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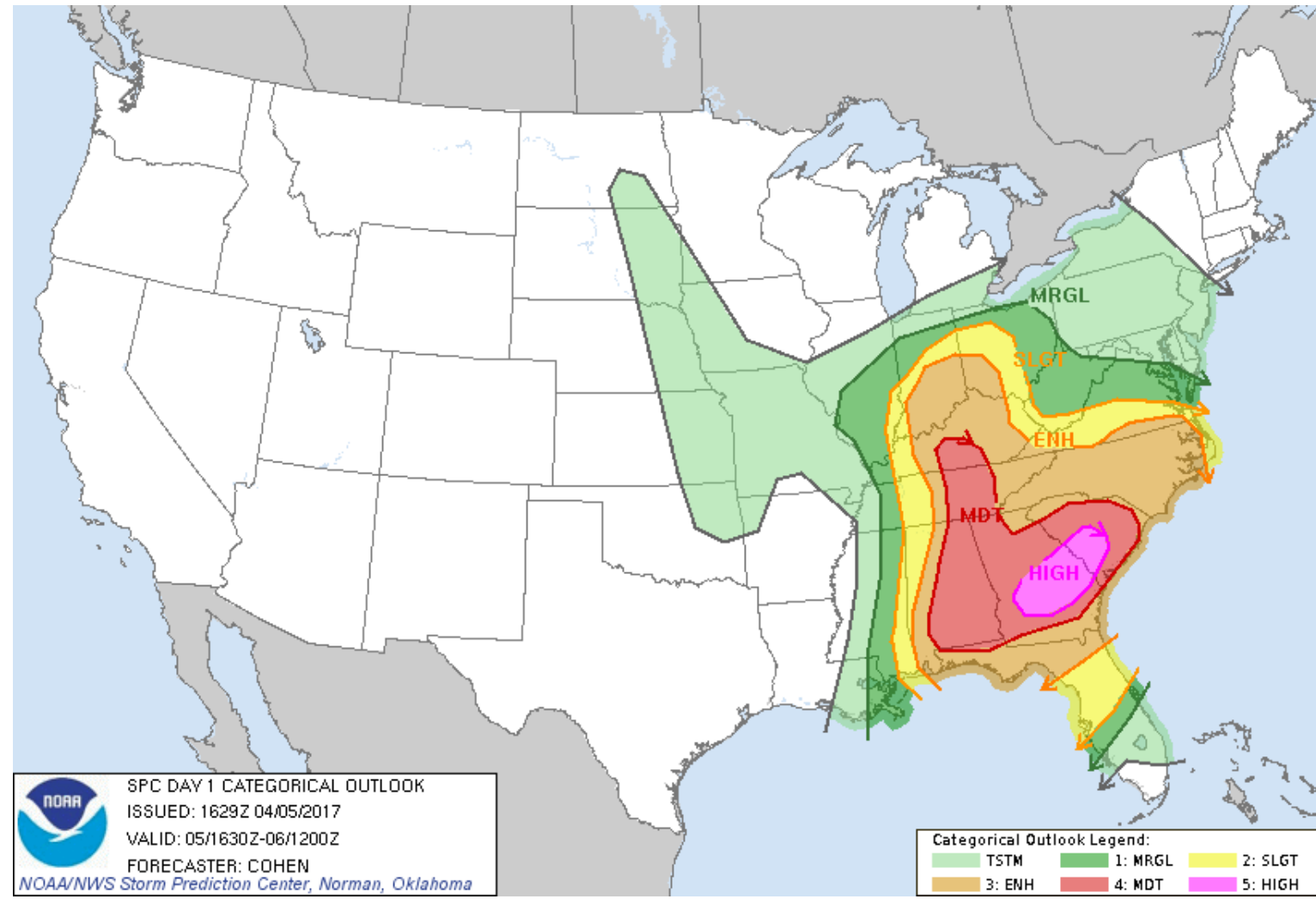
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## IOP3b (5 April 2017) Summary



- Moderate risk in place across far eastern portions of the VORTEX-SE domain with 15% (hatched) tornado probabilities

- Sfc cold pool from predawn elevated convection kept sfc layer stable for much of the day

- Late afternoon profiles exhibited weak-to-moderate sfc instability but maintained  $\sim 100 \text{ J kg}^{-1}$  of CIN

- Bulk shear near 80 kts with low-level hodographs that appeared favorable for supercells & tornadoes

- However, wind profiles were similar to the composite VORTEX2 nontornadic sounding (Coffer and Parker 2016)

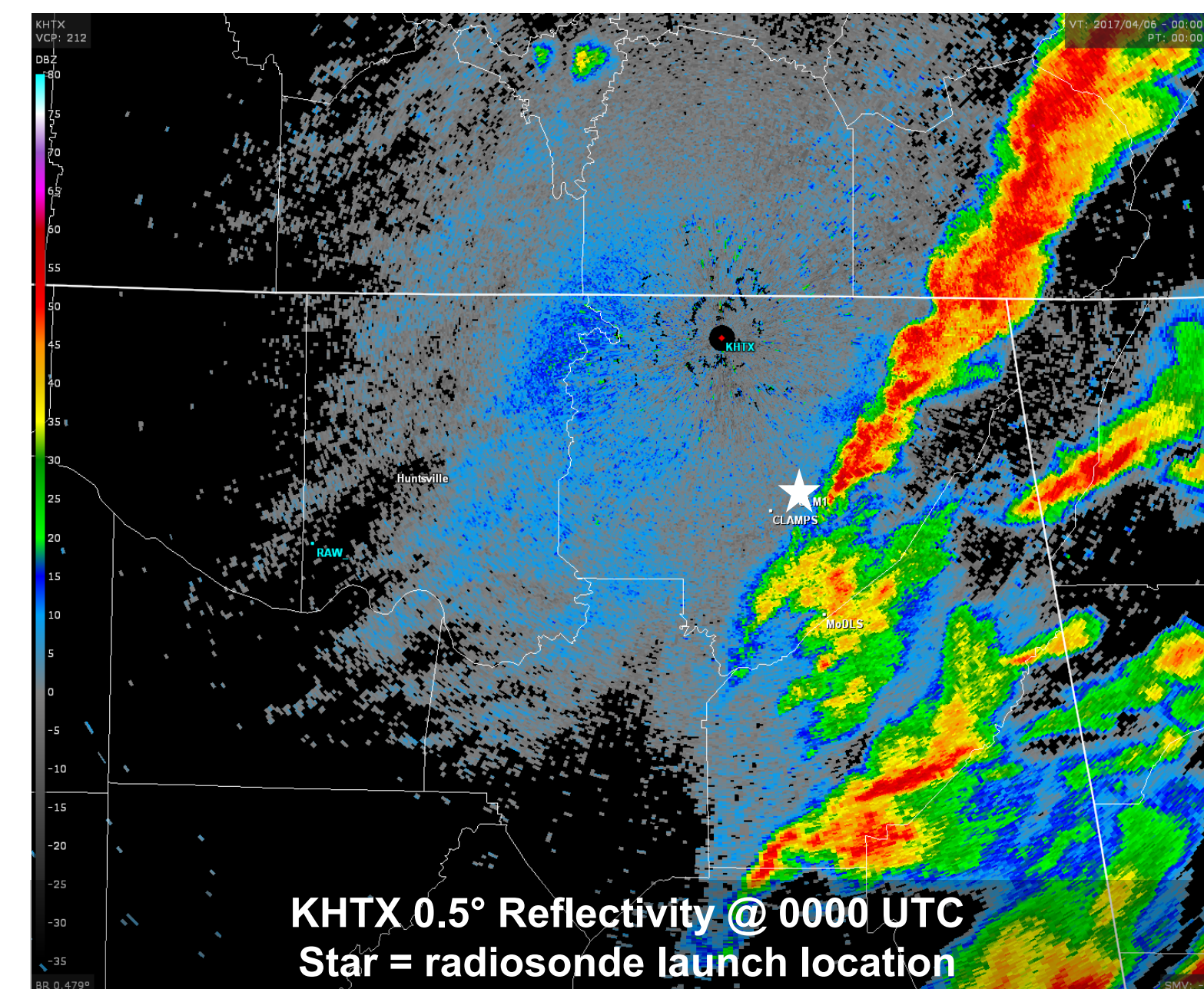
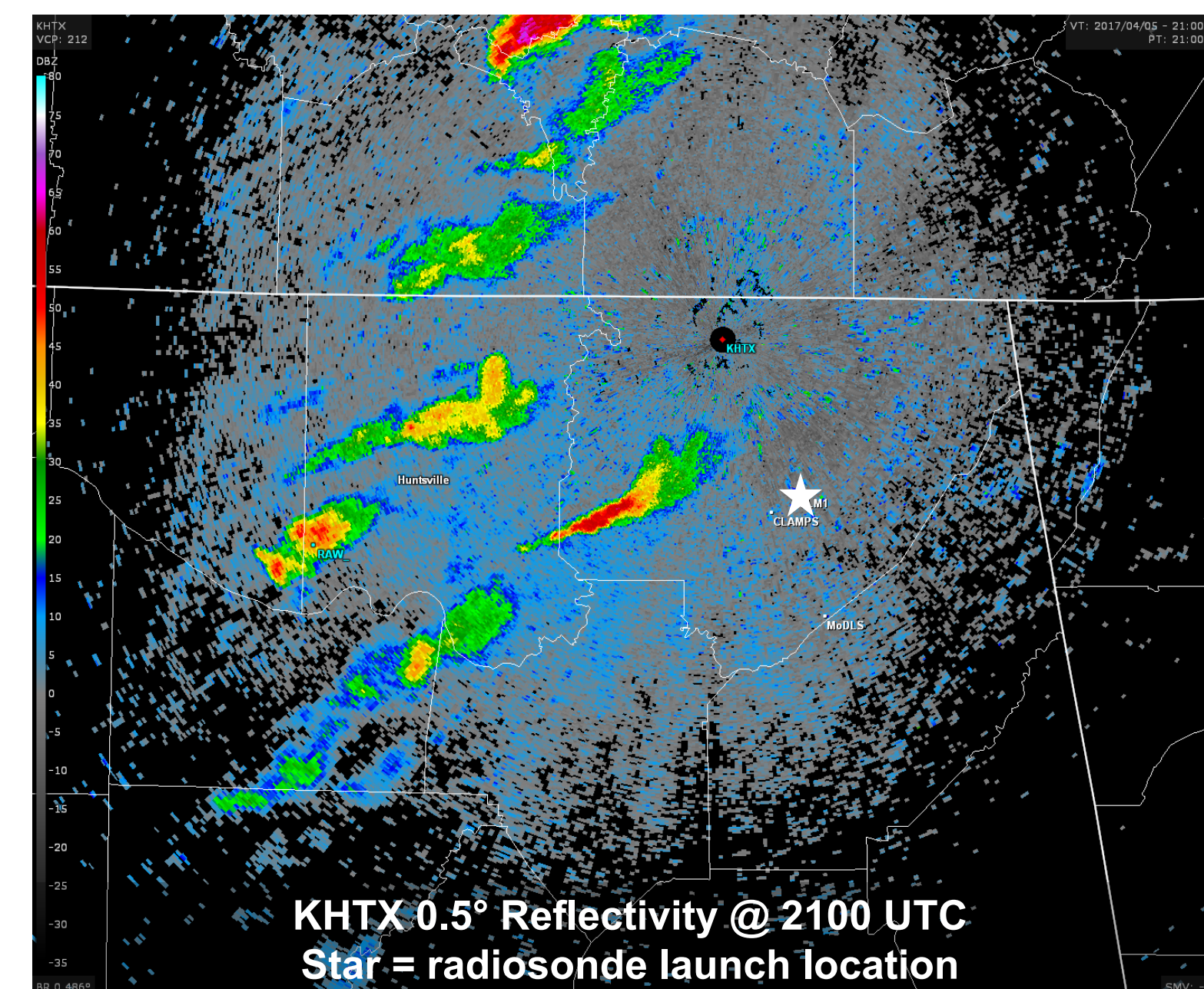
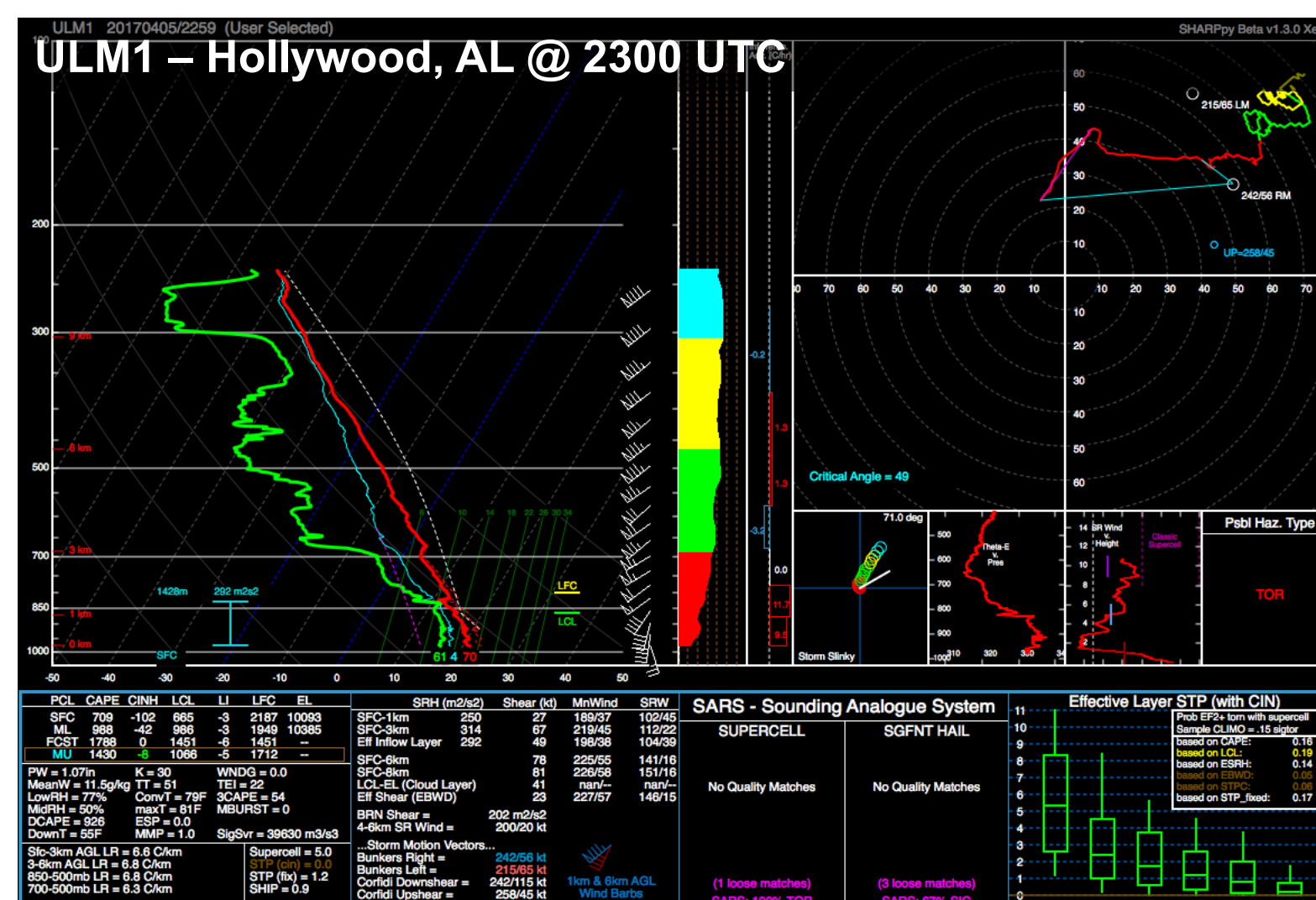
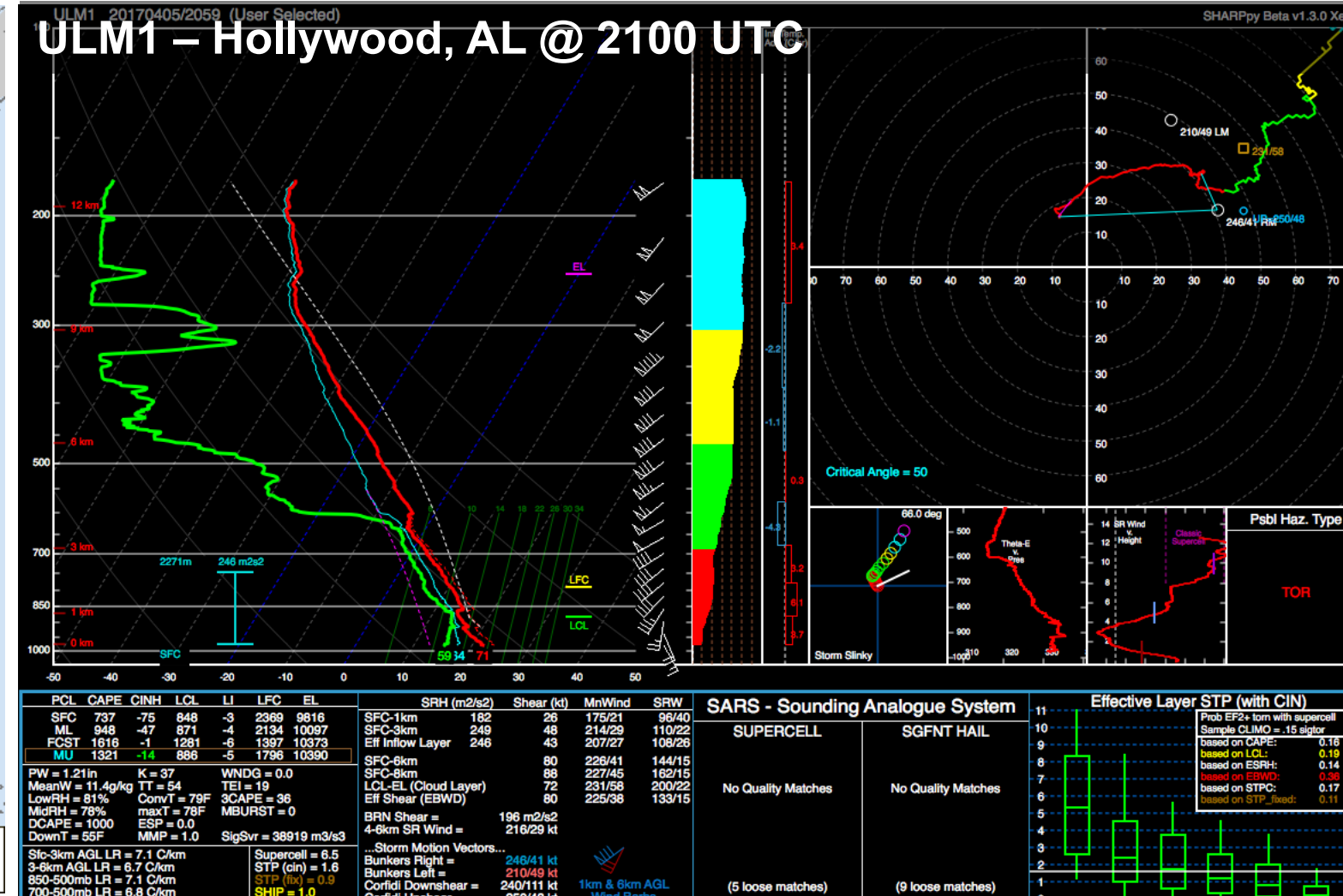
- A surface boundary combined with slightly unstable profiles lead to, scattered CI during the late afternoon

- Convective cells quickly became elongated as updrafts had difficulty organizing & sustaining under the weak instability and strong shear (KHTX reflectivity @ 2100 UTC)

- Convective coverage and intensity significantly increased near 00 UTC, as the surface boundary moved through the complex terrain of NE Alabama (KHTX reflectivity @ 0000 UTC)

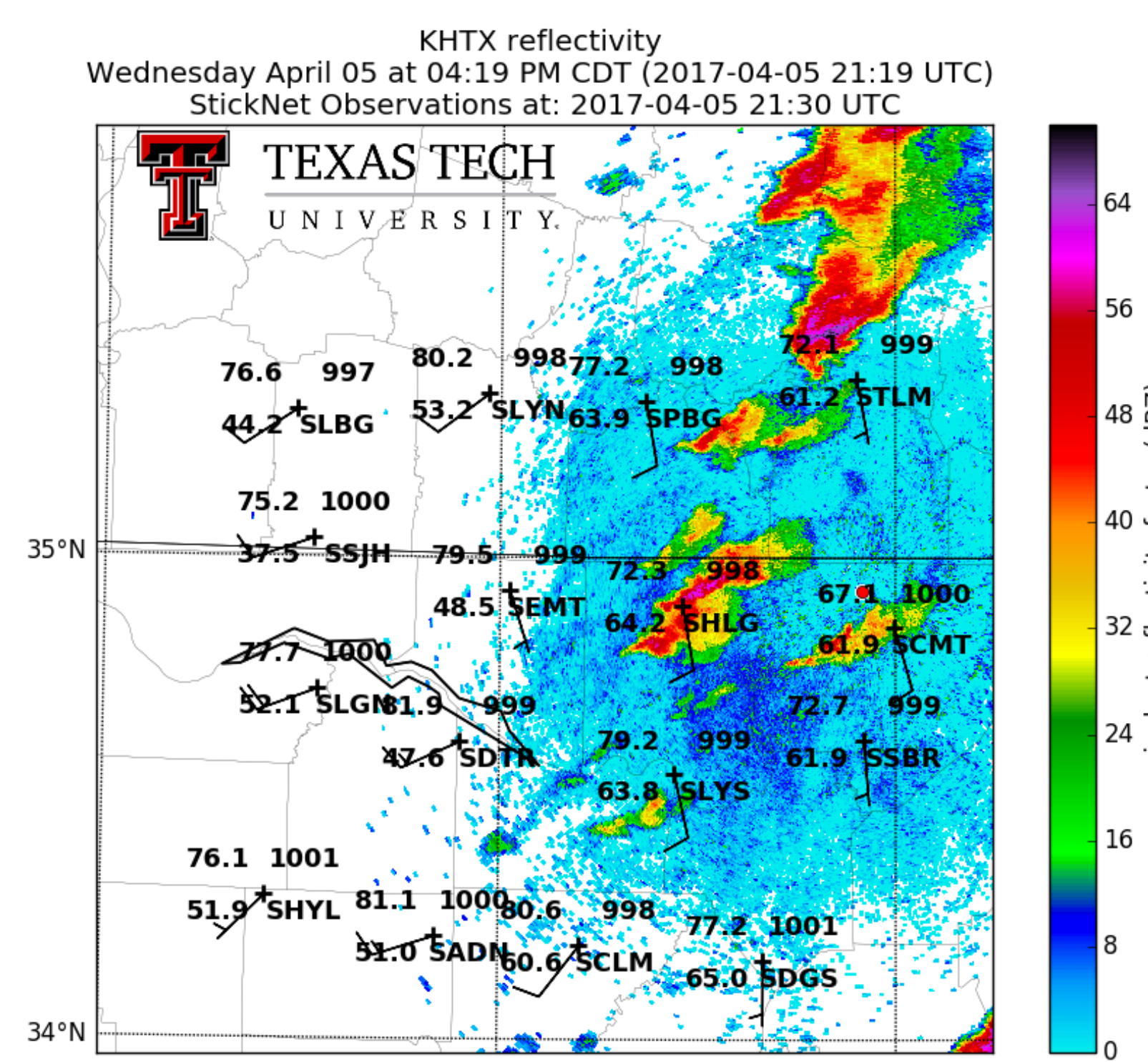
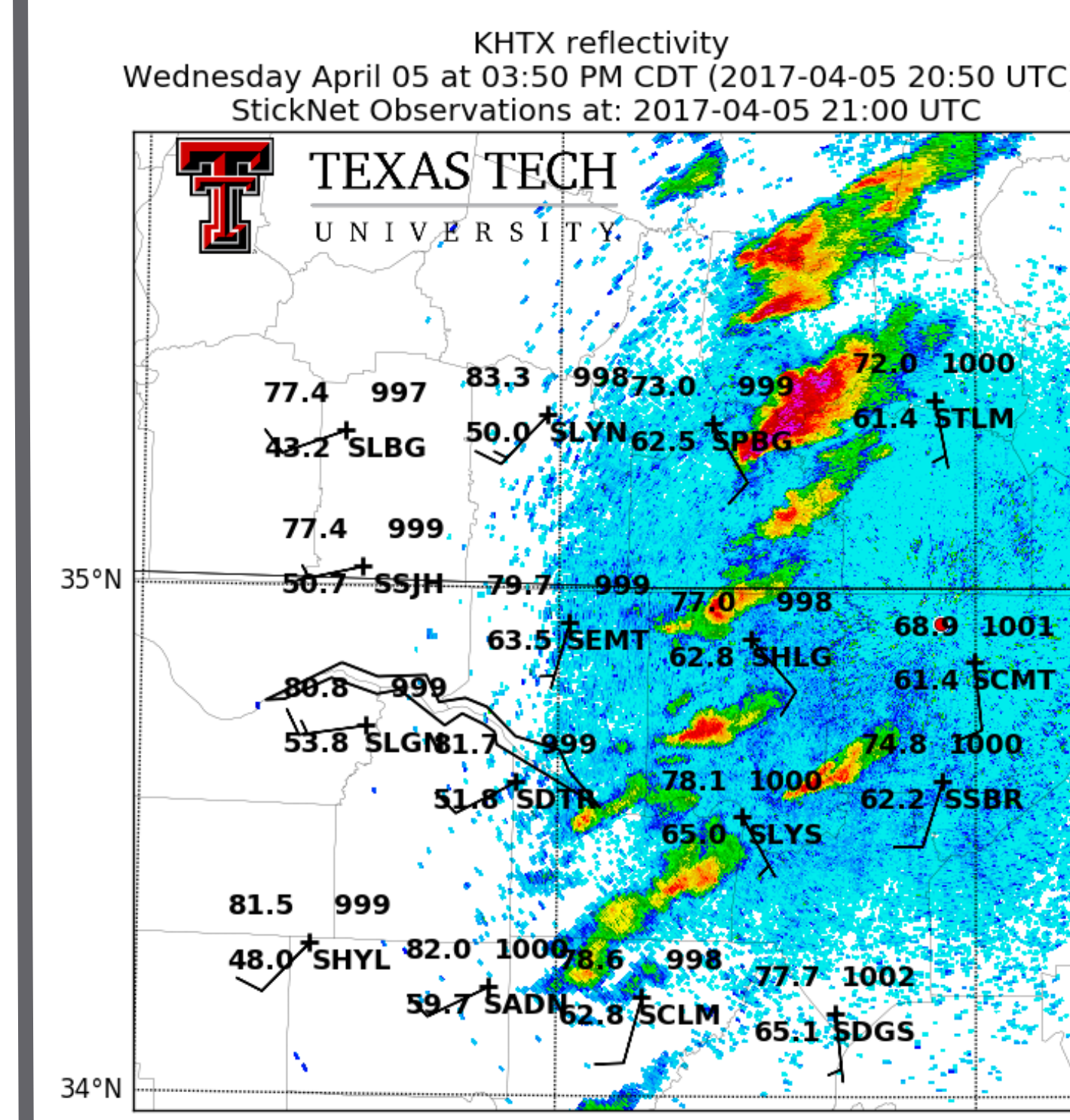
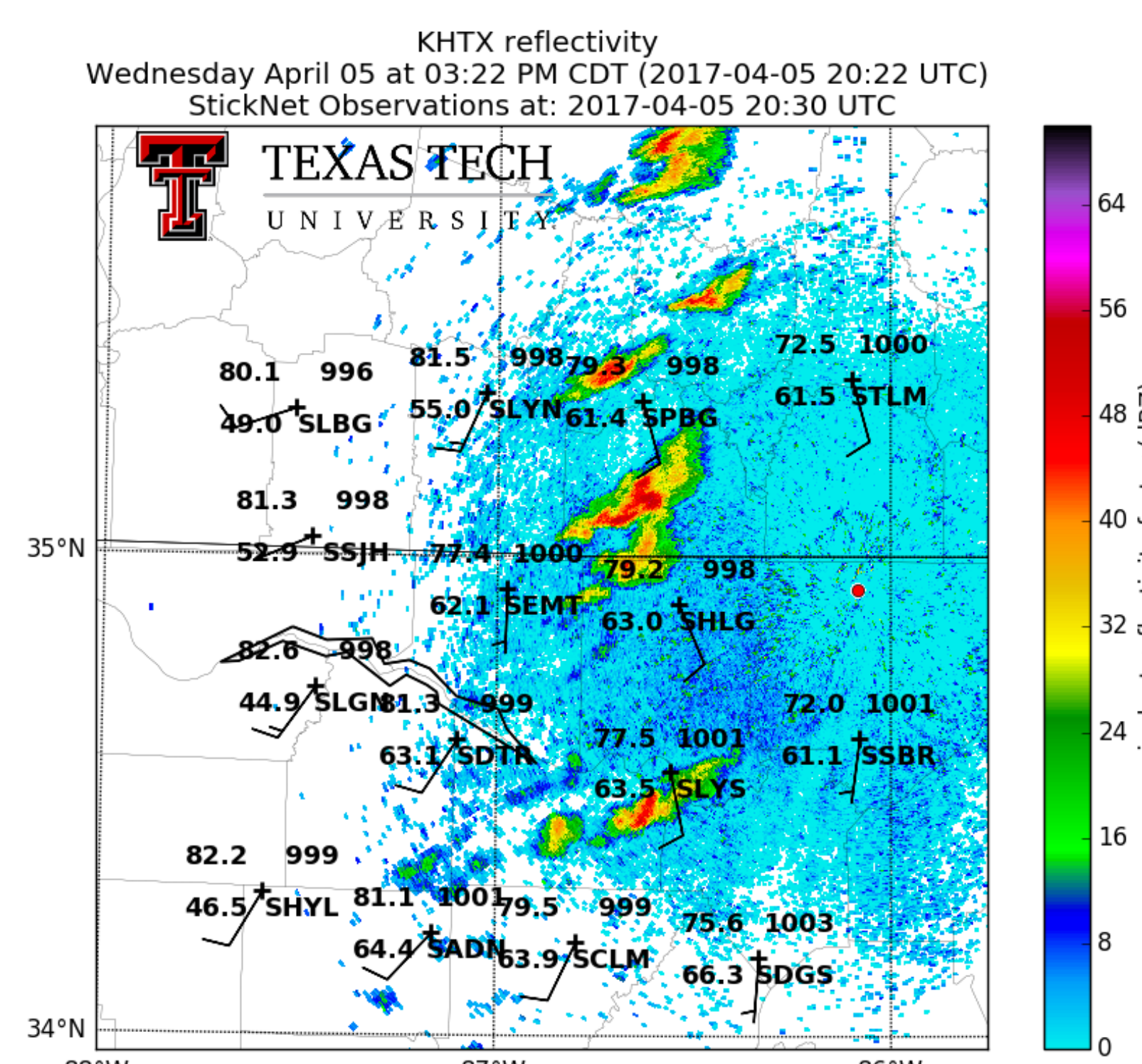
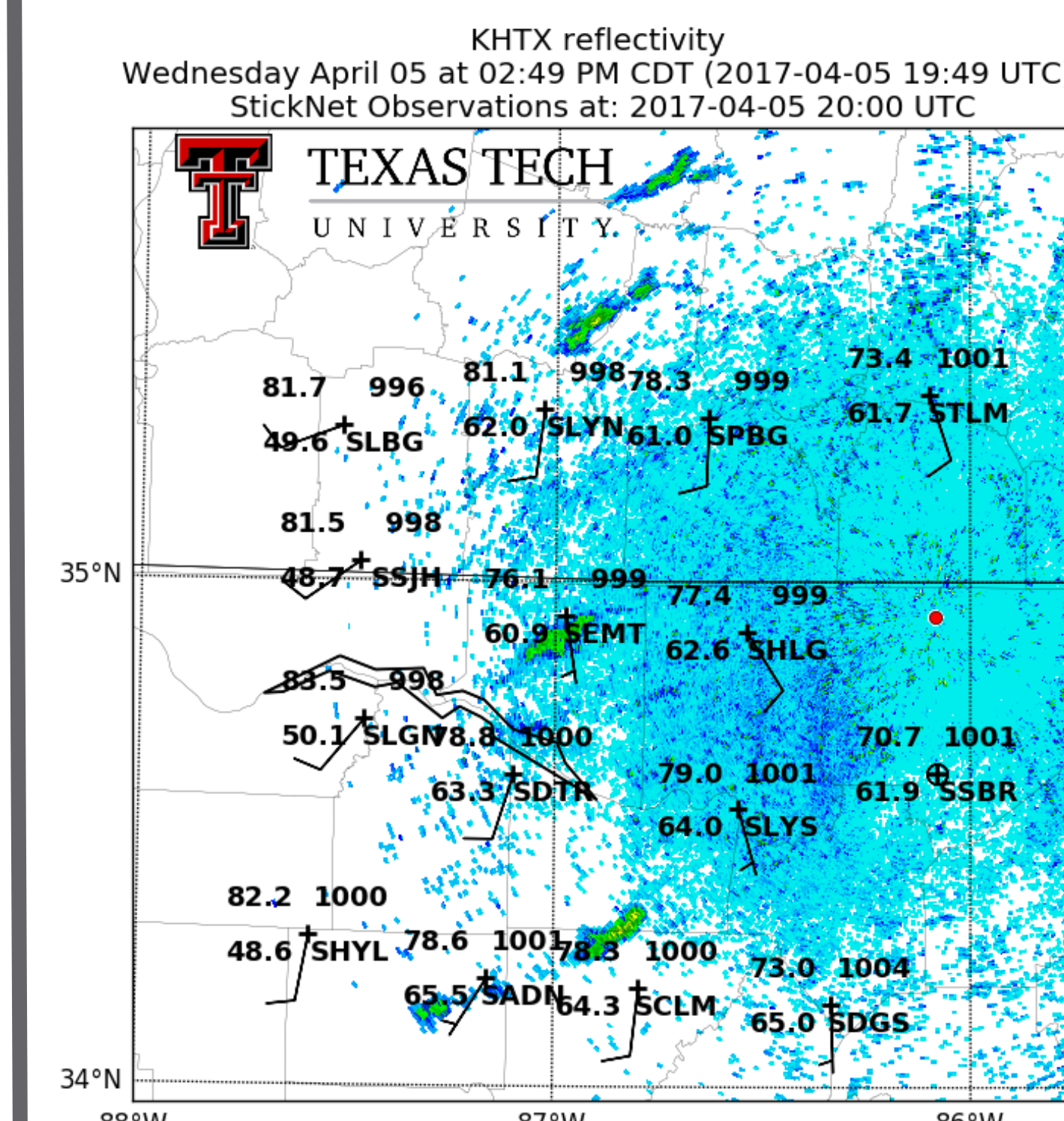
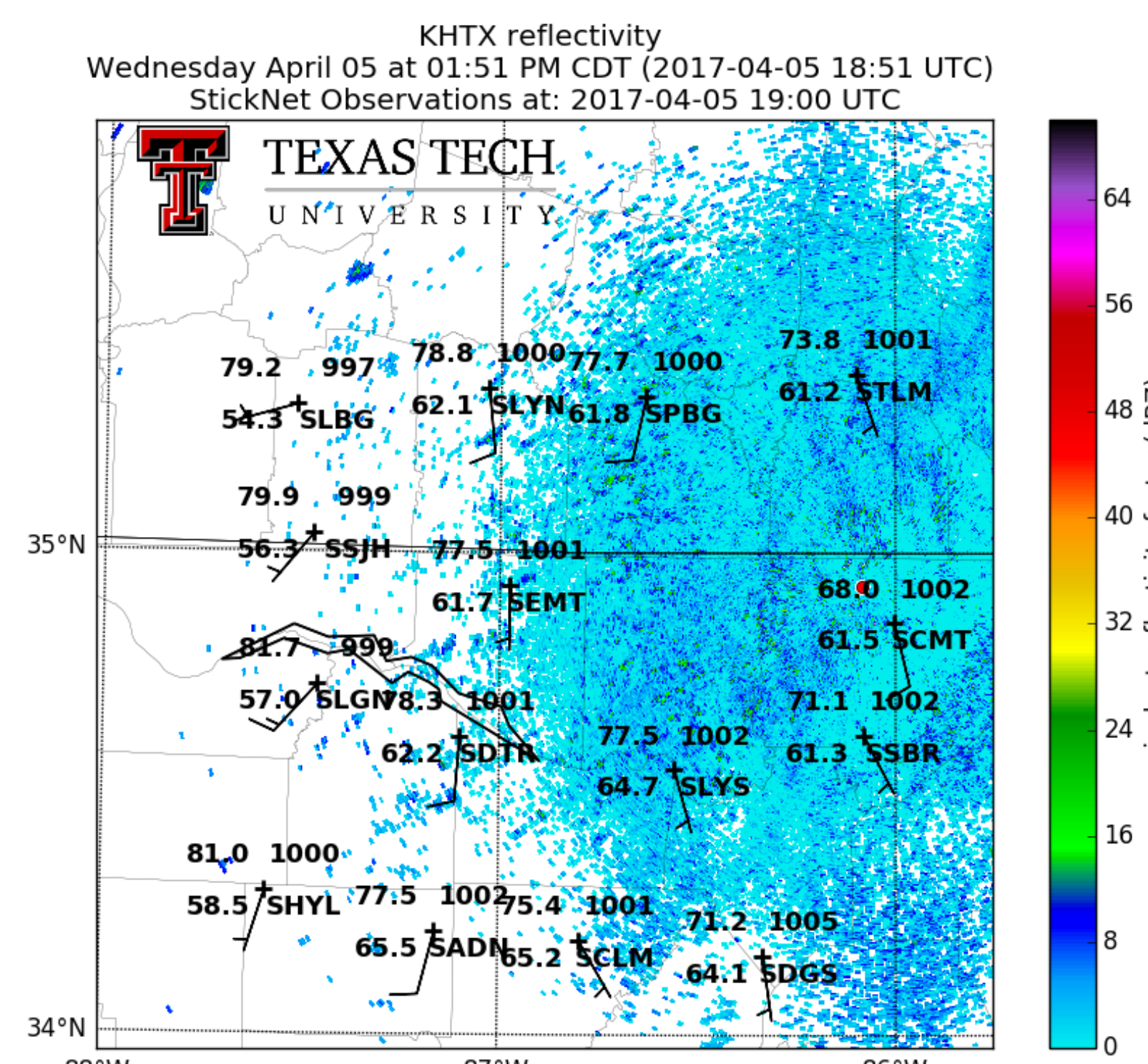
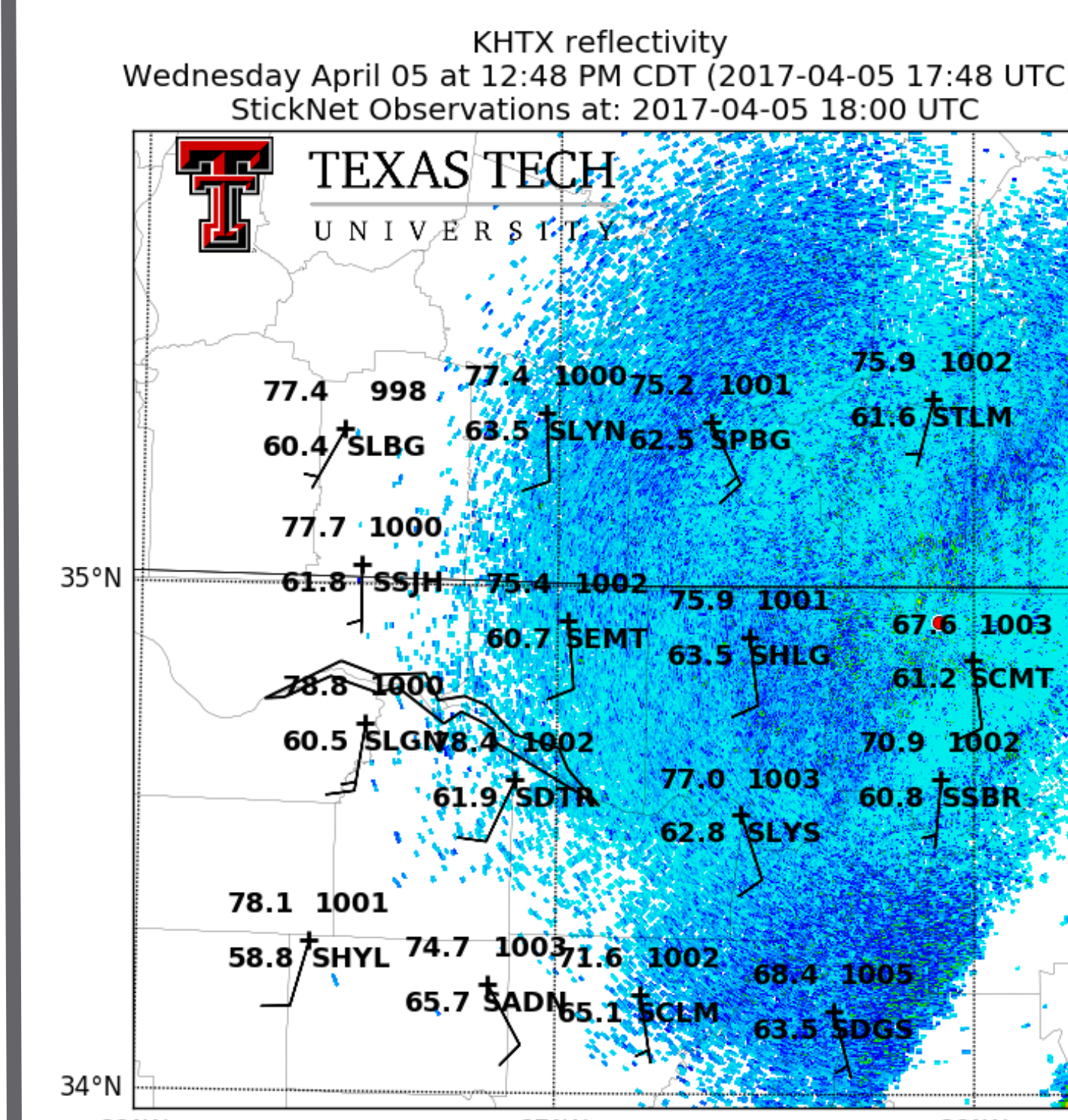
**Table 1:** Sounding Data from ULM Team 1 in Hollywood, AL

Time (UTC)	SBCAPE (CIN) ( $\text{J kg}^{-1}$ )	MLCAPE (CIN) ( $\text{J kg}^{-1}$ )	0-1 km SRH ( $\text{m}^2 \text{s}^{-2}$ )	0-3 km SRH ( $\text{m}^2 \text{s}^{-2}$ )	0-6 km Shear (kt)
1115	0 (0)	252 (-89)	237	292	80
1600	0 (0)	26 (-51)	84	203	68
1714	221 (-27)	272 (-18)	4	126	66
2000	153 (-120)	541 (-61)	156	237	75
2100	737 (-75)	948 (-47)	182	249	80
2200	828 (-83)	937 (-55)	185	270	84
2300	709 (-102)	988 (-42)	250	314	78



## Dryline in Alabama?

- Observations from the Texas Tech Stesonet suggests a weak dryline propagated through the VORTEX-SE domain ahead of a primary surface cold front
- Beginning near 1930 UTC,  $\sim 5\text{-}10^\circ \text{ F}$  of drying and  $\sim 3\text{-}5^\circ \text{ F}$  of warming was observed in the western domain along a pressure trough / wind shift line
- By 2030 UTC,  $\sim 20^\circ \text{ F}$  dew point gradient was observed across the domain; temperatures in the western domain remained steady or slightly increased
- Convection was initiated along this surface boundary around 1945 UTC



## Observations in Complex Terrain

- OU/NSSL CLAMPS (black star) located in TN River Valley
- UAH MoDLS (yellow star) located atop Sand Mountain Plateau
- $\sim 18 \text{ km}$  horizontal distance;  $\sim 200 \text{ m}$  elevation rise
- ULM Sounding Team 1 (white star)

- 0-1 km valley wind profiles  $5\text{-}10 \text{ m s}^{-1}$  stronger than those atop the plateau

- Winds in lowest 200 m noticeably stronger in the valley

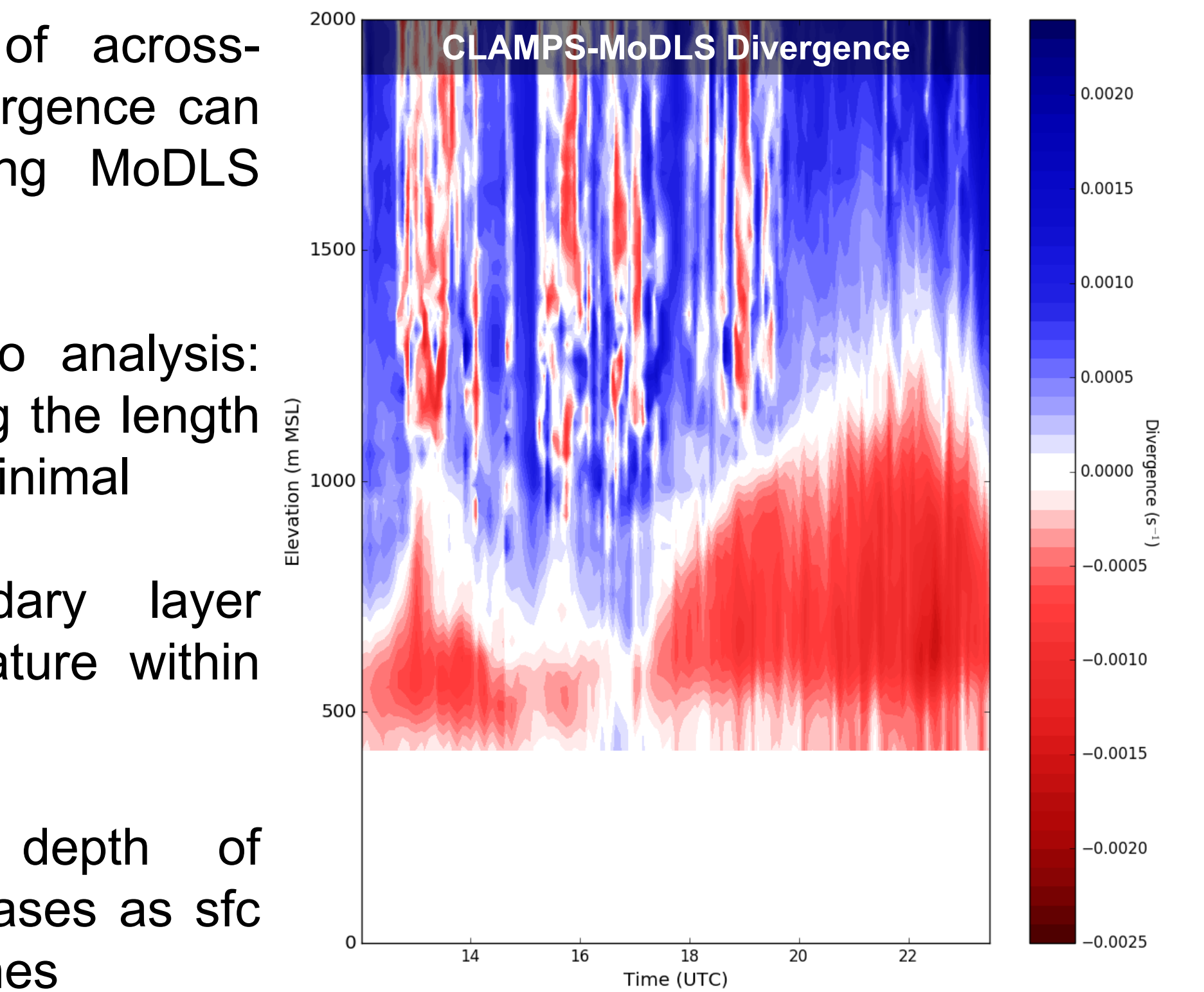
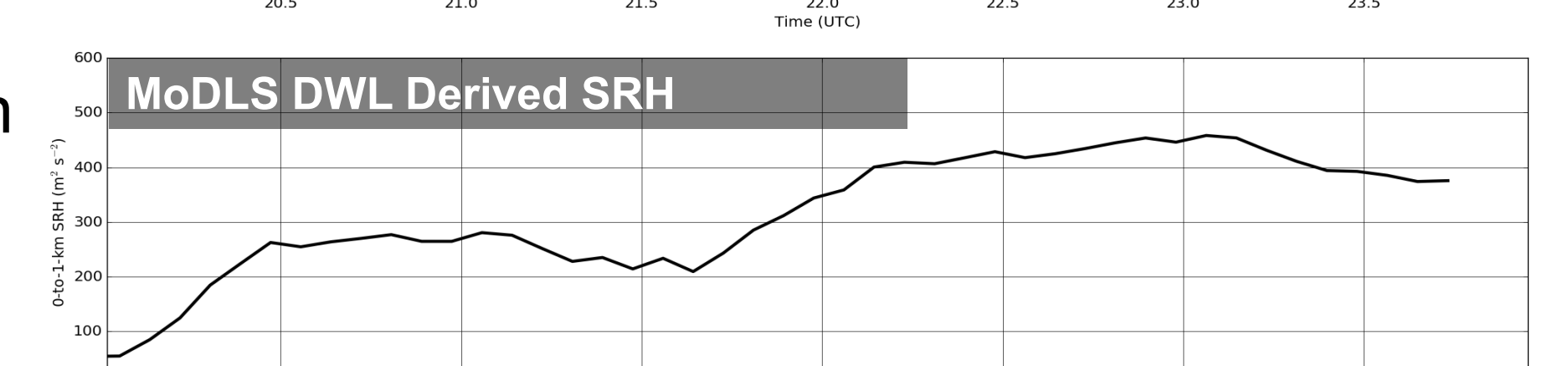
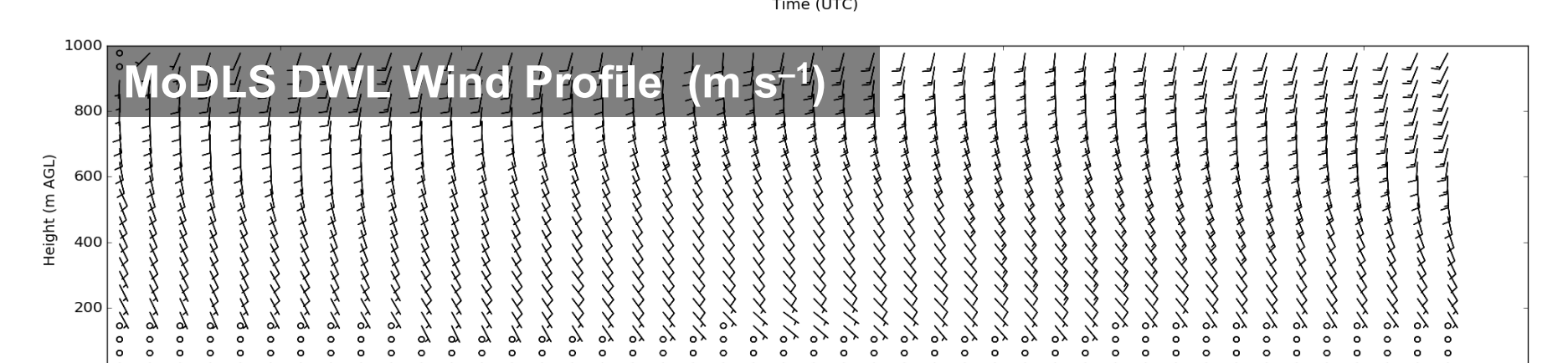
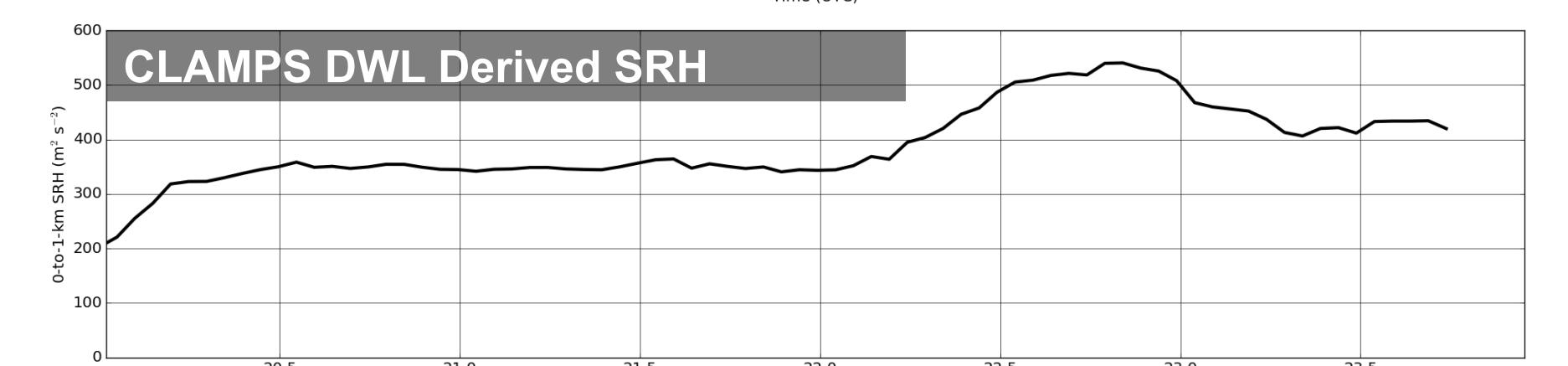
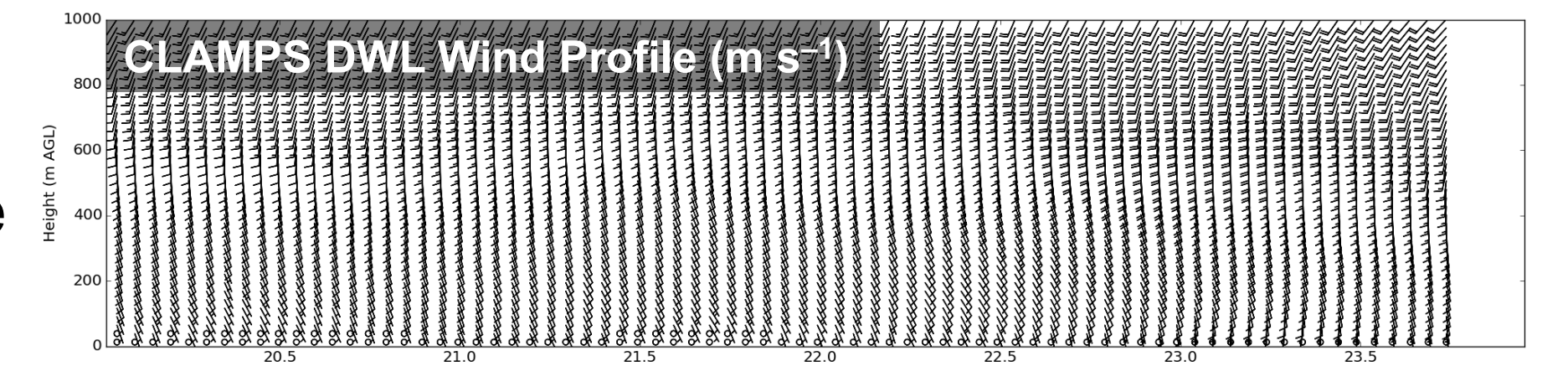
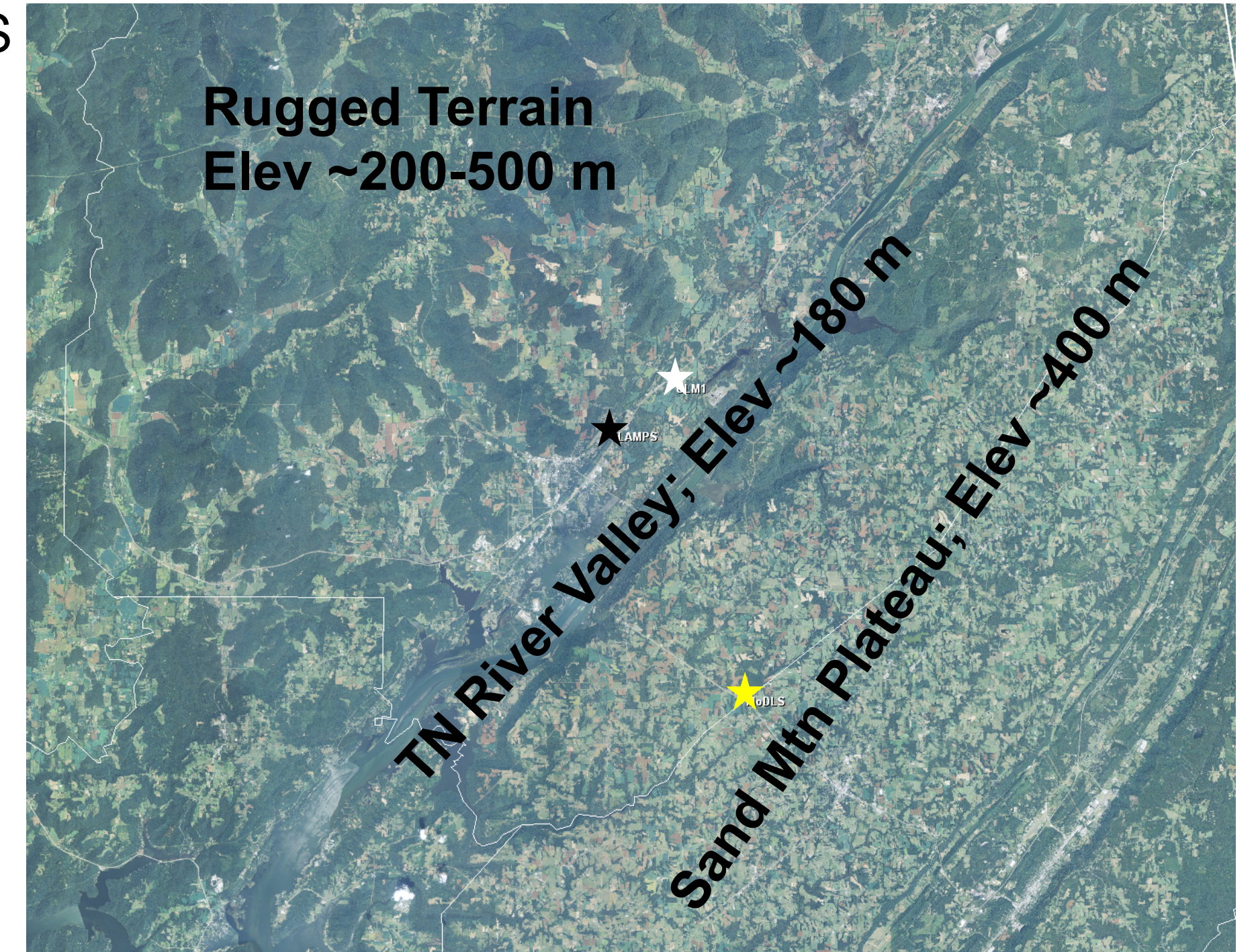
- As sfc boundary approached, valley winds veered from SSE to S, while plateau winds remained backed in the lowest 500 m

- The contribution of across-terrain flow to divergence can be estimated using MoDLS and CLAMPS data

- Key assumption to analysis: flow changes along the length of the terrain are minimal

- Persistent** boundary layer convergence signature within the terrain

- Magnitude and depth of convergence increases as sfc boundary approaches



## Future Work

- Analyses of divergence in western part of domain; compare with terrain analyses
- Multi-Doppler analyses of convection in the terrain
- Further analyses of radiosonde datasets; specifically differences in wind profiles in river valley vs plateau