



# A U.S. Climatology of Mesoscale Convective Systems: 1997–2013

Alex Haberlie and Walker S. Ashley

Meteorology Program, Department of Geography, Northern Illinois University, DeKalb, IL

Contact: ahaberlie@niu.edu



## Introduction

- Although previous research has examined the spatiotemporal distribution of MCSs in the U.S., few have examined the long-term climatology of these systems and even fewer have employed automated methods to detect and track MCSs.
- In general, prior investigations have used manual identification of MCSs and subsets of MCSs, limiting their scope and use in climatological investigations.
- We explore the utility of an automated MCS detection and tracking procedure and apply the algorithm using 17 years of national composite reflectivity data.

## Data

- 2-km, 5-minute WSI NOWrad® national radar composites
- Examined warm season (May – September) from 1997-2013
- Over 56,000 hours of national composite radar observations

## Detection and Tracking Method

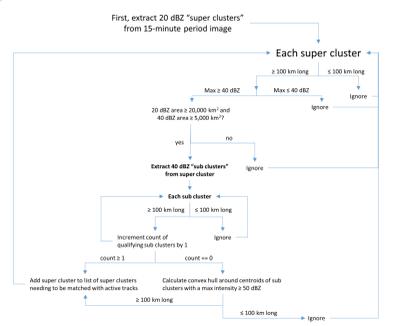


Fig. 1. Flow chart of the classification method employed by this study.

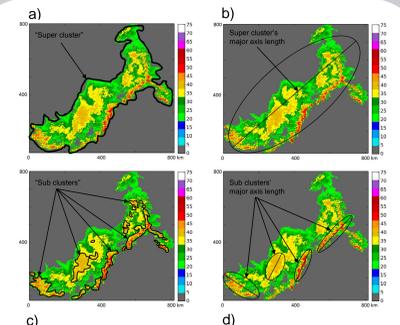


Fig. 2. Examples of cluster types and measurement approaches for super clusters (a,b) and sub clusters (c,d)

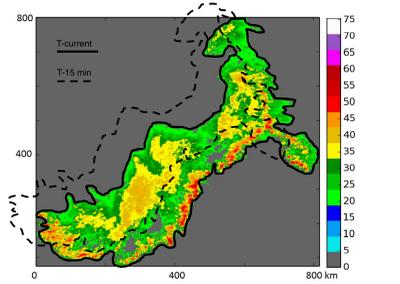


Fig. 3. Example of spatiotemporal overlap where the spatial extent of a super cluster from the previous scan (dotted line) overlaps the spatial extent of a super cluster from the current scan (solid line).

Our method (Figure 1) is based on MCS definition of Parker and Johnson (2000):

- Convective (40 dBZ) line  $\geq 100$  km in one dimension
- Meets this criteria for 3 hours

When a 40 dBZ threshold was applied, too many “convective systems” were found

- Many were related/within the same precipitating cluster

To improve tracking continuity between scans, our method identified all  $\geq 20$  dBZ clusters (i.e., “super clusters”; Figure 2.a) that met the minimum length requirement (Figure 2.b) in a given radar image. Further:

- Each super cluster must have at least 20,000 km<sup>2</sup> of stratiform ( $\geq 20$  dBZ) pixel coverage
- Each super cluster must have at least 5,000 km<sup>2</sup> of convective ( $\geq 40$  dBZ) pixel coverage
- Similar to the spatial requirements employed by Grams et al. 2006

Within each super cluster, find all  $\geq 40$  dBZ clusters (i.e., “sub clusters”; Figure 2.c)

- If at least one has a length  $\geq 100$  km (Figure 2.d)
  - Mark as MCS segment
- If not, draw convex hull around cells with  $\geq 50$  dBZ cores
- If convex hull length  $\geq 100$  km
  - Mark as MCS segment

Once MCS segments—qualifying super clusters—are identified, they are matched with existing, active super cluster tracks by testing for spatiotemporal overlap with the most recent segment (Figure 3).

If a new segment cannot be matched, it is labelled as the start of a new MCS track

After processing was completed, only tracks that spanned at least 3 hours were considered for the climatology

## Warm-season Climatology

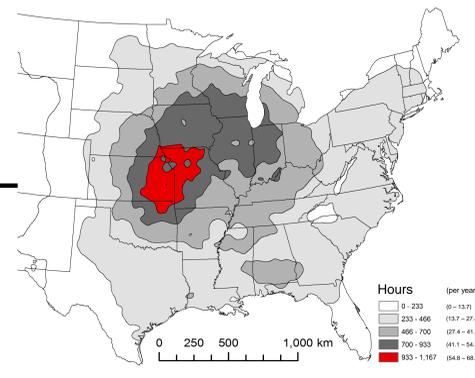


Fig. 4. Warm season MCS climatology for 1997-2013. The mapped values are the total time each 2-km grid is within the 20 dBZ shield of a qualifying MCS. The data were smoothed using a Gaussian filter.

## Monthly Climatology

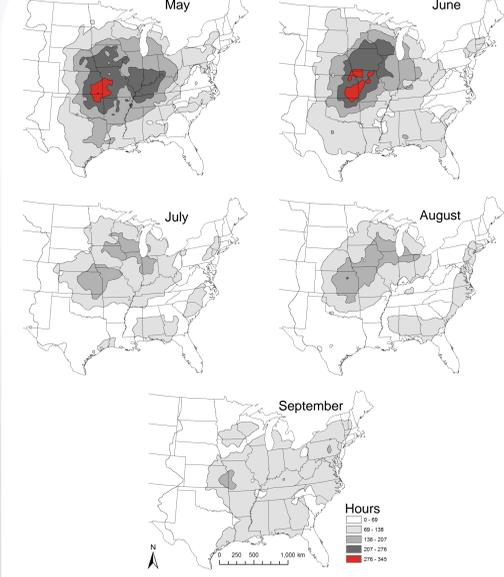


Fig. 6. As in Fig. 4., but for warm-season months. May and June produced the most hours of MCS coverage (i.e., within 20+ dBZ area) in the warm season.

## How are these values calculated?

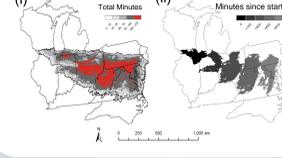


Fig. 5. Example of a “MCS swath” from the 29-30 June 2012 derecho. Panel (i) shows values that are added to the climatology and Panel (ii) shows the evolution of the MCS object through time.

## Yearly Warm-season Climatology

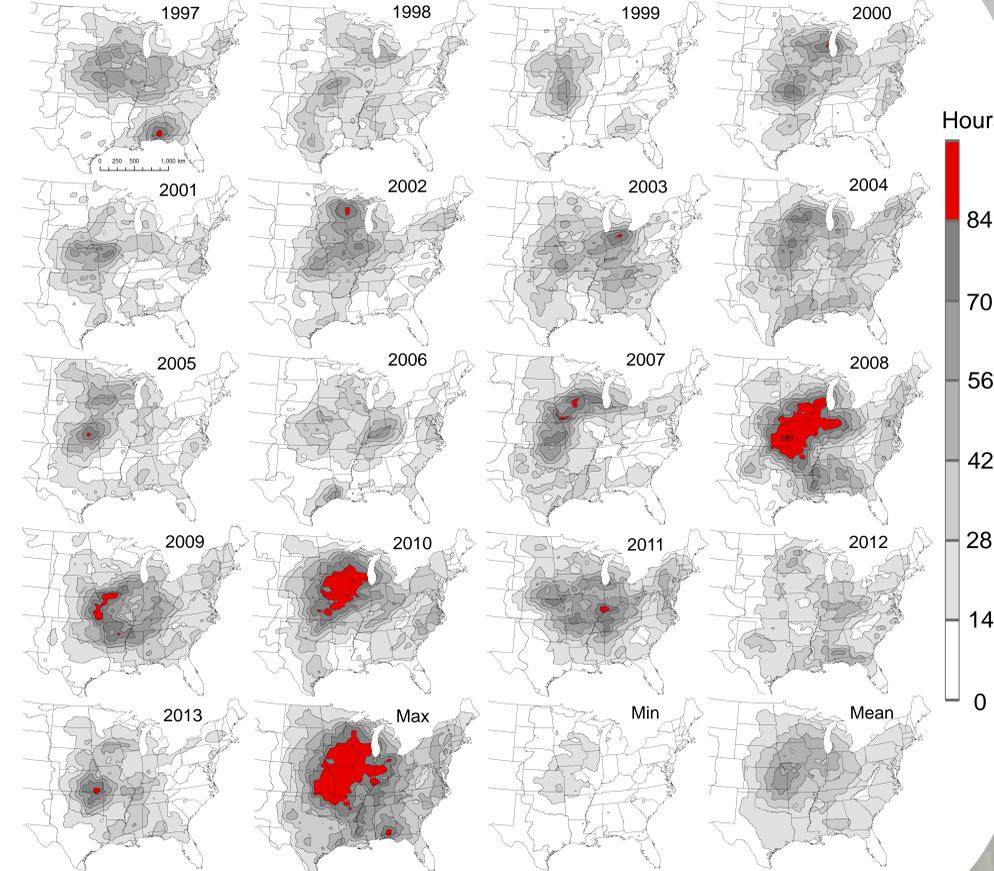


Fig. 7. As in Fig. 4., but by year, with max, min, and mean for each pixel included.

## Yearly MCS Areal Coverage

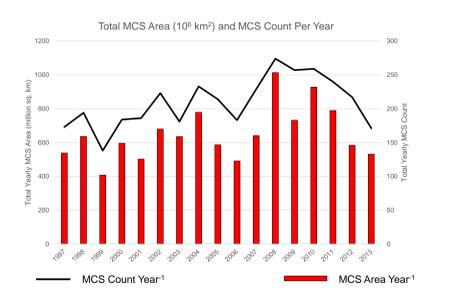


Fig. 8. Yearly total MCS area (left y-axis, red bars) and total MCS count (right y-axis, black line). Total MCS area was calculated by adding the total number of pixels associated with each  $\geq 20$  dBZ MCS shield every 15 minutes and multiplying this value by 4 to get the total square kilometers per year.

## Diurnal Climatology

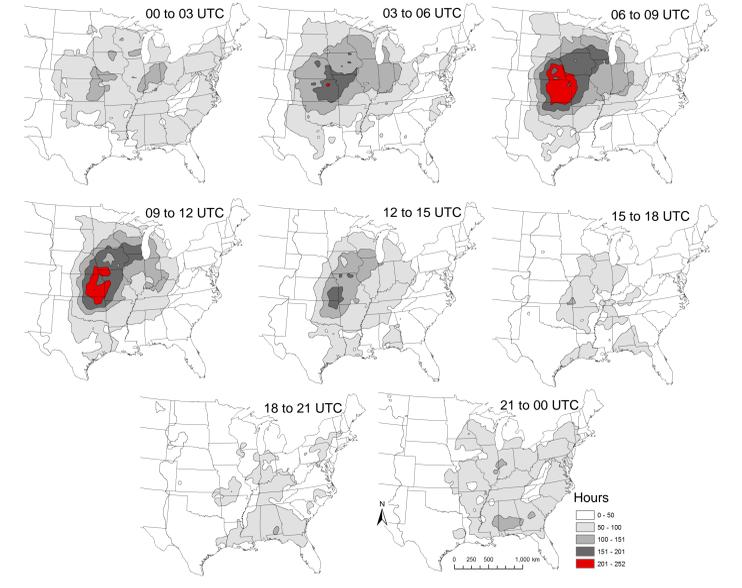


Fig. 9. As in Fig. 4., but for specific three hour periods.

## Conclusions

- The location of maximum MCS activity is consistent with results presented by similar studies
- Ashley et al. (2003) and Fritsch et al. (1986) preferred location of MCC rainfall matches reasonably well
  - Geerts (1998) estimated yearly MCS count for the Southeast U.S. in the warm-season (~220) same order of magnitude (171-274). Our values were lower due to a more strict MCS definition
- The procedure effectively distinguishes between convective system rainfall and isolated convective rainfall
- A convective rainfall climatology by Parker and Knievel (2005) and U.S. precipitation climatology show increases in precipitation nearer to the Gulf of Mexico, which is not evident in our results.
- The diurnal cycle of MCS occurrence and location matches well with previous radar climatologies that have inferred MCS occurrence:
- Carbone et al. (2002), Parker and Ahijevych (2007), Carbone and Tuttle (2008), and others show the west to east movement of convective systems and an overnight maximum in the Plains

Walker S. Ashley, Thomas L. Mote, P. Grady Dixon, Sharon L. Trotter, Emily J. Powell, Joshua D. Durkee, and Andrew J. Grundstein, 2003: Distribution of Mesoscale Convective Complex Rainfall in the United States. *Mon. Wea. Rev.*, 131, 3003–3017.  
 Matthew D. Parker and Jason C. Knievel, 2005: Do Meteorologists Suppress Thunderstorms? Radar-Derived Statistics and the Behavior of Moist Convection. *Bull. Amer. Meteor. Soc.*, 86, 341–359.  
 J. M. Fritsch, R. J. Kane, and C. R. Chelius, 1986: The Contribution of Mesoscale Convective Weather Systems to the Warm-Season Precipitation in the United States. *J. Climate Appl. Meteor.*, 25, 1333–1345.  
 Matthew D. Parker and Richard H. Johnson, 2000: Organizational Modes of Midlatitude Mesoscale Convective Systems. *Mon. Wea. Rev.*, 128, 3413–3436.  
 Matthew D. Parker and David A. Ahijevych, 2007: Convective Episodes in the East-Central United States. *Mon. Wea. Rev.*, 135, 3707–3727.  
 R. E. Carbone, J. D. Tuttle, D. A. Ahijevych, and S. B. Trier, 2002: Inferences of Predictability Associated with Warm Season Precipitation Episodes. *J. Atmos. Sci.*, 59, 2033–2056.  
 R. E. Carbone and J. D. Tuttle, 2008: Rainfall Occurrence in the U.S. Warm Season: The Diurnal Cycle. *J. Climate*, 21, 4132–4146.  
 Bart Geerts, 1998: Mesoscale Convective Systems in the Southeast United States during 1994–95: A Survey. *Wea. Forecasting*, 13, 860–869.  
 Jeremy S. Grams, William A. Gallus Jr., Steven E. Koch, Linda S. Wharton, Andrew Loughe, and Elizabeth E. Ebert, 2006: The Use of a Modified Ebert–McBride Technique to Evaluate Mesoscale Model QPF as a Function of Convective System Morphology during HPC02. *Wea. Forecasting*, 21, 288–306.