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1. Introduction

- Mixing refers to the homogenization of a scalar field. Although it is an important subject in fluid dynamics, mixing has not attracted much attention in the urban air quality literature, where processes such as dispersion and ventilation have been highlighted. For certain applications (e.g. the accidental release of toxic chemicals), however, knowledge of the mixing rate is extremely useful.
- This work investigates the mixing of a passive scalar in a unit-aspect-ratio street canyon. Using large-eddy simulation, the mixing rate is characterised using the decay rate of the variance. The mixing is inhomogeneous and strongly influenced by the mean circulation; the evolution is qualitatively distinct within (i) the central vortex; (ii) corners excepting the upper downwind one; (iii) remaining regions. For each region, separate regimes related to advection and mixing can be discerned.
- Following established theoretical predictions, the mixing rates are related to the divergence of Lagrangian trajectories (Lyapunov exponent) and Péclet number. On account of the open boundary and inhomogeneous mixing, sensitivity of the scalar field to the initial conditions persists in the long-time limit.

2. Stirring and Mixing

Pure stirring (advection)

$$\frac{\partial\theta}{\partial t} + \mathbf{u} \cdot \nabla\theta = \kappa \nabla^2 \theta^{\mathbf{v}}^{\mathbf{0}}$$

Stirring and mixing

$$\frac{\partial \theta}{\partial t} + \mathbf{u} \cdot \nabla \theta = \kappa \nabla^2 \theta$$



Fig.1 Creation of small-scale structures and associated homogenisation (left panel [1]; right panel [2]).

3. Theoretical Background: whether theories on mixing carry over to the literature of urban environment?

For a spatially smooth, large-scale velocity field, Large-scale nature of the canyon flow: e.g.

$$x_{n+1} = x_n + a\sin(y_n + \phi_n)$$

$$y_{n+1} = y_n + \cos(x_{n+1} + \phi_n)$$

the variance decay rate is predicted by the Lyapunov exponents in the absence of solid boundaries [3]. With no-slip boundary conditions, the long-time behaviour differs [4].

Theoretical predictions for the variance decay rate depend on the existence of a well-defined large-scale velocity field.



energy spectra inside the canyon at z=0.05H and 0.5H [5].

Mixing of a Passive Scalar in an Urban-Street Canyon

4. Model Configuration



5. Diagnostics

x (streamwise)

• Scalar variance is calculated with each pollutant block R_i . λ_i is the variance decay rate.

x/W=0

$$\langle \sigma^2 \rangle_{\mathcal{R}_i} = \frac{1}{V_i} \int\limits_{\mathcal{R}_i} \frac{\left(C(t, \vec{x}) - \langle C \rangle_{\mathcal{D}}\right)^2}{C_0^2} d\vec{x} \sim e^{-\frac{1}{2}}$$

• Divergence of Lagrangian trajectories. λ_0 is the largest Lyapunov exponent.

$$\|\delta \mathbf{x}(t)\| \approx e^{\lambda_0 t} \|\delta \mathbf{x}(t_0)\|$$

6. Results

On account of the open boundary and inhomogeneous mixing, sensitivity of the scalar field to the initial conditions persists in the long-time limit.

Fig.4 Long-time concentration fields (t=1000s) for different initial conditions.



 $S = C_0 \delta(t - t_0) \delta(\vec{x} - \vec{x}_0)$



Fig.3 (Top) Schematic diagram of the computational domain; (bottom) setup of initial conditions.

Initial conditions:

B: Bottom

C: Centre

T: Top

L: Left

e.g.

CL

R: Right



TR denotes the top-right one; **CC** located at the canyon centre.

(1)



The mean concentration decays exponentially for all initial conditions.





Fig.5 Decay of the mean concentration within the canyon.

7. Key Timescales

Initial variance decay rate for Set1 (central vortex) appears to be controlled by diffusion across streamlines. The mixing rate can be estimated from analytical predictions for cellular flow [6].



Fig.7 Illustration of a nomination vortex of radius R. TI tangential velocity V over R used on calculating the Péclet number: $Pe = (2R) V/\kappa$.

Following established theories, mixing rates of Set2 and Set3 appear to be controlled by the Lyapunov exponents.

Variance decay in the long-time limit is controlled by the escape of scalar across the open boundary at the roof level.

Table 1 Variance-decay timescales after Eq. (1)

										10-1	
Categories	CC	BL	BR	TL	$_{\rm BC}$	CL	\mathbf{CR}	TC	\mathbf{TR}	10^{-2}	
Set3					66	68	93	87	129	10^{-3} 10^{-4}	
Set2		32	51	42					$\mathcal{O}^{-}(\mathcal{O})^{2}$	10 ⁻⁵ - ී 10 ⁻⁶ -	
Set1	147									10 ⁻⁷ - 10 ⁻⁸ -	
	Set1									10^{-9}	

Table 2 Particle divergence timescales after Eq. (2)

								- 1-	(-)
Categories	CC	BL	BR	TL	BC	CL	\mathbf{CR}	TC	TR
Set3					35	27	31	21	12
Set2		26	26	25					
$\mathbf{Set1}$	-								
		1	Set2				Set3		

References

[1] Ngan, K. & Shepherd, T.G., 1997. J. Fluid Mech. [2] Rothstein, D., et al., 1999. Nature. [3] Haynes, P. H. & Vanneste, J., 2005. Phys. Fluids. [4] Salman, H. & Haynes, P.H., 2007. Phys. Fluids. [5] Lo, K.W. & Ngan, K., 2015. Boundary-Layer Meteoro. [6] Shraiman, B.I., 1987. Phys. Rev. A.



Initial variance decay depends on R_i. For large t (\gtrsim 800s), the variance decay rate is

Fig.6 Variance decay of the passive scalar for different initial conditions.

al	$u^* - (u) Po^{1/2}$
ne	$\kappa = \langle \kappa \rangle$ re
is	
et	$\tau^* = R^2 / \kappa^* \approx 177 \mathrm{s}$



8 Variance decay in the long-time limit. (Top) malisation is against the mean <C>; (bottom) malisation is over the initial value C_0 .