

Mechanism of Tropical Cyclone Frequency Changes Due to Global Warming

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Precipitation change

1. Introduction

Recent climate models (high-resolution AGCMs) consistently project that the global tropical cyclone (TC) frequency will decrease in the future greenhouse gas (GHG) warmed climate, but there is a large uncertainty in the regional TC frequency changes (Knutson et al. 2010). In the present study, the mechanism of global and regional TC frequency changes due to global warming is explored.

2. Mechanism of TC Frequency Changes (Hypothesis)

. The changes in global TC frequency are closely related to the changes in upward mass flux. • Relation between precipitation (P), upward mass flux (ω) and dry static stability (S):

 $\Delta P \quad \Delta S$

P S

ω

Thermodynamic equation: $\omega S \approx \alpha P$ (α is a scaling constant),

Effect		Global change					Regional change
		Radiative cooling	Precipitation	Stability	Upward mass flux	TC frequency	TC frequency
CO ₂	CO ₂	Decrease	Decrease		Decrease	Decrease	
	Other GHG	Increase	Increase		Increase	Increase	
SST	Uniform SSTA	Increase	Increase	Increase	Decrease	Decrease	
	Non-uniform SSTA						Shift

3. Model and Experiment

· Five 25year runs are conducted using the MRI-AGCM3.2 with 60km resolution. ·Observed SST (1979-2003) is prescribed for the present climate run, while CMIP3 ensemble mean SSTA (2075-2099) is added for the future climate run.





TC frequency change

4. Global Change and Regional Change (Uniform and Non-uniform SSTA effects)

the reduction of global TC frequency. · Most important remaining question is how the reduction of upward mass flux causes the reduction of TC frequency.

Upward mass flux change

Regional TC frequency change

- · Non-uniform SSTA effect (effect of SSTA relative to the tropical mean SSTA) is mainly responsible for the regional TC frequency.
- · Non-uniform SSTA effect causes shift of active convection area, leading to the shift of TC genesis area. It causes little change in global TC frequency.

Acknowledgement

This work was conducted under the framework of the "Projection of the Change in future Weather Extremes using Super-high-resolution Atmospheric Models" supported by the KAKUSHIN Program of the Ministry of Education, Culture, Sports, Science, and Technology (MEXT). The calculations were performed on the Earth Simulator.

References

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Upward mass flux Vertical p-velocity (500hPa)

Precipitation

Vertical shear Vector wind difference (200hPa and 850hPa)

Saturation deficit (600hPa)