

### Surface NO<sub>x</sub> Measurements in Northern Alabama During and After DC3

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**Abstract:** A residual gas nitrogen oxides (NO<sub>x</sub>) analyzer was employed for the purpose of measuring surface NO<sub>x</sub> in Huntsville, Alabama during the Deep Convective Clouds and Chemistry (DC3) campaign. Measurements were taken during a variety of conditions, including storm-free days and nights, non-thunderstorm outflow, and several thunderstorm outflows. In addition to the regular diurnal cycle in NO<sub>x</sub> caused by local anthropogenic activity, thunderstorm outflows produced small but measurable increases in surface NO<sub>x</sub>. The magnitude of NO<sub>x</sub> increase is supportive of the "backwards C" model profile of lightning NO<sub>x</sub>.

#### 1. INTRODUCTION

Nitrogen oxide production by lightning has been studied using a multitude of techniques. These

include aircraft, satellite, laboratory, chambers capturing rocket triggered lightning, and modeling studies (Table 1). Largely missing from these studies are surface observations of NO<sub>x</sub> before and after thunderstorms to measure the impact of lightning on surface NO<sub>x</sub>.

The Franzblau et al. (1989) study included measurements of surface NO<sub>x</sub>. They found that NO<sub>x</sub> from thunderstorm outflows was primarily nitric oxide. Since their spectroscopic measurements of lightning NO<sub>x</sub> were solely an NO<sub>2</sub> measurement, the surface measurements were used to apply a correction factor, thus increasing their estimates of NO<sub>x</sub> per lightning flash.

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First author	Year	Methodology	Molecules/flash	Moles/flash
Levine	1981	Laboratory	5.00E+24	8.30
Kumar	1995	Field study	5.00E+24	8.30
Dawson	1980	Theoretical	8.00E+24	13.28
Beirle	2010	Satellite	1.00E+25	16.61
Tuck	1976	Theoretical	1.10E+25	18.27
Hill	1980	Theoretical	1.20E+25	19.93
Koshak	2010	Theoretical	1.41E+25	23.40
Cooray	2009	Theoretical	2.00E+25	33.21
Lawrence	1995	Review	2.30E+25	38.19
Nesbitt	2000	Field study	2.67E+25	44.25
Huntrieser	2002	Field study	2.70E+25	44.84
Wang	1998	Laboratory	3.10E+25	51.48
Peyroux	1982	Laboratory	3.20E+25	53.14
Ridley	2004	Field study	3.20E+25	53.14
Beirle	2006	Satellite	5.40E+25	89.67
Koshak	2011a	Theoretical	6.09E+25	101.17
Koshak	2011b	Theoretical	7.23E+25	121.61
Sisterson	1990	Theoretical	8.20E+25	136.17
Noxon	1976	Field study	1.00E+26	166.06
Chameides	1977	Theoretical	1.00E+26	166.06
Kowalczyk	1982	Theoretical	1.00E+26	166.06
Bucsela	2010	Satellite	1.05E+26	174.36
Schumann	2007	Review	1.50E+26	249.09
Huntrieser	2011	Field study	1.51E+26	250.00
DeCaria	2000	Theoretical	1.56E+26	258.39
Fehr	2004	Field study	2.10E+26	348.72
Rahman	2007	Field study	2.40E+26	398.54
Chameides	1979	Theoretical	2.50E+26	415.14
DeCaria	2005	Theoretical	2.77E+26	460.00
Martini	2011	Theoretical	2.89E+26	480.00
Ott	2010	Theoretical	3.01E+26	500.00
Jourdain	2010	Theoretical	3.13E+26	520.00
Drapcho	1983	Field study	4.00E+26	664.23
Franzblau	1989	Field study	3.00E+27	4981.73

Table 1: A comparison of lightning NO<sub>x</sub> production from previous studies, adapted from Peterson and Beasley (2011).

## 2. METHODOLOGY

Nitrogen oxides were detected using a residual gas analyzer, the function of which is described in Peterson et al. (2009). The purpose of this method is not to reproduce the high speed reactions in lightning, but rather to obtain the overall contribution of lightning to NO<sub>x</sub> (Peterson et al. 2010). The residual gas analyzer was obtained on loan from the Tennessee Valley Authority, who also performed a calibration of the instrument prior to loaning it to us. Measurements were initially taken during the Deep Convective Clouds and Chemistry (DC3) field campaign (Carey et al. 2013), and continued through the rest of 2012.

## 3. RESULTS

To establish the role of lightning among all sources of NO<sub>x</sub> the background must first be known. Figure 1 shows a typical non-thunderstorm summer day in Huntsville. Here the primary source of NO<sub>x</sub> is vehicular traffic on Interstate 565 about 1 km south of the measurement site. There tends to be an increase in NO<sub>x</sub> around the time of the morning commute, and a smaller increase in NO<sub>x</sub> around the time of the afternoon commute. The morning peak is larger due to a shallower boundary layer in the morning.

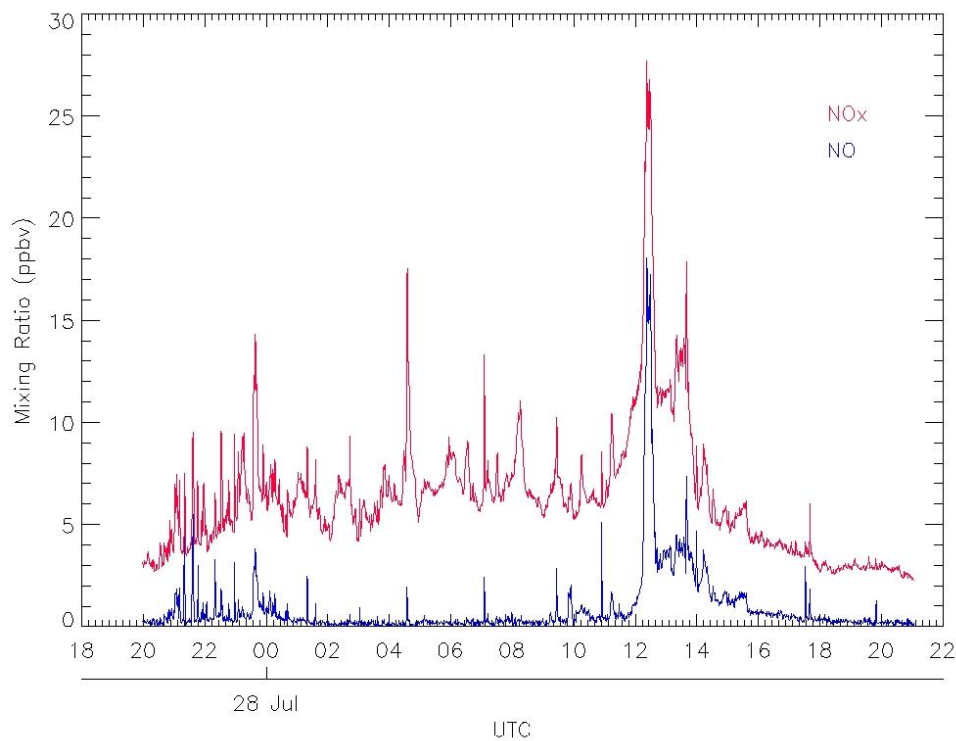


Figure 1: Summer background. Times on the bottom are UTC time. Concentrations on the left hand side are in ppb.

Figure 2 shows how the pattern changes for the cool season months. Two competing effects are at work: 1) a shallower boundary layer in the cooler months leads to higher surface concentrations of NO<sub>x</sub>, and 2) northerly winds due to passing cold fronts cuts off the I-565 source of NO<sub>x</sub>, leading to lower surface

concentrations of NO<sub>x</sub>. In this example southerly winds were present in the afternoon, with an overnight cold frontal passage leading to northerly winds in the morning. As a result the relative strength of NO<sub>x</sub> peaks was reversed, with a higher concentration in the afternoon.

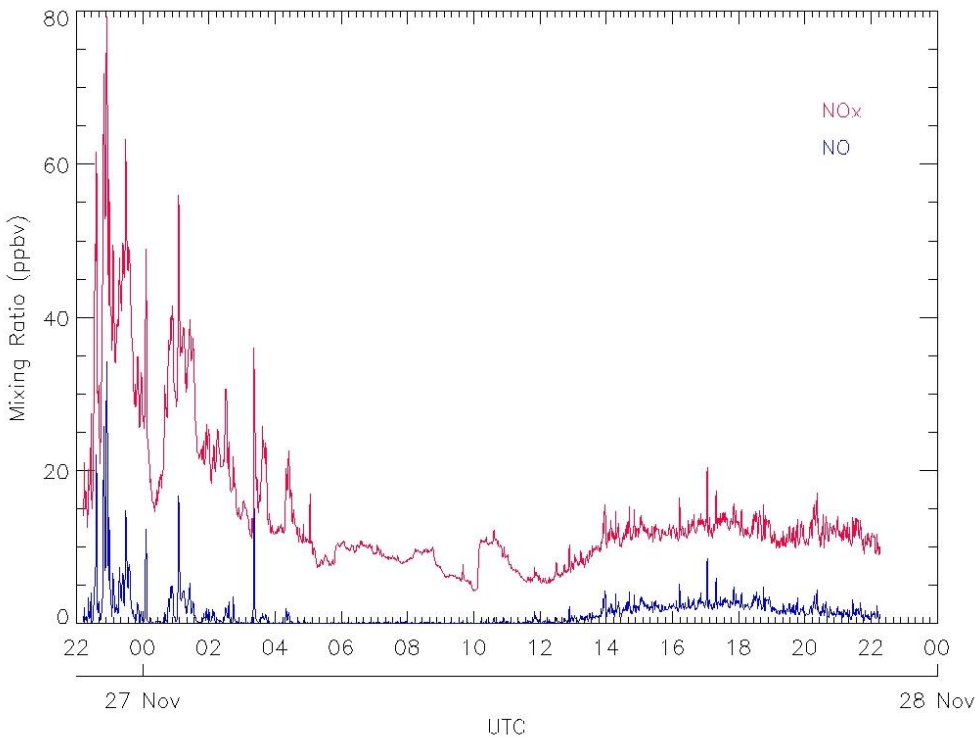
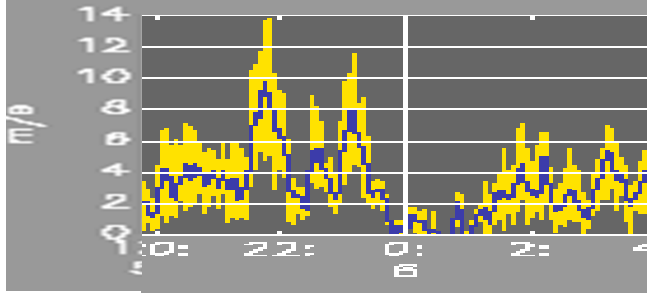
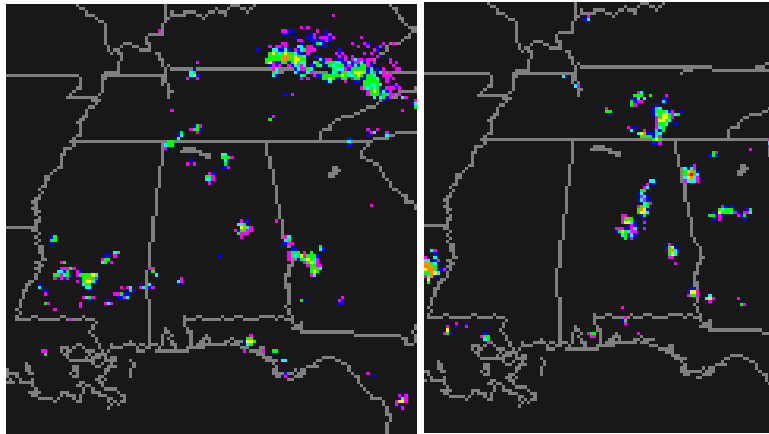
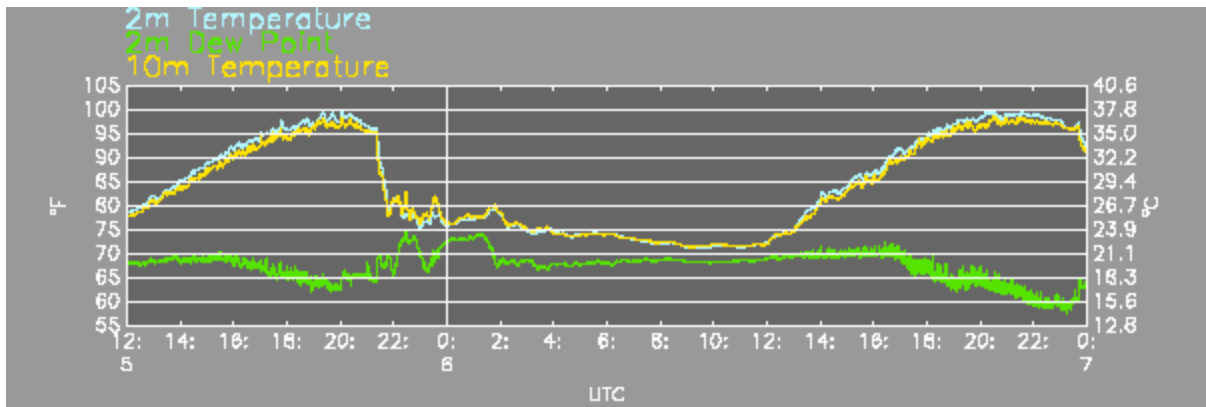


Figure 2: Winter background, with labeling the same as Figure 1.

The afternoon of July 5 (Figure 3) had numerous thunderstorms in north-central Alabama. Initial increases in NO<sub>x</sub> around 4:15 p.m. corresponded to the passage of an outflow boundary from a storm to the south. Wind speeds died down later in the afternoon/early evening and shifted to generally northerly, allowing additional advection of NO<sub>x</sub> from storms to the north of Huntsville. While not visible in the wind speed/direction data, a careful examination of temperature and

dewpoint data show the sudden dropoff in NO<sub>x</sub> after 8 p.m. corresponded to advection of non-thunderstorm air into the area, characterized by higher temperatures and lower dewpoints. Therefore the difference between the maximum and minimum NO<sub>x</sub> concentrations represents the total lightning NO<sub>x</sub> measured at the site, about 27 ppb. It should also be noted that air masses containing lightning NO<sub>x</sub> are primarily NO<sub>2</sub>, while NO<sub>x</sub> generated by traffic is a mix of NO and NO<sub>2</sub>.



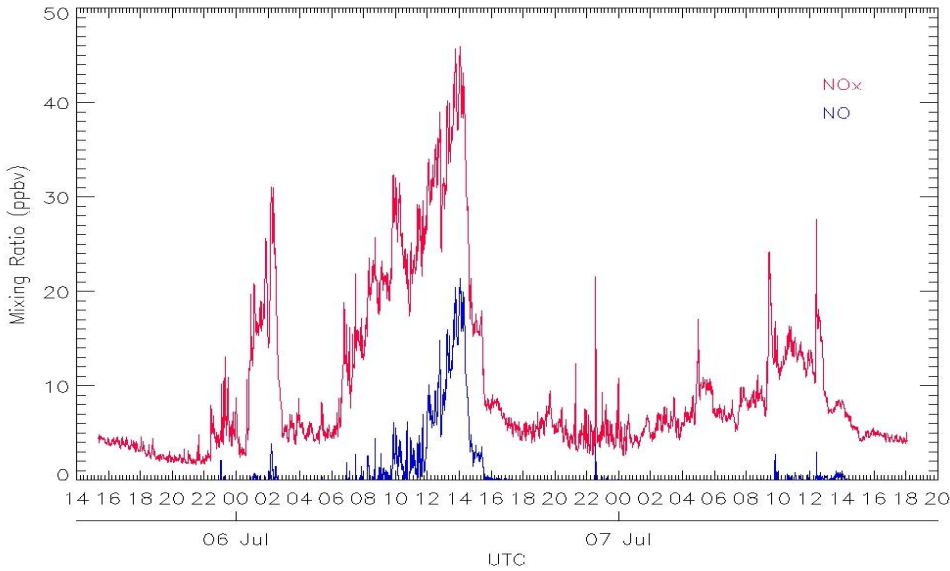
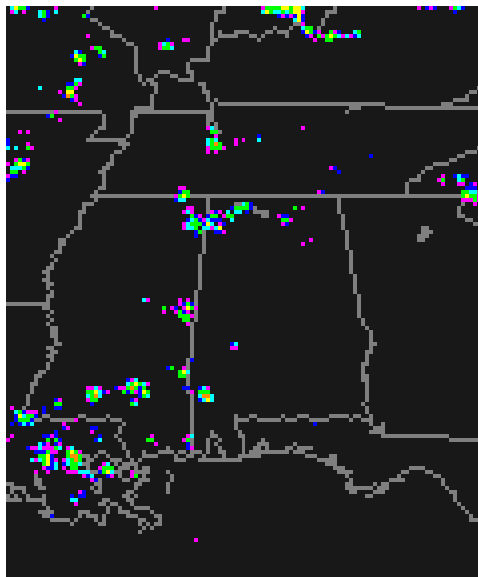
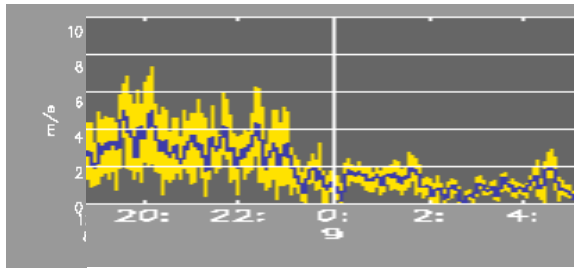
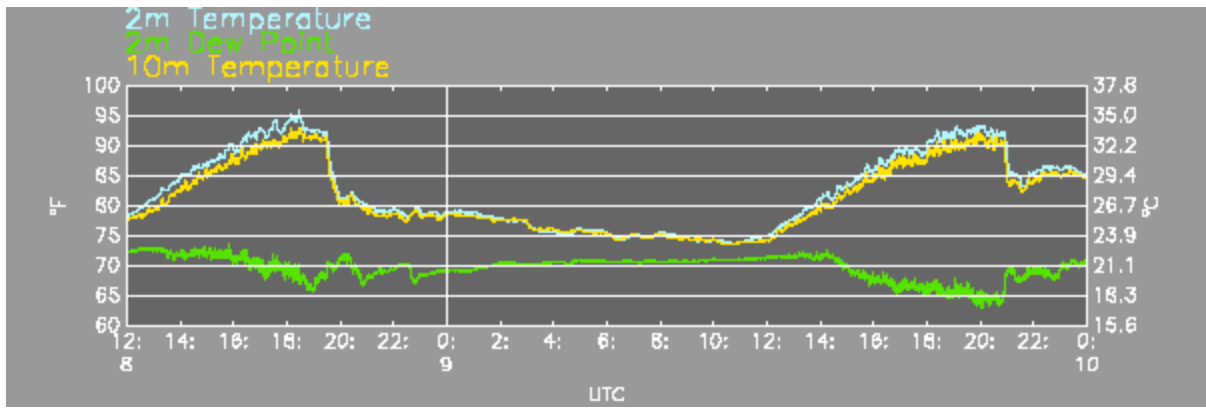


Figure 3: July 5-6. The top figure shows temperature and dewpoint, the second row shows NLDN lightning strikes from 4:15-4:30 p.m. (left) and 7:00-7:15 p.m. (right) local time, the next row shows wind speed for the evening of the 5<sup>th</sup>, with times in UTC, and the bottom row shows NO (blue) and NOx (red) concentration in UTC time.

For July 8 storms, temperature and dewpoint data are needed instead of wind data in order to determine thunderstorm vs. non-thunderstorm air masses. In this case (Figure 4) weak thunderstorms were present just south of Huntsville, which generated a drop in temperature of about five degrees Celsius but not an increase in wind speed. The result was a 5 ppb increase in NO<sub>2</sub>. Note that July 5 was a Sunday, so commuter NO<sub>x</sub> was much less of a factor than during the week.



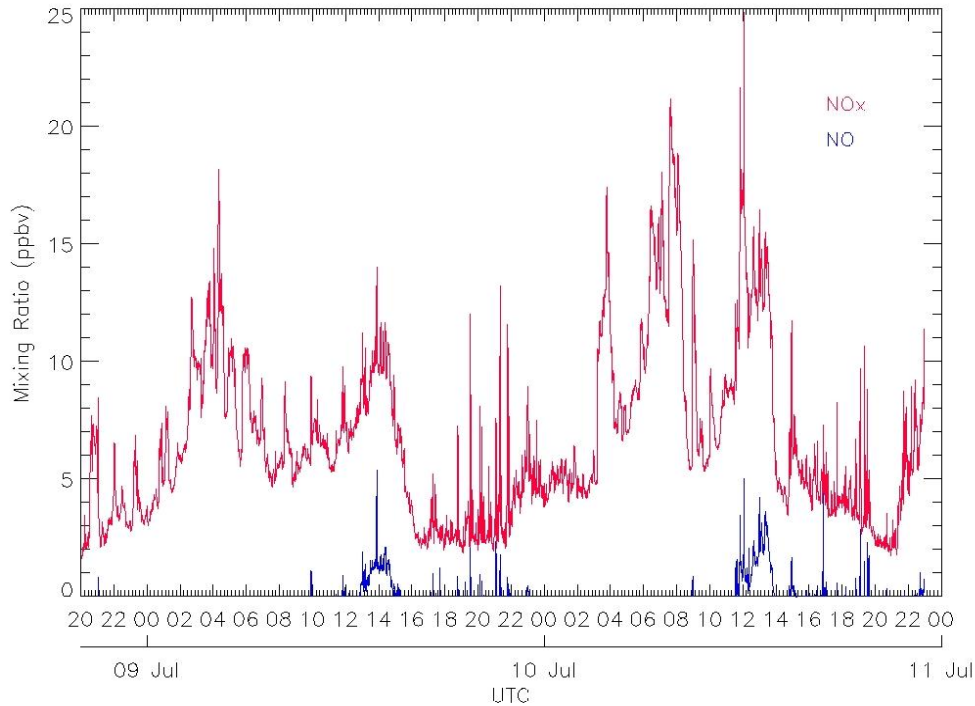
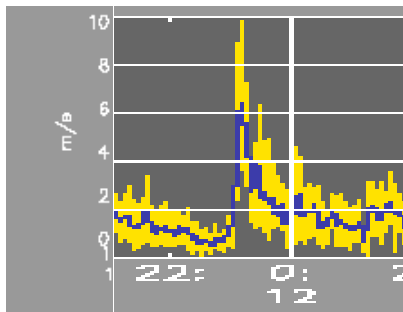
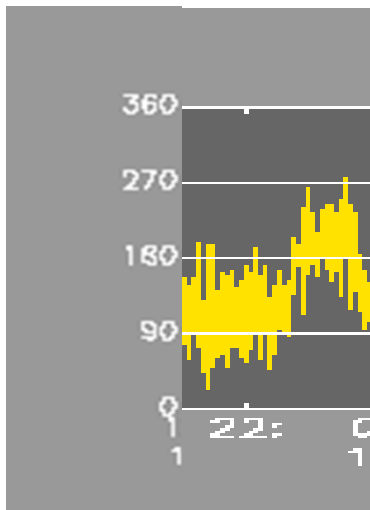
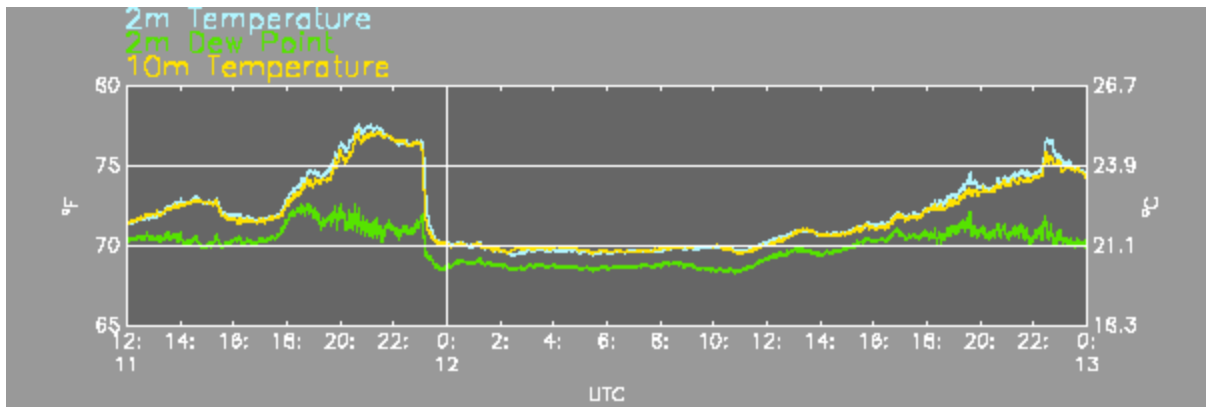


Figure 4: July 8-9. The top row gives temperature and dewpoint, the second row gives wind speed, the third row shows lightning from 2:30 to 2:45 p.m. CDT, and the bottom row shows NO<sub>x</sub> concentration.

The most noteworthy feature of July 11 (Figure 5) is the drop in NO<sub>x</sub> concentration at 6 p.m. While wind, temperature and dewpoint data all point to an outflow boundary passing at this time, NLDN data show the origin of the outflow is a region not producing lightning. This decrease may be considered to be the effect of replacing boundary layer air with free tropospheric air on NO<sub>x</sub> concentration, leading to a 15 ppb (about 75%) drop in total concentration. Since precipitation is shown to cause a large decrease in NO<sub>x</sub>, the level of increase in NO<sub>x</sub> shown in the previous cases is above and beyond the amount needed to offset this decrease.





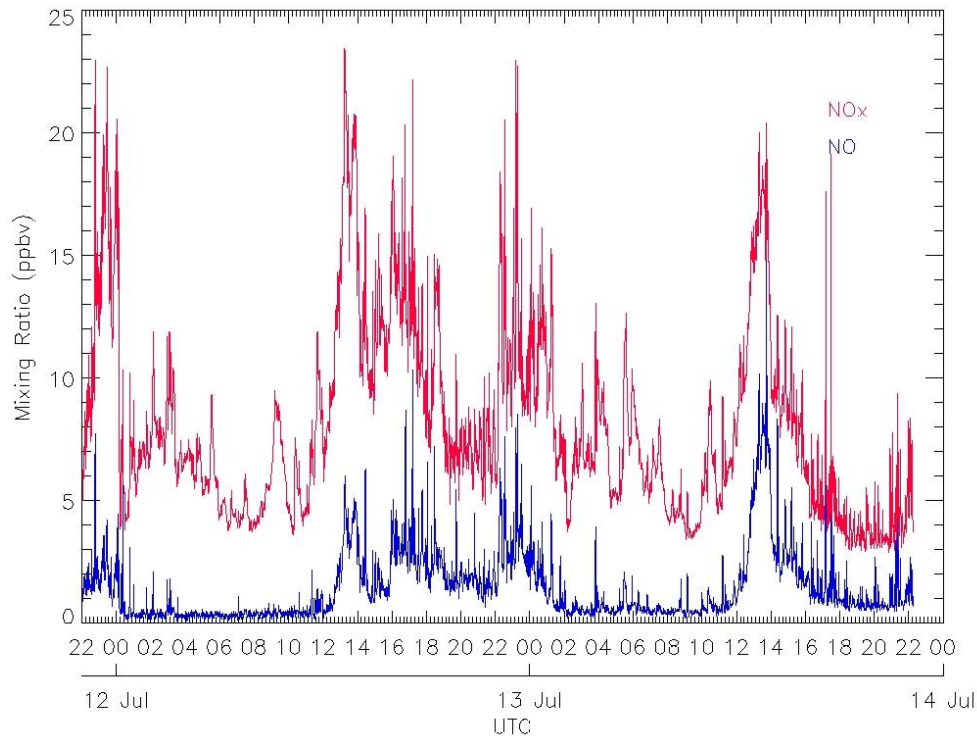
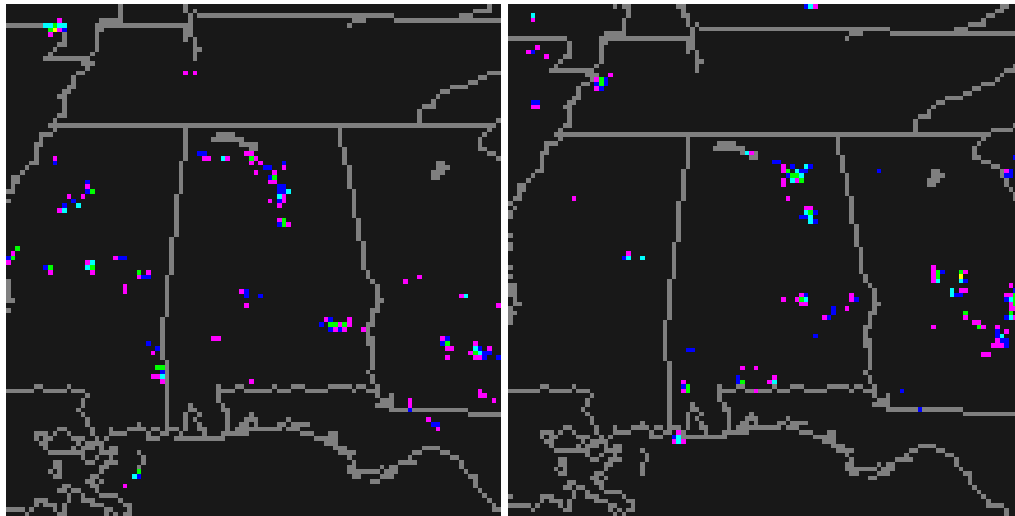


Figure 5: July 11-12. The top row shows temperature and dewpoint, the second row shows wind direction over UTC time, the third row shows wind speed over UTC time, the fourth row shows NLDN lightning at 5:00-5:15 (left) and 6:00-6:15 (right) p.m. CDT, and the bottom row shows NO (blue) and NO<sub>x</sub> (concentration over local time).

#### 4. DISCUSSION AND CONCLUSIONS

These three cases show that lightning generates a measurable amount of surface NO<sub>x</sub>. The chemical composition of the air masses measured containing lightning-generated NO<sub>x</sub> were clearly distinguishable from anthropogenic NO<sub>x</sub> due to a much higher ratio of NO<sub>2</sub> to NO; the cause of this higher ratio has not yet been determined. Winterrath et al. (1999) suggest ozone is introduced into

thunderstorms by stratospheric intrusions and non-lightning electrical discharges; this ozone could titrate NO into NO<sub>2</sub>. Surface NO<sub>2</sub> values compare favorably with Drapcho et al. (1983), who found a peak of 22 ppb following lightning discharges within 1.5 km; their NO<sub>2</sub> to NO ratio was similarly high. The meteorological signature of air masses containing lightning NO<sub>x</sub> is a faster than normal decrease in temperature accompanied by surface winds tracing back to a lightning active region. Conversely, advection of warmer, less humid air following the advection of thunderstorm outflow leads to a decrease in NO<sub>x</sub>. Similarly, advection of non-thunderstorm outflow also leads to a decrease in NO<sub>x</sub> as free tropospheric air will remove much of the preexisting NO<sub>x</sub> from the air in the absence of lightning generated NO<sub>x</sub>.

It is noted that the Ott et al. (2010) model only produces about 1 ppb NO<sub>x</sub> in simulated midlatitude storms, nearly an order of magnitude less than the amount shown in the above figures as produced above and beyond urban NO<sub>x</sub>. They state that vertical transport is too strong in their model, indicating the difference comes from the amount of NO<sub>x</sub> injection near the surface.

It is further noted that our results indicate no correction factor should have been applied to the Franzblau et al. (1989) measurements. The new value would be 1660 moles/flash, much closer to the value found in other studies.

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