Combined cloud radar and lidar aerosol observations in Potenza, southern Italy Pilar Gumà-Claramunt¹*, Fabio Madonna¹, Aldo Amodeo¹, Matthias Bauer-Pfundstein² ¹Istituto di Metodologie per l'Analisi Ambientale, Consiglio Nazionale delle Ricerche, Italy; ²METEK GmbH, Germany *Corresponding author: pilar.guma@imaa.cnr.it

INTRODUCTION

Which is the problem?

Giant and **ultragiant** aerosols ($r > 5 \mu m$) distribution and importance Combine different instruments aerosol measurements: for global meteorology and climate through warm rain processes and • Lidar and sun photometer are used to retrieve aerosol ice nucleation is not well known. They expedit warm rain processes microphysical properties between 100 nm and few μ m^[4] in absence of by acting as GCCN (Giant Cloud Condensation Nuclei^[1-3] and are thick clouds • Cloud radar can detect giant and ultragiant aerosols^[5] efficient **IN** (Ice Nuclei), increasing the ice formation temperature.

METHODOLOGY



RESULTS

Cloud radar aerosol dataset

Lidar simultaneous measurements analysis the identification of 150 aerosol and 576 insect lofted layers.

Some general statistics of these layers are shown in the next histograms: (a) the monthly distribution, (b) the daily distribution, and (c) the horizontal wind speed of the layers.









CONCLUSIONS

Clour radar aerosol observation

•Giant aerosols can be observed with a cloud radar

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- Entomology criteria are appropriate to discriminate aerosols and insects
- •Large number of layers detected (~ 20/year)
- The size range where aerosols can be observed is enlarged

- and

What can we do?

Lidar simultaneous measurements

Raman lidars (Nd-Yag) emit at 355, 532 and 1064 nm and receive the elastic backscattered light at the same wavelengths and the Raman backscattered light at 387 (N₂), 407 (H₂O₂) and 607 nm (N₂). By analysing the measured signals, the extensive and intensive parameters are obtained. The first depend on the particle concentration, while the second on the particle type.

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Giant aerosol effects

• The AOD seasonal evolution is in accordance with aerosol observations

• The **precipitation life cycle** is modified: lower accumulation can be retrieved by **inversion methods** preferentially $\sim 1\frac{1}{2}$ days after the observation

Synergy

•The **particles observed** by the two instruments are **different** •The size distribution (lidar) and the effective radius (radar) more probability of intense rainfall, occurring •Their merging is currently under investigation to obtain the enlarged size distributions

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How do we do it?

• Establishing an **aerosol detection** methodology with the **cloud radar**

- Analysing lidar simultaneous measurements and retrieving their size
- distributions by lidar inversion codes
- Creating an effective radius radar inversion code
- Using the lidar-radar synergy to obtain the aerosol effective radius

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Effective radius retrieval

jointly done. For the radar, a lookup reflectivity and LDR table was build with a T-matrix scattering code and is consulted to obtain the most probable particle size, shape and refractive index. For the lidar, an already existing inversion code is used. Then, the inverted parameters from both size instruments are combined.



Lidar and cloud radar synergy The application of the new methodology to the period March 2009 – June 2015 resulted in : The lidar measurements simultaneous to giant : The time-range evolutions of (a) the radar reflectivity and (b) the lidar Range Corrected Signal (RCS) at 1064 nm, (c) the aerosols observations were analyzed. Two examples is backscatter profiles, (d) the lidar size distribution and (e) the radar effective radius are the following: CASE A: dust IIRA36 05/08/2013 Tito Scalo (PZ), Italy MUSA 05/08/2013 Tito Scalo (PZ), Italy Lidar inversion: **Radar inversion:** (b) • $r_{eff} = 2.34 \pm 0.09$ ▶ r = 2.29 ± 0.09 μm and a start of the second of the second second start of the second second second second second second second s • $N = 31 \pm 1 \text{ cm}^{-3}$ 0.10 Radius [um ¹ • Axis ratio = $1.4 \pm$ 1.00 • $r_{eff} = 0.55 \pm 0.15 \ \mu m$ • Spheroid fraction = 0.2 ► RI = 2.27 - 0.57i r = 0.13 ± 0.015 μm 0% 10⁻¹⁴ 10⁻¹² 10⁻¹⁰ 10⁻⁸ 10⁻⁶ 10⁻³ 10⁻² 10⁻¹ 10⁰ 10¹ 19:10 Time UTC [hh:mm] 19:10 Time UTC [hh:mm] Backscatter [m⁻¹ sr⁻¹] 0 10 20 30 40 50 60 CASE B: smoke MUSA 19/06/2013 Tito Scalo (PZ), Italy MIRA36 19/06/2013 Tito Scalo (PZ), Italy **Radar inversion:** Lidar inversion: **(e)** (b) **(C)** (a) • $r_{eff} = 2.11 \pm 0.22$ Aerosols lofted I r = 1.87 ± 0.19 μm • $N = 90 \pm 10 \text{ cm}^{-3}$ incert 0.10 Radius [um ¹ 1.00 10.00 • Axis ratio = $1.3 \pm$ • $r_{eff} = 0.14 \pm 0.018 \ \mu m$ • Spheroid fraction = 0.2 ▶ RI = 2.22 - 0.51i 10⁻¹⁴ 10⁻¹² 10⁻¹⁰ 10⁻⁸ 10⁻⁶ ▶ r = 0.27 ± 0.05 μm 10⁻⁴ 10⁻³ 10⁻² 10⁻¹ 10⁰ 1 19:35 19:40 19:45 Time UTC [hh:mm] 19:50 19:40 Time UTC [hh:mm] Backscatter [m⁻¹ sr⁻¹] ▶ N = 420 ± 140 cm⁻³ RI = 1.53 - 0.005i 4•10⁶ 6•10⁶ 8•10⁶ RCS 1064 nm [a.u.]

⁷Depending on the system



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