# **PROCEDURE:** Based on Lee and Zawadzki (2006), Journal of Hydrology, p. 83-97.

- **RUC-2** forecasts provide initial indication of the 0° C isotherm height. No analysis is performed if < 2 km.
- A 5-point smoother is applied to the  $\Phi_{DP}$  and  $Z_H$  measurements.
- **Polarimetric info used to omit any path with non-liquid precipitation.**
- Measured and theoretical  $\Phi_{DP}$  differences are deduced for all rain paths > 20 km.
- Theoretical  $\Phi_{DP}$  is derived from  $Z_H = 5.7 \ge 10^4 K_{DP}^{1.075}$  (Brandes et al. 2005)
- A path along a given azimuth can contribute only one  $(\Phi_{DP-meas}, \Phi_{DP-theor})$  pair.
- A path may be composed of more than one segment.
- Rain paths are taken over 4 elevation angles (0.5°, 0.9°, 1.4° and 2.2°) and along azimuths free of ground echoes and/or shadows and at heights at least 0.5 km above ground and 0.5 km below the bright band.
- -Unlike Lee and Zawadzki (2006) who only considered  $\Phi_{DP-theor} > 3^{\circ}$ , all paths are kept regardless of the intensity, allowing its application to cases of light precipitation.  $\Phi_{DP-meas}$  differences may be negative in light precipitation due to significant noise caused by our fast scanning (6 rpm) radar antenna but are not rejected.
- Data from any 5-min radar cycle with fewer than 50 paths (selectable) are rejected.
- Data from cycles where the  $(\Phi_{\text{DP-meas}}, \Phi_{\text{DP-theor}})$  correlation  $\gamma < 0.4$  are rejected.
- -The accepted 5-min pairs are combined to produce a daily estimate at 1200 UTC.
- Calibration correction  $\varepsilon$  (to be added) is derived from the slope *m* of the least-square fit:  $\varepsilon = 10b\log_{10}(m)$  where b=1.075, the exponent of the  $Z_H - K_{DP}$  relationship used.
- The inclusion of a sufficiently large number (~10<sup>3</sup>) of ( $\Phi_{DP-meas}, \Phi_{DP-theor}$ ) pairs inevitably leads to a reliable estimate of the calibration.

# **OBSERVATIONS and CONCLUSIONS:**

-Polarimetric calibration with the selected  $Z_H$ - $K_{DP}$  relationship can be used for light and moderate stratiform precipitation if based on a sufficiently large number of paths ( $\sim 10^3$ ).

-The assumed  $Z_H$ - $K_{DP}$  relationship fails in heavy precipitation, in particular, when the average  $Z_{DR}$  over the path exceeds 2 dB.

-One of the  $Z_H$ - $K_{DP}$ - $Z_{DR}$  relationships proposed by Vivekanandan et al. (2003),  $K_{DP}$ =3.32x10<sup>-5</sup> $Z_H Z_{DR}^{-2.05}$  has been shown to overcome the limitation of the two-parameter relationship but any inconsistency found between the power and phase measurements cannot be attributed to only  $Z_{H}$  but to a combination of  $Z_H$  and of  $Z_{DR}$ .

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Distributions of observed  $K_{DP}$  vs observed  $Z_{H}$  for all paths within the indicated  $Z_{DR}$  interval, normalized to 100% at the  $(K_{DP}, Z_{H})$  pair of greatest occurrence. The grey pixels through the distribution represent the selected  $Z_H - K_{DP}$  relationship.  $K_{DP}$  is overestimated at high  $Z_{DR}$ .



Scatter plots of  $\Phi_{DP-meas}$  vs  $\Phi_{DP-theor}$  differences obtained with the  $Z_{H}-K_{DP}-Z_{DR}$  relationship derived by Vivekanandan et al. (2003) for the same 3-month data set stratified by the average  $Z_{DR}$  over the path.  $\Phi_{DP}$  differences are better predicted for all ranges of  $Z_{DR}$ .