

INTRODUCTION

- Davies-Jones and Stumpf (1977, 28th Conf. on Radar Meteorology, pp. 313-314) advocated using the detection and measurement of significant circulation around and areal contraction rate of a curve as a potential method for giving advanced warnings of tornadoes.
- Circulation (rate of areal contraction) is the line integral around a closed curve of velocity tangential (inward normal) to the curve.
- Circulation and areal contraction rate may be more useful than differential single-Doppler velocities in the characteristic velocity couplet for detecting and measuring the strength of convergent tornadic mesocyclones at low altitudes because of the following reasons:
 - a. Doppler rotational velocity (half the velocity difference between the two peaks in the velocity couplet) does not separate vortical and convergent flow,
 - b. circulation and areal contraction rate are (1) less scale dependent, (2) are tolerant of noisy Doppler velocity data, and (3) are relatively insensitive to range and azimuth.

3-D VELOCITY COMPONENTS OF THE RADAR TARGETS

- Since a Doppler radar scans the 3-D velocity components of a target with a volume scan, it is expeditious and computationally economical to determine the kinematic properties of the velocity field in the conical surfaces of constant elevation angle.
- Mathematically, the velocity components defined in a right-hand spherical coordinate system are

$$\mathbf{V} = V_R \hat{\mathbf{R}} + V_\alpha \hat{\boldsymbol{\alpha}} + V_\beta \hat{\boldsymbol{\beta}}, \quad (1)$$

where V_R is the radial component of the target in the radar viewing direction, V_α and V_β are both perpendicular to the radar viewing direction. V_α is along the upward inclined normal in the vertical plane containing the radar beam, and V_β is horizontal and perpendicular to the right of the beam in a constant α_o -surface.

- Here, $\hat{\mathbf{R}}, \hat{\boldsymbol{\alpha}}, \hat{\boldsymbol{\beta}}$ are the unit basis vectors of the spherical coordinate system.

CIRCULATION

- By Stokes' theorem, a circulation Γ around the boundary C of an area A in the α_o -surface is given by

$$\Gamma = \oint_C (V_R dR + V_\beta R \cos \alpha_o d\beta), \quad (2)$$

where the first term on the right-hand side of (2) is the observed circulation around C with V_β ignored.

- The so-called "cell circulation" Γ_{ob} around the i, k grid cell is computed from the first term on the right-hand side of (2) at the midpoint of each cell as $(d\Gamma_{i+1/2, k+1/2})_{ob} =$

$$\frac{\Delta R}{2} [V_R(R_i, \alpha_o, \beta_{k+1}) + V_R(R_{i+1}, \alpha_o, \beta_{k+1}) - V_R(R_{i+1}, \alpha_o, \beta_k) - V_R(R_i, \alpha_o, \beta_k)], \quad (3)$$

where i is the index in the range direction, k is the index in the azimuth direction, and $\Delta R = R_{i+1} - R_i$ is the range increment.

- Circulation in (3) is additive because the circulation around the outer perimeter of a union of contiguous grid cells is simply the sum of the circulations around the perimeter of each grid cell.
- Thus, the circulation around C by the line-integral method of Davies-Jones (1993, MWR, vol. 121, pp. 713-725) is computed, given by

$$\Gamma = \frac{\Lambda(V_R, R)}{2} - \cos \alpha_o \frac{\Lambda(RV_\beta, \beta)}{2}, \quad (4)$$

where $\Lambda(V_R, R)/2$ is the observed circulation, and $\Lambda(f, g) \equiv (f_1 g_2 - f_2 g_1 + f_2 g_3 - f_3 g_2 + \dots + f_{2M+3} g_{2M+4} - f_{2M+4} g_{2M+3} + f_{2M+4} g_1 - f_1 g_{2M+4})$.

- The method fits a piecewise-linear velocity field to the observation (linear between adjacent points on the circle).

AREAL CONTRACTION RATE

- The divergence (δ) in a constant α_o -surface is percentage areal expansion per unit time of the fluid element, given by

$$\delta = \frac{1}{A} \frac{dA}{dt}, \quad (5)$$

where A is the area of the fluid element, and $dA/dt \equiv A_t > 0 (< 0)$ is the rate at which the area of the curve expands (contracts).

- For a horizontal convergent flow, negative areal expansion rate ($A_t < 0$) corresponds to contraction.

- The element of area A in (5) on the α_o -surface (Hildebrand 1962, p. 304) is

$$dA = \cos \alpha_o R dR d\beta. \quad (6)$$

- Differentiating (6) with respect to time t and integrating by parts give the rate of areal expansion (the negative of the areal contraction rate) on the α_o -surface,

$$A_t = -\cos \alpha_o \oint_C V_R R d\beta + \oint_C V_\beta dR, \quad (7)$$

where the first and second terms on the right-hand side of (7) are the observed and unobserved effluxes.

- Using Davies-Jones' line-integral method, the expansion rate of the area A enclosed by C is

$$A_t = -\cos \alpha_o \frac{\Lambda(RV_R, \beta)}{2} + \frac{\Lambda(V_\beta, R)}{2}, \quad (8)$$

where the last term on the right-hand side of (8) is unobserved.

- Ignoring the last term on the right-hand side of (7), the so-called "cell areal expansion rate" represents the observed expansion rate of the i, k grid cell calculated at the midpoint of the cell and is therefore obtained by

$$\left(\frac{dA_{i+1/2, k+1/2}}{dt}\right)_{ob} = \frac{1}{2} \cos \alpha_o (\beta_{k+1} - \beta_k) \times \{R_{i+1} [V_R(R_{i+1}, \alpha_o, \beta_{k+1}) + V_R(R_{i+1}, \alpha_o, \beta_k)] - R_i [V_R(R_i, \alpha_o, \beta_{k+1}) + V_R(R_i, \alpha_o, \beta_k)]\}. \quad (9)$$

SIMULATED DOPPLER RADAR & RANKINE MODEL

- Simulated Doppler velocities are produced by scanning a radar across Rankine vortical and convergent flows.

- In (1), V_R at lowest elevation angle is computed from $V_R(R_i, \alpha_o, \beta_k) = (U \sin \gamma + V \cos \gamma) \cos \alpha_o$, (10)

where U is the radial component of an axisymmetric convergent flow, V is the tangential component of an axisymmetric cyclonic vortex, and γ is the angle between the tangential direction component V and the radar viewing direction β_k at a point.

- U and V can be modeled using the Rankine formulas, given by

$$U = U_x \left(\frac{\rho}{R_x}\right)^\epsilon$$

$$V = V_x \left(\frac{\rho}{R_x}\right)^\epsilon \quad (11)$$

where U_x, V_x are, respectively, radial and tangential velocity peaks, R_x is core radius, ρ is radius from center,

$\epsilon (= +1, -1)$ is the exponent that governs the inner velocity ($\rho \leq R_x$) and outer velocity profiles ($\rho > R_x$), respectively.

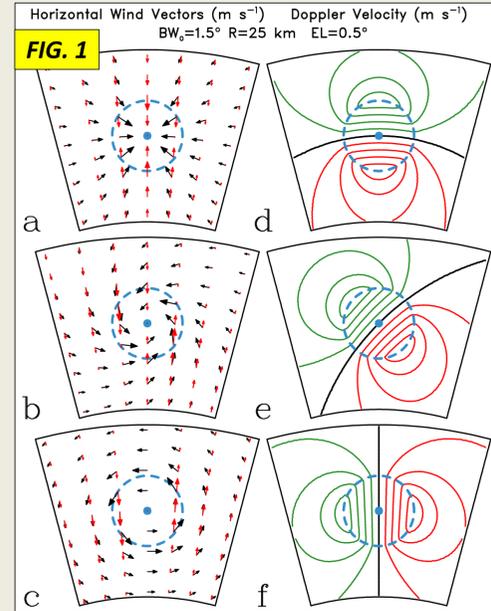


FIG. 1 Simulated, storm-relative horizontal (black) and Doppler (red) wind vectors with (a) convergent flow ($U_x = -25 \text{ m s}^{-1}, V_x = 0$), (b) combined convergence and rotation ($U_x = -17.8 \text{ m s}^{-1}, V_x = 17.8 \text{ m s}^{-1}$), and (c) rotational flow ($U_x = 0, V_x = 25 \text{ m s}^{-1}$) are plotted at 0.5° -elevation angle. Simulated, positive (negative) Doppler velocities represent flow away from (toward) the radar, shown by red (green) contours with contour interval of 5 m s^{-1} . Simulated, zero Doppler velocity contour (black) represents flow perpendicular to the radar viewing direction. Blue Doppler signature center is located at 25 km north of a simulated Doppler radar. Blue, dashed circle represents an axisymmetric Rankine core diameter ($2R_x$) of 5 km with its radial (U_x) and tangential (V_x) velocity peaks. Range increment is 240 m. The Doppler velocity pattern in (d)-(f) rotates counterclockwise, reflecting a change from (a) convergent flow through (b) combined convergence and rotation to (c) rotational flow.

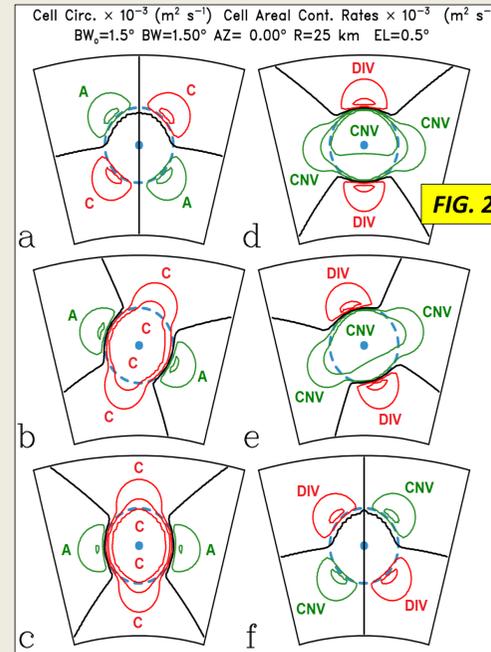


FIG. 2 Cell circulations in (a)-(c) and cell areal contraction rates in (d)-(f) are calculated. Positive (red) and negative (green) contours with zero (black) contour are indicated with contour interval of $0.5 \text{ m}^2 \text{ s}^{-1}$. **C** stands for Doppler cyclonic shear; **A** for Doppler anticyclonic shear; **CNV** for Doppler convergence; and **DIV** for Doppler divergence, as they correspond to red Doppler wind vectors (Figs. 1a-c).

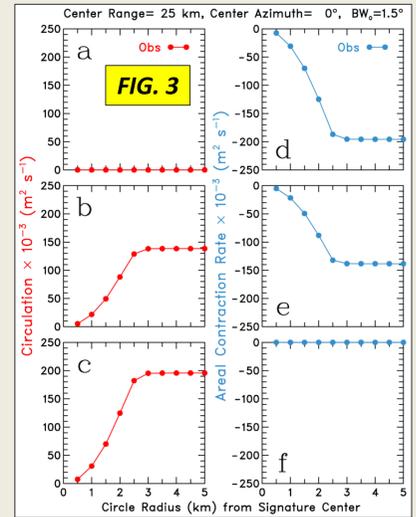


FIG. 3 Circulations [red dotted curves in (a)-(c)] and areal contraction rates [blue dotted curves in (d)-(f)] are functions of circle radius from the signature center, as are calculated from the first term on the right-hand side of (4) and (8), respectively.

CONCLUSIONS AND ON-GOING WORK

- Circulation and areal contraction rate are useful because (a) they help separate vortical and convergent flows, and (b) circulation around a circle is a better measure of vortex strength than Doppler rotational velocity.
- A potential vortex (the outer part of the Rankine vortex) and a sink each produce spurious "quadrupole" patterns in both observed vorticity and divergence. These patterns produce false readings in the potential vortex.
- For a circle centered on the vortex, the quadrupole contributions to circulation and areal contraction rate cancel.
- In our on-going work, we will demonstrate that circulation is much less range and azimuth dependent than Doppler rotational velocity is.



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 Contact: Vincent.Wood@noaa.gov

