



Characteristics Analysis and Influence of Environmental Variability on Idealized Simulations for a Long-lived Bow Echo with High Percentage of Positive Cloud-to-Ground Lightning

Jianhua Dai, Zhaohong Yuan, Lan Tao, and Min Sun CMA/SMS/Shanghai Central Meteorological Observatory, Shanghai, China 2017-01-23

The Eighth Conference on the Meteorological Application of Lightning Data Meteorological Society 97th ANNUAL MEETING | SEATTLE

A Long-lived MCS with a Bow Echo in mid- East China coast on Jun. 5, 2009



Synoptic analysis of 08:00(BJT) 500 hPa (trough with brown, previous trough with orange, and temperature change in last 24 hours with shaded sky-blue) on Jun 5,

2009. (Soundings: XZ-Xuzhou, BS-Baoshan) Behind a aloft trough, jets, lower FL, increasing lapse rate

On June 5, 2009, a long-lived squall line with a bow ech environment, and brought widely distributed large hail Jiangsu, Shanghai, and Zhejiang. (Yin et al., 2010; Dai et



Nantong Radar and ADTD lightning observations Phase I: 04Z-09Z, Jun. 5, 2009



Base reflectivity (in dBZ) from 1.5 degree scan of the Nantong CINRAD and 6-min cloud-to-ground (CG) lightning flashes (blue '+' denotes positive CG, black '-' negative CG) from the national ADTD lightning detection network 04Z-09Z of June 5, 2009

NanTong 0.5 deg Reflectivity and multi-point cross section display with the heights of -20°C and 0°C (1/5)



NanTong 0.5 deg Reflectivity and multi-point cross section display with the heights of -20°C and 0°C (2/5)



NanTong 0.5 deg Reflectivity and multi-point cross section display with the heights of -20°C and 0°C (3/5)



NanTong 0.5 deg Reflectivity and multi-point cross section display with the heights of -20°C and 0°C (4/5)



NanTong 0.5 deg Reflectivity and multi-point cross section display with the heights of -20°C and 0°C (5/5)



Storm structure and CG lightning in Phase I



Time series of 6-min (1) Nantong Doppler radar's VIL maximum (VILmax - kg/m², purple line) and Grids of VIL \ge 45 kg/m² (GridsVIL, light brown line), (2) the ADTD +CG lightning proportion (%PosCG - %, blue line), +CG lightning (CGs, flash, blue bars), -CG lightning (NegCGs, flash, green bars), and total CG lightning (flash, light orange bars), 0400-0900 Z, June 5, 2009

Positive CG in Storms/MCS

---(Rust et al. 1981; MacGroman 1988; Engholm et al. 1990; Saunders et al., 1991)

- *Upper levels in the anvil severe storms* by the advection of positive charge from the major convection regions
- *Adjacent to convective storms* associated with a local tilted electrical dipole
- Stratiform precipitation regions of MCSs
 - by *the advection of positive charge* from the upper levels of the convective portions of the storm
 - by stratiform precipitation processes, related to *ice-ice interactions with low supercooled liquid water contents*
- Inverted-polarity structure of severe storms
 - positive graupel charging occurs at temperatures above a "charge sign reversal temperature"
 - the reversal temperature moves to lower temperatures when the liquid water content is increased

Microphysics and Environmental Factors of Predominated +CG Storms (structure, dynamics)

- ----(Takahashi, 1978; Knapp, 1994; Williams et al., 2005; Carey and Buffalo, 2007)
- Microphysics
 - *larger cloud water concentrations in the mixed phase region* via the *non-inductive charging* (NIC) mechanism, a favorable condition for the *positive charging of large ice particles* that may result in thunderclouds with *a reversed polarity of the main cloud dipole*
- Environment conditions
 - An elevated cloud base height (CBH) /Low sub-cloud moisture / shallow warmcloud depth (WCD)
 - Broad and intense updrafts (CAPE -10°C~ -40 °C) and high liquid water contents (LWC) less entrainment of ambient air, transport a larger fraction of water content to the mixed-phase region
 - a drier low to mid-troposphere, higher cloud-base height, smaller *warm cloud depth**, stronger conditional instability, larger 0–3 km AGL wind shear, stronger 0–2 km AGL storm relative wind speed, and larger buoyancy in the mixed-phase zone. (Carey and Buffalo, 2007)

Sounding comparisons between MCS initiation and bow echo formation



00Z 58027-Xuzhou sounding + 06Z T/Td

Upstream to MCS initiation

06Z 58362-Baoshan Sounding Bow echo formation



A Descending Rear Inflow Jet after 07UTC



- •weak-moderate environmental shear
- •RIJ enhances the surface gust front

buoyancy gradient associated with the warm air in the ascending front to rear flow is less than the buoyancy gradient on the back edge of the cold pool, then the RIJ will descend to the ground

WRF simulation

Idealized simulation setup (Yuan 2015)
Sounding data: 20090605 06Z Baoshan 58362, Shanghai
H-Reso: 1x1km V-Reso: 51 levels, 400m
Microphysics scheme: Morrison double-moment (MP10)



No.	ID	Scheme	Sounding Modification
1	CTL	Control Run	none
2	LLR	Lower Lapse Rate	decreased T850hPa and increased T500hPa and T300hPa
3	MML	Moist Mid-level	increased dew point between 400 and 600 hPa
4	EMBL	Extremely Moist Boundary layer	dramatically increased dew point between surface and 925hPa ($+8^{\circ}C \& +6^{\circ}C$)
5	MBL	Moist Boundary layer	increased dew point between surface and 925hPa ($+5^{\circ}$ C & $+3^{\circ}$ C)
6	LHLW	Lower High-level winds	decreased wind speed between 300 and 500hPa





Reflectivity and wind vectors at 2km height of the WRF idealized model for the Jun 5 squall line

Model reflectivity and water content comparisons



Summary (1/3)

MCS structure

 vertical cross sections along the core of the bow echo showed a strong wind gust front, a mid-altitude radial convergence (MARC), and a downward-sloping rear inflow jet (RIJ).

Lightning

- The bow echo was dominated by *positive* CG lightning flashes (percentage over 66%) at the early stages.
- At its mature stage, a burst of negative CG flashes was found in the major convection areas.
- The predominant positive CG lightning flashes were associated with *rapid updraft intensification* stages of the bow echo, and predominant negative CG flashes with *updraft weakening*.
- Theme of the AMS2017 "Observations Lead the Way"
 - New instruments VHF lightning mapper, phased-array radar
 - High temporal resolution observations

Summary (2/3)

- Environmental factors
 - Vertical wind shear:
 - Along the squall line's motion axis, strong *surface to high-level vertical wind shear* interacted with the cold pool to enhance the lifting at the leading edge of the convective system, which made the squall line live longer.
 - The *medium-to-weak vertical wind shear* at low levels made a downward-sloping rear inflow jet (RIJ) which spread out along the surface and caused widely distributed strong wind gusts.
 - The vertical and horizontal size of the MCS were sensitive to the strong vertical shear created by the *high-level jets* (based on model simulations)

Summary (3/3)

- Environmental factors
 - High lapse rate
 - caused by the strong cold advection aloft over the warm advection at low levels was the crucial factor of a large convective energy or strong updrafts for this convective system.
 - *Rapid increase of boundary humidity* (along boundary layer convergence line)
 - may be the major cause of this deep convection case by providing a lower LFC and a larger CAPE
 - Dry layers at mid-levels and low-levels provided
 - evaporative cooling potential
 - shallow warm cloud depth
 - Warmer boundary ambient air
 - enlarged the temperature difference between downdrafts and environment.

References

- Bluestein H W and Jain M H. 1985. Formation of mesoscale lines of precipitation: Severe squall lines in Oklahoma during the spring. J. Atmos. Sci., 42: 1711–1732
- Carey, L. D., and K. M. Buffalo, 2007: Environmental control of cloud-to-ground lightning polarity in severe storms. *Mon. Wea. Rev.* 135, 1327-1353
- Cohen, A. E., M. C. Coniglio, S. F. Corfidi, and S. J. Corfidi, 2007. Discrimination of mesoscale convective system environments using sounding observations. Wea. Forecasting, 22:1045–1062
- Coniglio M C, Stensrud D J, and Richman M B. 2004. An observational study of derecho-producing convective systems. *Wea. Forecasting*, 19:320–337
- Dai, Jianhua, L. Tao, Y. Ding, et al. 2012. Case analysis of a large hail-producing severe supercell ahead of a squall line. *Acta Meteor. Sinica*, 70(4):609-627
- Engholm, C D, E R Williams, and R M Dole, 1990. Meteorological and electrical conditions associated with positive cloud-to-ground lightning. Mon. Wea. Rev., 118,470-487.
- Evans J S and Doswell C A III. 2001. Examination of derecho environments using proximity soundings. Wea. Forecasting, 16:329–342
- Fovell R G and Ogura Y. 1989. Effects of vertical wind shear on numerically simulated multicell storm structure. J. Atmos. Sci., 46: 3144–3176
- Fuquay, D. M. 1982. Positive cloud-to-ground lightning in summer thunderstorms, J. Geophys. Res., 87, 7131-7140
- James R P, Markowski P M, J. Fritsch M. 2006. Bow Echo Sensitivity to Ambient Moisture and Cold Pool Strength. *Mon. Wea. Rev.*, 134:950–964
- Johns, RH and Hirt WD. 1987. Derechos: Widespread convectively induced windstorms. Wea. Forecasting, 2:32-49
- Knapp, D. I., 1994. Using cloud-to-ground lightning data to identify tornadic thunderstorm signatures and nowcast severe weather. Natl. Wea. Digest, 19, 35–42.
- Makowski J A, D R MacGroman, M I Biggerstaff, and W H Beasly. 2013. Total lightning characteristics relative to radar and satellite observations of Oklahoma mesoscale convcetive systems. *Mon. Wea. Rev.*, 141:1593-1611
- MacGorman, D. R., D. W. Burgess, V. Mazur, et al., 1989. Lightning rates relative to tornadic storm evolution on 22 May 1981. J. Atmos. Sci., 46, 221-250
- MacGorman, D. R., and K. E. Nielsen. 1991. Cloud-to-ground lightning in a tornadic storm on 8 May 1986, Mon. Wea. Rev., 119, 1557-1574
- Morrison H, Thompson G, and Tatarskii V. 2009. Impact of cloud microphysics on the development of trailing stratiform precipitation in a simulated squall line: comparison of one- and two-moment cchemes. *Mon. Wea. Rev.*, 137:991-1007
- NCAR. User's Guide for the Advanced Research WRF (ARW) Modeling System Version 3.0., http://www.mmm.ucar.edu/wrf/users/docs/user_guide_V3.0/contents.html

References

- Parker, M. D., S. A. Rutledge, and R. H. Johnson, 2001. Cloud-to-ground lightning in linear mesoscale convective systems. *Mon. Wea. Rev.*, **129**, 1232–1242.
- Przybylinski R W. 1995. The bow echo: Observations, numerical simulations, and severe weather detection methods. *Wea. Forecasting*, 10:203–218
- Reap, R. M., and D. M. MacGorman, 1989. Cloud-to-ground lightning: Climatological characteristics and relationships to model fields, radar observations and severe local storms, *Mon. Wea. Rev.*, 117, 518-535
- Rotunno R J, Klemp B, and Weisman M L. 1988. A theory for strong, long-lived squall lines. J. Atmos. Sci., 45: 463–485
- Rust, W. D., D. R. MacGorman, and R. T. Arnold, 1981. Positive cloud-to-ground lightning flashes in severe storms, *Geophys. Res. Lett.*, 8, 791-794
- Rutledge, S. A., and D. R. MacGorman, 1988. Cloud-to-ground lightning activity in the 10-11 June 1985 mesoscale convective system observed during the Oklahoma-Kansas PRE- STORM Project, *Mon. Wea. Rev.*, 116, 1393-1408
- Saunders C P R, Keith W D, and Mitzeva R P. 1991. The effect of liquid water on thunderstorm charging. J. Geophys. Res., 96: 11007-11017
- Smull B F, and Houze R A. 1987. Rear inflow in squall lines with trailing stratiform precipitation. Mon. Wea. Rev., 115:2869-2889
- Takahashi T. 1978. Riming electrification as a charge generation mechanism in thunderstorms, J. Atmos. Sci., 35: 1536-1548
- Thorpe A J, M J Miller, and M W Moncrieff. 1982. Two-dimensional convection in non-constant shear: A model of midlatitude squall lines. *Quart. J. Roy. Meteor. Soc.*, 108: 739-762
- Weisman M L, Klemp J B, and Rotunno R. 1988. Structure and evolution of numerically simulated squall lines. J. Atmos. Sci., 45: 1990–2013
- Weisman M L. 1992. The role of convectively generated rear-inflow jets in the evolution of long-lived mesoconvective systems. *J. Atmos. Sci.*, 49(19):1826–1847
- Williams E R, Weber M E, and Orville R E. 1989. The relationship between lightning type and convective state of thunderclouds. *J. Geophys. Res.*, 94:13213-13220
- Williams E R, V Mushtak, D Rosenfeld, S Goodman, and D Boccippio, 2005. Thermodynamics conditions favorable to superlative thunderstorm updraft, mixed phase microphysics and lightning flash rate. *Atmos. Res.*, **76**, 288–306.
- Xue, M., 1990. Towards the environmental condition for long-lived squall lines: vorticity versus momentum. Preprints, *16th Conf. on Severe Local Storms*, Kananaskis Provincial Park, Alberta, Canada, Oct. 22-26, Amer. Meteor. Soc., 24-29
- Ying, D, H. Wu, B. Zhang, et al. 2010. Analysis on a severe convective weather triggered by sea breeze front. *Plateau Meteor*.,29(5):1261-1269.
- Yuan,Z. 2015. Study of the influence of the different horizontal resolutions and microphysical setups on the idealized simulation of a squall line. *Acta Meteor. Sinica*, 73(4):648-666

Questions or Comments?