

The Impact of Radar Data Assimilation for Simulate a Supercell in Brazil Southern.

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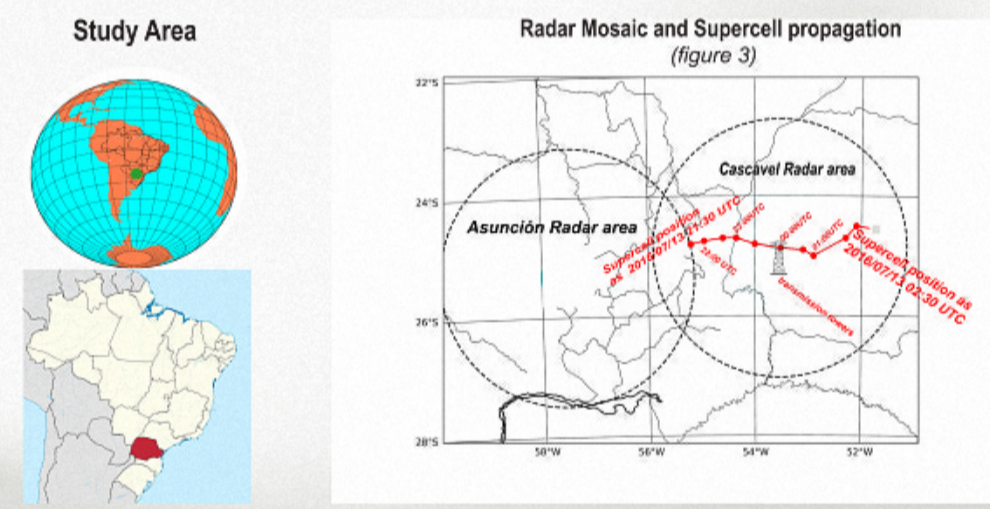
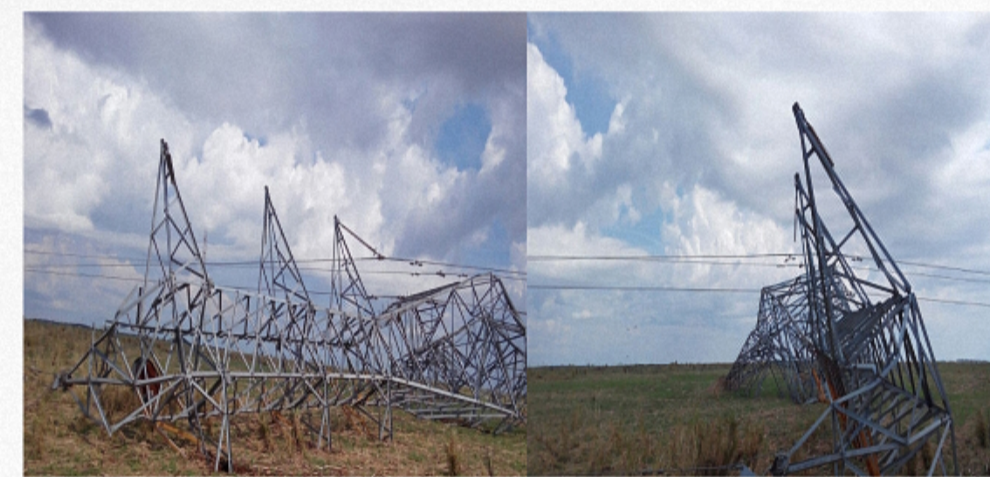
Introduction

Doppler radar measurements has a significant advantage due its temporal and spatial high resolution, allowing a good identification of patterns and signatures of mesoscale systems.

On 12 July 2016, a supercell was registered by the S-Band Dual-Pol Weather Radar of Cascavel in Southern Brazil (Calvetti et al. 2016) this supercell produced a downburst that blown down a high-voltage transmission tower in the region (figure 1,2), supercell developed and propagate over Paraguay and western Parana State in Brazil during 5 hours at 18 ms⁻¹ (figure 3), this supercell produced a downburst that blown down a high-voltage transmission tower in Cascavel City, was been registered wind gust of 32.6 ms⁻¹ at 40m in a tower installed 300m from the tower that blown down.

The purpose of this study is analyzing the impact of radar data assimilation on the numerical weather model with different assimilation methods to simulate the genesis and the propagation of the supercell.

Figure 1,2 The fallen high voltage transmission tower



Methodology

The data is provided from Cascavel S-Band Radar (Operated by SIMEPAR) localized in the city of Cascavel-PR Brazil (Radar location: Lon = -53.829, Lat = -24.870), and also Asunción S-Band Radar localized Asunción, Paraguay (Lon = -57.523, Lat = -25.333).

Radial velocity is assimilated, the indirect methods, as used to retrieved rainwater and water vapor from reflectivity data (Xiao et al. 2007)

In this study it has been used the WRF model version 3.9 (Skamarock, J. et al. 2008) for simulations, and it was tested the in four data assimilation methods: 3DVAR, (Xiao and Sun 2007), 4DVAR (Melloul et al. 2005) the variational hybrid 3DEnVar (Gao et al. 2015) and 4DEnVar (Lu et al. 2008), all assimilation methods, from the WRF Data Assimilation version 3.9 (WRFDA) package (Barker et al. 2004).

The boundary were obtained from the Global Forecast System (GFS) with a resolution 0.25 degrees. For calculate the background error is used one month backward the event of analyze x forecast. In hybrids, the physical and dynamical ensemble with 40 members for ensemble-estimated error covariances

Configuration of the experiments

Assimilation cycles = 2;
Start time of first cycle = 19UTC;
Start time of the second cycle = 22UTC;
First cycle data time windows = 3 hours forward and 3 hours backward;
Second cycle data time windows = 1 hours forward and 2 hours backward;
cv_options=5;

Minimization options
4DVAR, 4DEnVar, 3DEnVar, 3DVAR;
EPS = 1*-3, 1*-7, 1*-10, 1*-25;
mmax = 200, 400, 400, 400;

Figure 4 - Domains configuration

Physics Options	
9km	3km
mp_physics = Thompson	Thompson
ra_lw_physics = RRTMG	RRTMG
ra_sw_physics = RRTMG	RRTMG
radt = 5	5
sf_sfcly_physics = Monin-Obukhov	Monin-Obukhov
sf_surface_physics = Noah	Noah
bl_pbl_physics = MYJ	MYJ
blndt = 0	0
cu_physics = 0	0
pltp_requested = 5000	5000
e_vrt = 60	60
time_step = 30s	15s

One Way nest option is used.

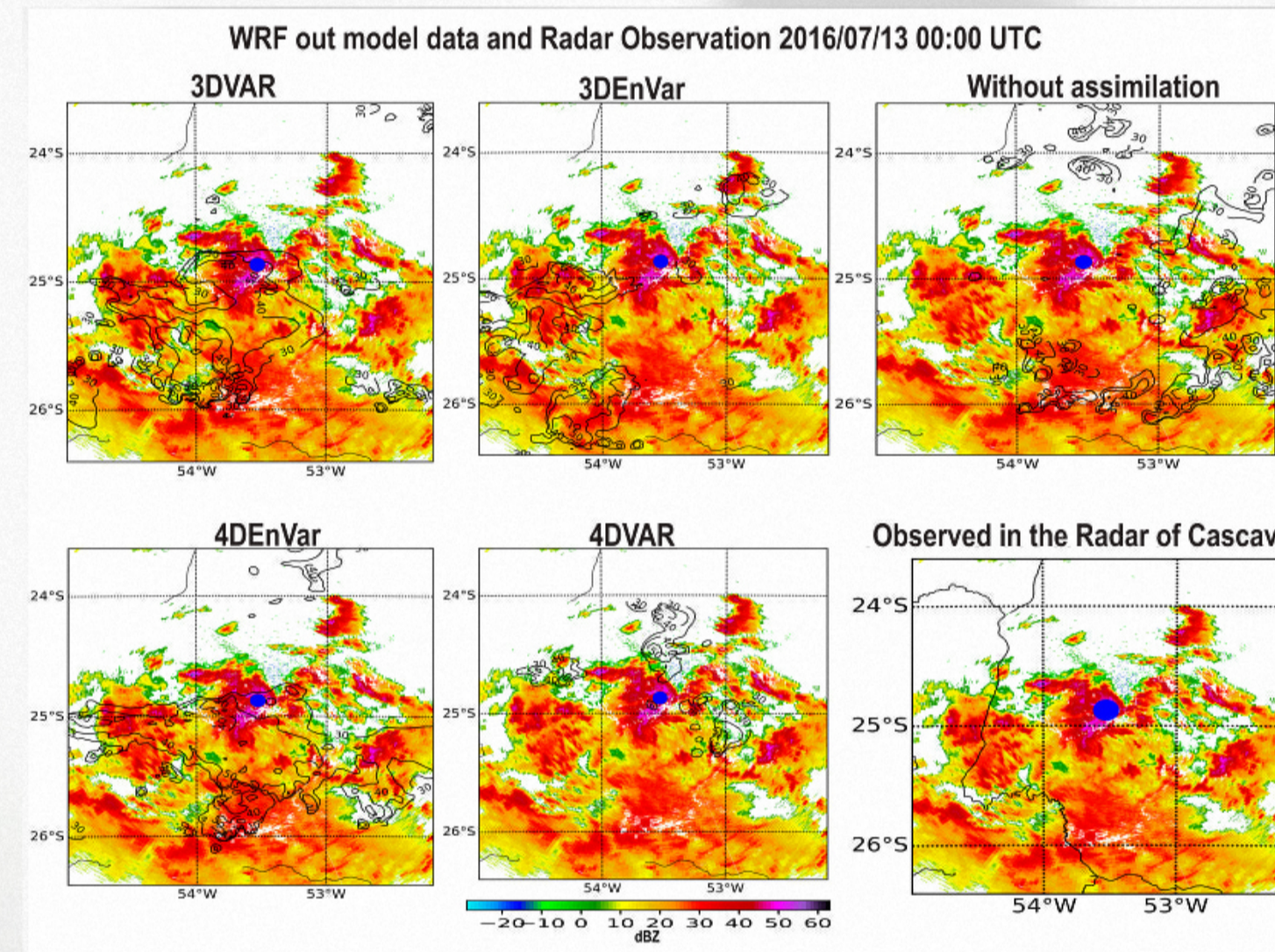


Figure 5. Contour = dBZ > 30 from WRF Simulation shaded dBZ from Cascavel Weather Radar

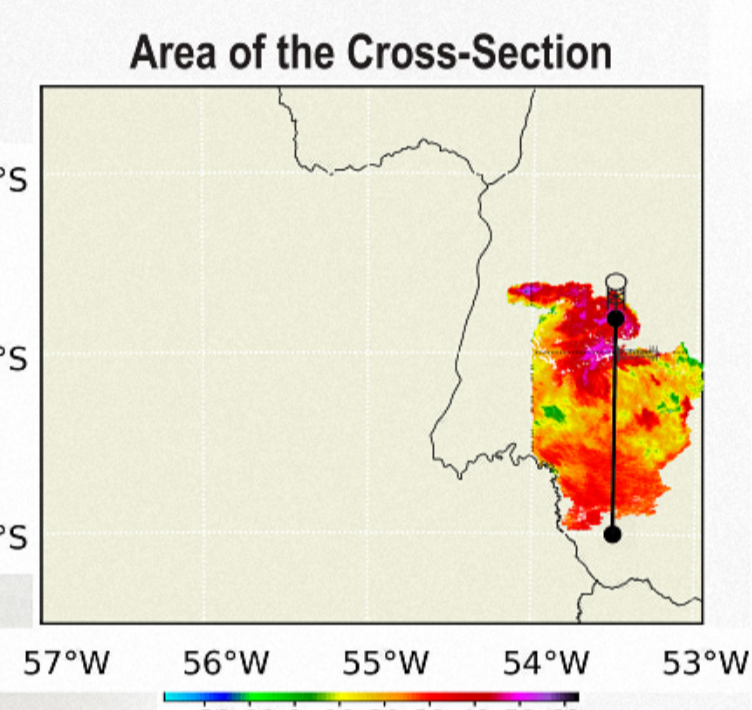


Figure 6. In this figure, shows the location of the vertical cross-section in figure 9, 10, 11, 12, 13, 14, 15. Shaded dBZ from Cascavel Weather Radar

To make the figures, some modules of the python programming language were used: for the figures with radar data the Pyart module (Helmus, J.J. et al. 2016), and for the geolocalized figures the Basemap, and the figures with model data the module wrf-python (UGAR, 2016) all 2D figures we used the graphic libraries of matplotlib (Hunter, J. et al. 2012). For 3D figures the VAPOR software is used (Clyne, J. et al. 2007).

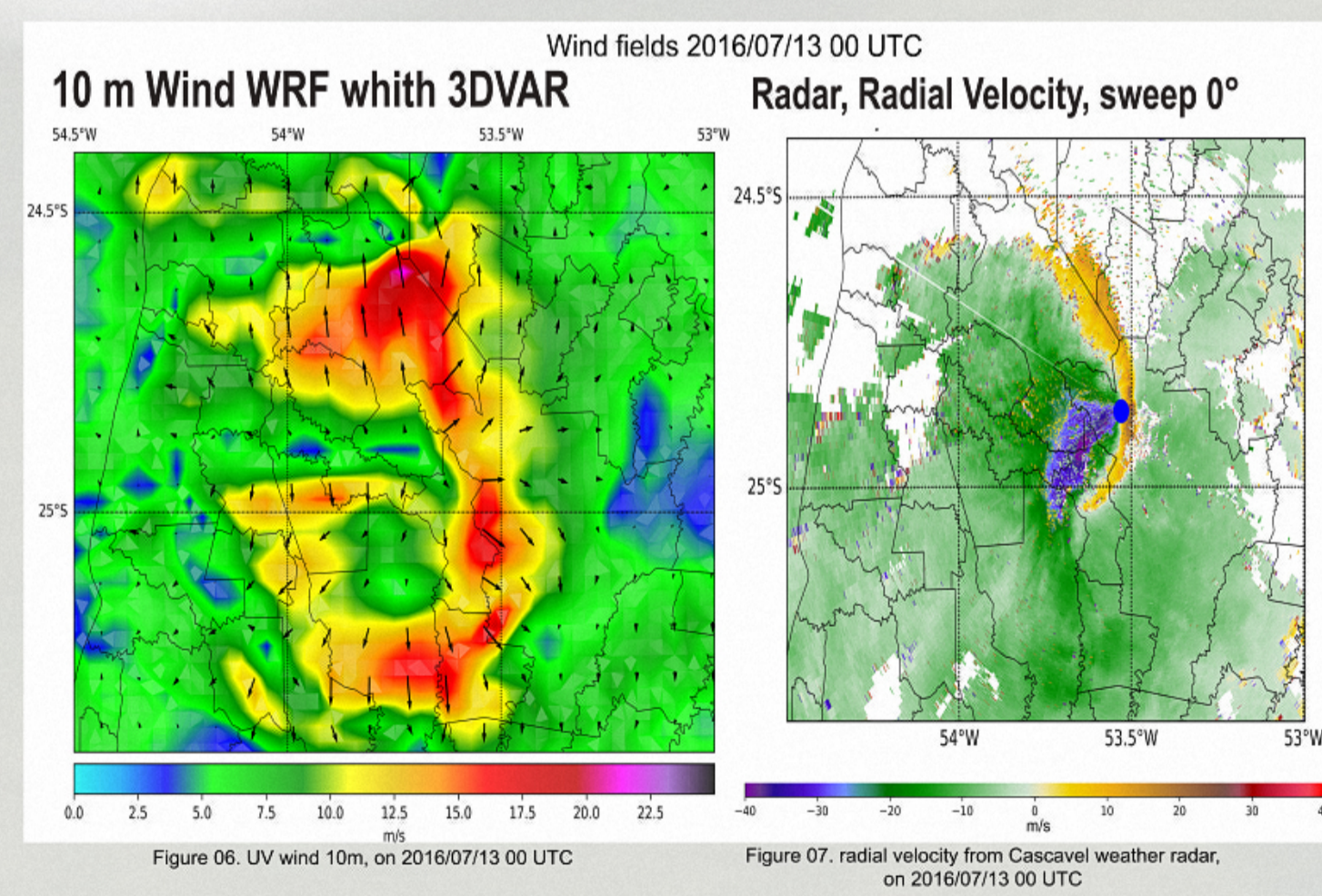


Figure 06. UV Wind 10m, on 2016/07/13 00 UTC

Figure 07. radial velocity from Cascavel weather radar, on 2016/07/13 00 UTC

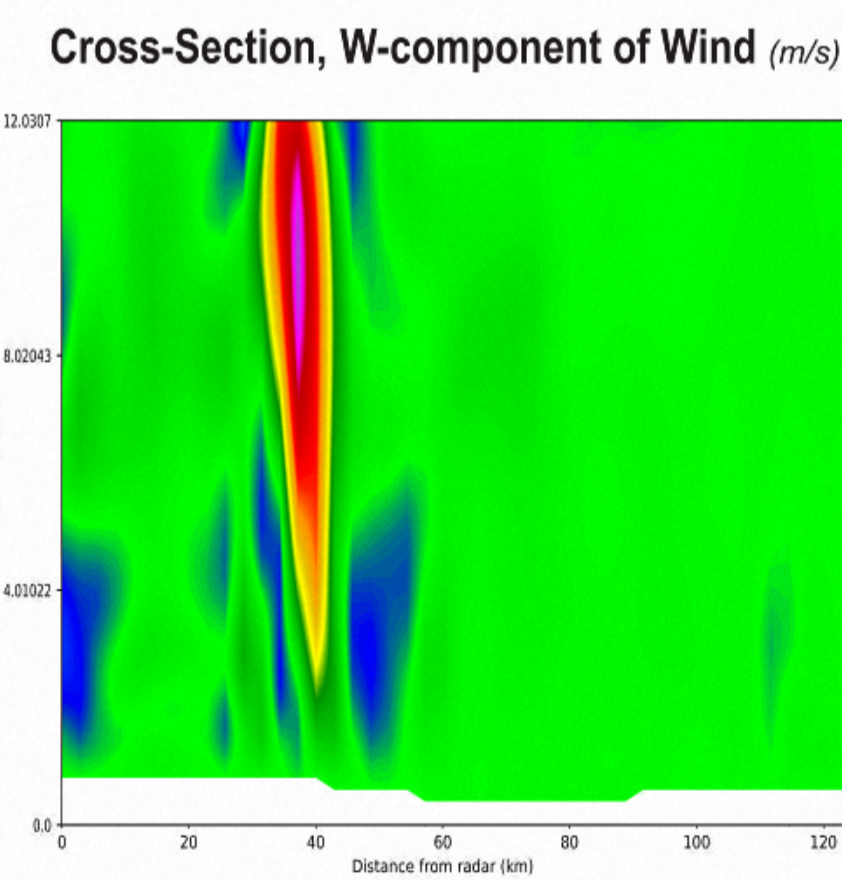


Figure 12. Vertical Wind speed from WRF simulation with 3DVAR analysis, on 2016/07/13 00 UTC

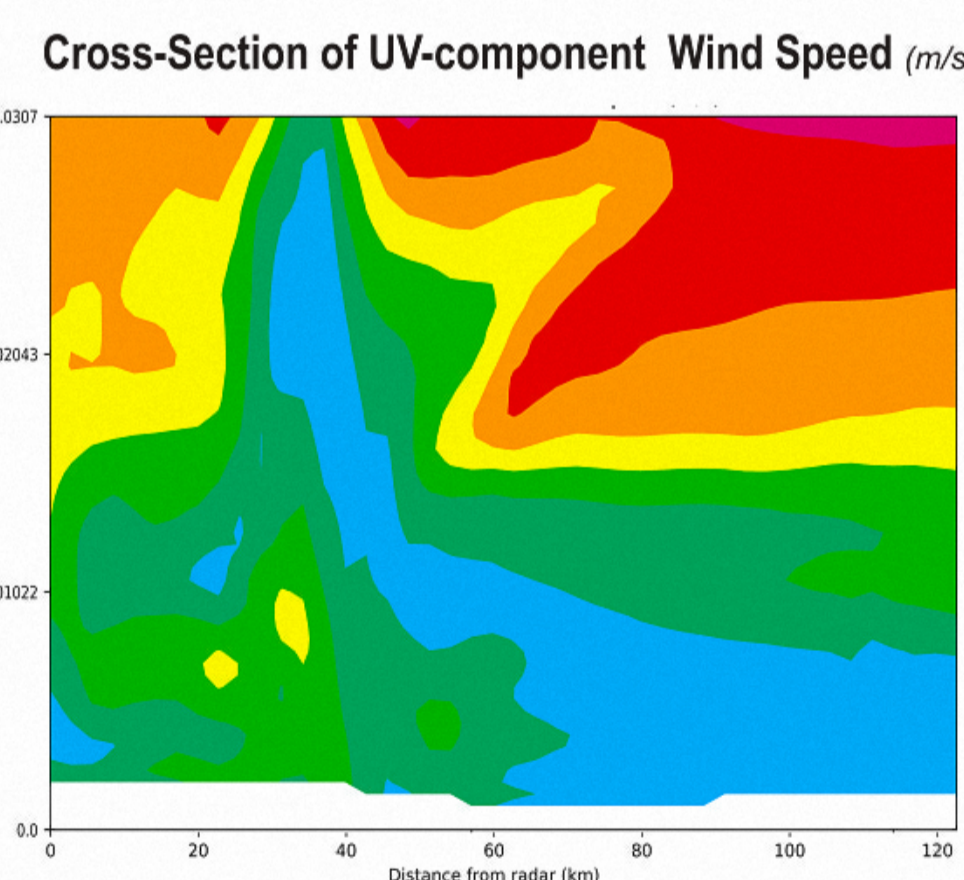


Figure 13. Wind speed from WRF simulation with 3DVAR analysis, on 2016/07/13 00 UTC

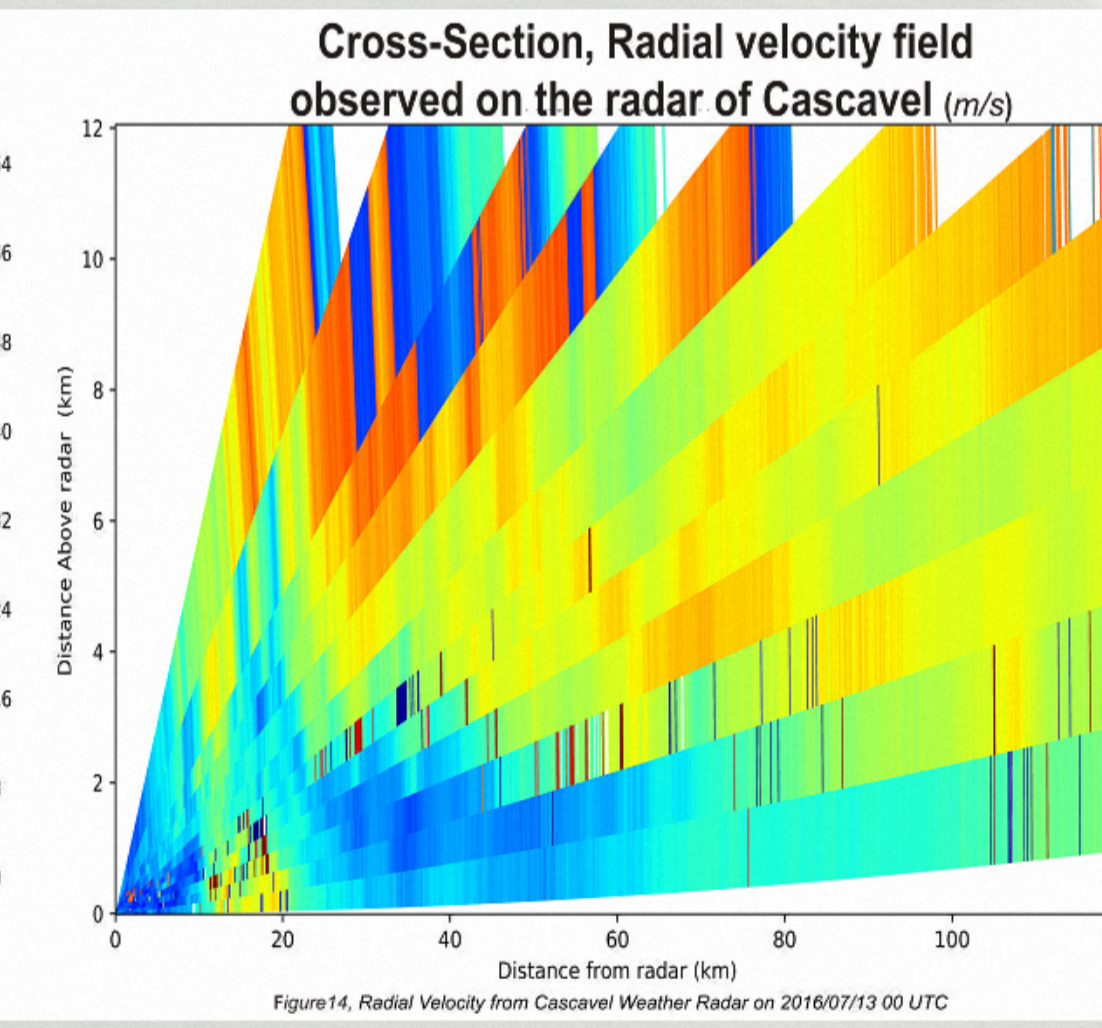


Figure 14. Radial Velocity from Cascavel Weather Radar on 2016/07/13 00 UTC

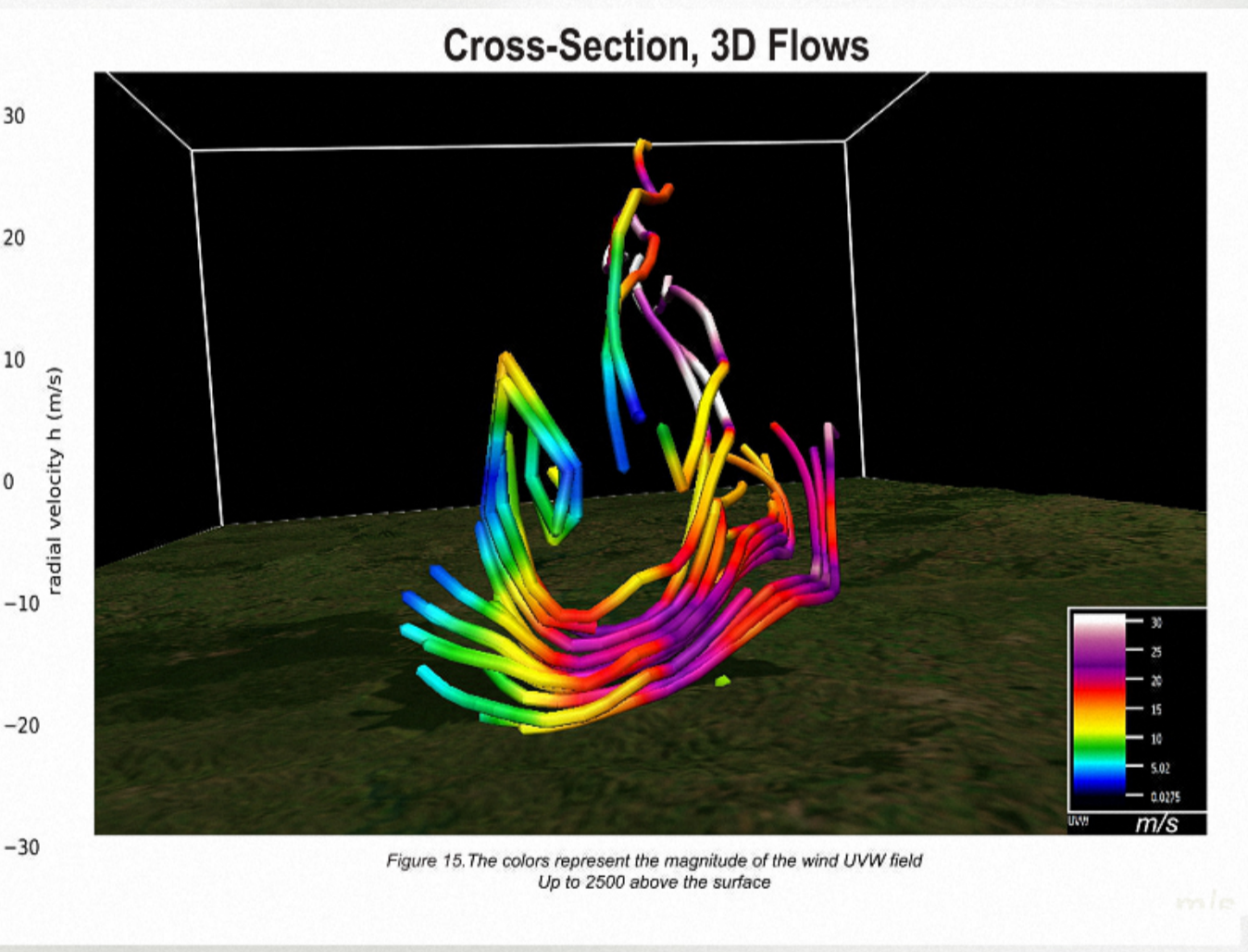


Figure 15. The colors represent the magnitude of the wind U/V/W field Up to 2500 above the surface

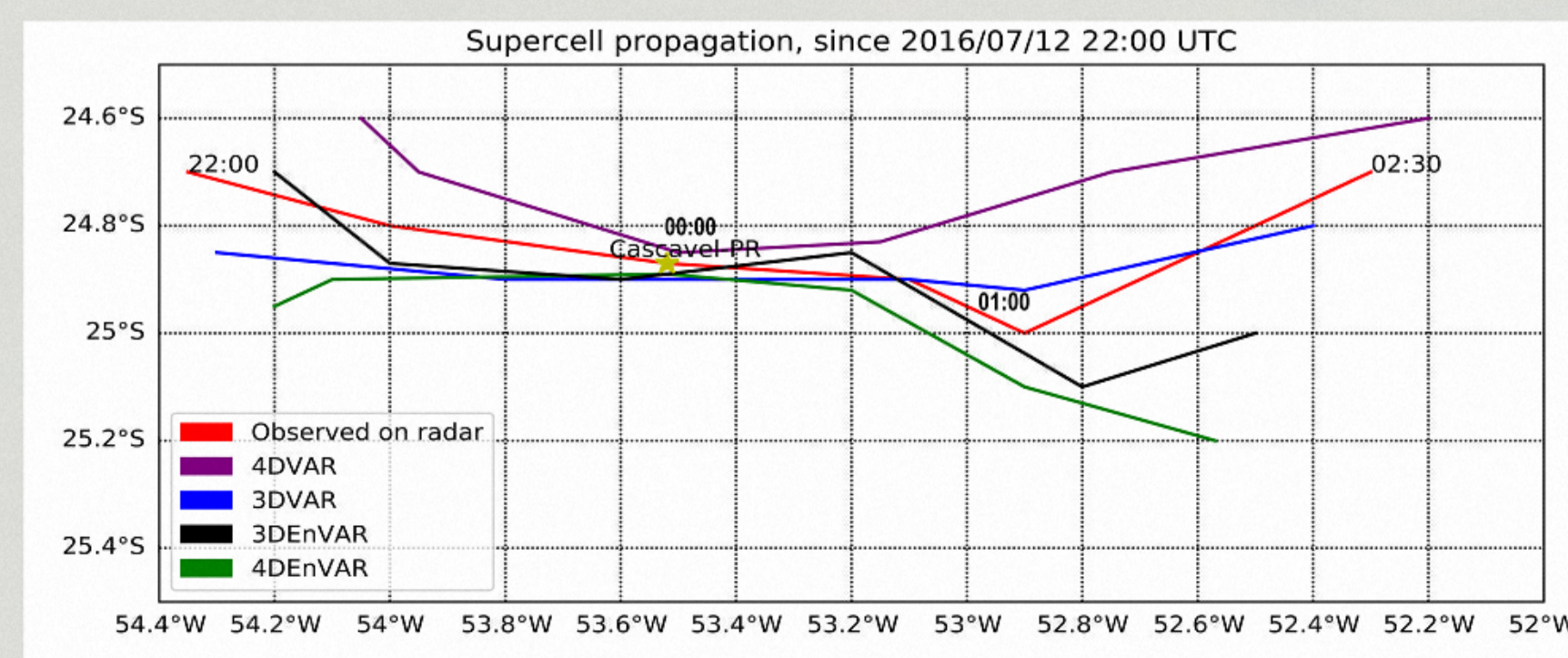


Figure 15. Propagation of the supercell between 22 UTC on day 12 to 02:30 UTC on the day 13, method by method

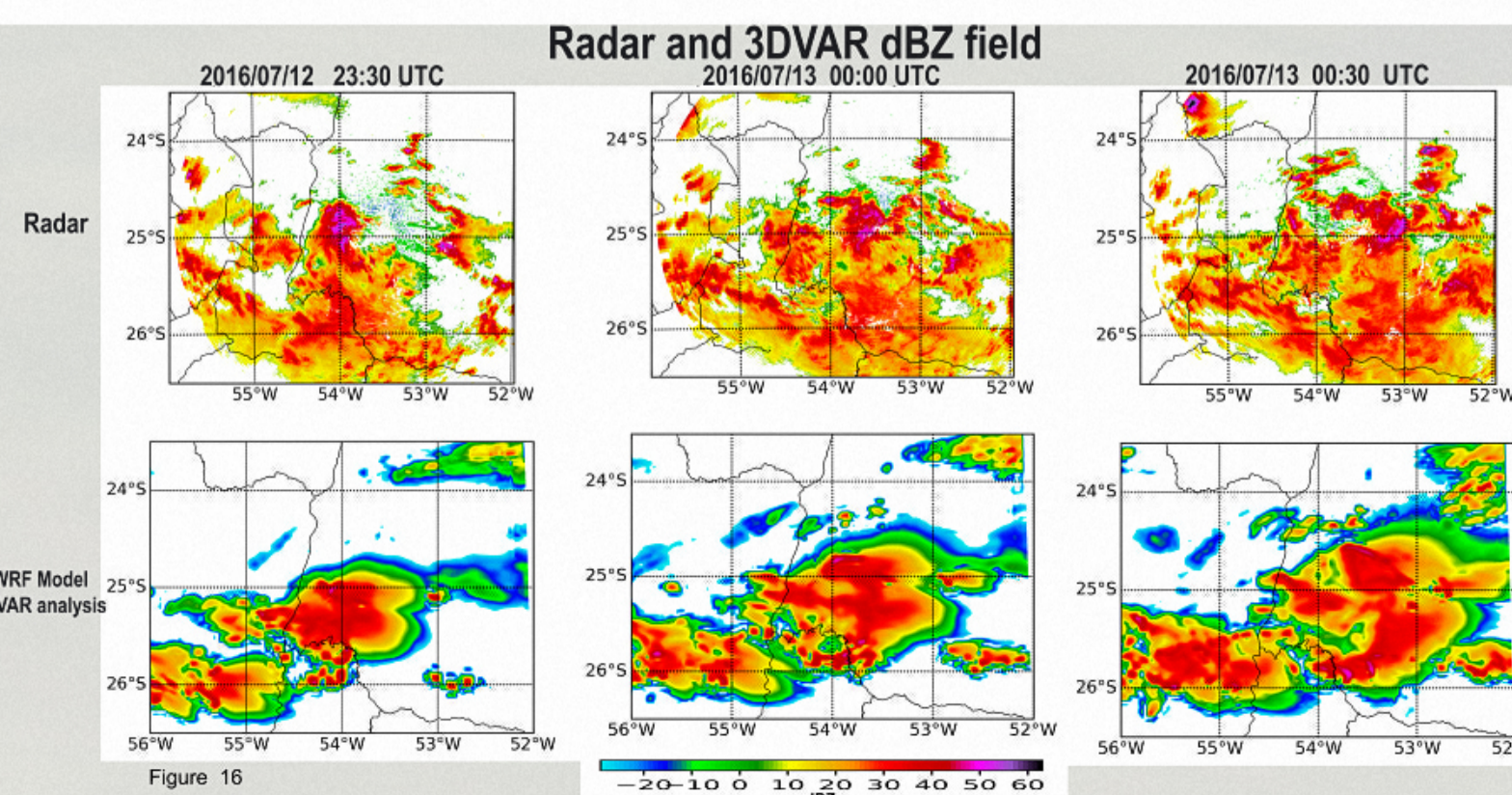


Figure 16

Results

In the figure 5, without assimilation the radar data, was not able to simulate the supercell, which yielding a scattered convection. In the experiments with assimilation, there was an improvement, approaching observed radar data, can reproduce characteristics more similar to those observed in radars data, but the results vary greatly, for each of the assimilation methods, this is extremely visible in figure 5 and 16, when when intercepting a moment, the dBZ field in the model is different for each assimilation method (figure 5), and over time, the propagation varies (figure 16).

With wind as 10 meters with the 3DVAR it is possible to detect much similarity in the wind direction, as well as its intensity, when compared with the radar data in figure 7, in the figures 12, 13, 14 with vertical data on the region where the tower fell, showing a burst of propagation with very similar intensity and generating a wind flows turbulent in the atmosphere (figure 15), it is also possible to visualize a good spatial accuracy if we compare with the radar radial velocity in figure 14.

Another characteristic visible in figure 9, if compared with figure 10, the model data also presents an accuracy placing high reflectivity values in a region very close to the observation by the radar.

In this case, the 3DVAR among the methods studied (figure 5) was the only one that presented these signs and features for forecast in 1 hour (figure 16), because the other methods of assimilation in the analysis did not present these signals and characteristics that can be observed in the radar data, and with a smaller spatial precision when compared including propagation over time (figure 16).

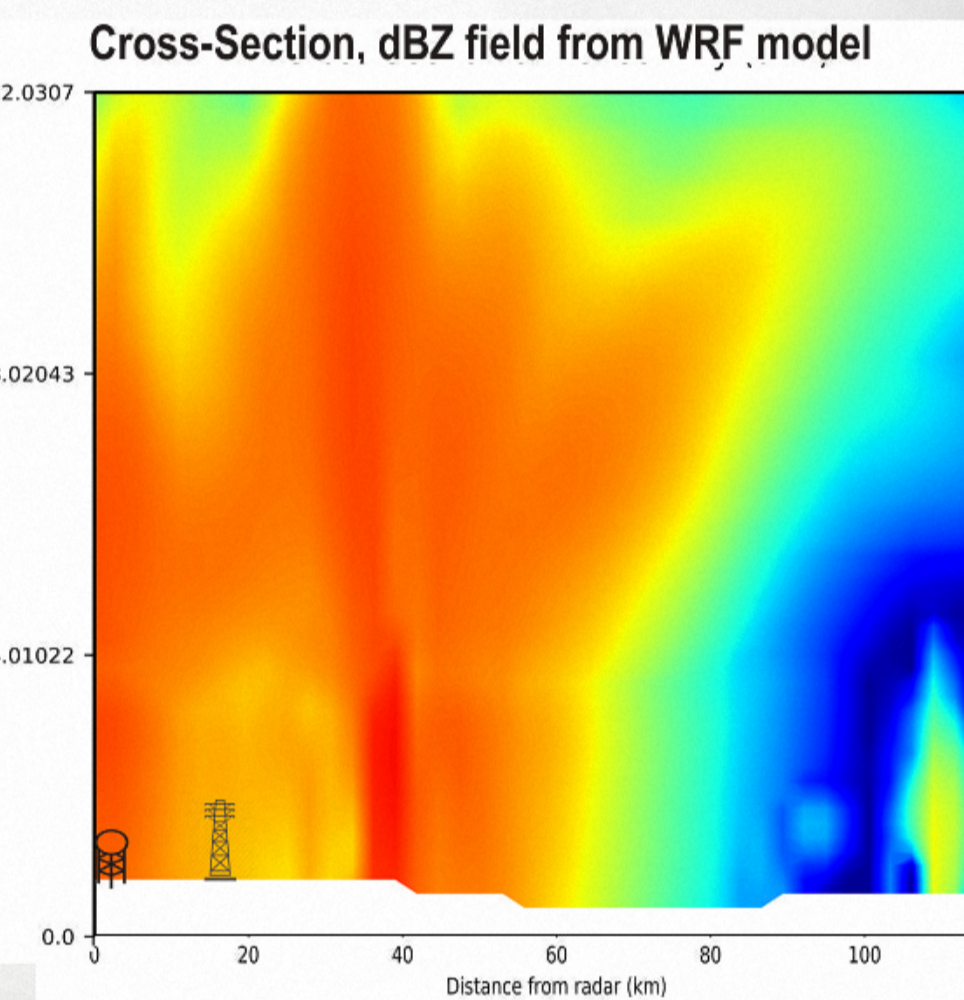


Figure 9. The figure above represents the dBZ field from WRF model with 3DVAR method on 2017/07/13 00UTC, the icons represent the approximate position of the radar, and the tower of high transmission

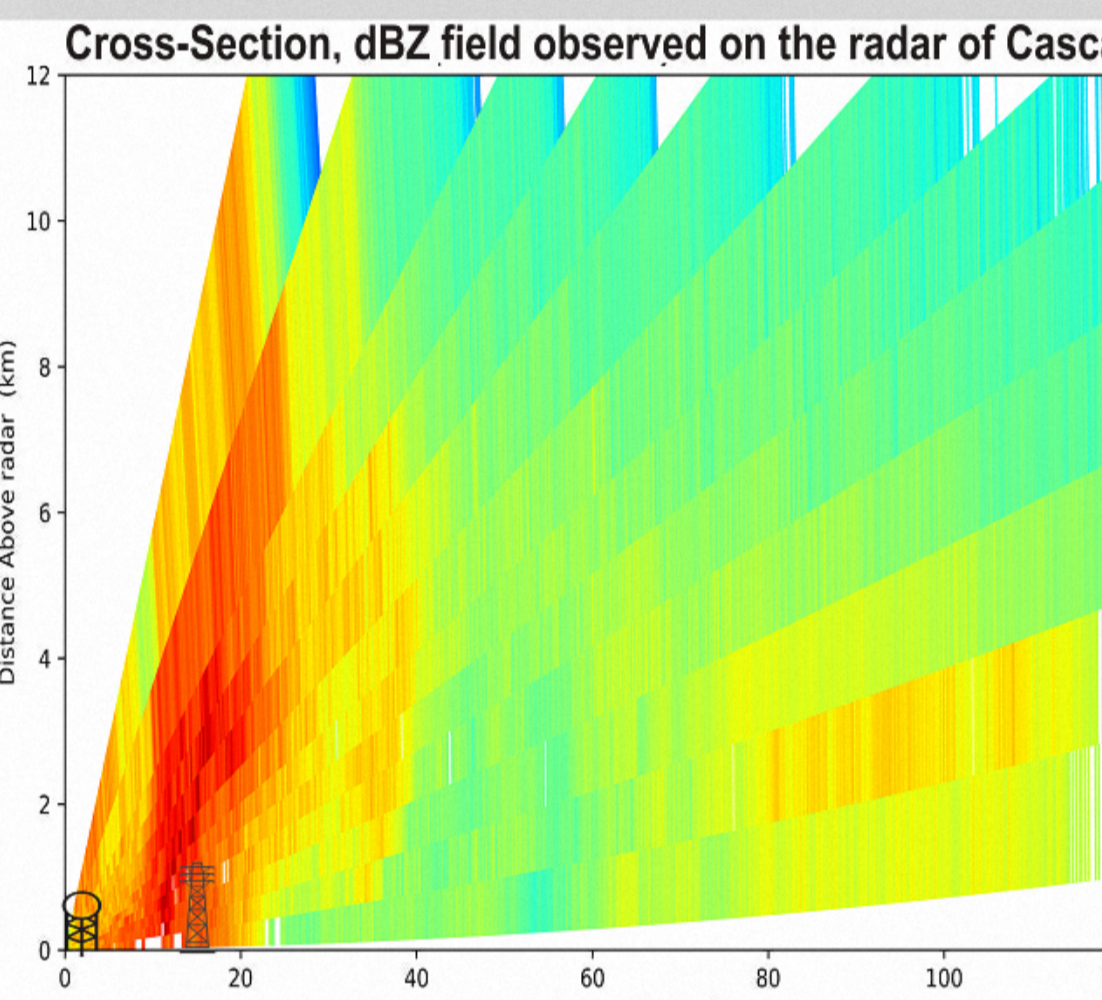


Figure 10. The figure above represents the dBZ field of the Cascavel radar on 2017/07/13 00UTC, the icons represent the approximate position of the radar, and the tower of high transmission

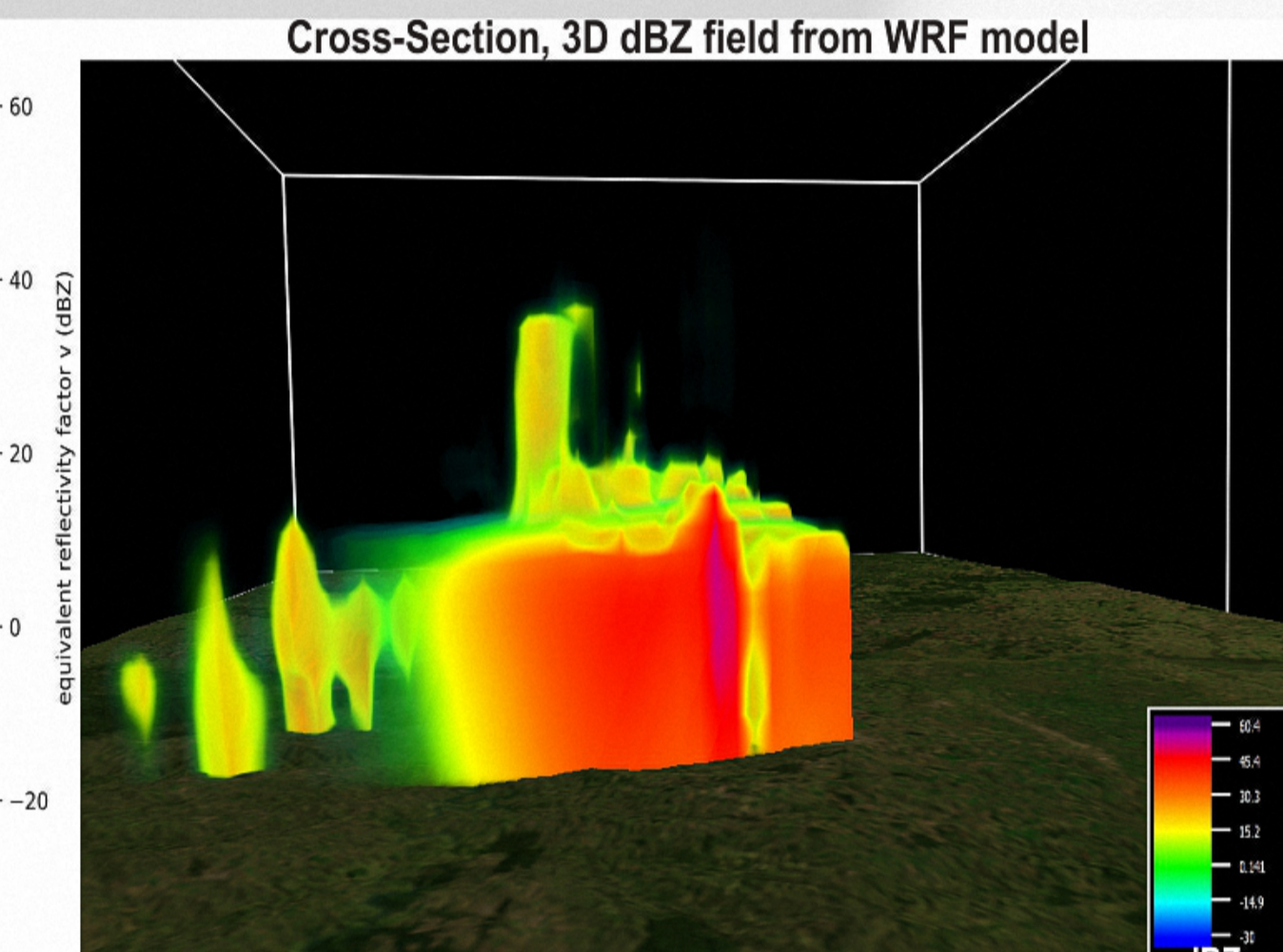


Figure 11. The shaded is the dBZ field from the model, with the method 3DVAR, on 2016/07/13 00 UTC, the area forward with a cut is the same indicated area of the figure 6.

Conclusions

To gain better insight into the evolution of the storms the radar data assimilation shows essentials in this experiment allows simulations and detect more accurately dynamical and thermodynamical structures inside the supercell, including finding signatures of the phenomenon, and is apparent from the results with assimilation, they have created an environment on a temporal and spatial scale very close to that observed (especially with 3DVAR) but with different results between the experiments.

Another point that is observable was the propagation of the supercell (figure 16), which can be represented with high precision in forecast, especially with 3DVAR analyze, but with a very high variation from method to method.

Future studies are needed to explore the impact different variational methods mainly for nowcasting, and the impact of change of convergence criterion (EPS) between the methods.

References

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