

A Dissection of the Surface Temperature Biases in the Community Earth System Model

Tae-Won Park¹, Yi Deng¹, Ming Cai², and Jee-Hoon Jeong³

¹*School of Earth and Atmospheric Science, Georgia Institute of Technology, Atlanta, Georgia, USA*

²*Department of Meteorology, Florida State University, Tallahassee, Florida, USA*

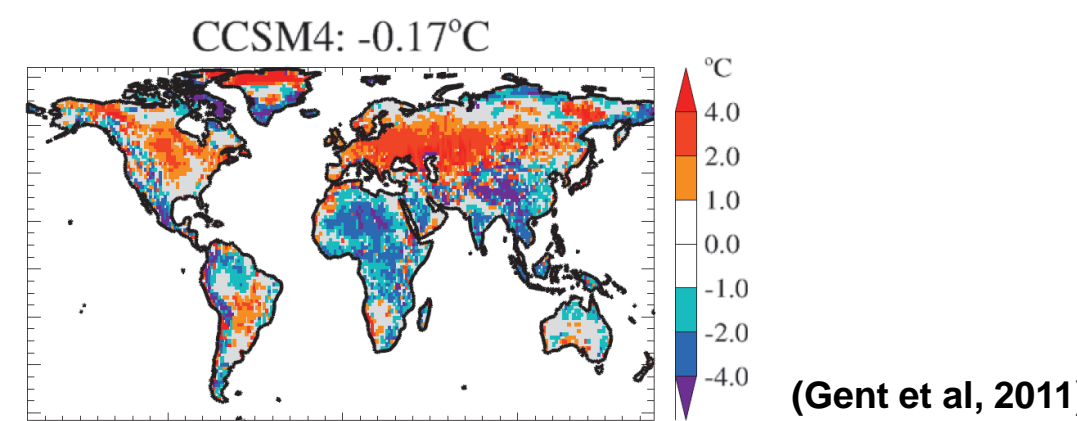
³*Faculty of Earth Systems and Environmental Sciences, Chonnam National University, Gwangju, South Korea*



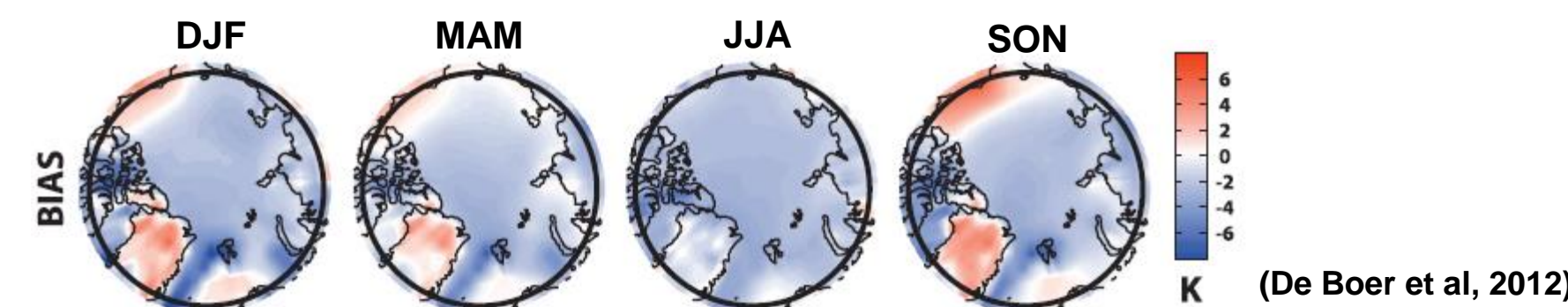
Introduction

Surface Temperature Biases in Climate Models

- Land temperature bias in CCSM4



- Arctic temperature bias in CCSM4



The present study aims to ...

Quantify the relative contributions of seven radiative (physical: albedo, water vapor, and cloud) and non-radiative (dynamical: sensible/latent heat flux, surface dynamics, and atmospheric dynamics).

Data

- Observation** : ERA-Interim (1979-present)
- Model** : CESM1 historical run (1850-2005)
- Used Variables**
 - Solar insolation, surface pressure, surface temperature, surface latent/sensible heat flux, surface downward/upward SW, air temperature, specific humidity, cloud amount, and cloud liquid/ice water
- Analysis period**: Annual mean of 1979-2005

Methods

CFRAM Formulation

- The **total energy balance** at M atmospheric layers and one surface (M+1)th layer

$$\vec{R} = \vec{S} + \vec{Q} \leftarrow \text{Energy due to non-radiative dynamical processes}$$

↑
SW radiation flux
LW radiation flux

- The difference between two climate states

$$\Delta \vec{E} = \frac{\partial \vec{E}}{\partial t} = \Delta \vec{S} - \Delta \vec{R} + \Delta \vec{Q}$$

Change in energy storage
→ negligible

$$\Delta \vec{S} \approx \Delta \vec{S}^{(a)} + \Delta \vec{S}^{(w)} + \Delta \vec{S}^{(c)}$$

$$\Delta \vec{R} \approx \Delta \vec{R}^{(w)} + \Delta \vec{R}^{(c)} + \frac{\partial \vec{R}}{\partial T} \Delta \vec{T}$$

$$\Delta \vec{Q} \approx \Delta \vec{Q}^{(SH)} + \Delta \vec{Q}^{(LH)} + \Delta \vec{Q}^{(sfc_dyn)} + \Delta \vec{Q}^{(atmos_dyn)}$$

Planck feedback matrix $\left(\frac{\partial \vec{R}}{\partial T} \right) = \begin{pmatrix} \frac{\partial R_1}{\partial T_1} & \dots & \frac{\partial R_1}{\partial T_{M+1}} \\ \vdots & \ddots & \vdots \\ \frac{\partial R_{M+1}}{\partial T_1} & \dots & \frac{\partial R_{M+1}}{\partial T_{M+1}} \end{pmatrix}$

- Rearranging the terms ...

$$\Delta \vec{T} = \left(\frac{\partial \vec{R}}{\partial T} \right)^{-1} \left\{ \begin{array}{l} \Delta \vec{S}^{(a)} \text{ Albedo} \\ \Delta \vec{S}^{(w)} \text{ Water vapor} \\ \Delta \vec{S}^{(c)} \text{ Cloud} \\ \Delta \vec{R}^{(w)} \text{ Sensible heat flux} \\ \Delta \vec{R}^{(c)} \text{ Latent heat flux} \\ \Delta \vec{Q}^{(SH)} \text{ Surface dynamics} \\ \Delta \vec{Q}^{(LH)} \text{ Atmospheric dynamics} \\ \Delta \vec{Q}^{(sfc_dyn)} \\ \Delta \vec{Q}^{(atmos_dyn)} \end{array} \right\}$$

Decomposition Procedure

Define Surface Temperature Bias

Model: CESM1 Surface Temperature
Observation: ERA-Interim Surface Temperature
→ Bias in CESM1 = Model - Observation

Input for Radiative transfer model

Surface	Multi-layer
Solar insolation	Air temperature
surface pressure	Specific humidity
surface temperature	Cloud amount
surface latent/sensible heat flux	Cloud liquid water
surface downward/upward SW	Cloud ice water

Energy perturbation terms

$$\Delta \vec{S}^{(a)} \quad \Delta \vec{S}^{(w)} \quad \Delta \vec{S}^{(c)} \\ \Delta \vec{R}^{(w)} \quad \Delta \vec{R}^{(c)} \quad \Delta \vec{Q}^{(sfc_dyn)} \quad \Delta \vec{Q}^{(atmos_dyn)}$$

Partial temperature changes

$$\Delta T^{albedo} \quad \Delta T^{water\ vapor} \quad \Delta T^{cloud} \\ \Delta T^{SH} \quad \Delta T^{LH} \quad \Delta T^{sfc_dyn} \quad \Delta T^{atmos_dyn}$$

Annual-mean

Fu-Liou radiative transfer model
(Fu and Liou, 1992; 1993)

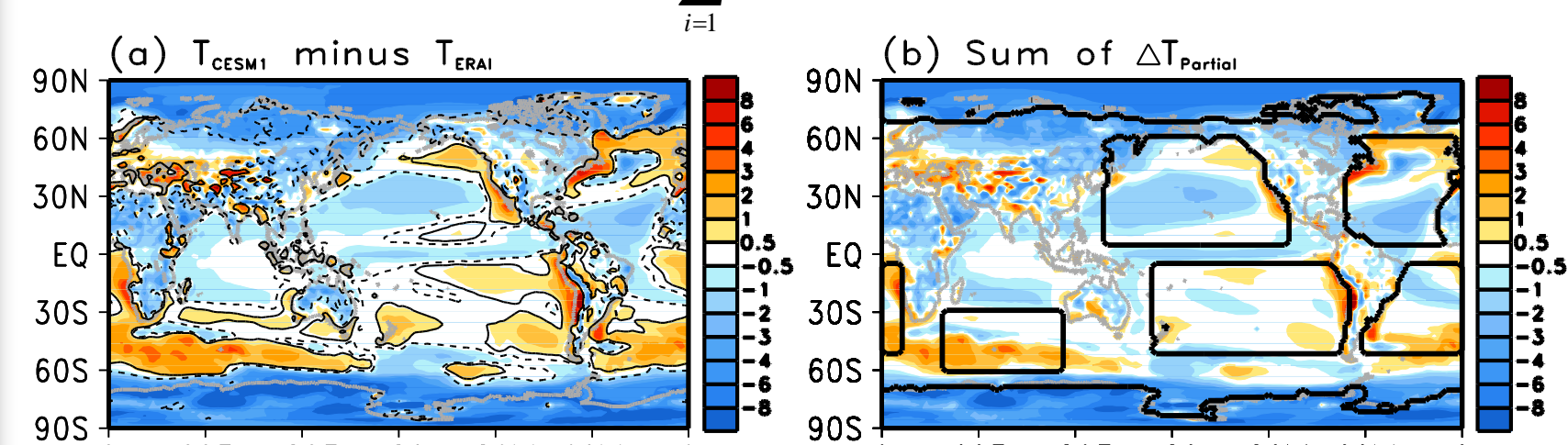
CFRAM
(Lu and Cai, 2009)

Partial Temperature Biases

Difference of TS between CESM1 and ERAI vs. Total TS changes

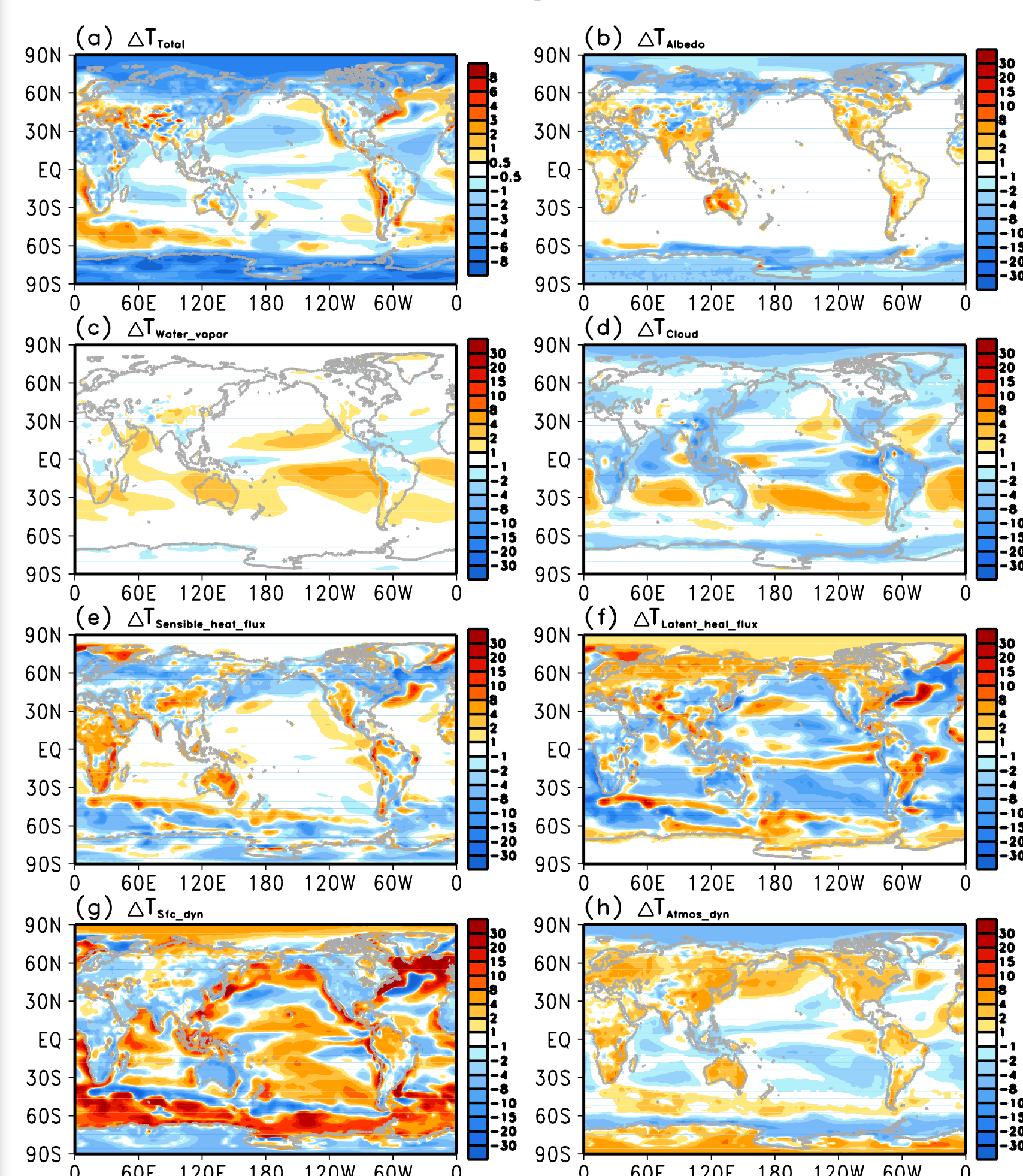
Difference of TS in original data: $\Delta \vec{T} = \vec{T}(\text{CESM1}) - \vec{T}(\text{ERAI})$

CFRAM-calculated $\Delta \vec{T}$: $\Delta \vec{T} = \sum_{i=1}^7 \Delta \vec{T}_i$



$$\vec{T}(\text{CESM1}) - \vec{T}(\text{ERAI}) \approx \sum_{i=1}^7 \Delta \vec{T}_i$$

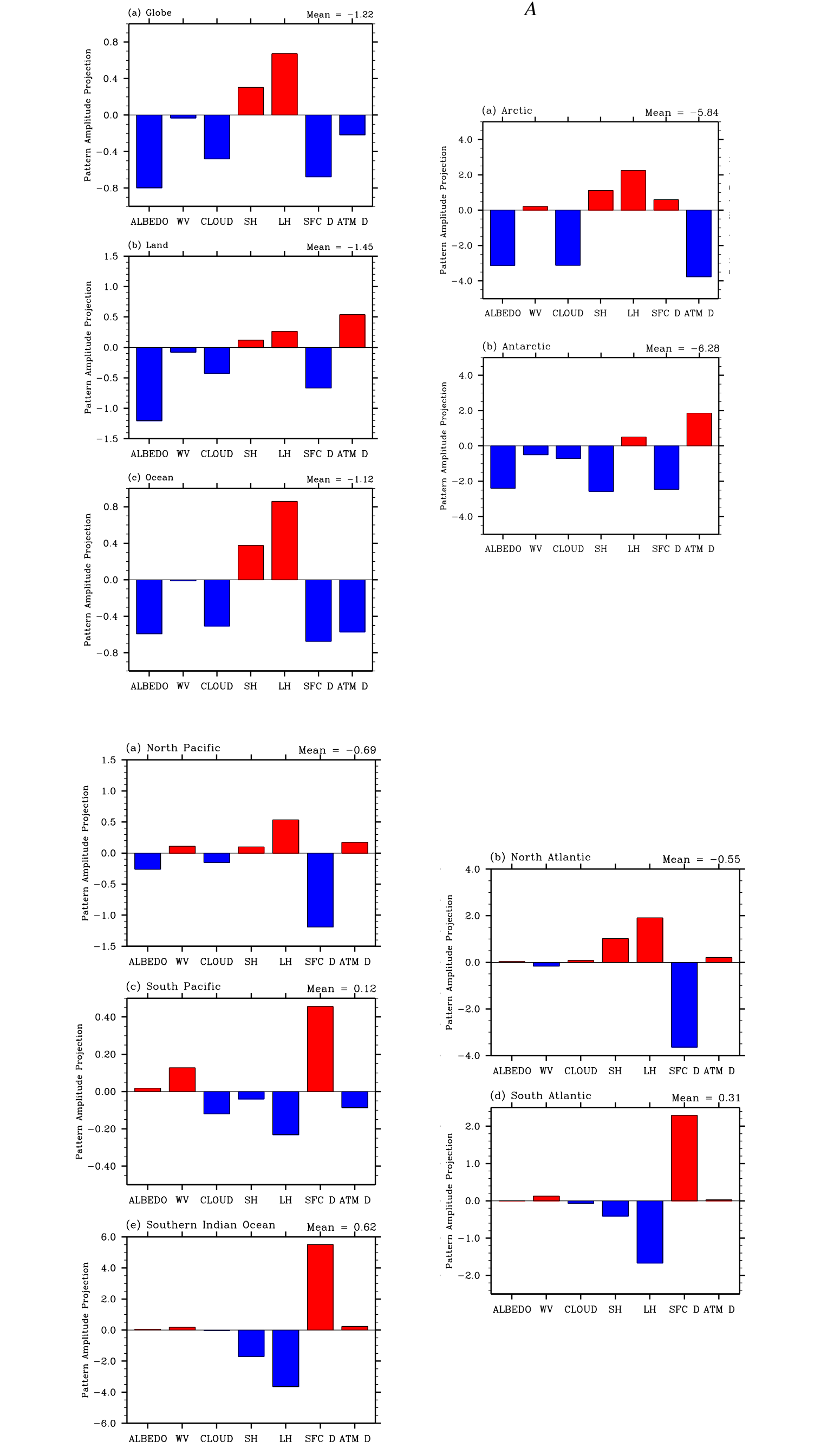
Partial Surface Temperature Biases



Quantification of the relative contributions

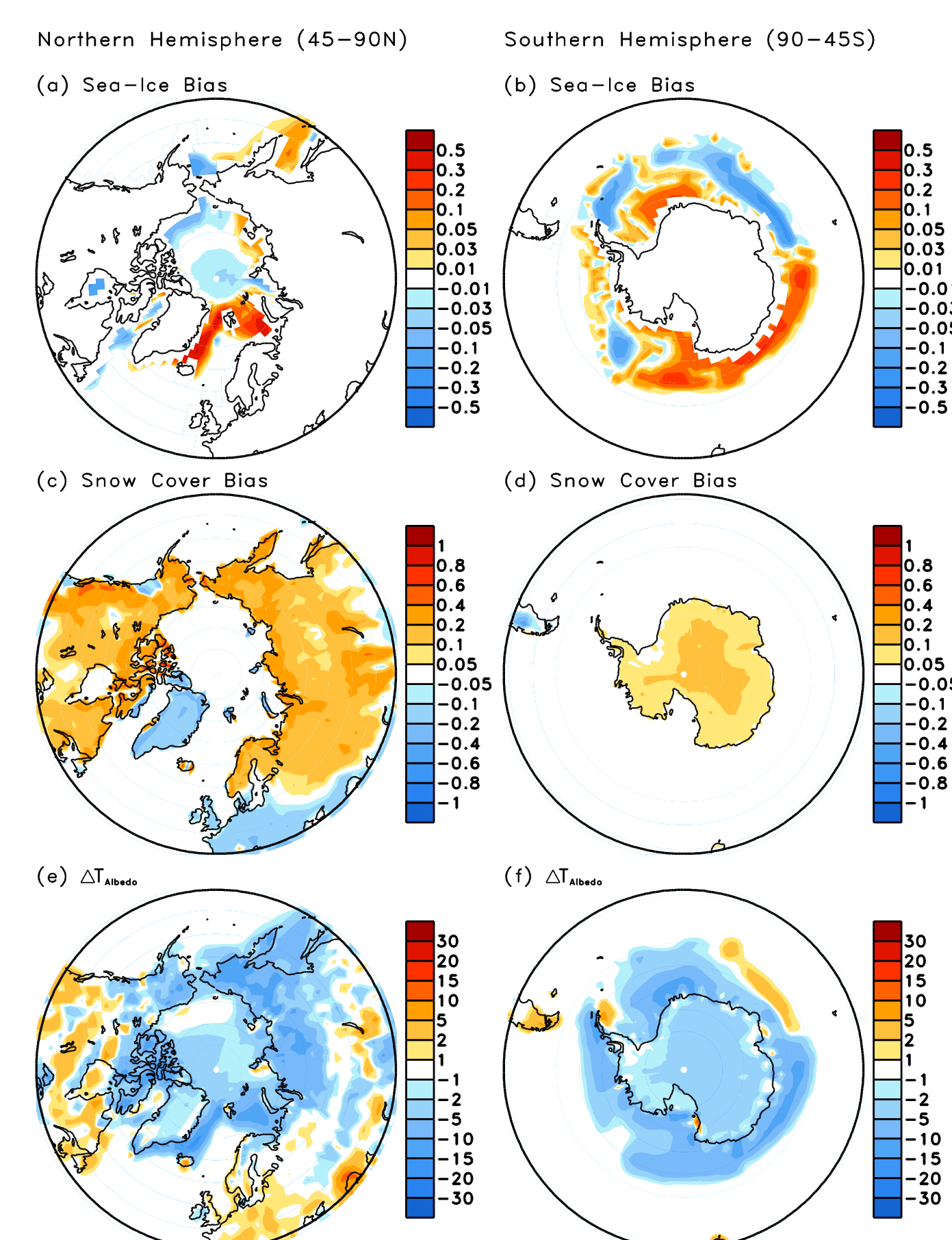
- Pattern Amplitude Projection (PAP) coefficients

$$\text{PAP}_i = A^{-1} \int_A \Delta \vec{T} \cos \phi d\lambda d\phi \cdot \frac{A^{-1} \int_A \Delta \vec{T}_i \cos \phi d\lambda d\phi}{A^{-1} \int_A (\Delta \vec{T})^2 \cos \phi d\lambda d\phi}$$

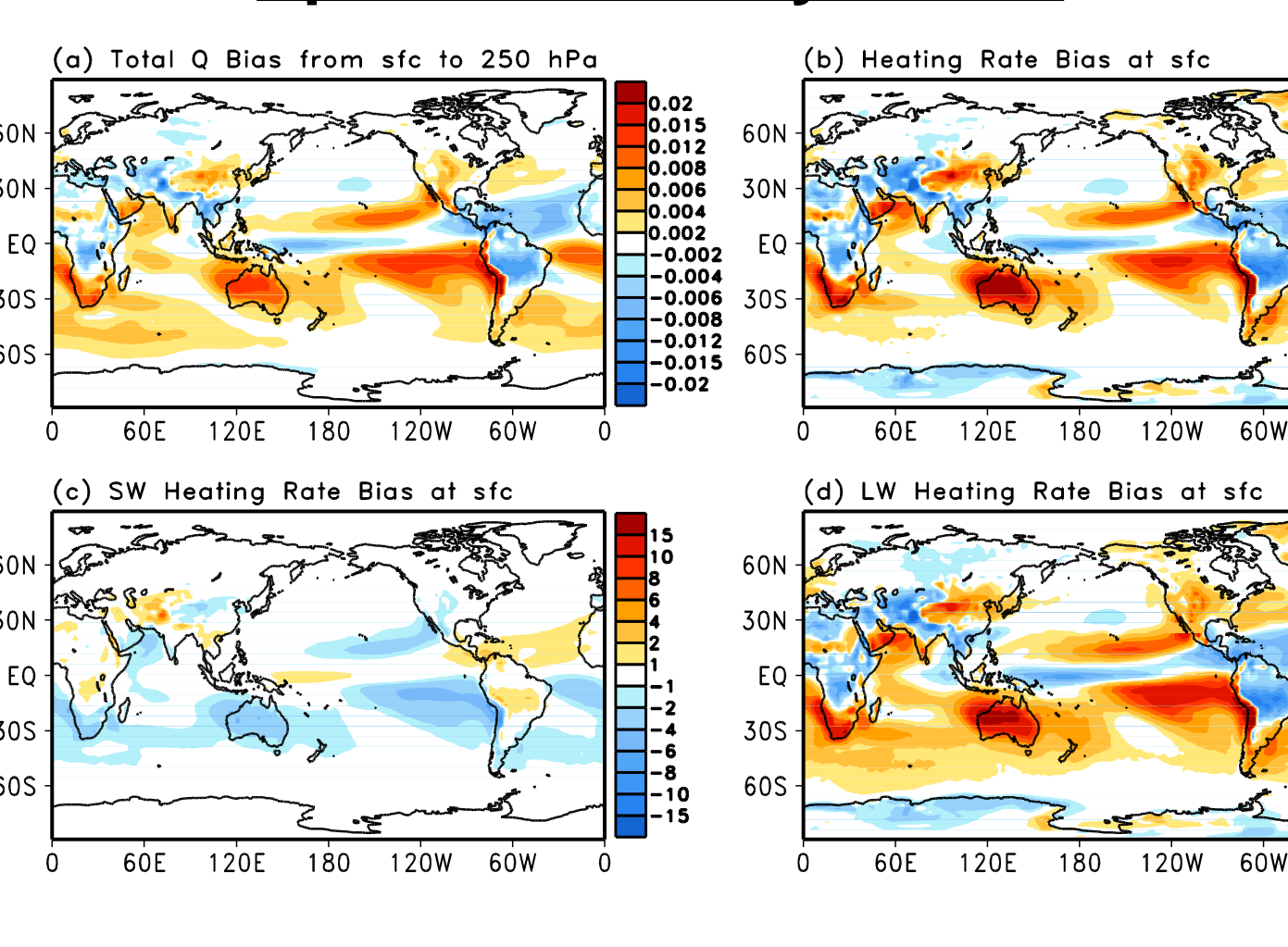


Possible Causes of Partial Temperature Biases

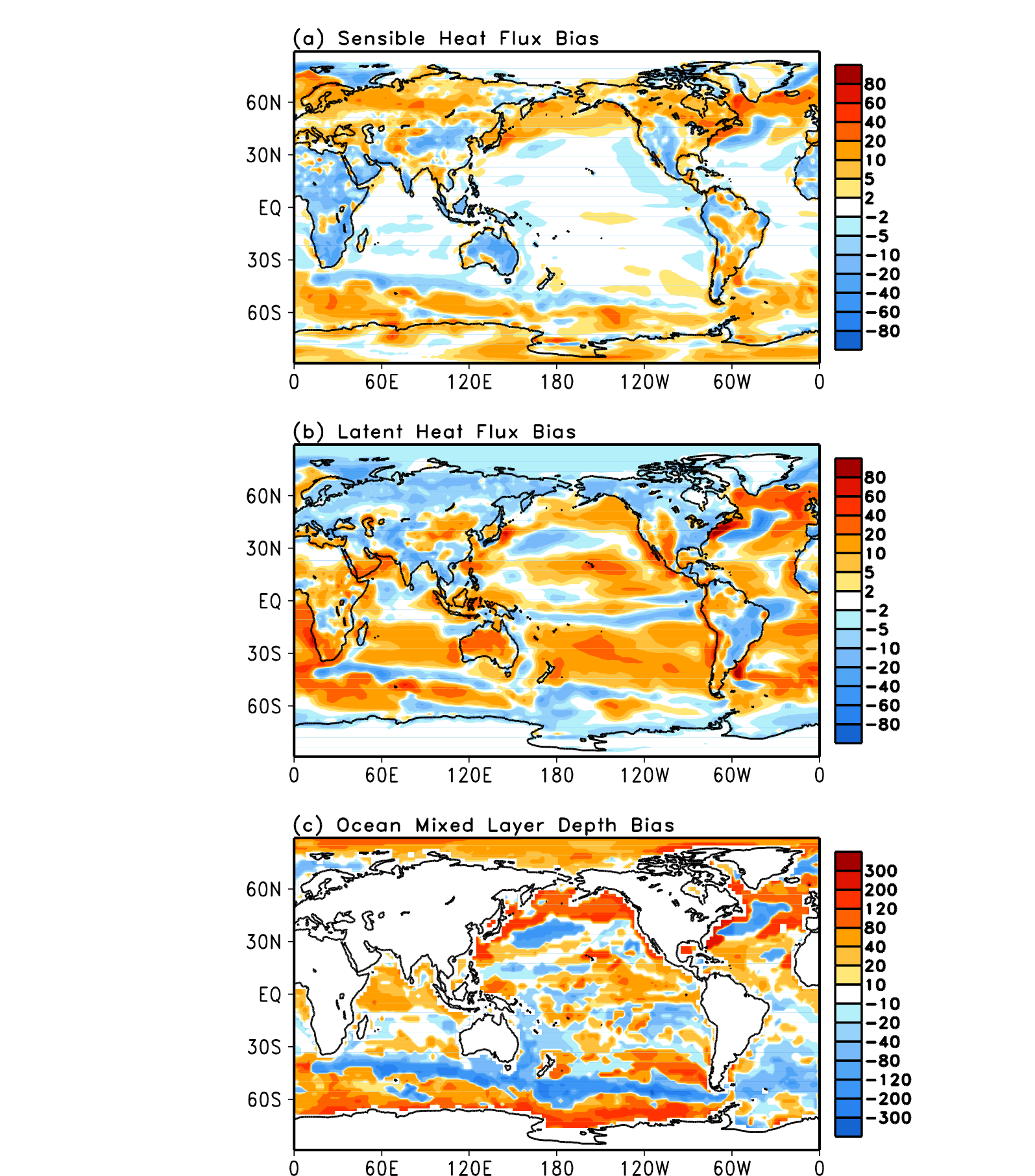
Sea-ice and snow cover biases



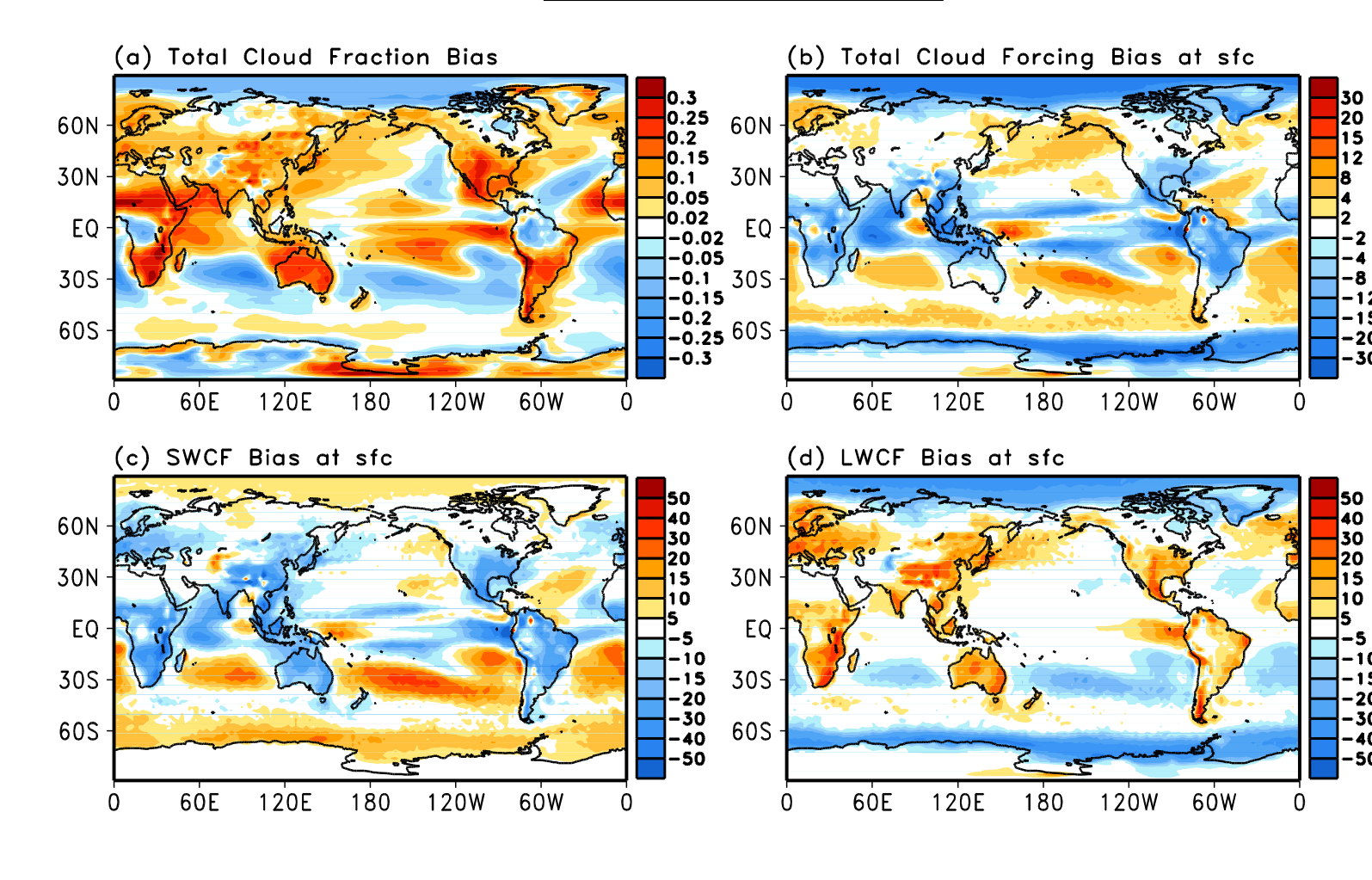
Specific humidity biases



SH/LH and ocean mixed layer depth biases



Cloud biases



Physics vs Dynamics

- Radiative (physically-induced) energy biases**
 - albedo, water vapor, and cloud
- Non-radiative (dynamically-induced) energy biases**
 - sensible/latent heat flux, surface dynamics, and atmospheric dynamics

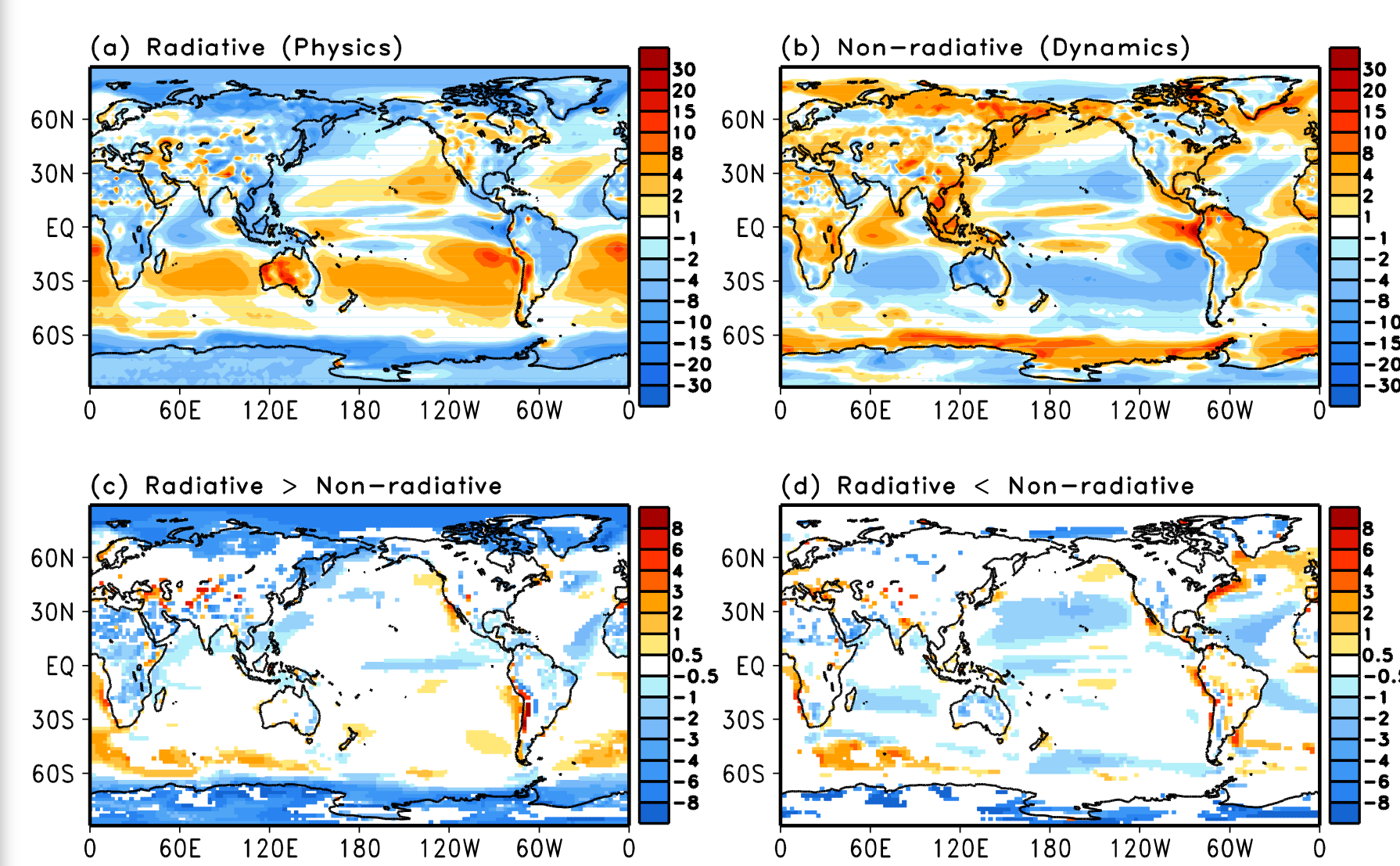


Table 1. The number of total, radiative-origin, and non-radiative-origin grid-points for globe, land, ocean, Arctic, and Antarctica

	Total	Radiative > non-radiative	Radiative < non-radiative
Globe	9,852	6,306 (64.0%)	3,546 (36.0%)
Land	2,969	2,090 (70.4%)	879 (29.6%)
Ocean	6,883	4,216 (61.2%)	2,667 (38.8%)
Arctic	1,194	1,109 (92.9%)	85 (7.1%)
Antarctic	1,203	897 (74.5%)	306 (25.5%)

Acknowledgments: The Georgia Tech authors (Deng and Park) and FSU author (Cai) were supported by DOE Office of Science Regional and Global Climate Modeling (RGCM) program DE-SC0005596 and Grant DE-SC0004974, respectively. Deng is also supported by NSF grant AGS 114760. Jee-Hoon Jeong is supported by the Korea Meteorological Administration Research and Development Program under Grant CATER 2012-3066.