

# **Impact of Eurasian spring snow decrement on East Asian summer precipitation**

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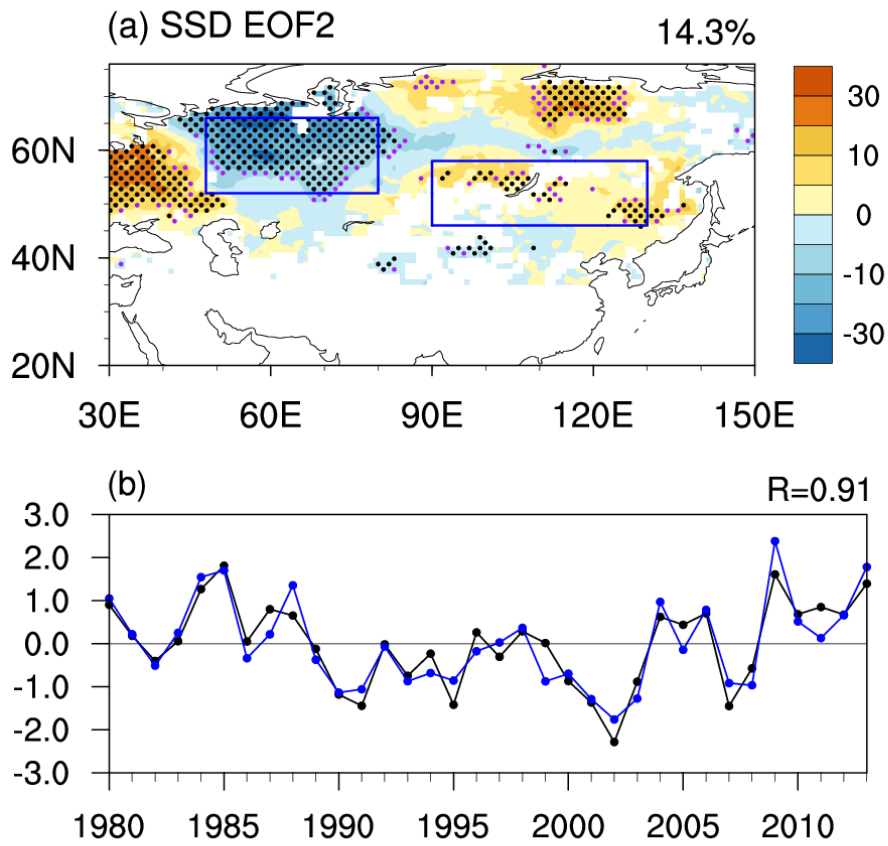
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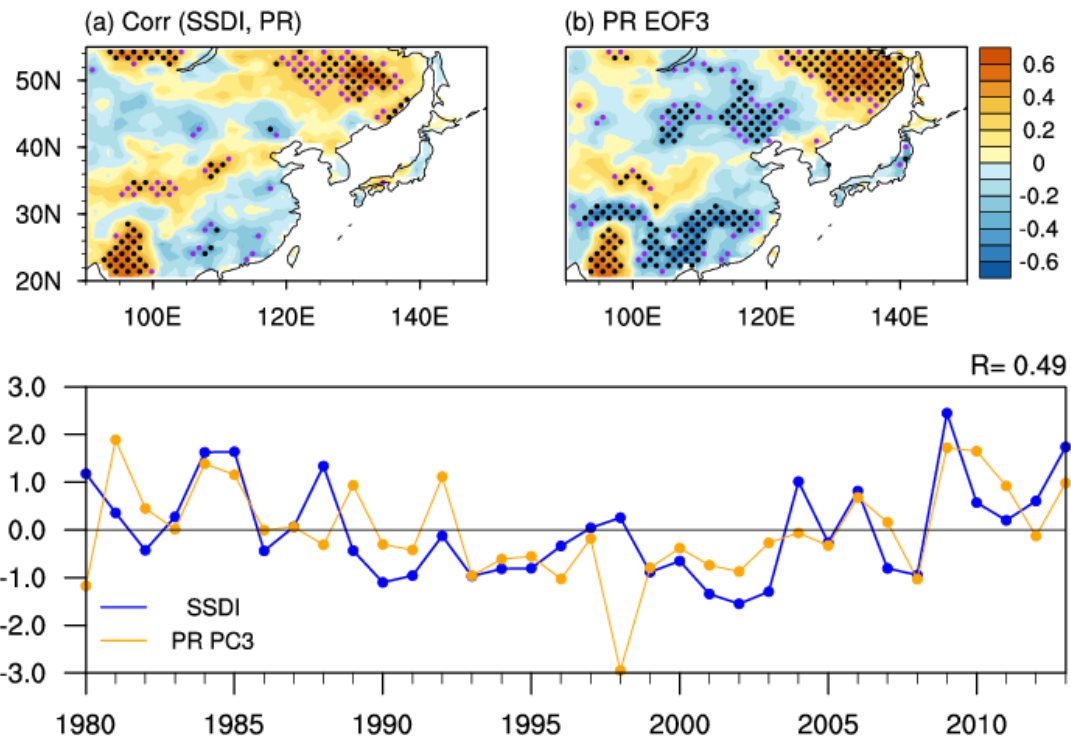
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 2 **Figure 1.** (a) Spatial pattern (shadings, units: mm) of the second EOF (EOF2) and (b) detrended  
 3 time series of the second principal component (PC2, black line) of SSD over the Eurasian  
 4 continent during 1980–2013. The blue line in (b) shows the SSDI index defined as the difference  
 5 of SSD between the rectangles with positive values (46°–58°N, 96°–138°E) and negative values  
 6 (54°–68°N, 48°–84°E) shown in (a). Small and large dotted areas in (a) denote values with  
 7 statistical significance exceeding the 90% and 95% confidence levels, respectively.

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11 **Figure 2.** (a) Correlation coefficients between the SSDI index and the subsequent summer  
 12 precipitation, (b) spatial pattern of the third EOF mode of East Asian summer precipitation, and (c)  
 13 detrended time series of SSDI and summer precipitation (PC3). R represents the correlation  
 14 coefficient between the two indices. Small and large dotted areas denote values with statistical  
 15 significance exceeding the 90% and 95% confidence levels, respectively.

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17 In the present study, we investigated the relationship between Eurasian spring snow decrement  
 18 (SSD) and East Asian summer precipitation and the related thermodynamic and dynamic  
 19 mechanisms using both observational data and the CAM3.1 model. The results show that the  
 20 second EOF mode of Eurasian SSD exhibits a west-east dipole pattern, with a negative center  
 21 located over eastern Europe and the West Siberia Plain (EEWSP) and a positive center over the  
 22 area around Baikal Lake (BL). This anomalous SSD pattern is significantly associated with the  
 23 third EOF mode of East Asian summer rainfall through triggering an anomalous mid-latitude  
 24 Eurasian wave train. The reduced SSD over EEWSP tends to decrease the local soil moisture from

25 spring to the following summer, thereby increase the surface heat flux and near-surface  
26 temperatures. Similarly, the increase in SSD over BL is accompanied by anomalously low  
27 near-surface temperatures. Changes in near-surface temperatures further intensify the meridional  
28 temperature gradient and lower-level baroclinicity, leading to the acceleration of the upper-level  
29 subtropical westerly jet stream. At the 500 hPa geopotential heights, an anomalous cyclone and  
30 anomalous anticyclone emerge to the north and south, respectively, of the exit region of East  
31 Asian subtropical jet stream. Meanwhile, the changed surface thermal conditions enhance the  
32 local 1000–500 hPa thickness over EEWSP while decreasing it over BL. These factors both create  
33 favorable physical conditions for the maintenance and enhancement of the anomalous  
34 mid-latitude Eurasian wave train prevailing over the regions from eastern Europe eastward to the  
35 Northwest Pacific. We further explored the origin of the Eurasian wave train and found that there  
36 are zonally oriented WAFs spreading from eastern Europe eastward to East Asia. This finding  
37 demonstrates the role of anomalous SSD in triggering the Eurasian wave train. These circulation  
38 patterns ultimately significantly influence the precipitation over East Asia, especially over China,  
39 with excessive precipitation over regions west of BL, northeastern China and the Yellow River  
40 valley and deficient precipitation over Inner Mongolia and southern China. Therefore, Our study  
41 confirms the significant role of Eurasian SSD in influencing East Asian summer precipitation.

42 Our model results demonstrate that the CAM3.1 can reproduce the positive summer  
43 precipitation anomalies over most of northern East Asia and negative anomalies over southern  
44 China. Corresponding to the imposed anomalous Eurasian SSD forcings in CAM3.1, positive  
45 surface air temperature and atmospheric thickness responses over EEWSP and negative responses  
46 over BL are primarily simulated. In the simulation the subtropical jet stream over East Asia is  
47 strengthened, the Eurasian mid-latitude wave train is formed, and reinforced WAFs propagate

48 from eastern Europe to East Asia. These simulated dynamic and thermodynamic processes result  
49 in according precipitation anomalies to occur over East Asia.