

EQUILIBRIUM TRANSLATION MODEL – A KEY TO PREDICTION OF TROPICAL HURRICANE INTENSITY

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1. INTRODUCTION

As it is commonly accepted generation and development of tropical hurricane (TH) depends from the following three key conditions: wide area with sea-surface temperature (SST) more than 26 °C, environmental wind with low vertical shear and the Coriolis parameter sufficient for vorticity generation (Palmen and Newton 1969, Emanuel 1999, Goldenberg et al. 2001). Below it is shown that additionally certain type of conformity between dynamical and thermal environmental fields also is decisive, specifically, through rapid intensification of the most powerful hurricanes.

The spectrum of characteristic scales of tropical hurricane (TH) begins with sprayed by wind water droplet and ends by comparable to the planetary sizes inflowing atmospheric currents. Description of such a complex phenomenon requires accounting of great variety of interrelated irreversible thermo-hydrodynamic processes. That is why numerical methods of modeling in combination with wide field observations become main instrument of TH research during last decades. By now this "numerical" attack is in progress with certain achievements, for instance, in TH track forecasting (Webber 2005a).

At the same time the problem of forecasting of rapid changes of TH intensity still remains as the most critical (Krishnamurti et al. 2005, Webber 2005). According to review of forecast advisories (UNISYS Weather: Hurricane/Tropical Data/Archive) during 2004-2005 regular predictions have overlooked or significantly underestimated practically all cases of rapid intensification.

Pressing character of the problem is reflected also by wide front of research of potential mechanisms of TH intensification (Emanuel 2004, Zhu et al. 2004, Wu and Braun

2004, Persing and Montgomery 2005, Lin et al. 2005, Shapiro and Moller 2005, Corbosiero et al. 2005, Wu et al. 2005, Möller and Shapiro 2005). High emphasis is placed also on vital problem of clarification of potential influence of global warming on TH intensity (Trenberth 2005, Kerr 2005, Emanuel 2005).

Aforementioned principal challenges highlight pressing contemporary problem of proper combination of numerical methods with adequate qualitative physical models.

In certain degree this gap is bridged by equilibrium translation model (ETM) (Shekrladze 2004a). ETM links rapid development of tropical hurricane to conformity of dynamical and thermal environmental fields. According ETM TH achieves peak efficiency in conversion of ocean heat potential into cyclonic motion of atmosphere when its translation is found to be in certain accordance with intensity of heat removal from ocean upper layer (alignment effect).

In the framework of ETM the role of the main characteristic of TH development is gained by non-dimensional alignment number incorporating main thermal and dynamical parameters of the system TH-ocean. Alignment number, representing a measure of conformity dynamical and thermal fields, by its part ultimately directly is linked to so-called heat involvement factor.

Fundamental character of the effect is confirmed by wide field data on the cases of rapid intensification observed in 2004-2005 (including tropical hurricanes Charley, Katrina and Wilma). Potential is created for qualitative improvement of tropical hurricane intensity prediction

2. EQUILIBRIUM TRANSLATION MODEL

Certain initial conformity of dynamical and thermal environmental fields almost always holds in the zones of a TH development. It is caused

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by formation of background SST field by background dynamical field itself. In main zones of TH development direction of background SST elevation, as a rule, coincides with environmental wind. In trade winds zone SST elevation rate is of order of $10^{-6} \text{ }^{\circ}\text{C m}^{-1}$ (McGraw-Hill Encyclopedia of Ocean and Atmospheric Sciences 1980).

Another level of conformity of environmental fields is formed by longitudinal SST jump induced by TH itself. TH translation necessarily leads to considerable lowering of SST on its rear. The resultant SST jump forms local SST elevation rate of order of $10^{-5} \text{ }^{\circ}\text{C m}^{-1}$. In this connection the influence of SST jump mentioned assumes prevailing significance.

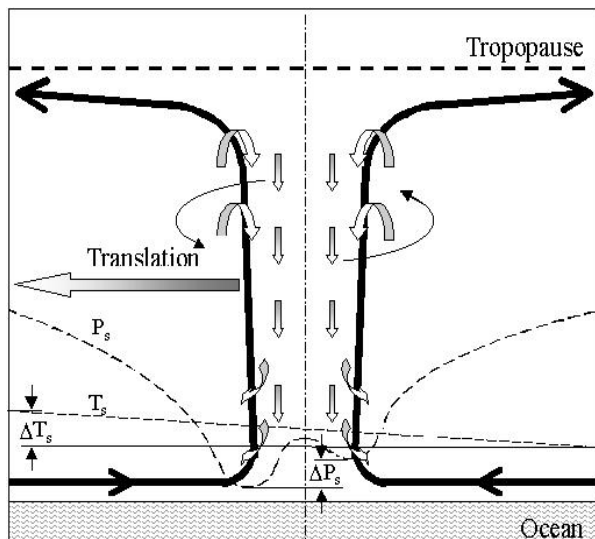


Fig. 1. Scheme of TH thermal driving mechanism.

Interaction of environmental fields may be considered based at the scheme of TH thermal driving mechanism presented in the Fig. 1.

SST jump (ΔT_s) caused by asymmetry of energy inflow in TH leads to more intense ascending flow in frontal part of the eye wall cloud with corresponding minimum of air pressure. Generated by thermal asymmetry pressure drop between frontal and back parts of the eye wall cloud (ΔP_s) just creates driving force.

Intensity of heat and mass transfer from a sea surface to TH is little affected by translation speed. Here main role is played by much higher air tangent velocity (for instance, beginning from outer boundary with tangent velocity 15 ms^{-1}). In this connection reduction of translation speed (leading to prolonging of TH passage above given area of a sea surface) steps up the share of heat removed from upper ocean layer, and vice versa,

gathering of translation reduces cooling of the layer mentioned, all other things being the same.

Inverse dependence mentioned introduces rather strong feedback into thermal driving mechanism. Reduction of translation speed, alongside with increasing ΔT_s , leads to elevation of pressure drop ΔP_s causing, for its part, TH acceleration.

In such a manner, TH not only prefers to shift toward SST elevation direction, but it also tends to establish certain equilibrium between translation speed and integral heat flux.

Generation of thermal driving force depends from TH translation speed and hurricane thermal potential (HHP) (Leipper and Volgenau 1972) of given area. Besides, generation of the same driving force needs more slow translation of TH at high value of HHP and vice-versa. This inverse dependence roughly corresponds to constancy of heat involvement factor (equal to the share of HHP removed by TH from an ocean through passage of given area) during equilibrium translation. The last supposition is supported by analysis of the field data on TH Opal (1995) presented in Fig. 2.

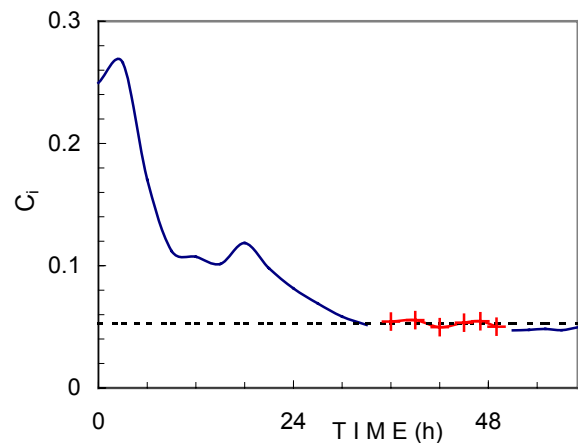


Fig. 2. Variation of heat involvement factor (C_i) during development of TH Opal (1995): crosses – TH rapid intensification stage.

At the same time establishment of equilibrium translation is possible only through it's favoring by dynamical field. ETM assumes that Just conformity of dynamical and thermal fields triggers intensive TH development (alignment effect).

2.1 The First Approximation

Realization of ETM in the first approximation (FAP) (Shekrladze 2004a) is built upon

consideration of uniform straightforward translation of circular TH. Within accepted assumptions heat balance of an ocean upper layer inside an area with leading role of tangent winds in heat removal from a sea surface is determined in following form:

$$2C_i Q R_{15} \left(U_{tr} - U_{dr} - C_\phi \frac{dR_{15}}{d\tau} \right) = \pi R_{15}^2 q \quad (1)$$

Here C_i is heat involvement factor; Q is averaged inside R_{15} HHP; R_{15} is radius of a TH at tangent velocity 15 m s^{-1} ; q is averaged inside R_{15} integral (sensitive and latent) heat flow from sea surface to TH; U_{dr} is component of sea surface drift speed parallel to translation; τ is time.

Assuming in addition insignificance of a sea surface drift, relationship between main dynamical and thermal parameters of combined ocean-atmosphere system during equilibrium translation (to say, at the stage of TH intensification) is reduced to following equation for non-dimensional alignment number:

$$N_{al} = \frac{U_{tr} Q}{q R_{15}} = 25 \quad (2)$$

For TH holding roughly constant size and other parameters near to reference case (TH Kenna, 2000), production qR_{15} is determined by following empirical equation:

$$qR_{15} = 2.125 \cdot 10^8 \left(\frac{U_{\max}}{145} \right)^{0.5} \text{ Wm}^{-1}, \quad (3)$$

Here U_{\max} is maximum tangent wind velocity at given instance of time in knots.

2.2 The Second Approximation

Analysis of ETM in the second approximation (SAP) covers general case of TH of non-circular geometry (Fig. 3). By the goal of utilization of TH parameters reflected in regular forecast advisories TH radius at tangent winds 34 knots

(17.5 m s^{-1}) is accepted as TH outer boundary. Integral heat flow (sensitive and latent) removed from left behind sea strip may be written in following form:

$$A_{34} q = C_i Q W_{34} (U_{bb} - U_{dr}) \quad (4)$$

Here the nomenclature corresponds to Fig. 3; q is averaged inside A_{34} integral heat flux to TH; Q is averaged inside A_{34} HHP; U_{bb} is translation speed of TH back boundary center; W_{34} ($U_{bb} - U_{dr}$), to a certain approximation, determines increment of cooled sea surface.

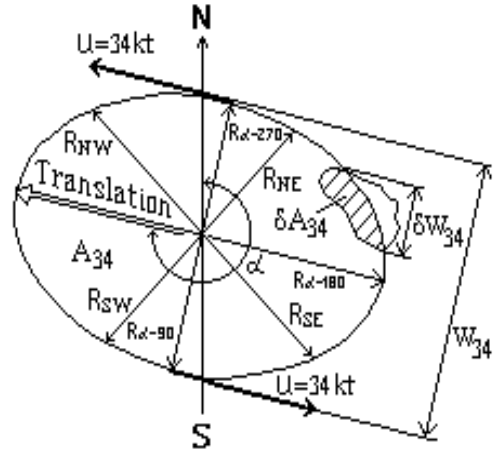


Fig. 3. Scheme of non-circular TH translation: A_{34} is area inside tangent velocity 34 knots (17.5 m s^{-1}); R_{NE} , R_{SE} , R_{SW} and R_{NW} are TH radii at tangent velocity 34 knots in Northeast, Southeast, Southwest and Northwest quadrants, respectively; α is TH translation azimuth; W_{34} is TH transverse width at tangent velocity 34 knot; δA_{34} is an area of land surface covered by A_{34} ; δW_{34} is projection onto W_{34} of the part of the back boundary of A_{34} covered by land surface.

Slightly refined value of heat involvement factor also is introduced through SAP based at reanalysis of the field data on TH Opal:

$$C_i \approx \text{Const} = 1/19 \quad (5)$$

Smooth (linear with the angle) distribution of intermediate TH radius is assumed through

calculation of $R_{\alpha-90}$, $R_{\alpha-180}$ and $R_{\alpha-270}$ serving further for determination of TH transverse width and back boundary center coordinates. The area A_{34} is determined as sum of the quadrants using aforementioned 4 values of TH radius. Further, with regard to insignificance of sea surface drift, corresponding to value $C_i \approx 1/19$ condition of establishment of equilibrium translation is written:

$$N_{al} = \frac{\pi W_{34} U_{bb} Q}{2A_{34} q} = \frac{U_{bb} Q}{q R_{ef}} \approx 30 \quad (6)$$

Here R_{ef} is effective radius of TH:

$$R_{ef} = 2A_{34} / \pi W_{34} \quad (7)$$

In such a manner, in the framework of SAP, characteristic linear size of non-circular TH turns to be a function of the square and transverse width of area inside tangent velocity 34 knot.

Further it is necessary to evaluate influence of land surface covering. According to equations (6-7) land surface influences heat balance through reduction of sea surface inside A_{34} and TH active back boundary. Besides, as the first type of influence works through any location of land surface inside A_{34} , reduction of back boundary may take place only if it immediately is crossed by land surface.

Reasoning from these deductions accounting of influence of land surface is carried out using correction factors accounting the shares of δA_{34} and δW_{34} through determination of alignment number. Besides, such a simple approach is applicable only if covered by TH land surface makes comparatively small part of A_{34} . In the case of significant share of land surface more detailed additional analysis is necessary.

Finally, it should be noted also that assumption about decisive role of ocean upper layer heat potential naturally restricts applicability of ETM in the case of small TH radius when contribution of initial energy content of environmental air also may become valuable.

2.3 Three-zone model of integral heat transfer

Evaluation of average integral heat flux from sea surface to TH turns as the most complicated problem through analysis of individual cases in

the framework of ETM. Available results of analysis and field data (Bortkovskiy 1983, Grigorkina 1986, Jacob et al. 2000, D'Asaro 2003, Shekrladze 2004b) on integral heat fluxes have rather wide dispersion (from 0.25 kW m^{-2} to around 4 kW m^{-2}). In the context of ETM it represents exceptional importance the records (Shay et al. 2000) of immediate measurement of sensitive and latent heat fluxes made during extremely rapid intensification (to say, during equilibrium translation) of TH Opal in 1995.

Averaging of heat flux within the area inside A_{34} attaches to this parameter certain peculiarities. As area covered by tangent winds in the range $17-25 \text{ m s}^{-1}$ makes around 80% of A_{34} , this parameter to greater extent depends from heat flux at outer boundary. Besides, within the assumptions made, the last is independent from the most intense tangent winds. At the same time the zone with intense tangent winds also contributes in formation of average heat flux compensating in certain degree the zone's smallness by high values of local fluxes. In this connection adequate evaluation of average heat flux needs accounting of real tangent wind velocity distribution inside A_{34} .

In general integral heat flux from ocean upper layer to TH is function of temperature, humidity and velocity fields. Besides, in contrast to single-phase turbulent flow, here decisive role is played by intensity of seawater spraying by tangent winds. As it follows from corresponding analysis general solution of the problem needs further analytical and experimental studies. Below approximate three-zone model is developed focused to equilibrium translation mode.

As it is known (Bortkovskiy 1983, Grigorkina 1986, Shay et al. 2000) dynamical impact of TH on an ocean includes intensive vertical mixing of seawater. Besides, resulted by vertical mixing downward heat flow exceeds sought upward integral heat flow at an order of magnitude (Shay et al. 2000). In this connection vertical mixing becomes decisive factor through transformation of initial SST field. Besides, as it is shown by Goni and Trinanes (2003), this process mainly is restricted by initial HHP. Consequently, these two factors: intensity of vertical mixing and initial HHP level are main factors determining real SST field under TH.

At the same time constancy of involvement factor causes another important peculiarity of heat transfer specifically during equilibrium translation. In general, removal by TH of constant share of HHP needs more time at high

HHP and vice-versa (other things being the same). In this connection equilibrium translation is comparatively slow at high HHP and comparatively fast at low HHP. Accordingly, the area with high initial HHP undergoes more intensive cooling than the area with low initial HHP.

This circumstance allows assuming that SST field under TH is roughly the same irrespective of initial HHP value if translation corresponds to equilibrium mode (other things, to say, tangent wind distribution, being the same). Correspondingly, to a certain approximation, average integral heat flux during equilibrium translation also turns to be independent from initial HHP. It is evident also that the last conclusion, assuming inter-compensating influences of intensity of vertical mixing and initial HHP field, is valid only for equilibrium translation. In such a case average heat flux becomes single-valued function of tangent wind distribution that essentially simplifies further analysis.

Simultaneously the subject matter reduces to near-equilibrium mode of translation restricting applicability of the approach outside of this mode. However, at this stage such a specializing is quite allowable in respect to our focusing to equilibrium translation.

Further, having in account field data structure of regular forecast advisories, three-zone model of heat transfer is introduced. The first (outer) zone covers the area between tangent winds 34 knot (17.5 m s^{-1}) and 50 knot (25.75 m s^{-1}). The second (intermediate) zone covers the area between tangent winds 50 knot and 64 knot (33 m s^{-1}). The third (central) zone covers the area inside tangent wind 64 knot. Within used approach maximum tangent wind influences integral heat transfer only in central zone. Besides, influence of central zone on average integral heat flux is smoothed by comparatively small share of this zone in A_{34} (5 - 15 %).

Further, reasoning from all above and using available field data, following empirical equation is developed for average integral heat flux from sea surface to TH:

$$q = [375(R_1^2 - R_2^2) + 600(R_2^2 - R_3^2) + 1600(U_{\max}/155)R_3^2]R_1^{-2} \text{ W m}^{-2} \quad (8)$$

Here R_1 , R_2 and R_3 are average outer radii in meters of first, second and third zones determined as a quarter of square root from sum

of squares of corresponding radii in aforementioned four quadrants; U_{\max} is maximum tangent wind velocity inside TH at given position in knots. Characteristic for the first and second zones values 375 W m^{-2} and 600 W m^{-2} are determined through analysis and rounding the data recorded during rapid intensification of TH Opal (Shay et al. 2000). The value 1600 W m^{-2} is determined through equating of arithmetic mean of 8 values of alignment number of super-typhoons Dianmu (2004) and Chaba (2004) at intensity 155 knot to 30.

Finally, it should be noted also that revealed by equation (8) strong dependence of integral heat flux from tangent wind velocity (in comparison with single-phase turbulent heat transfer) looks quite reasonable in the context of decisive role of sea water spraying.

3. CORRELATION OF THE FIELD DATA

The main objective of correlation of the field data in the framework of equations (6-8) is verification of predicted by ETM linkage of TH rapid intensification to alignment number (alignment effect).

TH translation speed (U_{bb}) is determined through geographical coordinates of TH back boundary center as arithmetic mean of average values for straight translation at two lengths of standard track data prior and subsequent to given position. In certain degree accuracy of this procedure may be assessed using two other alternative procedures of determination of U_{bb} through average values only for one length of standard track data prior or subsequent to given position.

Comparison of aforementioned three alternative values allows to revealing TH translation with significantly variable speed or through curved or zigzagging trajectory when received procedure may turn to be a priori inapplicable. In any case unavailability of local values of translation speed affects accuracy of correlation although in majority of cases situation is mitigated by rather straight and uniform translation of TH just at the stage of rapid intensification.

The problem with initial data precision manifests itself also with HHP. Following from the results of comparative consideration of different HHP maps (NRL hurricane heat potential maps 2004-2005, Goni and Trinanes - hurricane heat potential maps, 2004-2005, Cook - hurricane heat potential maps, 2004-2005), the maps Goni and Trinanes hurricane heat potential

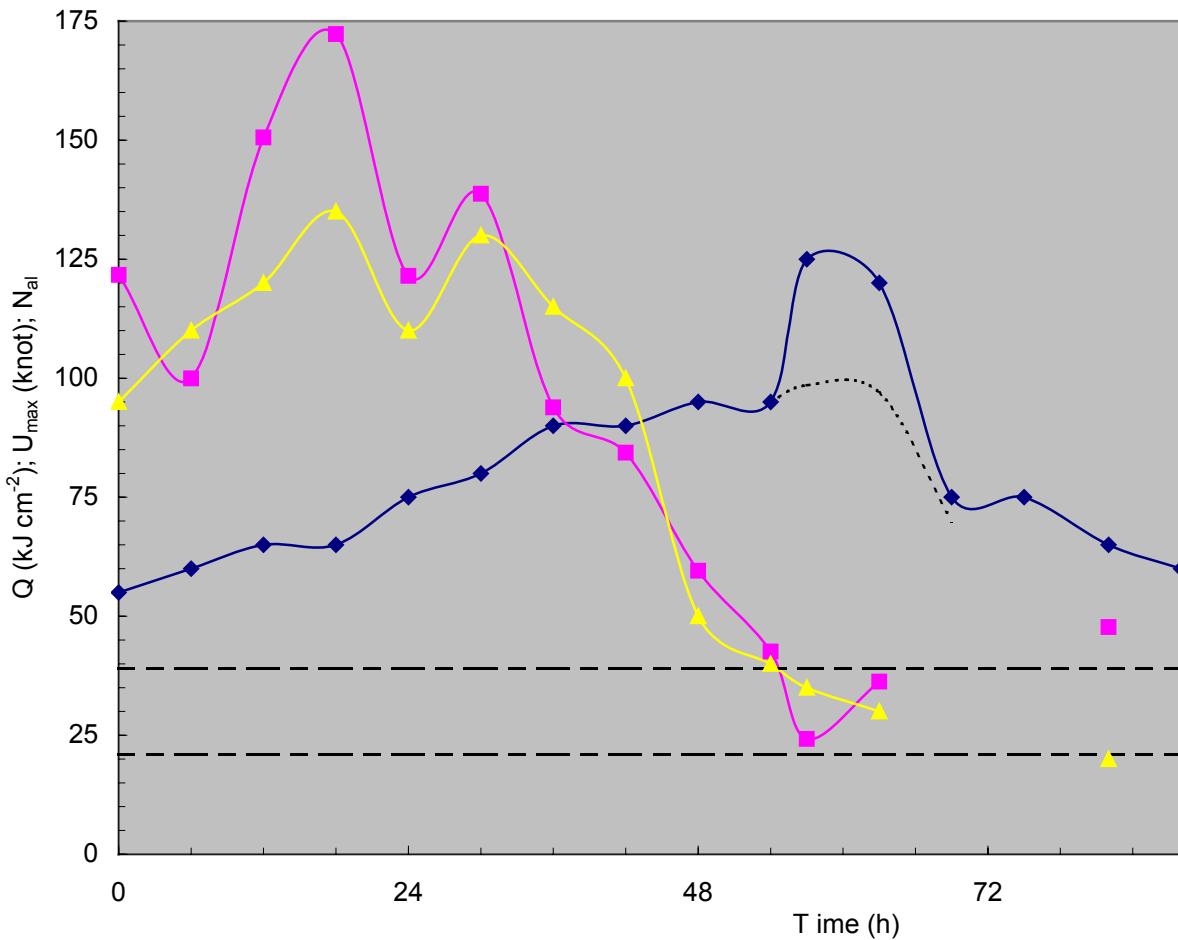


Fig. 4. Correlation of the field data on TH Charley (North West Atlantic, August, 2004) with equations (6-8): triangles - Q (kJ cm^{-2}); rhomboids - U_{max} (knot); dotted curve - regular forecast issued 15: 00 (UT) 13.08.2004; squares - N_{ai} ; horizontal lines - the range $N_{\text{ai}}=30\pm 30\%$; blank spaces at 69-75th hours correspond to crossing Florida Peninsula; the point 0 corresponds to 09:00 (UT) 11.08. 2004.

maps (2004-2005) are selected as a basis for consideration.

The stage of rapid intensification is specified by intensification rate no less than 5,0 knot/h and achievement of minimum 4th category of intensity. Following 8 THs are identified as satisfying such a condition among all THs observed in 2004-2005 in all relevant zones of Tropical Ocean: hurricanes Charley, Katrina and Wilma (North West Atlantic), super typhoons Dianmu and Chaba (North West Pacific), cyclones Nancy and Olaf (South Pacific) and Ingrid (South Pacific and South Indian).

Aforementioned correction factors to A_{34} and W_{34} are used through accounting of influence of partial land covering only if corresponding reduction is no more than a third of either of these two parameters. At TH positions, when greater reduction of A_{34} or W_{34} takes place, alignment

number is considered as uncertain in developed framework quantity that is reflected in corresponding figures by blank spaces. The results of correlation of some important cases are presented in Figures 4-8.

The first among considered cases is TH Charley, which relates to the number of the most disastrous cases of unexpected rapid intensification. Springing at east boundary of Caribbean Sea Charley passes Northwest approaching west end of Cuba. As it follows from Fig. 4, despite very high values of HHP (100-130 kJ cm^{-2}), Charley develops quite moderately at this stage that should be linked to significant excursion of its real translation from equilibrium mode ($N_{\text{ai}}\approx 80-150$).

Crossing Cuba and turning north to Florida Peninsula Charley is found in the zone with significantly lower HHP (30-40 kJ cm^{-2}).

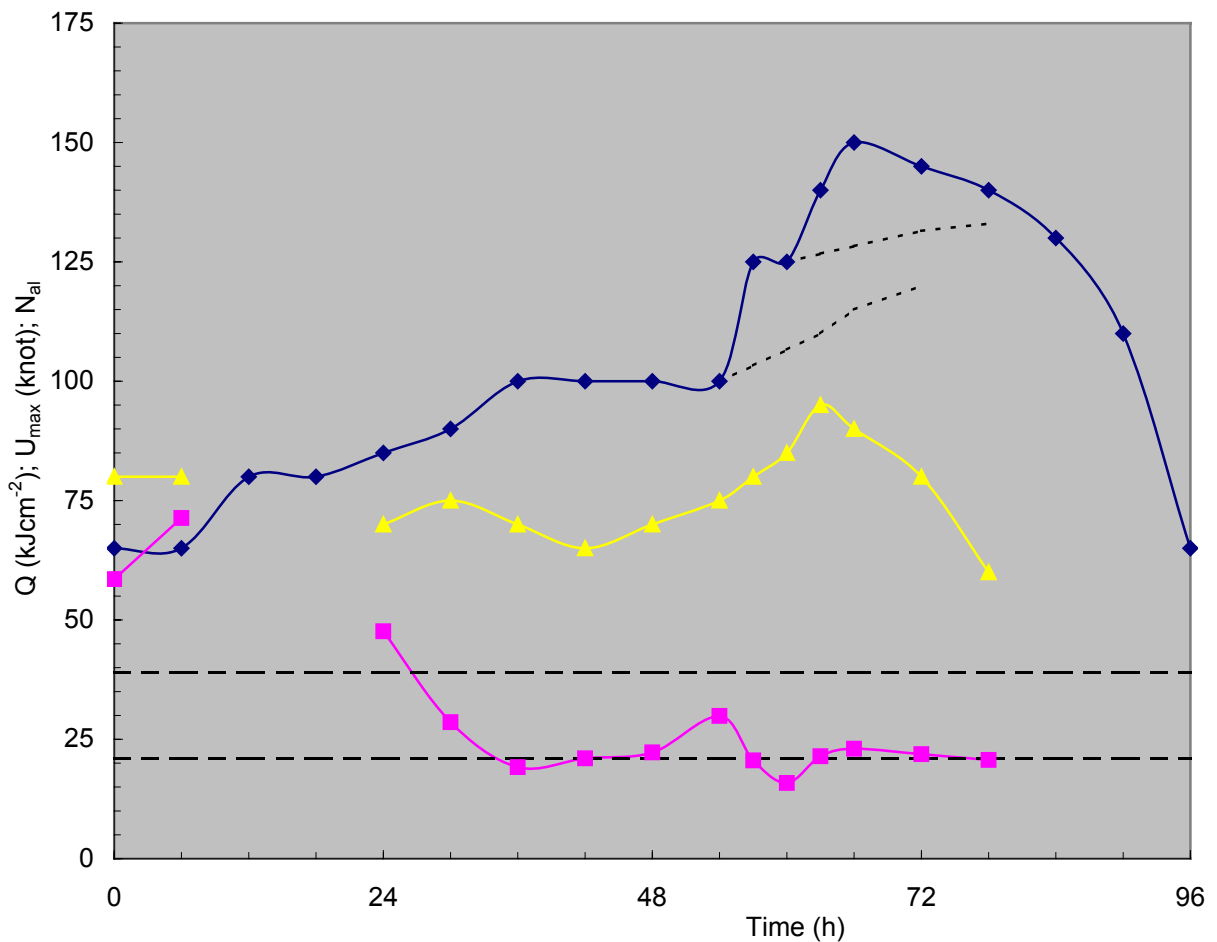


Fig.5. Correlation of the field data on TH Katrina (North West Atlantic, August, 2005) with equations (6-8): triangles - Q (kJ cm^{-2}); rhomboids - U_{\max} (knot); dotted curves - regular forecasts issued 03:00 (UT) and 09:00 (UT) 28.08. 2005; squares - N_{al} ; horizontal lines - the range $N_{al}=30\pm 30\%$; blank spaces at 12-18th hours correspond to crossing Florida Peninsula, at 84-96th hours - to entering continental part of USA; the point 0 corresponds to 21:00 (UT) 25.08. 2005.

Alignment number significantly reduces (from around 80 to around 25). Charley's translation immediately approaches equilibrium translation mode and Charley strengthens very rapidly (from 95 to 125 knot) during only 3 hours.

Unfortunately, Charley's unexpected intensification took place just prior to landfall to Florida Peninsula. Besides, intensification rate at this stage (10 knot/h) is second high among all selected cases of rapid intensification. Correspondingly, as Charley's rapid intensification is observed against rather sharp reduction of HHP, considered sequence of events clearly demonstrates crucial role of alignment effect in the observed phenomenon. At the same time the effect evidently is overlooked by numerical models serving through preparation of regular forecasts.

Differing sequence of events is observed during development of TH Katrina (Fig. 5). Springing in the zone of Bahamas Islands Katrina crosses Florida Peninsula and Gulf of Mexico making final landfall in the zone of New Orleans. Consequent to establishment of extremely favorable for alignment effect conditions ($N_{al}\approx 20-30$; $Q\approx 70-80 \text{ kJ cm}^{-2}$) Katrina strengthens in central part of the Gulf from 100 to 125 knot during only 3 hours (intensification rate 8.33) achieving consequent to short-time "respite" intensity 150 knot. Further Katrina weakens alongside with reduction of both of these parameters although it retains extremely dangerous level of intensity up to landfall. (Katrina made landfall as TH of 4th category).

As alignment number holds favorable for TH intensification range during almost all life cycle of

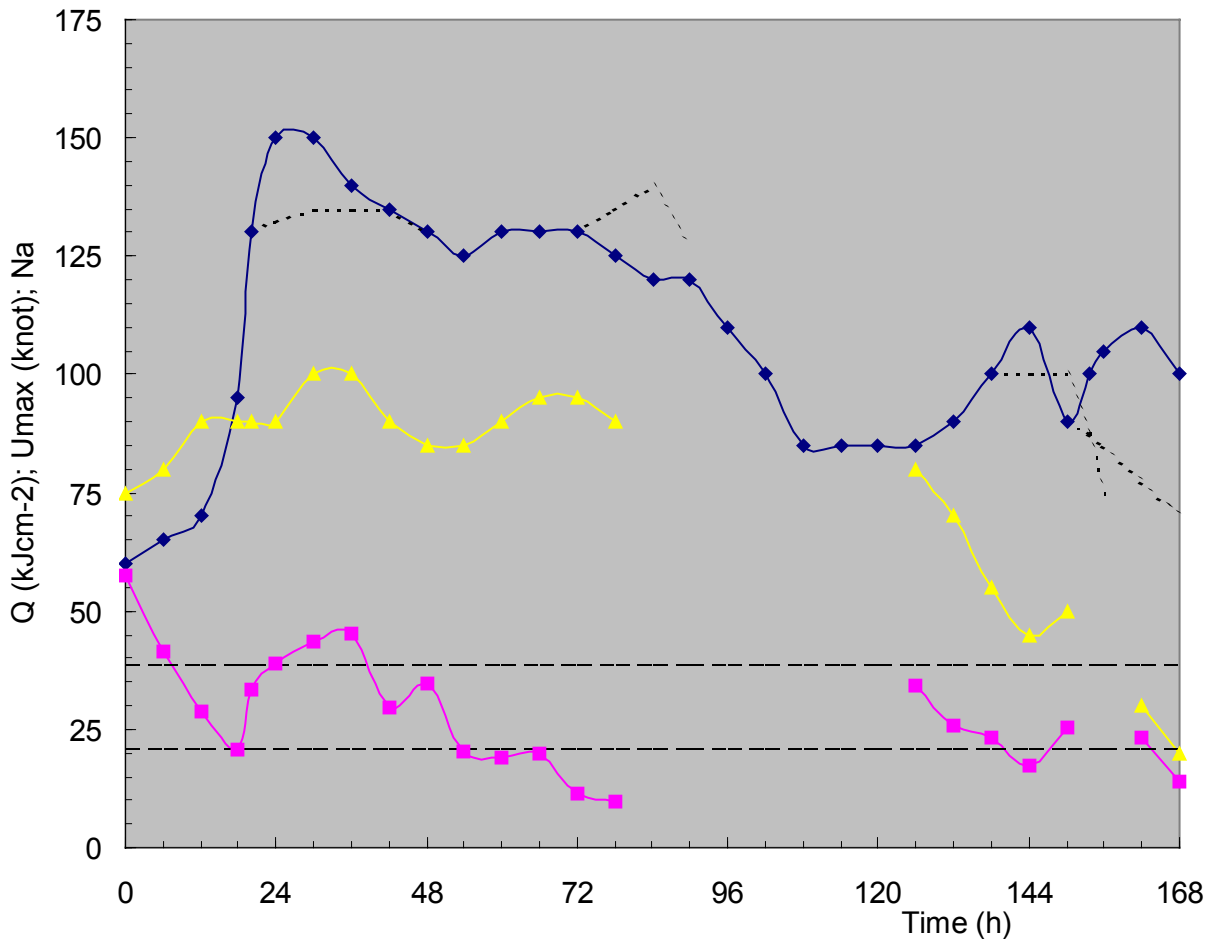


Fig. 6. Correlation of the field data on TH Wilma (North West Atlantic, October, 2005) with equations (6-8): triangles - Q (kJ cm^{-2}); rhomboids - U_{\max} (knot); dotted curves - regular forecasts issued 05:00 (UT) 19.10.2005, 09:00 (UT) 21.10.2005, 03:00 (UT) and 21:00 (UT) 24.10.2005; squares - N_{a1} ; horizontal lines - $N_{a1}=30\pm 30\%$; blank spaces at 84-120 hours correspond to covering of Yucatan Peninsula, at 153-156 hours - to crossing Florida Peninsula; the point 0 corresponds to 09:00 (UT) 18.10.2005.

Katrina, in contrast to Charley, variation of HHP naturally gains leading role in TH development. This circumstance clearly is reflected in Fig. 5 by similarity of TH intensity and HHP curves. Correspondingly, forecasting misses also are comparatively insignificant.

The last is TH Wilma (Fig. 6) among cases of rapid intensification observed in 2004-2005 in North West Atlantic. Springing in central part of Caribbean Sea with high values of HHP ($70\text{-}95 \text{ kJ cm}^{-2}$), Wilma passes northwest, hits Yucatan Peninsula and turns northeast to Florida.

Further Wilma, touching firstly west end of Cuba, crosses Florida Peninsula. Although Wilma starts at rather high alignment numbers ($N_{a1}\approx 55$) its translation promptly approaches equilibrium mode explosively intensifying from 95 to 130 knot

with record-breaking intensification rate 17.5 knot/h during 2 hours. For 4 more hours Wilma achieves 150 knot.

Alternating stages of intensity attenuation and re-intensification, further Wilma remains as rather strong TH of 3rd category through crossing Florida Peninsula.

As it follows from Fig. 6 Wilma's intensity curve adequately correlates with alignment number. At the same time, except underestimation of forthcoming rapid intensification, regular forecasts repeatedly fail in prediction of intensity changes, including overlooking of continuation of Wilma's intensification through landfall to Florida Peninsula and re-intensification through entering North Atlantic.

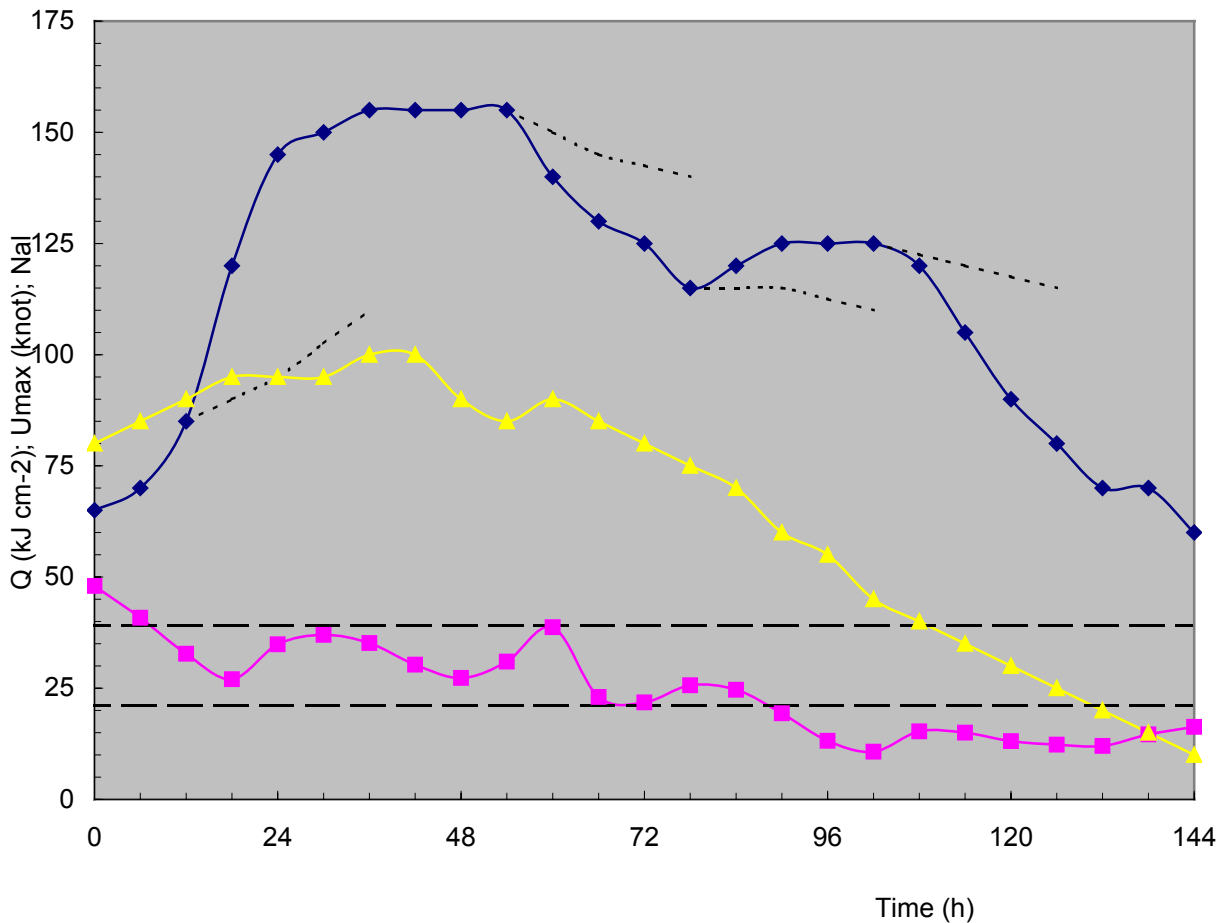


Fig. 7. Correlation of the field data on super typhoon Dianmu (North West Pacific, June, 2004) with equations (6-8): triangles - Q (kJ cm^{-2}); rhomboids - U_{max} (knot); dotted curves - regular forecasts issued 06:00 (UT) 16.06.2004, 00:00 (UT) 17.06.2004, 00:00 (UT) 18.06.2004 and 00:00 (UT) 19.06.2004; squares - N_{al} ; horizontal lines - $N_{\text{al}}=30\pm 30\%$; the point 0 corresponds to 18:00 (UT) 14.06. 2004.

Combined influence of alignment number and HHP clearly is revealed also through correlation of the filed data on super typhoon Dianmu (Fig. 7) that has developed up to record high for 2004 155-knot intensity in North West Pacific.

As it follows from Fig. 7, consequent to reduction of alignment number to around 30 and corresponding approach to equilibrium translation mode Dianmu rapidly strengthens from 85 to 120 knot during 6 hours with intensification rate equal to 5.83.

Subsequent to Dianmu super typhoon Chaba (Fig. 8) repeats record high for 2004 155-knot intensity in North West Pacific. Starting with rather high values of alignment number in the zone with rather high HHP ($100\text{-}115 \text{ kJ cm}^{-2}$), translation of Chaba promptly approaches equilibrium mode. Simultaneously Chaba rapidly strengthens from 125 to 155 knot during 6 hours with intensification

rate 5.0. Besides, correlation between TH intensity and N_{al} is observed during almost all life cycle of Chaba. In this connection Chaba keeps very high intensity even through tangible reduction of HHP.

Crucial role of alignment effect through TH rapid intensification is clearly demonstrated also through correlation of the field data on cyclone Nancy (Fig. 9). Besides, in full contrast to regular forecast, TH rapid intensification with intensification rate 6.25 starts from the very beginning within the range $N_{\text{al}}\approx 30$. During 12 hours intensity of Nancy increases from 45 to 120 knot achieving further maximum level 125 knot. At 24th hour Nancy begins to weaken also in full contrast with regular forecast.

Similar situation is observed also through rapid intensification of cyclone Ingrid in the Gulf of Carpentaria (Fig. 10). Ingrid strengthens from 80 to 120 knot with intensification rate 6.67

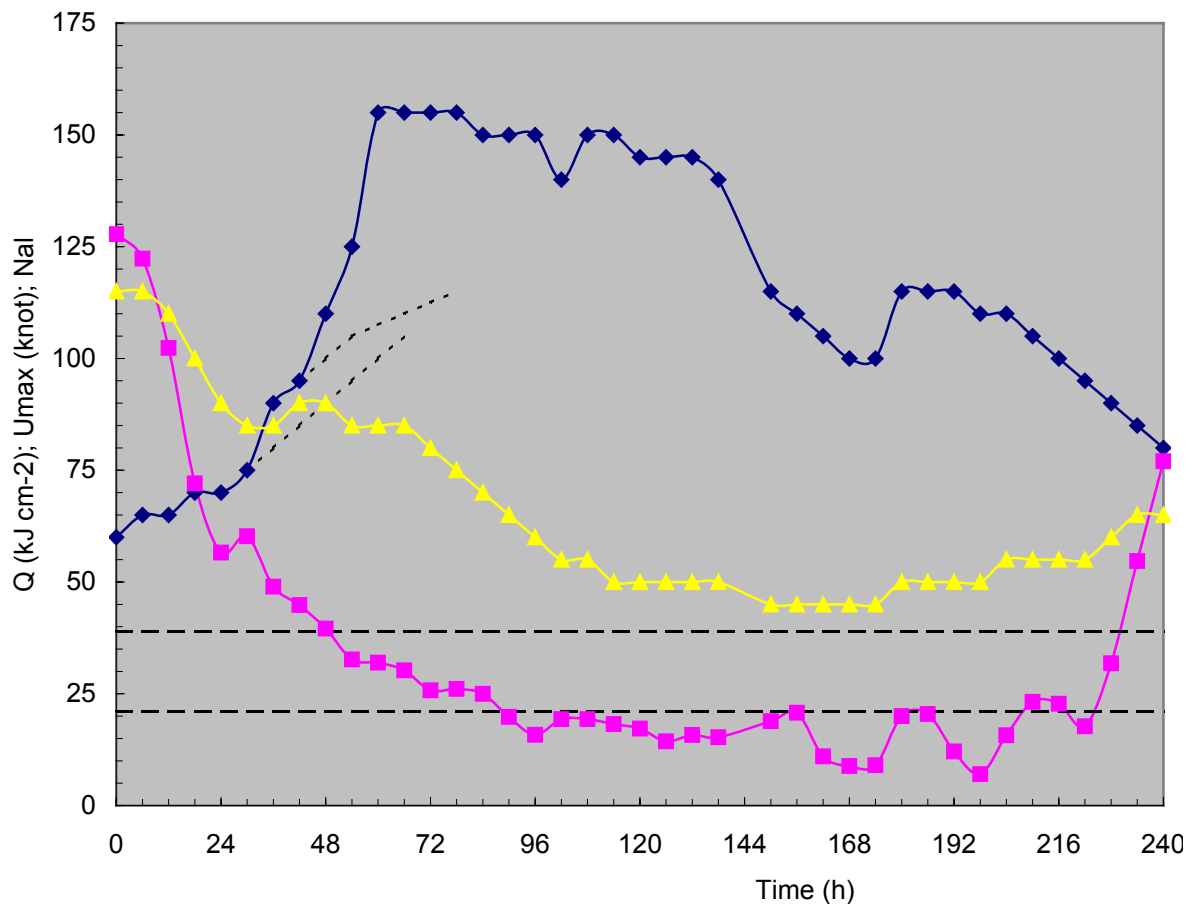


Fig. 8. Correlation of the field data on super typhoon Chaba (North West Pacific, August, 2004) with equations (6-8): triangles - Q (kJ cm^{-2}); rhomboids - U_{\max} (knot); dotted curves - regular forecasts issued 12:00 (UT) 21.08.2004; 00:00 (UT) 22.08.2004; squares - N_{al} ; horizontal lines - $N_{al}=30\pm 30\%$; the point 0 corresponds to 06:00 (UT) 20.08. 2004.

within the range $N_{al}=20-30$. It presents interesting peculiarity of Ingrid's development existence of 3 phases of intensification during its life cycle. Besides, among all considered cases Ingrid represents exclusive case when the rapid intensification takes place in the mode of re-intensification.

As a whole, the conclusion on leading role of alignment effect through TH rapid intensification equally is supported by correlation of all selected cases except cyclone Olaf (Fig. 11) that deserves of special consideration.

As it follows from Fig. 11, rapid intensification of Olaf (from 90 to 125 knot during 6 hours with intensification rate 5.83) takes place at rather high values of alignment number ($N_{al}\approx 200-90$).

According to corresponding analysis, this miscorrelation may be linked to small radius of Olaf. As it was mentioned above, small TH radii restrict applicability of ETM when assumption

about decisive role of ocean upper layer heat content may lose its force and contribution of initial energy content of environmental air also may become valuable.

Really, Olaf has the smallest average radius at tangent velocity 50 knot to the end of rapid intensification ($R_2=56$ km) among all cases (the radius R_2 is selected here as TH parameter most suitable for evaluation of the role of oceanic heat removal). Among considered cases maximum value of R_2 takes place during rapid intensification of super- typhoon Chaba ($R_2=142$ km). Neighbor to Olaf values of this parameter are observed in the cases of cyclone Ingrid ($R_2=71$ km) and TH Charley ($R_2=84$ km). At that leading role of alignment effect clearly manifests itself in both of these cases.

Such a contrast between TH Olaf, on the one hand, and THs Ingrid and Charley, on another, partly may be caused also by high

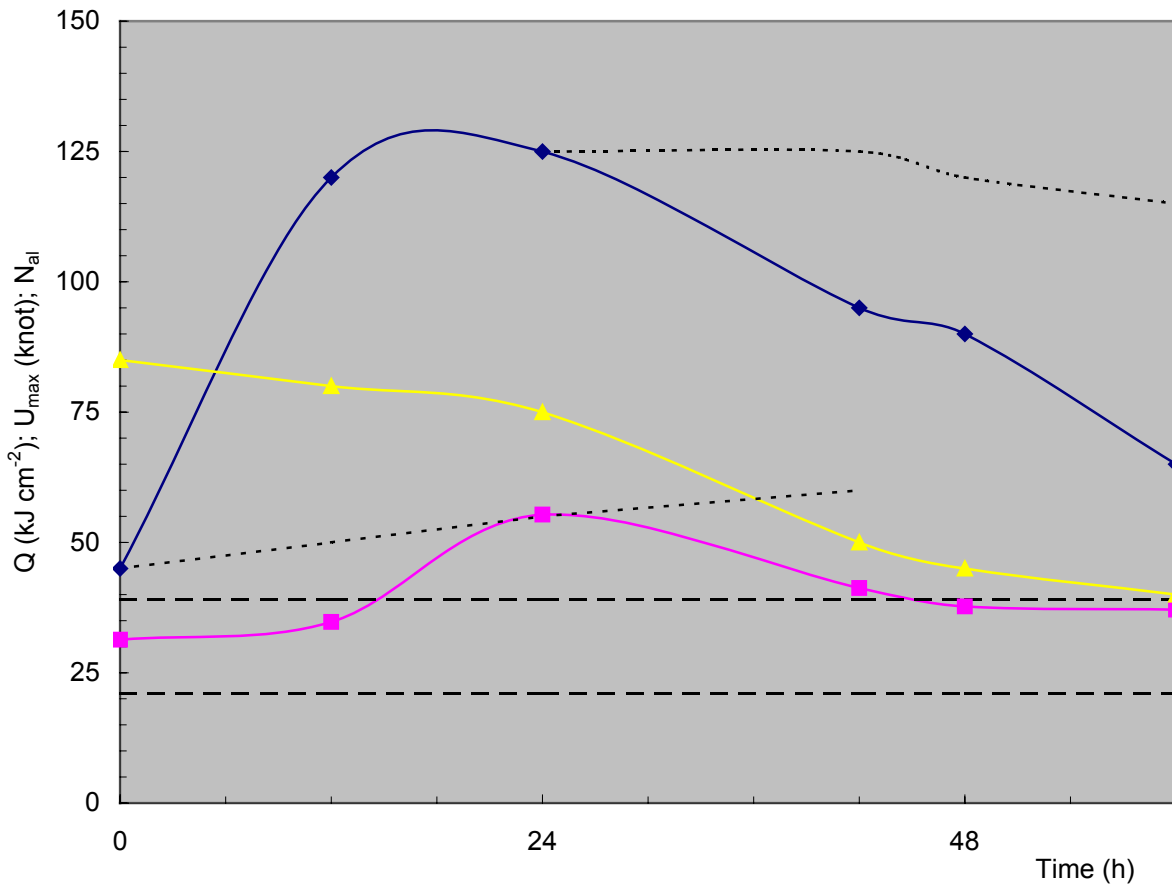


Fig.9. Correlation of the field data on cyclone Nancy (South Pacific, February, 2005) with equations (6-8): triangles - Q (kJ cm^{-2}); rhomboids - U_{max} (knot); dotted curves - regular forecasts issued 18:00 (UT) 13.02. 2005 and 18:00 (UT) 14.02. 2005; squares - N_{al} ; horizontal lines - the range $N_{\text{al}}=30\pm 30\%$; the point 0 corresponds to 18:00 (UT) 13.02. 2005.

humidity of Olaf's inflow formed above extent open sea area of South Pacific with very high HHP ($120\text{-}140 \text{ kJ cm}^{-2}$).

In contrast to it contribution was significant of comparatively dry continental air through inflow formation in the cases of rapid intensification of TH Ingrid in the region of the Gulf of Carpentaria and TH Charley in the region of Florida Strait.

However, relevance range of ETM requires further detailed study.

4. CONCLUDING REMARKS

Predicted by ETM alignment effect manifests itself in overwhelming majority of cases of TH rapid intensification observed in 2004-2005 in all relevant zones of Tropical Ocean (in 7 cases from 8 cases in total). Combined influence of alignment effect and HHP on TH intensity also is established although in some cases alignment effect takes

place even against significant reduction of HHP.

Preliminary rough evaluation also is made of limitations of ETM linked to initial energy content of air inflow in the range of small radii of TH.

Comprehensive essence is revealed of Non-dimensional alignment number (incorporating main integral parameters of the system TH-ocean) as a measure of conformity of dynamical and thermal fields.

Unified algorithm is developed describing wide field data without use of any fitting coefficient variable from one TH to another.

The potential of qualitative physical models once again is demonstrated to discover basic features of any subject matter, on this occasion, by the example of exclusively complex and powerful multi-scale natural phenomenon.

Further development of MET and incorporation of alignment effect into numerical models are necessary next steps toward

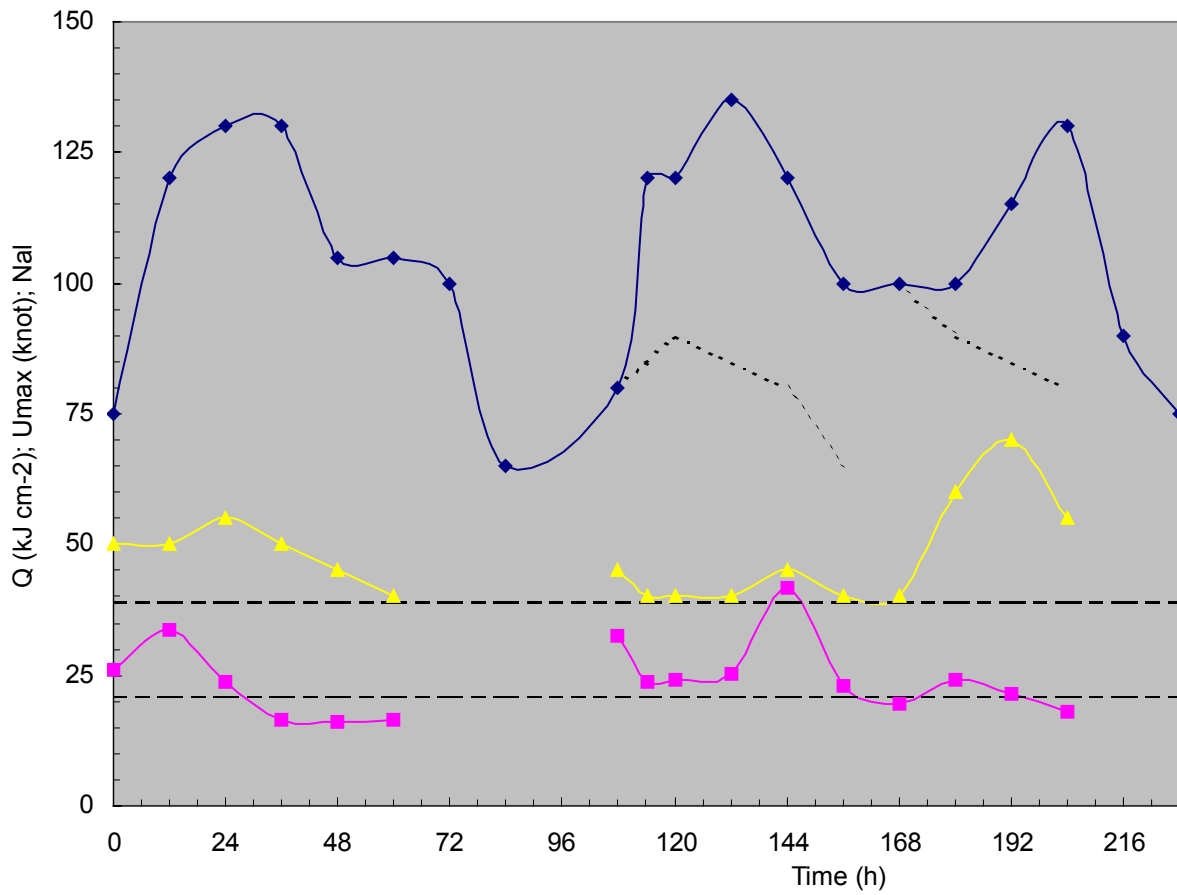


Fig. 10. Correlation of the field data on cyclone Ingrid (South Pacific and South Indian, March, 2005) with equations (6-8): triangles - Q (kJ cm^{-2}); rhomboids - U_{\max} (knot); dotted curves - regular forecasts issued 06:00 (UT) 11.03. 2005 and 18:00 (UT) 13.03. 2005; squares - N_{al} ; horizontal lines - the range $N_{al}=30\pm 30\%$; blank spaces at 72-84th hours correspond to crossing Cape Peninsula, at 216-218th hours - to covering of north coast of Australia; the point 0 corresponds to 18:00 (UT) 06.03.2005.

qualitative improvement of TH intensity forecasting. At the same time developed here algorithm already allows evaluation of current regular forecasts in real-time mode.

Investigation of the effect may contribute also in clarification of vital problem of potential influence of global warming on regional features of development of tropical hurricanes.

As correlation is decisive of TH equilibrium translation with prevailing regional winds, in some cases elevation of HHP may lead even to hindering of TH development (by mistuning of TH translation and heat removal). Otherwise, reduction of HHP may cause opposite effect.

In this context, the most unwanted scenarios may be realized in the regions where long-term elevation of HHP is accompanied by favorable for alignment effect transformation of correlation between equilibrium mode of TH translation and prevailing regional winds.

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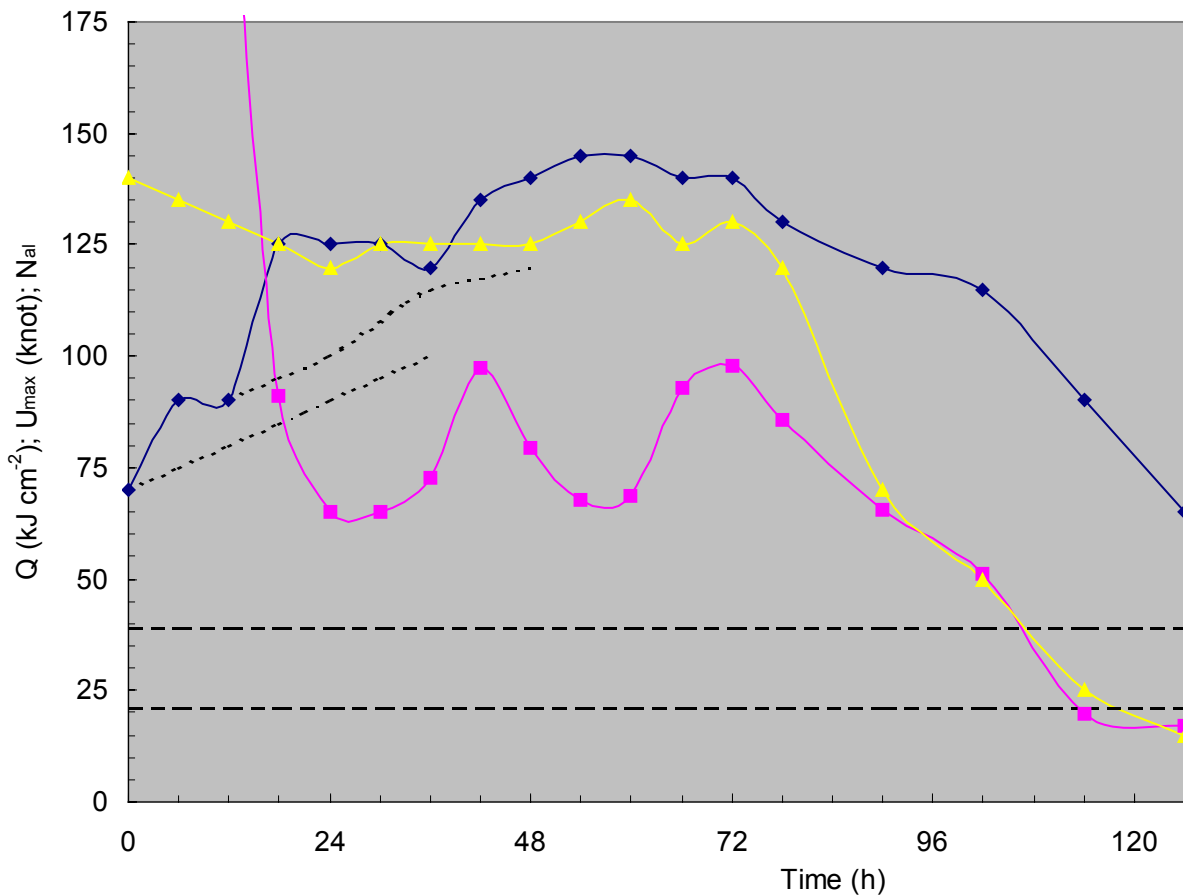


Fig. 11. Correlation of the field data on cyclone Olaf (South Pacific, February, 2005) with equations (6-8): triangles - Q (kJ cm^{-2}); rhomboids - U_{\max} (knot); dotted curves - regular forecasts issued 00:00 (UT) and 12:00 (UT) 14.02.2005; squares - N_{al} ; horizontal lines - $N_{al}=30\pm 30\%$; the point 0 corresponds to 00:00 (UT) 14.02. 2005.

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**PROCEDURE AND TABLES
OF CORRELATION OF THE FIELD DATA ON THE
CASES OF TH RAPID INTENSIFICATION
(2004-2005)**

The appendix to the Paper “**EQUILIBRIUM TRANSLATION MODEL – A KEY TO PREDICTION OF TROPICAL HURRICANE INTENSITY**” contains detailed description of the algorithm of correlation of the field data on tropical hurricane (TH) rapid intensification in the framework of the second approximation (SAP) of equilibrium translation model (ETM), results of correlation of all 8 cases of rapid intensification observed in 2004-2005 in all relevant zones of Tropical Ocean (including tables of correlation). Numbering of equations and figures continues corresponding numbering of the Paper. Made in the appendix references correspond to the chapter of the Paper “REFERENCES”.

1. TH geometrics

Presented in Fig.3 idealized non-circular contour is based at 4 values of TH radius at tangent velocity 34 knot in Northeast, Southeast, Southwest and Northwest quadrants (R_{NE} , R_{SE} , R_{SW} , R_{NW}) specified in regular advisories (45° , 135° , 225° and 315° from North, respectively) according to UNISYS Weather.

At first stage four radii R_{α} , $R_{\alpha-90}$, $R_{\alpha-180}$ and $R_{\alpha-270}$ are determined through simple smooth trigonometric interpolation of aforementioned initial data:

$$R_{\alpha} = R_{SW} - \frac{R_{SW} - R_{NW}}{90}(\alpha - 315) \quad (9)$$

$$R_{\alpha-90} = R_{SE} - \frac{R_{SE} - R_{SW}}{90}(\alpha - 90 - 135) \quad (10)$$

$$R_{\alpha-180} = R_{NE} - \frac{R_{NE} - R_{SE}}{90}(\alpha - 180 - 45) \quad (11)$$

$$R_{\alpha-270} = R_{NW} - \frac{R_{NW} - R_{NE}}{90}(\alpha - 270 + 45) \quad (12)$$

Further TH transverse size (W_{34}), average radii at tangent wind with speeds 34 knot, 50 knot and 64 knot (R_1 , R_2 and R_3 , respectively), an area inside tangent velocity 34 knot (A_{34}), coordinates (latitude - $L_{t_{bb}}$, longitude - Ln_{bb}) and translation speed (U_{bb}) of TH back boundary center are determined:

$$W_{34} = R_{90} + R_{270} \quad (13)$$

$$R_1 = 0.25 \left(R_{NE,34}^2 + R_{SE,34}^2 + R_{SW,34}^2 + R_{NW,34}^2 \right)^{0.5} \quad (14)$$

$$R_2 = 0.25(R_{NE,50}^2 + R_{SE,50}^2 + R_{SW,50}^2 + R_{NW,50}^2)^{0.5} \quad (15)$$

$$R_3 = 0.25(R_{NE,64}^2 + R_{SE,64}^2 + R_{SW,64}^2 + R_{NW,64}^2)^{0.5} \quad (16)$$

$$A_{34} = (\pi / 4)(R_{NE,34}^2 + R_{SE,34}^2 + R_{SW,34}^2 + R_{NW,34}^2) \quad (17)$$

$$Lt_{bb} = Lt_c \pm (R_{\alpha-180} / (1.11 \cdot 10^5)) \cos \alpha \quad (18)$$

$$Ln_{bb} = Ln_c \pm (R_{\alpha-180} / (1.11 \cdot 10^5 \cdot \cos Lt_{bb})) \sin \alpha \quad (19)$$

Here Lt_c and Ln_c are corresponding coordinates of TH center specified in corresponding regular advisory; the length of a degree of latitude is accepted as constant and equal to $1.11 \cdot 10^5$ m; the length of a degree of longitude is accepted as equal to $\cos Lt_{bb} \cdot 1.11 \cdot 10^5$ m;

2. TH back boundary translation

As the field data specified in regular forecast advisories do not allow determining of local translation speed, U_{bb} at current position is determined through geographical coordinates of TH back boundary center as arithmetic mean of average values for straight translation at two lengths of standard track data prior and subsequent to current position:

$$U_{bb} = \frac{U_{bb-1} + U_{bb+1}}{2} \quad (20)$$

$$U_{bb-1} = \frac{\left((Lt_{bb} - Lt_{bb-1}) \cdot 1.11 \cdot 10^5 \right)^2 + \left((Ln_{bb} - Ln_{bb-1}) \cdot 1.11 \cdot 10^5 \cos Lt_{bb} \right)^2}{\tau - \tau_{-1}} \quad (21)$$

$$U_{bb+1} = \frac{\left((Lt_{bb+1} - Lt_{bb}) \cdot 1.11 \cdot 10^5 \right)^2 + \left((Ln_{bb+1} - Ln_{bb}) \cdot 1.11 \cdot 10^5 \cos Lt_{bb} \right)^2}{\tau_{+1} - \tau} \quad (22)$$

Here subscripts -1 and +1 correspond to TH positions and lengths of standard track data prior and consequent to current position according regular advisories, respectively.

Unavailability of local values of translation speed affects accuracy of correlation in any case although in majority of cases situation is mitigated by rather straight and uniform translation of TH just at the stage of rapid intensification.

In certain degree accuracy of determination through equation (20) may be assessed using alternative procedures of determination of U_{bb} through average values only for one length of standard track data prior or subsequent to given position (through equation (21) or equation (22)). Comparison of aforementioned three values allows to revealing TH translation with significantly variable speed or through curved or zigzagging trajectory.

3. Determination of average HHP

The problem with initial data precision manifests itself also with HHP. Among different HHP maps (NRL maps, Goni and Trinanes maps, Cook maps) only NRL maps and the maps of Goni and Trinanes cover all relevant zones of Tropical Ocean. Besides, NRL maps present the most important range of HHP (40-90 kJcm^{-2}) by hardly distinguishable colors. In this connection the maps of Goni and Trinanes, characterized by comparatively high color resolvability, are adopted as a basis for our consideration.

HHP is determined as average weighted value for sea surface inside A_{34} for the time point just prior to entering of TH front in given area. Conversion of the color into figures imports human factor affecting accuracy of determination. As TH translation is determined through regular advisories by lag 6 hours, the lag 24 hours between the HHP maps also causes some additional error.

4. Integral heat flux

Average inside A_{34} integral heat flux is determined based at three-zone model of heat transfer introduced in the paper. The first zone covers the area between tangent winds 34 knot (17.5 ms^{-1}) and 50 knot (25.75 ms^{-1}). The second zone covers the area between tangent winds 50 knot and 64 knot (33 ms^{-1}). The third zone covers the area inside tangent wind 64 knot. Average integral heat flux is determined through empirical equation (8) of the Paper.

5. Correction factors on partial land covering

Accounting of influence of land surface covered by TH is carried out using following correction factors to W_{34} and A_{34} through determination of effective radius according equation (7) and alignment number according equation (6):

$$C_W = 1 - \frac{\delta W_{34}}{W_{34}} \quad (23)$$

$$C_A = 1 - \frac{\delta A_{34}}{A_{34}} \quad (24)$$

6. Accuracy of correlation of the field data

Accuracy of correlation of the field data in the framework of equations (6-8) is determined by the procedures of determination of HHP using HHP maps and calculation of U_{bb} and qR_{ef} . Errors in determination of these parameters may vary in rather wide range depending from values of parameters themselves. For instance, HHP and U_{bb} determination errors are high at these low absolute values (around $\pm 30\%$ and even more).

An error connected with neglecting of a sea surface drift is rather low in majority of cases (no more than $\pm 5\%$). Accuracy of equation (8) may be evaluated as $\pm 20\%$. Some specific errors are connected with accepted way of determination of U_{bb} . Firstly, it is intersection by alignment number curve of the zone of alignment effect. As uninterrupted data on U_{bb} are not available, really sharp transition also may take place without realization of any regime inside $N_{ai}=30\pm 30\%$. So, here we have uncertainty that may be eliminated only using more detailed field data. Secondly, it is connected with mentioned above underestimation of U_{bb} at TH indirect translation.

In such a complicated situation accuracy of correlation in the best way may be characterized by dispersion of the field data. In rather good accordance with ETM dispersion $N_{ai}=30\pm 30\%$ is typical for the stage of rapid intensification of almost all TH involved in correlation. Corresponding level of accuracy of correlation may be accepted for majority of considered cases having in mind that in separate situations correlation may be much less accurate.

7. Correlation of the field data

7.1. Composition of initial field data involved in correlation

Calculation of alignment number through equations (6-24), in combination with HHP determined using corresponding HHP map, involves following parameters specified in regular advisories:

- Time points and corresponding TH center coordinates at current position and at positions prior and subsequent to current position;
- Maximum tangent wind velocity at current position;
- TH radii at tangent wind velocity 34 knot in Northeast, Southeast, Southwest and Northwest quadrants at current position;
- TH radii at tangent wind velocity 50 knot in Northeast, Southeast, Southwest and Northwest quadrants at current position;
- TH radii at tangent wind velocity 64 knot in Northeast, Southeast, Southwest and Northwest quadrants at current position;
- TH translation azimuth at the length of standard track data prior to current position.

Simultaneously additional parameters are involved such as an area of land surface covered by A_{34} (δA_{34}) and a projection onto W_{34} of the part of the back boundary of A_{34} covered by land (δW_{34}) both determined using TH geometrics at current position and geographical map. It should be stressed also that equations (6-24) form unified algorithm designed for description of the field data on different TH using abovementioned initial field data without using of any fitting coefficient variable from one TH to another.

7.2. Selection of the cases

The case of rapid intensification is specified by availability of the stage with intensification factor no less than 5.0 knot/h and achievement to the end of the stage of minimum 4th category of intensity. Consideration covers all 8 THs of 4th and 5th category identified as cases of rapid intensification among all tropical hurricanes observed in 2004-2005 in all relevant zones of Tropical Ocean. Selected in such a manner cases are: hurricanes Charley, Katrina and Wilma (North West Atlantic), super typhoons Dianmu and Chaba (North West Pacific), cyclones Nancy (South Pacific), Olaf (South Indian) and Ingrid (South Pacific and South Indian).

7.3. Comparison between FAP and SAP

Previous realization of ETM (FAP) is based at assuming of uniform straightforward translation of circular TH of constant size at open sea surface. Heat balance of underplaying sea surface is considered inside TH boundary with tangent wind speed 15 m s^{-1} .

In the case of straightforward translation of circular TH of constant size translation speed U_{bb} (determining increment of cooled sea surface) turns to be equal to TH center translation speed (U_{trc}). In addition, uniform translation allows determining U_{trc} "as average for straight line covering 2 time intervals of standard track data prior and subsequent to this instance" (Shekrladze 2004a). Accordingly, conditions of alignment effect are reduced to equation (3) of the paper.

Production qR_{15} is determined by empirical equation (4) of the paper based at parameters of TH Kenna. Expansion of the equation linking heat flux only to maximum tangent velocity on all cases refers to assumption in the framework of FAP of similar radial distribution of tangent wind speeds in all THs.

Transition in the SAP to consideration of heat balance inside R_{34} allows using the field data on tangent wind distribution specified in regular forecast advisories. Determination of heat balance based at U_{bb} taking in account TH non-circular geometry allows description of TH curved translation with variable TH's shape. Three-zone model of heat transfer and corresponding empirical equation (7) allow accounting of real radial distribution of tangent wind speed. Introduction of correction factors (23) and (24) also improves description of TH development. At the same time determination of U_{bb} through equation (15) only slightly improves description of curved translation and unavailability of local current values of U_{bb} lowers accuracy of analysis.

As it follows from above, results of FAP and SAP may differ or coincide depending from a number of parameters. As it follows from analysis of such a comparison all considered cases of rapid intensification might be subdivided into three groups.

The first group includes the cases with practically identical correlation of the field data by previous and refined realizations of ETM (TH Katrina, super typhoons Dianmu and Chaba). Rather close coincidence of FAP and SAP may be explained by validity of the approximation underlying FAP of ETM in these cases.

The second group includes the cases with clearly different but rather close curves of correlation of the field data (THs Charley and Wilma, cyclone Nancy). As it follows from corresponding analysis, despite some quantitative differences, FAP adequately describes the process of Charley's development. Besides, it also predicts the stage of rapid intensification although the stage itself is described by SAP evidently better. However, taking in account in addition related sizes of Charley and Kenna (FAPs reference case), FAP of ETM may be considered as an instrument that might be used for prediction of high probability of Charley's rapid intensification through landfall to Florida Peninsula 13th of August 2004. As a whole, more relevant description of TH development by SAP of ETM manifests itself in all cases of this group.

The smallest in number third group includes the cases with significantly differing curves of correlation of the field data by FAP and SAP (cyclones Ingrid and Olaf). Besides, FAP evidently leads to miscorrelation.

Finally, results of comparison may be evaluated as quite valuable. On the one hand, these results demonstrate principal limitations of the FAP (by the way, together with demonstration of sizable area of its applicability as well); on the other hand, qualitative widening becomes evident of covering of different cases of TH rapid intensification by ETM through SAP.

7.4. Tables of correlation of the field data in the framework of SAP

Detailed tables of correlation of all considered cases of rapid intensification in the framework of equations (6-24) are presented below. Each row of the table describes certain TH position. The row includes 17 parameters borrowed from regular forecast advisory and value of HHP determined through Goni and Trinanes HHP map. The same row presents the main result of correlation for given TH position: current value of alignment number. Correction factors to A_{34} and W_{34} on partial land covering also are presented. These factors are used only if corresponding reduction is no more than a third of either of these two parameters. Besides, alignment number is considered as uncertain in developed framework quantity at TH positions with greater reduction of A_{34} or W_{34} . Such a situation is reflected by blank spaces in corresponding tables. Dimensions of initial parameters in tables correspond to regular advisories and HHP maps. Each of tables is provided by necessary comments.

TABLES OF CORRELATION OF TH RAPID INTENSIFICATION CASES OBSERVED IN 2004-2005

Table 1.

Correlation of the field data on TH CHARLEY (West North Atlantic)

TIME	LAT	LON	R ₆₄				R ₅₀				R ₃₄				α	C _W	C _A	Q	U _m	N _{ai}
			NE	SE	SW	NW	NE	SE	SW	NW	NE	SE	SW	NW						
(h)	(deg)	(deg)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	[grad]			(kJ cm ⁻²)	(knot)	
-6	16.0	72.8																		
0	16.9	74.7					20	0	0	0	100	50	0	75	295	1.00	1.00	95	55	121.7
6	16.5	76.1					40	0	0	40	100	75	0	75	285	1.00	1.00	110	60	99.9
12	17.0	77.5	25	0	0	25	40	0	0	40	100	75	0	75	290	0.69	0.84	120	65	150.6
18	17.8	78.7	25	0	0	25	40	0	0	40	100	75	30	75	300	0.80	0.89	135	65	172.3
24	18.6	79.9	20	0	0	20	50	20	0	40	100	75	30	75	305	1.00	1.00	110	75	121.5
30	19.7	81.2	20	15	10	20	50	30	10	40	100	90	30	75	310	1.00	1.00	130	80	138.7
36	21.2	81.9	25	20	10	20	60	50	20	40	110	100	40	90	335	1.00	0.94	115	90	93.9
42	22.2	82.4	25	20	10	20	60	50	20	40	110	100	45	90	340	0.84	0.75	100	90	84.3
48	23.9	82.9	25	20	10	20	60	50	20	40	110	100	45	90	345	0.74	0.89	50	95	59.6
54	25.2	82.8	25	25	15	15	50	50	40	40	90	90	75	75	360	1.00	1.00	40	95	42.6
57	26.0	82.4	25	25	15	15	50	50	40	40	90	90	75	75	5	1.00	0.73	35	125	24.2
63	26.9	82.2	20	20	0	0	30	40	40	25	40	75	75	50	15	1.00	0.68	30	120	36.3
69	29.1	81.1	20	20	20	10	30	40	30	25	40	75	40	40	20	0.00	0.08		75	
75	31.2	80.5	25	50	20	10	50	75	30	25	75	100	40	40	20	0.34	0.97		75	
82	33.2	79.0	25	40	0	0	50	75	0	0	75	100	40	40	25	0.74	0.76	20	65	47.7
88	36.0	77.0																		

The point 0 corresponds to 09:00 UT, August 11, 2004. Rapid intensification (with intensification rate 10 knot/h) is observed at 54-57th hours. Blank spaces at 69-75th hours correspond to crossing Florida Peninsula. R₂=84 km at 57th hour.

Table 2.

Correlation of the field data on TH KATRINA (West North Atlantics)

TIME	LAT	LON	R ₆₄ NE	R ₆₄ SE	R ₆₄ SW	R ₆₄ NW	R ₅₀ NE	R ₅₀ SE	R ₅₀ SW	R ₅₀ NW	R ₃₄ NE	R ₃₄ SE	R ₃₄ SW	R ₃₄ NW	a [grad]	C _w	C _A	Q (kJ cm ⁻²)	U _{in} (knot)	N _{bl}	
(h)	(deg)	(deg)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	[grad]			(kJ cm ⁻²)	(knot)		
-6	26.2	79.3																			
0	26.1	79.9	15	15	10	10	25	25	20	20	70	70	50	60	270	1.00	0.88	80	65	58.6	
6	25.5	80.7	10	10	0	10	20	20	0	20	60	60	30	60	270	1.00	0.66	80	65	71.3	
12	25.3	81.5	20	20	0	0	60	60	20	0	75	75	40	30	270	0.34	0.80		80		
18	25.1	82.2	20	20	0	0	60	60	20	0	75	75	40	30	270	0.45	0.97		80		
24	24.8	82.9	20	20	15	10	60	60	35	20	75	75	55	35	270	0.79	1.00	70	85	47.6	
30	24.6	83.6	20	20	15	10	60	60	35	20	75	75	55	35	270	1.00	1.00	75	90	28.6	
36	24.4	84.4	35	30	30	25	60	60	45	60	130	90	90	130	270	1.00	1.00	70	100	19.2	
42	24.5	85.0	35	30	30	25	60	60	45	60	130	90	90	130	270	1.00	1.00	65	100	21.0	
48	24.6	85.6	40	30	30	30	60	60	60	60	140	90	90	130	270	1.00	1.00	70	100	22.3	
54	25.0	86.2	60	30	30	30	75	60	60	60	140	90	90	130	270	1.00	1.00	75	100	29.9	
57	25.1	86.8	60	45	45	45	80	80	65	80	140	100	100	140	270	1.00	1.00	80	125	20.6	
60	25.4	87.4	75	75	50	75	100	100	75	100	160	160	125	140	270	1.00	1.00	85	125	15.9	
63	25.7	87.7	75	75	50	75	100	100	75	100	160	160	125	140	270	1.00	1.00	95	140	21.5	
66	26.0	88.1	90	90	50	75	120	120	75	100	180	180	125	140	270	1.00	1.00	90	150	23.0	
72	26.9	89.0	90	90	50	90	120	120	75	120	200	180	125	180	270	1.00	0.98	80	145	21.9	
78	27.6	89.4	90	90	50	80	110	100	75	100	200	200	150	180	270	1.00	0.92	60	140	20.7	
84	28.8	89.6	90	90	50	80	110	100	75	100	200	200	150	180	270	1.00	0.59		130		
90	30.2	89.6	110	110	60	60	140	140	75	75	200	200	150	100	270	1.00	0.47		110		
96	31.9	89.6	50	50	30	30	75	100	75	75	100	180	100	100	270	0.38	0.19		65		
102	33.5	88.5																			

The point 0 corresponds to 21:00 UT, August 25, 2005. Rapid intensification (with intensification rate 8.33 knot/h) is observed at 54-57th hours. Blank spaces at 12-18th hours correspond to crossing Florida Peninsula, at 84-96th hours - to entering continental part of USA. R₂=142 km at 57th hour.

Table 3.

Correlation of the field data on TH WILMA (West North Atlantics)

TIME	LAT	LON	R ₆₄	R ₆₄	R ₆₄	R ₆₄	R ₅₀	R ₅₀	R ₅₀	R ₅₀	R ₃₄	R ₃₄	R ₃₄	R ₃₄	α	C _W	C _A	Q	U _{in}	N _{ai}	
			NE	SE	SW	NW	NE	SE	SW	NW	NE	SE	SW	NW							
(h)	(deg)	(deg)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	[grad]			(kJ cm ⁻²)	(knot)		
	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21
-6	15.8	80.2																			
0	15.7	80.0					20	20	0	20	60	60	30	60	300	1.00	1.00	75	60	57.5	
6	16.5	80.6	15	15	0	15	30	20	20	30	105	75	50	105	320	1.00	1.00	80	65	41.3	
12	16.7	81.5	15	15	15	15	50	30	30	50	120	75	60	120	290	1.00	1.00	90	70	28.7	
18	16.8	82.1	15	15	15	15	60	30	30	60	135	90	90	135	285	1.00	1.00	90	95	20.7	
20	16.9	82.0	15	15	15	15	60	30	30	60	135	90	90	135	295	1.00	1.00	90	130	33.6	
24	17.2	82.5	15	15	15	15	60	30	30	60	140	90	90	140	295	1.00	1.00	90	150	38.9	
30	17.4	83.2	45	20	20	45	70	45	45	70	140	90	90	140	300	1.00	1.00	100	150	43.5	
36	17.7	83.7	50	40	40	50	75	60	60	75	200	200	100	150	295	1.00	1.00	100	140	45.2	
42	18.1	84.3	60	40	40	60	75	75	60	75	200	200	110	150	300	1.00	1.00	90	135	29.7	
48	18.3	85.0	60	40	40	60	75	75	60	75	200	200	110	150	295	1.00	1.00	85	130	34.7	
54	18.4	85.5	80	75	45	75	110	85	60	110	225	150	110	225	295	1.00	1.00	85	125	20.3	
60	18.9	85.7	75	75	60	75	110	90	75	110	175	150	120	175	310	1.00	1.00	90	130	19.2	
66	19.3	86.0	75	75	60	75	110	90	75	110	175	150	120	175	325	1.00	1.00	95	130	19.8	
72	20.0	86.2	75	75	60	75	110	90	75	110	175	150	120	175	330	1.00	0.97	95	130	11.3	
78	20.2	86.5	75	75	60	75	100	90	75	100	175	175	120	150	325	0.92	0.90	90	125	9.7	
84	20.6	86.9	75	75	60	75	100	100	75	100	175	175	120	150	320	0.64	0.77		120		
90	20.8	86.9	75	75	60	75	100	100	75	100	175	175	120	150	330	0.64	0.79		120		
96	20.9	87.2	75	75	60	75	100	100	75	100	175	175	120	150	300	0.62	0.77		110		
102	21.3	87.0	75	75	60	75	100	100	75	100	175	175	120	150	0	0.65	0.79		100		
108	21.4	87.1	75	75	50	75	120	100	75	100	175	175	120	150	0	0.63	0.79		85		
114	21.8	86.9	60	60	50	60	100	90	75	90	175	175	120	150	10	0.62	0.81		85		
120	22.1	86.6	60	60	40	40	100	100	75	90	175	175	125	175	45	0.62	0.85		85		
126	22.7	85.8	60	60	40	40	100	100	75	90	175	175	125	175	45	0.71	0.93	80	85	34.0	
132	23.5	84.9	65	75	50	50	125	125	90	90	200	200	150	150	45	0.88	0.96	70	90	25.7	
138	24.4	83.7	65	75	50	50	125	125	90	90	200	200	150	150	50	1.00	0.92	55	100	23.4	
144	25.5	82.4	75	80	75	50	125	125	125	90	200	200	175	150	50	0.91	0.76	45	110	17.3	
150	26.9	80.0	75	85	75	50	125	125	100	90	200	225	200	150	45	1.00	0.79	50	90	25.4	
153.5	28.1	78.8	75	90	75	40	125	150	100	90	200	225	200	150	45	0.30	0.81		100		

1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21
156	29.0	77.4	75	90	75	40	125	150	100	90	200	225	200	150	45	0.60	0.94		105	
162	31.6	74.3	75	90	75	40	125	150	100	90	200	225	200	150	45	1.00	1.00	30	110	23.0
168	34.8	70.0	75	90	60	20	125	150	100	40	150	275	375	100	50	1.00	1.00	20	100	13.9
174	34.8	70.0																		

The point 0 corresponds to 21:00 UT, September 19, 2005. Rapid intensification (with intensification rate 17.5 knot/h) is observed at 18-20th hours. Blank spaces at 84-120th hours correspond to covering of Yucatan Peninsula, at 153-156th hours - to crossing Florida Peninsula. $R_2=88$ km at 20th hour.

Table 4.

Correlation of the field data on super typhoon DIAMNU (North West Pacific)

TIME	LAT	LON	R ₆₄ NE	R ₆₄ SE	R ₆₄ SW	R ₆₄ NW	R ₅₀ NE	R ₅₀ SE	R ₅₀ SW	R ₅₀ NW	R ₃₄ NE	R ₃₄ SE	R ₃₄ SW	R ₃₄ NW	α [grad]	C _W	C _A	Q (kJ cm ⁻²)	U _{in} (knot)	N _{ai}	
(h)	(deg)	(deg)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	[grad]			(kJ cm ⁻²)	(knot)		
-6	9.8	137.0																			
0	10.7	137.4	20	20	20	20	25	25	30	25	100	100	110	100	25	1.00	1.00	80	65	48.0	
6	11.2	137.5	20	20	20	20	25	25	30	25	100	100	110	100	10	1.00	1.00	85	70	40.8	
12	11.9	137.2	20	20	20	20	25	25	30	25	105	105	115	105	350	1.00	1.00	90	85	32.7	
18	12.7	136.9	35	35	35	35	50	50	60	50	140	140	160	140	340	1.00	1.00	95	120	27.0	
24	13.4	136.8	35	35	35	35	70	70	75	70	130	140	130	130	350	1.00	1.00	95	145	34.9	
30	14.3	136.6	40	40	40	40	75	75	75	75	145	140	140	130	350	1.00	1.00	95	150	37.0	
36	15.1	136.0	40	40	40	40	70	75	70	70	140	150	140	140	325	1.00	1.00	100	155	35.2	
42	15.7	135.6	40	40	40	40	70	75	70	70	140	150	140	140	325	1.00	1.00	100	155	30.3	
48	16.5	135.2	40	40	40	40	70	75	70	70	140	150	140	140	330	1.00	1.00	90	155	27.3	
54	17.1	134.8	40	40	40	40	70	75	70	70	140	150	140	140	330	1.00	1.00	85	155	31.0	
60	17.6	133.8	40	40	40	40	70	70	60	60	150	150	150	150	305	1.00	1.00	90	140	38.7	
66	17.9	132.8	60	60	60	60	90	90	90	90	150	150	170	170	295	1.00	1.00	85	130	23.0	
72	18.0	132.1	70	70	70	70	90	120	120	90	150	150	170	170	285	1.00	1.00	80	125	21.8	
78	18.4	131.8	70	70	70	70	100	100	100	100	150	150	170	170	315	1.00	1.00	75	115	25.7	
84	19.3	131.3	70	70	70	70	100	100	100	100	150	150	160	170	330	1.00	1.00	70	120	24.6	
90	20.3	130.4	70	70	70	70	100	100	100	100	150	150	160	170	326	1.00	1.00	60	125	19.4	
96	21.2	129.7	70	70	70	70	100	100	100	100	160	160	180	180	325	1.00	1.00	55	125	13.2	
102	22.4	129.3	70	70	50	50	150	150	120	120	220	220	180	180	340	1.00	1.00	45	125	10.7	
108	23.3	129.1	50	50	50	50	120	150	80	80	200	220	180	180	350	1.00	1.00	40	120	15.3	
114	24.4	129.2	40	40	40	40	120	100	100	100	220	200	180	180	5	1.00	1.00	35	105	15.0	
120	25.6	129.3	40	40	40	30	120	100	100	50	200	200	180	120	10	1.00	1.00	30	90	13.1	
126	26.9	129.7	40	40	40	30	100	120	80	50	180	200	180	120	15	1.00	1.00	25	80	12.3	
132	28.0	130.4	30	20	20	20	90	90	70	40	200	200	180	120	30	1.00	0.99	20	70	12.0	
138	29.5	131.3	0	0	0	0	80	80	60	60	180	180	160	160	30	1.00	0.96	15	70	14.6	
144	31.0	132.1																			

The point 0 corresponds to 18:00 UT, June 14, 2004. Rapid intensification (with intensification factor 17.5 knot/h) is observed at 12-18th hours. R₂=98 km at 18th hour.

Table 5.

Correlation of the field data on super typhoon CHABA (North West Pacific)

TIME	LAT	LON	R ₆₄	R ₆₄	R ₆₄	R ₆₄	R ₅₀	R ₅₀	R ₅₀	R ₅₀	R ₃₄	R ₃₄	R ₃₄	R ₃₄	α	C _w	C _A	Q	U _{in}	N _{ai}
			NE	SE	SW	NW	NE	SE	SW	NW	NE	SE	SW	NW						
(h)	(deg)	(deg)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	[grad]			(kJ cm ⁻²)	(knot)	
1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21
-6	13.7	158.0																		
0	14.5	157.0	0	0	0	0	10	10	10	10	75	80	80	85	300	1.00	1.00	115	60	127.8
6	14.6	155.4	0	0	0	0	15	15	15	15	90	90	90	85	280	1.00	1.00	115	65	122.3
12	14.8	154.1	0	0	0	0	15	15	15	15	90	90	90	90	280	1.00	1.00	110	65	102.4
18	14.7	153.2	10	10	10	10	30	30	30	30	70	90	90	70	270	1.00	1.00	100	70	72.1
24	14.3	151.9	10	10	10	10	40	40	40	40	110	125	125	110	265	1.00	1.00	90	70	56.6
30	14.1	150.3	20	20	20	20	40	45	45	40	110	140	140	110	265	1.00	1.00	85	75	60.2
36	14.4	148.8	20	30	30	20	50	60	60	50	120	160	160	120	280	1.00	1.00	85	90	49.0
42	14.0	147.4	30	40	40	30	60	80	80	60	140	170	170	140	260	1.00	1.00	90	95	44.8
48	14.4	146.3	35	45	45	35	50	80	80	50	140	180	180	140	285	1.00	1.00	90	110	39.6
54	14.4	145.4	40	50	50	40	60	90	90	60	140	180	180	140	270	1.00	1.00	85	125	32.7
60	14.8	144.8	30	50	50	30	75	90	90	75	180	200	200	180	295	1.00	1.00	85	155	31.9
66	15.2	143.6	30	50	50	30	75	90	90	75	180	200	200	180	285	1.00	1.00	85	155	30.3
72	15.6	142.9	30	50	50	30	75	90	90	75	180	220	220	180	300	1.00	1.00	80	155	25.8
78	16.3	142.1	40	60	60	40	80	120	120	80	160	210	210	160	310	1.00	1.00	75	155	26.1
84	16.9	141.2	40	60	60	40	80	120	120	80	160	200	220	160	310	1.00	1.00	70	150	25.0
90	17.4	140.2	40	60	60	40	80	120	120	80	160	200	220	160	300	1.00	1.00	65	150	19.8
96	18.0	139.6	40	60	60	40	80	120	120	80	160	220	200	160	315	1.00	1.00	60	150	15.8
102	18.5	139.1	40	60	60	40	80	120	120	80	160	200	200	160	315	1.00	1.00	55	140	19.3
108	19.4	138.7	40	60	60	40	80	120	120	80	160	200	200	160	335	1.00	1.00	55	150	19.3
114	20.3	138.2	40	40	40	40	60	80	80	60	130	200	200	130	330	1.00	1.00	50	150	18.2
120	21.2	137.6	40	40	40	40	60	80	80	60	130	200	200	130	330	1.00	1.00	50	145	17.2
126	22.0	137.1	40	40	40	40	80	90	90	80	160	200	200	160	330	1.00	1.00	50	145	14.4
132	22.7	136.6	40	40	40	40	80	90	90	80	170	200	200	170	330	1.00	1.00	50	145	15.8
138	23.4	136.2	40	40	40	40	80	90	90	80	150	180	180	170	330	1.00	1.00	50	140	15.3
150	24.8	135.2	50	50	45	45	90	90	70	70	200	200	170	170	325	1.00	1.00	45	115	18.9
156	25.7	135.1	50	50	50	50	90	90	70	70	150	150	130	130	345	1.00	1.00	45	110	20.8
162	26.1	134.9	50	50	50	50	90	90	65	65	150	150	125	125	335	1.00	1.00	45	105	11.0
168	26.6	134.5	50	50	50	50	90	90	65	65	150	150	125	125	325	1.00	1.00	45	100	8.8

1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21
174	26.8	134.3	50	50	50	50	90	90	80	80	180	180	130	130	320	1.00	1.00	45	100	9.0
180	27.1	134.2	50	50	50	50	100	90	90	90	200	180	130	130	335	1.00	1.00	50	115	20.0
186	27.2	133.8	60	60	60	60	130	110	80	80	220	220	180	180	305	1.00	1.00	50	115	20.5
192	27.3	133.6	80	60	60	60	100	80	80	80	220	140	140	200	300	1.00	1.00	50	115	12.1
198	27.7	132.7	80	60	60	60	100	80	80	80	220	200	180	200	295	1.00	1.00	50	110	7.0
204	27.8	132.1	70	60	60	60	90	80	80	80	220	200	200	180	285	1.00	1.00	55	110	15.7
210	28.2	131.3	70	60	60	60	90	80	80	80	220	200	190	180	300	1.00	1.00	55	105	23.2
216	28.8	130.5	50	50	50	50	80	80	80	80	200	180	180	180	315	1.00	1.00	55	100	22.8
222	29.4	130.0	50	50	50	50	80	80	80	80	200	180	180	180	325	1.00	0.97	55	95	17.7
228	30.2	129.7	50	50	50	50	85	85	85	85	200	180	180	180	340	1.00	0.96	60	90	31.8
234	31.5	130.2	50	50	50	50	80	80	80	80	180	170	160	160	20	1.00	0.93	65	85	54.7
240	33.3	131.3	40	40	40	40	60	60	60	60	180	140	110	110	25	1.00	0.79	65	80	77.0
246	35.0	133.0	40																	

The point 0 corresponds to 06:00 UT, August 20, 2004. Rapid intensification (with intensification rate 5.0 knot/h) is observed at 54-60th hours. $R_2=154$ km at 18th hour.

Table 6.

Correlation of the field data on cyclone NANCY (South Pacific)

TIME	LAT	LON	R ₆₄	R ₆₄	R ₆₄	R ₆₄	R ₅₀	R ₅₀	R ₅₀	R ₅₀	R ₃₄	R ₃₄	R ₃₄	R ₃₄	α	C _W	C _A	Q	U _{in}	N _{ai}
			NE	SE	SW	NW	NE	SE	SW	NW	NE	SE	SW	NW	[grad]			(kJ cm ⁻²)	(knot)	
(h)	(deg)	(deg)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)						
0	13.6	164.2	0	0	0	0	0	0	0	0	50	60	60	50	165	1.00	1.00	85	45	31.3
12	14.0	163.5	30	30	30	30	60	60	60	60	100	120	120	100	125	1.00	1.00	80	120	34.7
24	15.3	160.9	30	30	30	30	60	60	60	60	100	120	120	100	125	1.00	1.00	75	125	55.3
42	19.2	158.7	30	30	25	25	50	50	25	25	120	120	100	100	170	1.00	1.00	50	95	41.2
48	20.3	158.3	25	25	25	25	40	40	25	25	110	110	90	90	165	1.00	1.00	45	90	37.7
60	22.0	159.9	0	0	0	0	30	30	30	30	80	100	100	80	165	1.00	1.00	40	65	37.1

The point 0 corresponds to 18:00 UT, February 13, 2005. Rapid intensification (with intensification rate 5.0 knot/h) is observed at 0-12th hours. R₂=111 km at 12th hour. In connection with small number of positions TH back boundary translation speeds at 0 and 60 hours are determined through equations (17) and (16), respectively.

Table 7.

Correlation of the field data on cyclone OLAF (South Pacific)

TIME	LAT	LON	R ₆₄	R ₆₄	R ₆₄	R ₆₄	R ₅₀	R ₅₀	R ₅₀	R ₅₀	R ₃₄	R ₃₄	R ₃₄	R ₃₄	α	C _w	C _A	Q	U _{in}	N _{ai}
			NE	SE	SW	NW	NE	SE	SW	NW	NE	SE	SW	NW	[grad]			(kJ cm ⁻²)	(knot)	
(h)	(deg)	(deg)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)	(knmi)						
-6	9.6	177.4																		
0	9.4	178.0	15	15	15	15	25	25	25	25	80	90	90	80	285	1.00	1.00	140	70	254.8
6	9.2	177.9	15	15	15	15	30	30	30	30	80	90	90	80	0	1.00	1.00	135	90	238.9
12	9.4	177.1	15	15	15	15	30	30	30	30	80	90	90	80	110	1.00	1.00	130	90	208.7
18	9.8	175.9	20	20	20	20	30	30	30	30	80	90	90	80	105	1.00	1.00	125	125	91.0
24	10.7	175.1	20	20	20	20	30	30	30	30	85	95	95	85	130	1.00	1.00	120	125	65.1
30	11.5	174.5	20	20	20	20	30	30	30	30	80	90	90	80	135	1.00	1.00	125	125	65.2
36	12.0	174.1	25	25	25	25	35	35	35	35	95	85	85	95	140	1.00	1.00	125	120	72.6
42	12.2	173.6	25	25	25	25	40	40	40	40	115	105	105	115	120	1.00	1.00	125	135	97.3
48	12.2	172.6	25	25	25	25	45	45	45	45	120	105	120	120	95	1.00	1.00	125	140	79.3
54	12.6	171.3	35	35	35	35	45	45	45	45	95	110	110	95	110	1.00	1.00	130	145	67.7
60	13.3	170.1	35	35	35	35	45	45	45	45	100	120	120	100	120	1.00	1.00	135	145	68.4
66	14.4	169.2	20	20	20	20	40	40	40	40	95	105	105	95	140	1.00	1.00	125	140	93.0
72	15.4	168.2	25	25	25	25	40	40	40	40	95	105	105	95	135	1.00	1.00	130	140	97.6
78	16.3	167.3	25	25	25	25	40	40	40	40	95	105	105	95	145	1.00	1.00	120	130	85.9
90	18.5	165.7	25	25	25	25	45	40	40	45	115	105	105	115	140	1.00	1.00	70	120	65.5
102	21.2	162.9	25	25	25	25	45	40	40	45	120	110	110	120	140	1.00	1.00	50	115	51.2
114	24.9	160.8	40	40	40	40	60	50	50	60	120	140	120	140	155	1.00	1.00	25	90	19.9
126	28.0	160.6	0	0	0	0	35	35	25	25	105	105	105	110	175	1.00	1.00	15	65	17.2
138	30.7	161.3																		

The point 0 corresponds to 00:00 UT, February 14, 2005. Rapid intensification (with intensification rate 5.83 knot/h) is observed at 12-18th hours. R₂=56 km at 18th hour.

Table 8.

Correlation of the field data on cyclone INGRID (South Pacific and South Indian)

TIME	LAT	LON	R ₆₄ NE	R ₆₄ SE	R ₆₄ SW	R ₆₄ NW	R ₅₀ NE	R ₅₀ SE	R ₅₀ SW	R ₅₀ NW	R ₃₄ NE	R ₃₄ SE	R ₃₄ SW	R ₃₄ NW	α [grad]	C _W	C _A	Q (kJ cm ⁻²)	U _m (knot)	N _{al}	
-12	13.2	148.3																			
0	13.7	149.0	15	15	15	15	45	45	40	45	110	110	110	110	115	1.00	1.00	50	75	26.1	
12	14.2	148.6	25	25	25	25	55	55	45	50	105	95	80	90	245	1.00	1.00	50	120	33.8	
24	13.9	148.0	25	25	25	25	45	45	45	45	100	100	100	100	305	1.00	1.00	55	130	23.8	
36	13.9	146.9	35	35	35	35	45	45	45	45	90	90	90	90	270	1.00	1.00	50	130	16.5	
48	13.8	146.0	20	20	20	20	35	35	35	35	85	85	80	85	280	1.00	0.93	45	105	16.0	
60	13.6	145.0	20	20	20	20	35	35	35	35	85	85	75	85	270	1.00	0.89	40	105	16.4	
72	13.2	143.7	25	25	25	25	35	35	35	35	80	80	70	70	285	1.00	0.48		100		
84	13.5	141.7	0	0	0	0	30	30	30	30	60	60	60	60	255	0.35	0.51		65		
108	12.3	138.7	10	10	10	10	20	20	20	20	60	70	70	70	290	1.00	1.00	45	80	32.5	
114	12.2	137.8	25	25	20	30	40	40	25	45	80	80	65	90	275	1.00	0.98	40	120	23.6	
120	11.9	136.8	25	25	20	30	40	40	25	45	80	80	65	90	280	1.00	0.84	40	120	24.2	
132	11.6	134.9	25	20	20	30	35	25	25	45	70	65	65	90	275	1.00	0.80	40	135	25.1	
144	11.3	132.6	25	25	25	25	35	35	35	35	80	65	65	80	275	0.69	0.93	45	120	41.9	
156	11.4	131.1	25	25	25	25	35	35	35	35	80	65	65	80	255	0.77	0.91	40	100	22.9	
168	11.6	130.2	25	25	25	25	35	35	35	35	80	55	55	80	250	0.75	0.93	40	100	19.5	
180	11.7	129.2	25	25	25	25	35	35	35	35	80	55	55	80	255	0.85	0.98	60	100	24.1	
192	12.3	128.3	25	25	25	25	35	35	35	35	100	65	65	100	235	1.00	1.00	70	115	21.4	
204	13.2	127.5	35	35	35	35	50	40	40	50	105	85	85	105	230	1.00	0.97	55	130	18.2	
216	14.5	126.8	35	35	35	35	50	40	40	50	105	85	85	105	210	0.63	0.48		90		
228	15.5	126.4	35	35	35	35	50	40	40	50	105	85	85	105	170	0.37	0.30		75		
240	16.2	126.5																			

The point 0 corresponds to 18:00 UT, March 06, 2005. Rapid intensification (with intensification rate 6.67 knot/h) is observed at 108-114th hours. Blank spaces at 72-84th hours correspond to crossing Cape Peninsula, at 216-218th hours - to covering of north coast of Australia. R₂=71km at 114th hour.