

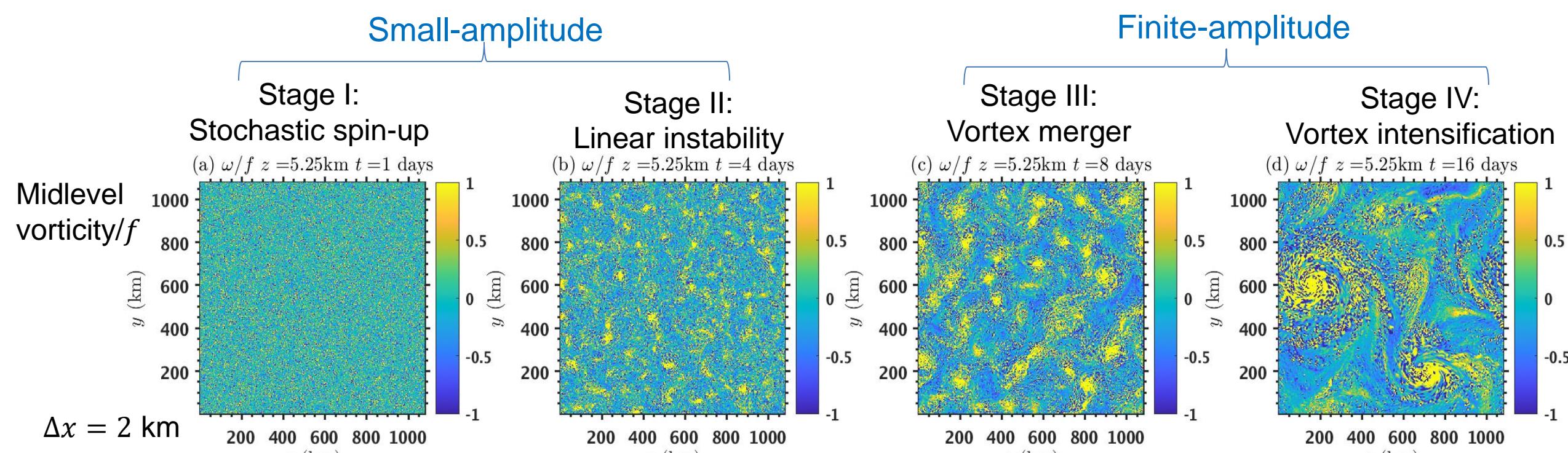
The stochastic spin-up of vorticity in spontaneous tropical cyclogenesis [1]

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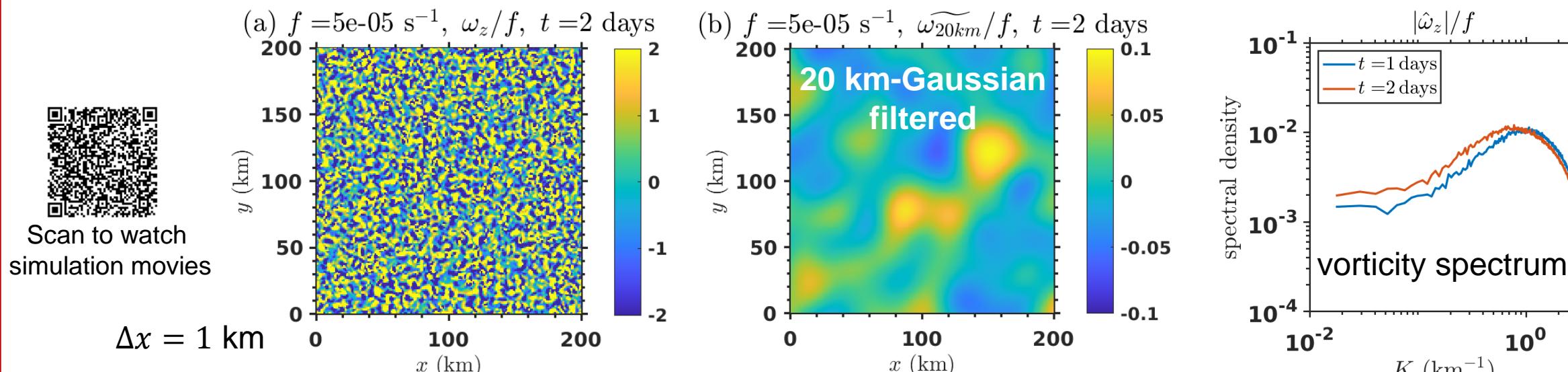
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Background: spontaneous TC genesis [2]

- Doubly periodic domain, uniform SST, Bryan Cloud Model 1, SST=300 K, $f = 10^{-4} \text{ s}^{-1}$, no background shear.
- Doubling longwave radiative feedback to detect the small-amplitude instability [3].

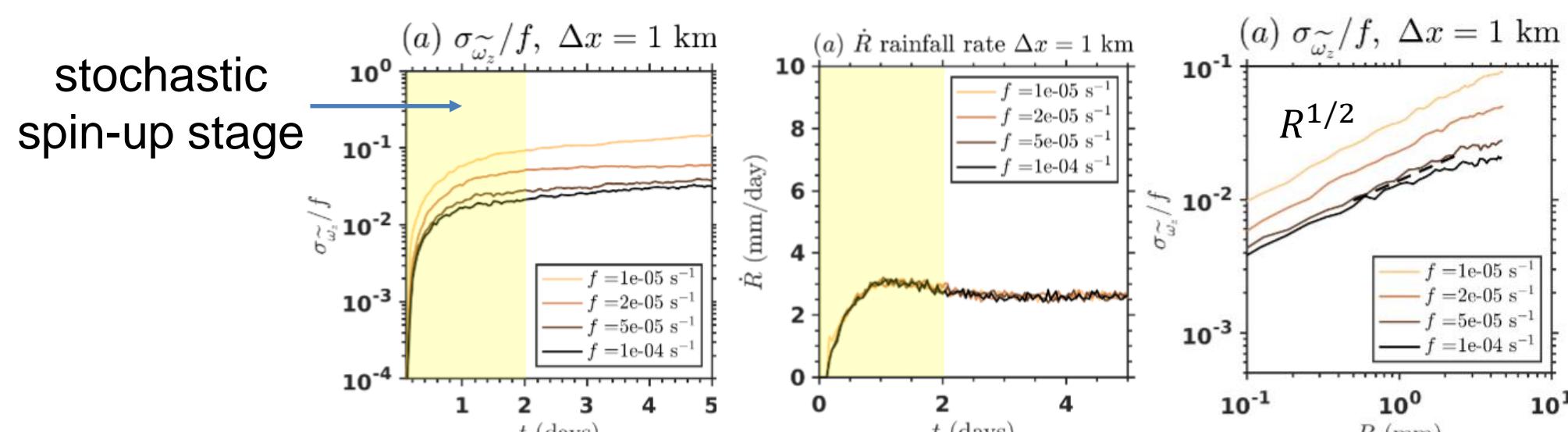


Key quantity: mesoscale midlevel vorticity



- Convection is NOT a small-scale noise – it projects onto larger scales [4].
- The **mesoscale vorticity** and vapor fluctuation [5] are **initial seeds** for mesoscale feedback.
- How to quantify the mesoscale imprint of the “convectively generated vorticity noise”?

Experimental result: power law



- The standard deviation of mesoscale vorticity $\tilde{w_z}$, $\sigma_{\tilde{w_z}}$, grows rapidly within the first two days.
- Using the accumulated rainfall rate R as a rescaled time coordinate, there is $\sigma_{\tilde{w_z}} \propto R^{1/2}$.
- The variable R represents the accumulated convective stirring effect.

Main References

- [1] Fu H, O'Neill M. The Stochastic Spin-up of Vorticity in Spontaneous Tropical Cyclogenesis[J]. *J. Atmos. Sci.*, 2024, 81(5): 855-870.
- [2] Bretherton C S, Blossey P N, Khairoutdinov M. An Energy-Balance Analysis of Deep Convective Self-Aggregation above Uniform SST[J]. *J. Atmos. Sci.*, 2005, 62(12): 4273-4292.
- [3] Fu H, O'Neill M. The Small-Amplitude Dynamics of Spontaneous Tropical Cyclogenesis. Part I: Experiments with Amplified Longwave Radiative Feedback[J]. *J. Atmos. Sci.*, 2024, 81(2): 381-399.
- [4] Mapes B E. Equilibrium vs. Activation Control of Large-Scale Variations of Tropical Deep Convection[M]//The Physics and Parameterization of Moist Atmospheric Convection. Dordrecht: Springer Netherlands, 1997: 321-358.
- [5] Biagioli G, Tompkins A M. A Dimensionless Parameter for Predicting Convective Self-Aggregation Onset in a Stochastic Reaction-Diffusion Model of Tropical Radiative-Convective Equilibrium[J]. *J. Adv. Model. Earth Syst.*, 2023, 15(5): e2022MS003231.

A stochastic model based on the vorticity equation

Convection: random local pulses

Background wind: random vertical shear

$$w \sim \sum_{n=1}^{\infty} \exp\left(-\frac{|\mathbf{x}-\mathbf{x}_n|^2}{r_c^2}\right) \delta(t - t_n)$$

$$\boldsymbol{\omega}_h \cdot \nabla_h w \sim \frac{\Delta U}{H} \mathbf{n} \cdot \nabla_h w$$

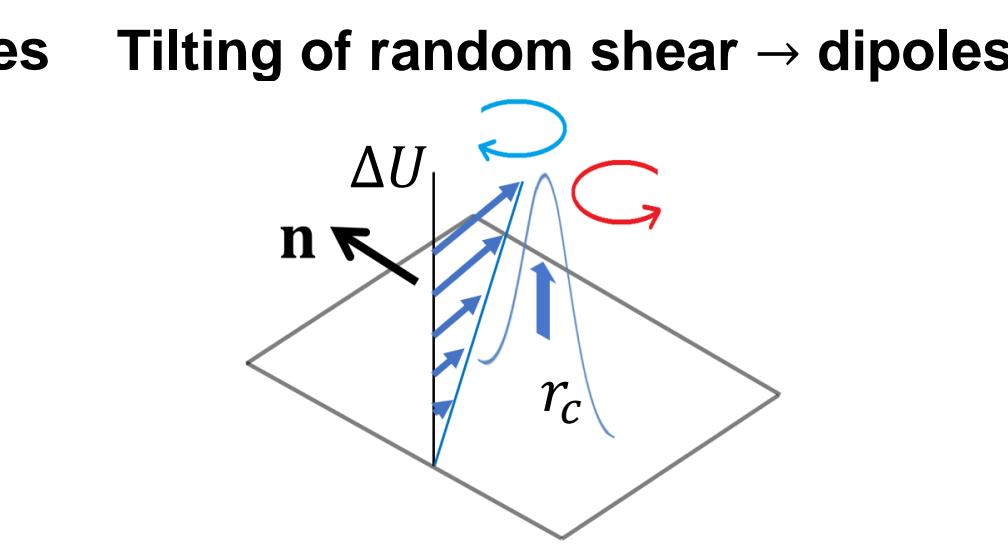
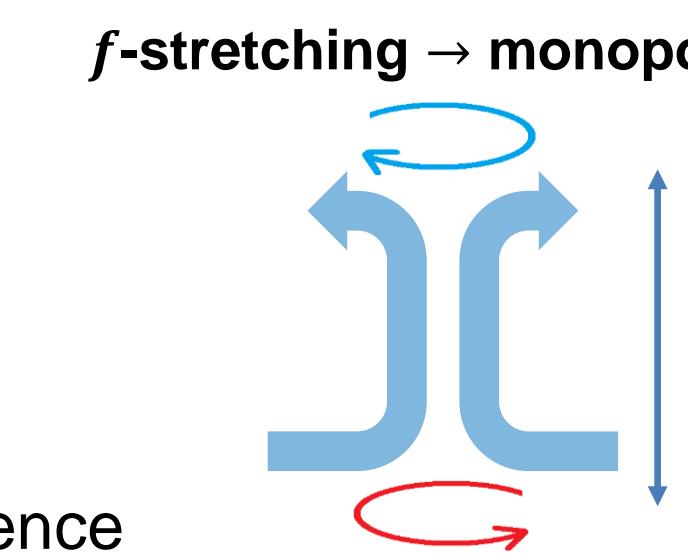
Dirac-Delta function

\mathbf{x}_n : random location vector

\mathbf{n} : a horizontal random unit vector

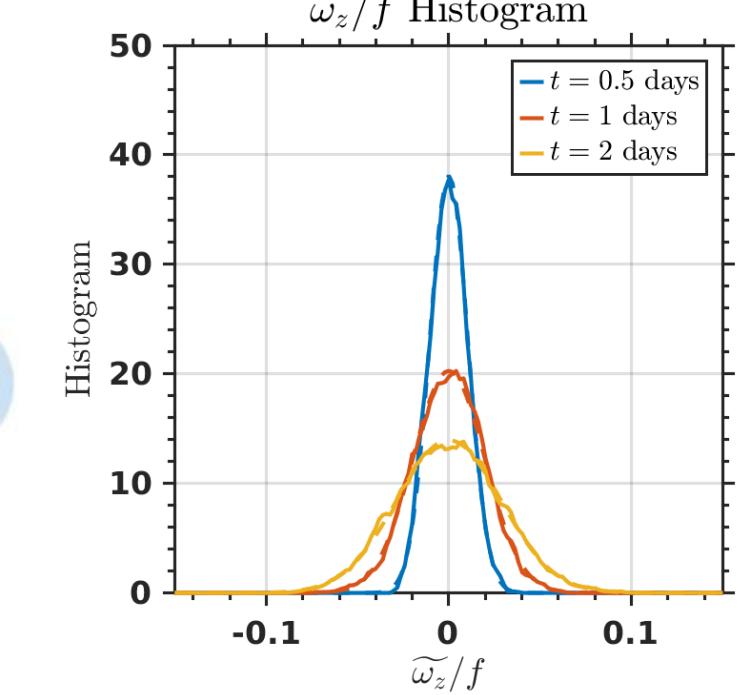
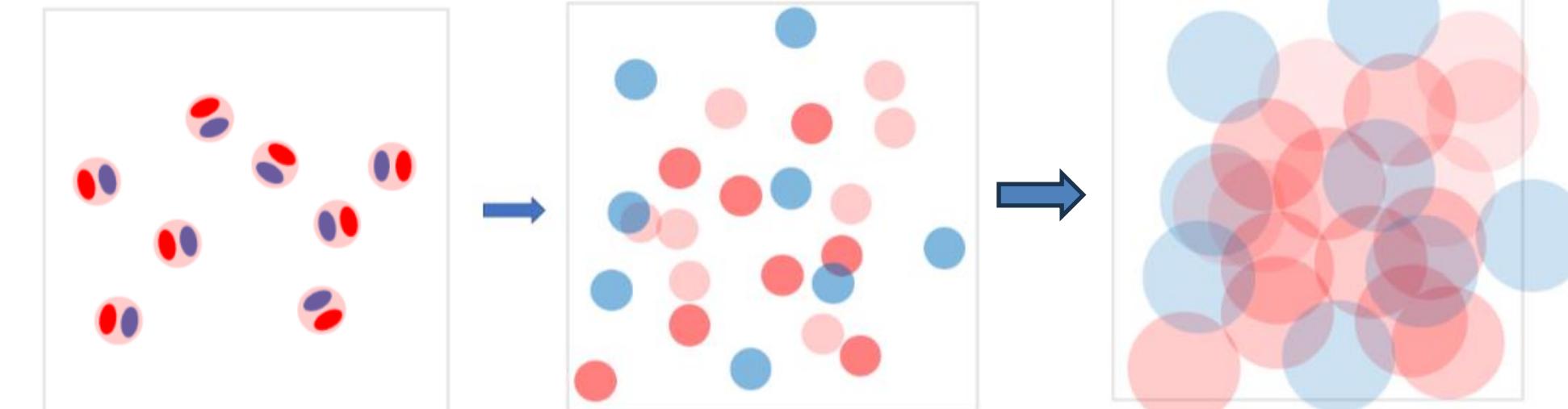
ΔU : magnitude of eddy vertical wind difference

$$\frac{\partial \omega_z}{\partial t} = \underbrace{-f \nabla_h \cdot \mathbf{u}_h}_{f \text{ stretching}} + \underbrace{\nabla_h \cdot (w \boldsymbol{\omega}_h)}_{\text{tilting flux}} - \underbrace{\nabla_h \cdot (\mathbf{u}_h \omega_z)}_{\text{advective flux}},$$



Modeling the histogram of mesoscale vorticity

Vorticity monopoles + dipoles Redistributed by 2D flow Mesoscale vorticity



cyclones by stretching + cyclones by tilting + anticyclones by tilting:
normal distribution(+) normal distribution(+) normal distribution(-)
= normal distribution

Solution of the stochastic model

$$\text{Standard deviation } \sigma_{\tilde{w_z}} = (\sigma_x^2 + \sigma_y^2 + \sigma_z^2)^{1/2} = f(\Omega t_R)^{1/2} (1 + Ro_c^2)^{1/2}$$

Convective Rossby number: $Ro_c = \frac{f_*}{f}$ growth timescale
the relative importance of tilting to f -stretching

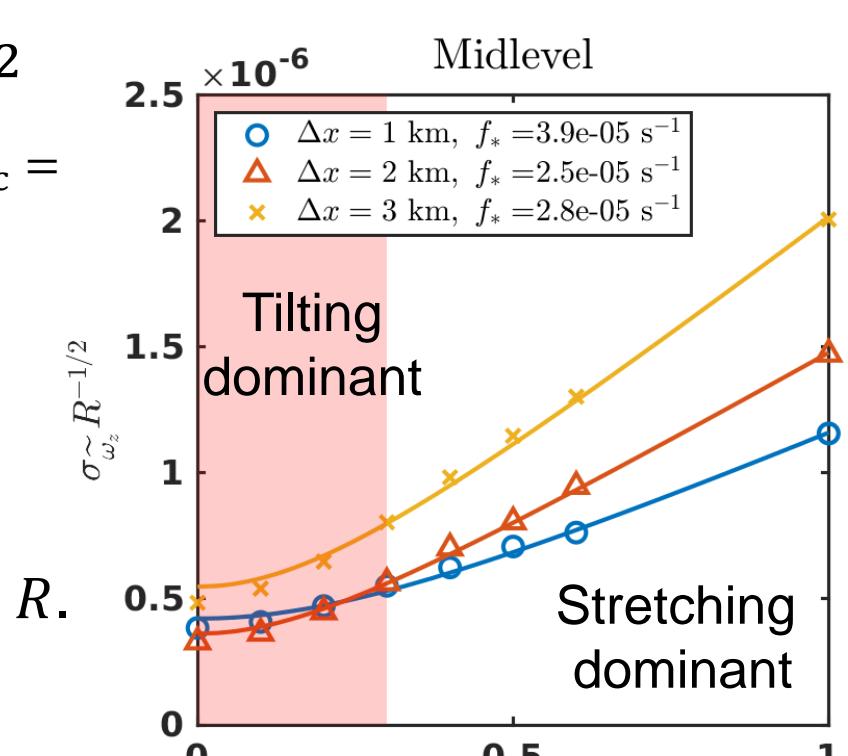
$$Ro_c \equiv \frac{f_*}{f} = \left(\frac{2}{\pi}\right)^{1/2} \frac{\Delta U}{f r_c}$$

ΔU : eddy vertical wind difference r_c : cloud radius

t_R : time coordinate rescaled by the domain-averaged accumulated rainfall R .

$f \downarrow$, $Ro_c \uparrow$, tilting more dominant

- $Ro_c \sim 1 \Leftrightarrow f_* \sim 3 \times 10^{-5} \text{ s}^{-1} \Leftrightarrow 12^\circ \text{N}$, transition between tilting & stretching regimes.
- $\sigma_{\tilde{w_z}}$ is sensitive to the grid spacing Δx , which influences the simulated cloud radius r_c .
- Future work: include background vertical shear.



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