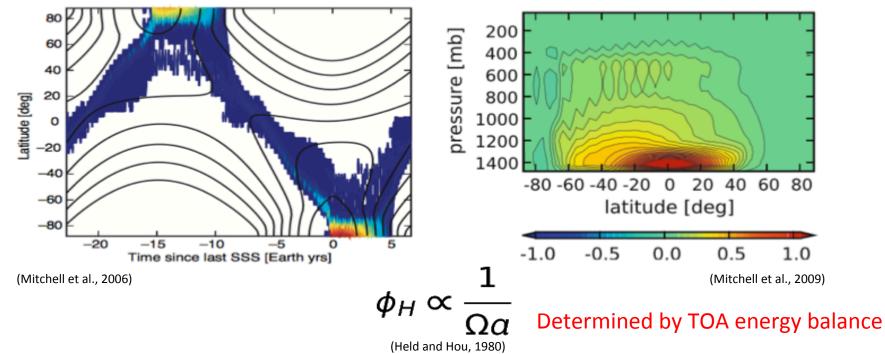
How do dynamics affect ITCZ position?

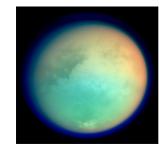
ITCZ on Titan

Titan ITCZ reaches pole during solstice, as part of a global Hadley cell

Zonal mean precipitation

Zonal and time mean streamfunction





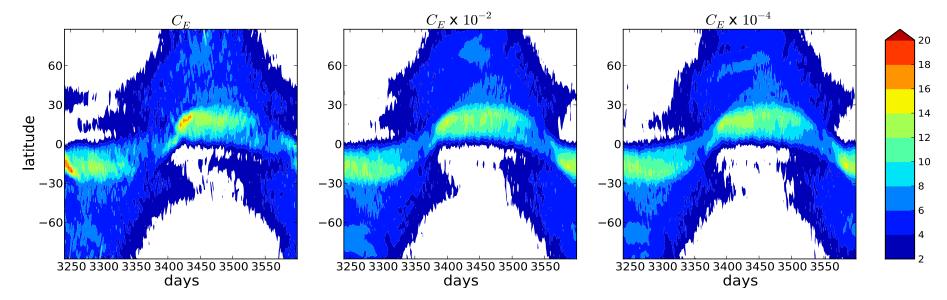
Use a moist convection aquaplanet GCM with a seasonal cycle

- Model developed by Frierson (2006) and Frierson (2007) using FMS spectral core
 - zonally symmetric, slab mixed layer ocean
 - no clouds or water vapor feedbacks
 - moist convection
- 3-D; 64 x 128 horizontal resolution; 25 vertical levels

- 1) Simulations run for decreased heat capacity keeping everything else Earth-like
 - 10-year seasonal runs where top-of-atmosphere insolation is empirical fit to Earth's current insolation (Hartmann, 1994)
- 2) Simulations run under **eternal solstice** forcing for 10 years
- 3) Simulations run with $\Omega = \alpha \Omega_E$ for $\alpha = 4, 2, 1, 1/2, 1/4, 1/6, 1/8, 1/16, 1/32$

Decreased heat capacity has minimal effect on ITCZ migration

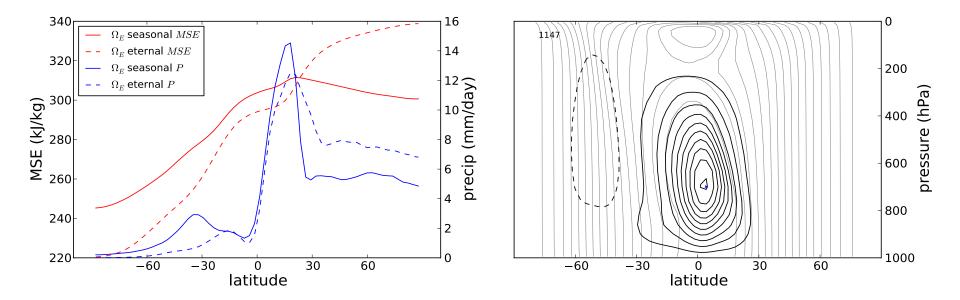
Zonal mean precipitation (mm/day)

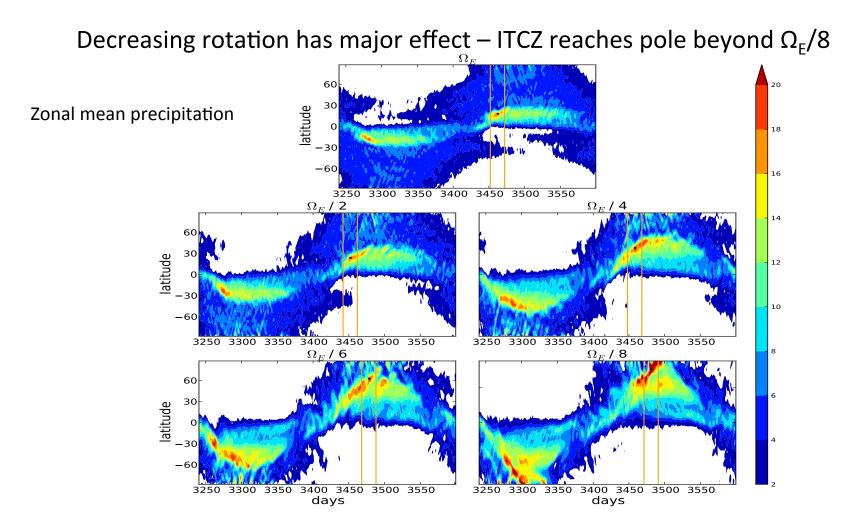


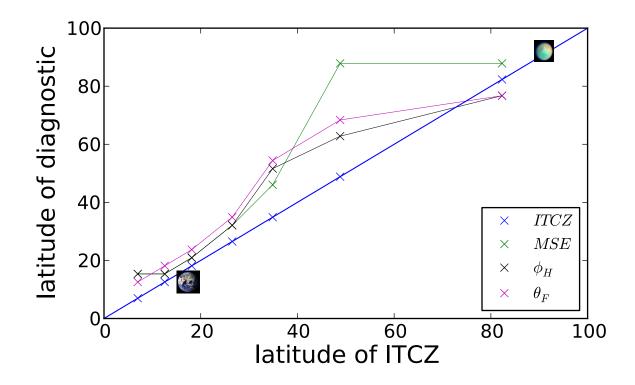
Maximum MSE condition fails in eternal solstice case

Zonal/time mean precipitation and MSE

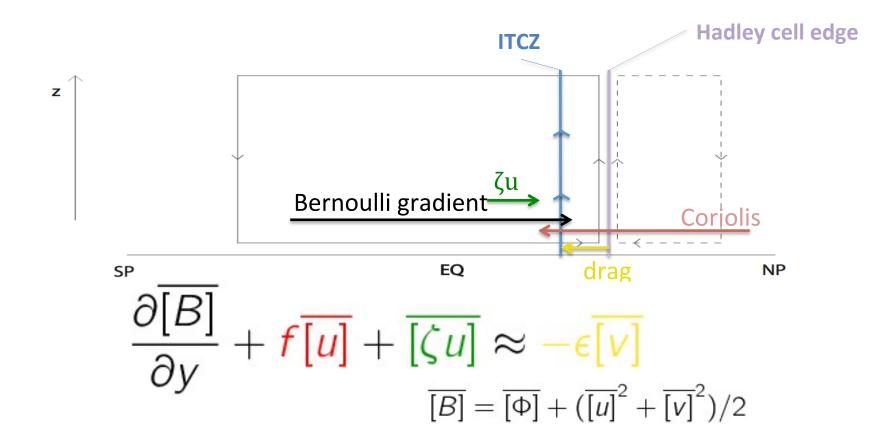
Zonal and time mean streamfunction



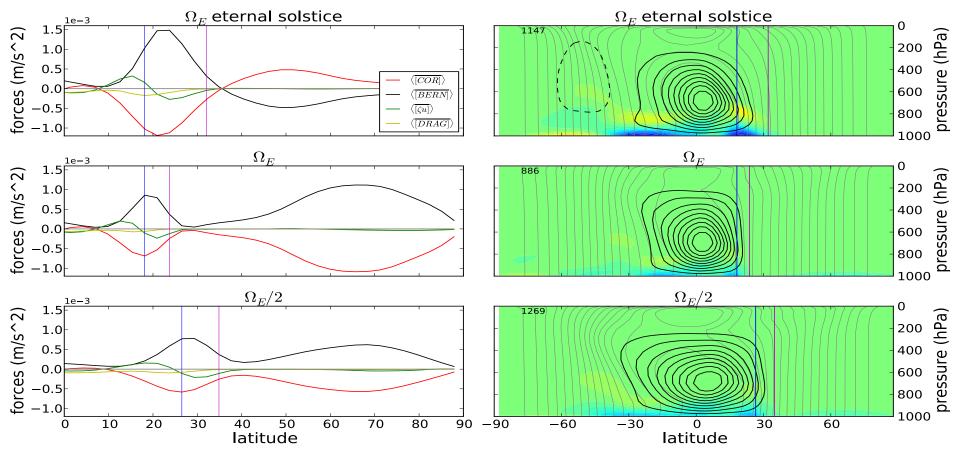




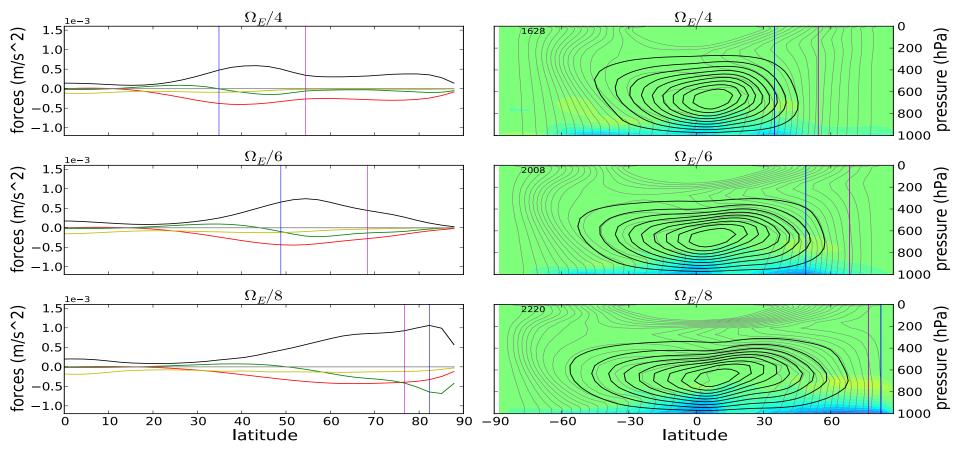
Boundary layer forces



Boundary layer forces



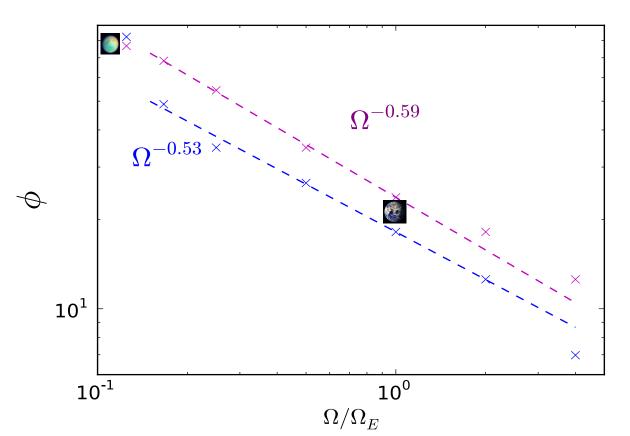
Boundary layer forces



Summary

- Controls on Earth's ITCZ are still unclear and most of the literature has focused on thermodynamic mechanisms as well as on the planet Earth, so we have approached the problem from a planetary perspective and with an emphasis on dynamics
- We find that surface temperature/MSE and heat capacity do not always control ITCZ position, and that rotation rate and/or radius are the dominant controls in an aquaplanet context
- We examine boundary layer dynamics and find the ITCZ migration to be well correlated with friction and other forces.

ITCZ and energy flux equator scale according to power law



 $\phi_H \propto \Omega^{-rac{2}{3}}$

(Caballero et al., 2008)