

Optimized polarimetric radar relations for snow estimation

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#### **Current state of radar snow estimates**



Source	Z(S) relation for dry snow
Gunn and Marshall (1958)	$Z = 448S^2$
Sekhon and Srivastava (1970)	$Z = 399S^{2.21}$
Ohtake and Henmi (1970)	$Z = 739S^{1.7}$
Puhakka (1975)	$Z = 235S^2$
Koistinen et al. (2003)	$Z = 400S^2$
Huang et al. (2010)	$Z = (106-305)S^{(1.11-1.92)}$
Szyrmer and Zawadzki (2010)	$Z = 494S^{1.44}$
Wolfe and Snider (2012)	$Z = 110S^{2}$
WSR-88D, Northeast	$Z = 120S^{2}$
WSR-88D, north plains-upper Midwest	$Z = 180S^{2}$
WSR-88D, high plains	$Z = 130S^{2}$
WSR-88D, Intermountain West	$Z = 40S^{2}$
WSR-88D, Sierra Nevada	$Z = 222S^2$

The multiplier in the power-law relations varies by an order of magnitude, no clear path how to chose optimal S(Z)



Multitude of S(Z) relations (Bukovčić et al. 2018)

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#### **Basic formulas**



$$S = 610^{-4} \pi \int_{0}^{D_{\text{max}}} \frac{\rho_{s}(D)}{\rho_{w}} D^{3} V_{t}^{(s)}(D) N(D) dD \sim M_{2+\gamma}$$

#### **Radar reflectivity**

$$Z = \frac{|K_{\rm i}|^2}{|K_{\rm w}|^2} \int_0^{D_{\rm max}} \frac{\rho_{\rm s}^2(D)}{\rho_{\rm i}^2} D^6 N(D) dD \sim M_4$$

$$M_n = \int D^n N(D) dD$$

$$S \sim f_{rim}^{0.12} N_{0s}^{0.35} Z^{0.62}$$

The multiplier in the S(Z) relation changes more than an order of magnitude because N0s varies by 4 orders of magnitude

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Snow size distribution

 $N(D) = N_{0s}D^{\mu} \exp(-\Lambda_s D)$ 

#### **Snow density**

$$\rho_{\rm s}(D) = \alpha_{\rm u} f_{\rm rim} D^{-1}$$

(frim is the degree of riming)

Snow fall velocity  $V_t^{(s)} \sim D^{\gamma}$ 

#### Analysis of snow disdrometer data







**Specific differential phase** 

$$K_{\rm DP} = \frac{0.27\pi}{\lambda \rho_i^2} \left(\frac{\varepsilon_{\rm i} - 1}{\varepsilon_{\rm i} + 2}\right)^2 F_{shape} F_{orinet} \int \rho_s^2(D) D^3 N(D) dD \sim M_{\rm F}$$

Z is proportional to the 4<sup>th</sup> moment of snow SD whereas KDP is proportional to its 1<sup>st</sup> moment

Number concentration

$$N_{\rm t} = \int N(D) dD \sim M_0$$

Z is proportional to the 4<sup>th</sup> moment of snow SD whereas N<sub>t</sub> is proportional to its 0<sup>th</sup> moment

Shape factor





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 $f_{rim}$  – riming degree,  $F_0 \& F_s$  – particle orientation & shape parameters,  $\mu$  – PSD shape parameter,  $N_{\rm t}$  – number concentration, p – atmospheric pressure,  $\lambda$  – radar wavelength



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 $f_{rim}$ ,  $F_{o}F_{s}$ , and  $N_{t}$  can be retrieved at ASOS stations – snowfall rate + extinction coefficient  $\sigma_{e}$  needed!

$$f_{\text{rimA}}(S_{\text{ASOS}}, \sigma_{\text{eASOS}}, Z, \mu) = const \times f_1(\mu) \frac{S_{\text{ASOS}}^a}{\sigma_{\text{eASOS}}^b Z^c}$$

Z, Kdp – radar 
$$F_{o}F_{s} = F_{osA}(K_{DP}, Z, \sigma_{eASOS}, f_{rimA}, \mu) = const \times f_{2}(\mu) \frac{Z^{g}(K_{DP}\lambda)^{h}}{f_{rimA}^{i}\sigma_{eASOS}^{j}}$$

 $D_{\rm m} = f(f_{\rm rim}, \, \mu, \, \sigma_{\rm e}, \, Z) - mean \, volume \, diameter$ 

 $N_{\rm t} = 2f_3(\mu)\sigma_{\rm e}/(\pi D_{\rm m}^2)$ 

S,  $\sigma_e - ASOS$ 

We can use ASOS measurements to directly evaluate "goodness" of our relations – how much the best possible radar estimate deviates from the ASOS measurements of S



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### **Example of f<sub>rim</sub> and F<sub>o</sub>\*F<sub>s</sub> retrieval**





Constant  $F_0^*F_s$  factor deviates significantly from the retrieved  $F_0^*F_s$ , potentially introducing large biases.



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Addition of dynamically estimated  $N_{t}$ ,  $f_{rim}$ ,  $F_{o}$ , and  $F_{s}$  introduces large improvement. Adjusting  $\mu$  from 0 to -0.6 in Sopt.(Kdp, Z) provides even better results for this event.



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Addition of dynamically estimated *N*<sub>t</sub>, f<sub>rim</sub>, *F*<sub>o</sub>, and *F*<sub>s</sub> introduces large improvement. Adjusting μ from 0 to -0.4 in Sopt.(Kdp, Z) provides even better results for this event.



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Addition of dynamically estimated  $N_{\rm t}$ ,  $f_{\rm rim}$ ,  $F_{\rm o}$ , and  $F_{\rm s}$  introduces large improvement.



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Addition of dynamically estimated  $N_{t}$ ,  $f_{rim}$ ,  $F_{o}$ , and  $F_{s}$  introduces large improvement. Adjusting  $\mu$  from 0 to -1 in Sopt.(Kdp, Z) provides even better results for this event.



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Addition of dynamically estimated  $N_{\rm t}$ ,  $f_{\rm rim}$ ,  $F_{\rm o}$ , and  $F_{\rm s}$  introduces large improvement.



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### **Moving forward**



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- Projection of polarimetric S estimates from DGL (-20°C to -10°C), where we have more reliable radar measurements of Kdp and estimates of N<sub>t</sub> to the ground, using particle trajectories
- Using ASOS estimates of  $f_{rim}$ ,  $F_oF_s$ , and  $N_t$  from S and extinction to learn how to optimally adjust the multipliers of proposed relations, and expand information spatially
- To use optimal S(Z) relation according to polarimetric classification (physically-based) instead of current regional dependence



#### References





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