Joint interpretation of surface-based meteorological and forecast reanalysis data from West Antarctica using a suite of nonlinear tools provides an improved climatology for this poorly known region and reveals important changes in spatial coherency between inland and ice-shelf sites

1. INTRODUCTION

A deeper understanding of regional meteorology in the Antarctic is required for improved interpretations of the ever-growing body of ice-core-based paleoclimate records from this region. Benefits would also accrue in related areas such as operational forecasting at McMurdo and global change research. Artificial neural network (ANN) techniques offer new approaches to improving our record of surface observations and our understanding of the regional atmospheric circulation, two keys to this important problem.

Apart from a small number of mostly coastal manned stations, the only source of continuous, direct measurements of near-surface weather on the West Antarctic ice sheet and Ross Ice Shelf is the network of automatic weather stations (AWS) begun in 1980. Unfortunately, the AWS are widely and non-uniformly distributed, record only a few basic measurements, and are subject to instrument failures that may go for months without repair due to limited opportunities for service. The records are also relatively short with many stations in place only since the late 1980’s/early 1990’s. This makes it difficult to use these highly valuable records for comprehensive climatological studies. Recent work with ANN-based methods has shown a way around some of these problems.

With 6-hourly GCM-scale upper air data from the ECMWF 15-year reanalysis (ERA-15) as predictors (e.g., 500 mb geopotential height, 850 mb temperature advection) and the available AWS temperature and pressure data as targets, multilayer feed-forward ANNs were trained to predict the missing AWS observations from the forecast data (Reusch and Alley 2002). Using this technique, we have developed complete 15-year temperature and pressure records (1979-93) for six West Antarctic AWS: Siple, Byrd, Lettau, Marilyn, Elaine and Ferrell (Table 1). Errors for temperature prediction are approximately equal to those from a satellite-based methodology but with no exposure to problems from surface melt events or sensor changes (Reusch and Alley 2002; Shuman and Stearns 2001).

With complete 15-year records from six sites covering ~90° of longitude, it is credible to do detailed climatological studies of, for example, interannual and spatial variability. The spatial pattern of temperature anomalies and principal component analysis suggests significant differences both between the ice shelf and ice sheet sites and between the two ice sheet sites. The records also allow us to put these AWS sites into context with respect to ENSO.

Self-organizing maps (SOMs), a second ANN technique, have proved useful for analysis of synoptic-scale circulation in temperate latitudes (Hewitson and Crane 2002). The use of SOMs allows development of synoptic climatologies with an arbitrary number of smoothly transitioning climate states, in contrast to traditional synoptic classification techniques. Results from SOM analyses are applicable both to the ice-core interpretation problem and to studies of global change. SOM-derived maps of synoptic variables such as temperature and geopotential height can be compared to ice core data to examine the relationship between the proxy and the atmosphere. Climate change can be studied by looking at how state transitions evolve over

### Table 1. AWS Locations

<table>
<thead>
<tr>
<th>Station</th>
<th>Latitude</th>
<th>Longitude</th>
<th>Elevation (m)</th>
<th>Date Installed</th>
<th>Distance¹ (km)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Byrd</td>
<td>80.01° S</td>
<td>119.40° W</td>
<td>1530</td>
<td>February 1980</td>
<td>11.5</td>
</tr>
<tr>
<td>Elaine</td>
<td>83.13° S</td>
<td>174.17° E</td>
<td>60</td>
<td>January 1986</td>
<td>71</td>
</tr>
<tr>
<td>Ferrell</td>
<td>77.91° S</td>
<td>170.82° E</td>
<td>55</td>
<td>December 1980</td>
<td>49.5</td>
</tr>
<tr>
<td>Lettau</td>
<td>82.52° S</td>
<td>174.45° W</td>
<td>45</td>
<td>January 1986</td>
<td>8.3</td>
</tr>
<tr>
<td>Marilyn</td>
<td>79.95° S</td>
<td>165.13° E</td>
<td>75</td>
<td>January 1987</td>
<td>6.1</td>
</tr>
<tr>
<td>Siple</td>
<td>75.90° S</td>
<td>84.00° W</td>
<td>1054</td>
<td>January 1982¹</td>
<td>103.9</td>
</tr>
</tbody>
</table>

¹ Distance to the nearest ERA-15 gridpoint.
² Siple AWS was removed in April 1992.

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time. We anticipate new insights and improved ice-core interpretations from application of SOMs to the West Antarctic atmosphere.

Predictive ANNs can also be used to directly relate the atmosphere to the ice core proxies by training the ANN to predict ice core data from atmospheric data. The relative contributions of each predictor can provide new insight to the atmosphere-ice core relationship. This is an area of ongoing and future research.

2. METHODOLOGY

2.1 Data

Automatic Weather Station Data: The main source of direct meteorological data in West Antarctica is the network of AWS maintained by the University of Wisconsin-Madison since 1980 (Lazzara 2000). All stations provide air temperature and pressure, wind speed and direction; some stations also have relative humidity and the 3.0-0.5 m (above nominal snow surface) temperature difference. Data were selected to match ECMWF time-steps from the three hourly quality-controlled archive for each AWS. Two geographic categories are present: ice sheet (Byrd and Siple) and Ross Ice Shelf (the rest). All sites in this study are within the south/southeast Pacific sector of West Antarctica (Table 1).

Grided Meteorological Data: ERA-15 is our source for GCM-scale meteorological data at 2.5° horizontal resolution for the period 1979-1993 (ECMWF 2000). Potential problems have been noted with the ECMWF (re)analysis data over Antarctica, stemming in part from the flawed surface elevations used in these models (Genthon and Braun 1995). Elevation errors exceeding 1000 m exist in some areas of Queen Maud Land and the Antarctic Peninsula (e.g., Figure 3, Genthon and Braun 1995). Topography in West Antarctica is generally much better but errors from outside our study area will still have an influence on the reanalysis data. Evaluations of several operational products (e.g., Bromwich et al. 1995; 1998; 2000; Cullather et al. 1998) suggest that the ECMWF analyses are still the best available data sets for Antarctica despite these issues.

Ice-core Data: Four shallow cores from central West Antarctica (Reusch et al. 1999) provided high-resolution glaciochemical and annual accumulation rate data for comparison with the ERA-15 data.

2.2 Artificial Neural Networks

At the simplest level, artificial neural networks (ANNs) are a computer-based problem solving tool inspired by the original, biological neural network – the brain. Because of their ability to generate non-linear mappings during training, ANNs are particularly well-suited to complex, real-world problems such as understanding climate (Elsner and Tsonis 1992; Tarassenko 1998). Meteorological examples include an improved understanding of controls on precipitation in southern Mexico (Hewitson and Crane 1994), prediction of summer rainfall over South Africa (Hastenrath et al. 1995) and northeast Brazil (Hastenrath and Greischar 1993), and extreme event analysis in the Texas/Mexico border region (Cavazos 1999).

The ANN-based methodology used to develop the records for the six AWS sites is essentially the same as described in Reusch and Alley (2002). Two refinements (using more input data and adding insolation as a predictor) were tried but had no appreciable affect on predictive skill. Due to different elevations (Table 1), the same pressure levels could not be used at all six sites, but this did not affect the ANN training methodology. We have used the MATLAB® Neural Network Toolbox (Haykin 1999; Demuth and Beale 2000) for the AWS predictions.

The SOM-based analyses required that the ERA-15 data be resampled to an equal area grid before making annual and seasonal averages. We used both 250 km and 125 km versions of the National Snow and Ice Data Center EASE-Grid (Armstrong and Brodzik 1995). The SOM analyses were then performed with SOM-PAK software (Kohonen et al. 1996).

3. RESULTS

3.1 Seasonal Anomalies

Seasonal temperature anomalies can be quite large and range from −6.1 °C up to 7.8 °C. The extreme minima tend to occur in fall while extreme maxima occur in winter (creating an issue for annual anomaly calculations). The 1980 annual warm anomaly is predominantly due to abnormal warmth in the austral winter of that year with additional contributions from a warm fall. The evolution of this anomaly is fairly similar at all sites. The 1988 annual warm anomaly is also related to abnormal winter warmth but in this case the additional contributions come from a warm spring. Only Siple does not show a strong warm anomaly in this timeframe. The four ice shelf sites also have a strong single-season (fall) warm anomaly in 1983. It does not appear in the annual anomalies due to cold anomaly offsets in other seasons. Cold anomalies appear at all sites in fall (1982, 1987 – weak at Ferrell), winter (1986), and spring (1981 – weak at Byrd, 1986 – except Siple). The ice sheet sites and Marilyn also have cold anomalies in summer 1981 and winter 1983. Thus the 1982 annual cold anomaly appears to be at both the ice shelf and ice sheet sites but is hidden at the latter by offsetting warm/neutral seasons. In contrast, the 1983 ice sheet annual cold anomaly is distinct and only present at those sites. This implies different behaviors, with respect to temperature, between the ice sheet and ice shelf sites. Marilyn also appears to have characteristics of both domains (also suggested by the ERA-15 prediction statistics).

3.2 Multivariate Analysis

Principal component analysis (PCA) of the joined temperature records readily shows the dominance of the annual cycle (nearly 80% of the variance is found in the first component, PC1). A PCA after removing the
annual cycle has two significant components with 51% and 17% of the variance, respectively. The ice shelf sites are most influenced in PC1 with 60-65% of their variance. PC1 also affects the ice shelf sites to a lesser extent (Siple at 20% and Byrd at 35%). PC2 is dominated by the ice shelf sites suggesting a separate behavior for Siple and Byrd relative to the ice shelf sites. This may be related to a differing ENSO influence between the ice shelf and Ross Ice Shelf.

PCA of the joined pressure records (annual cycle removed) also has two significant components. In this case, PC1 has nearly 80% of the variance. Byrd (70%) joins the ice shelf sites (86-94%) in PC1 with a smaller loading on Siple (35%). PC2 (15%) is dominated by Siple (58%) and some influence from Byrd (16%). Siple thus appears to behave somewhat differently for pressure. Otherwise, the pressure pattern resembles the associations seen in the temperature PCA.

3.3 Preliminary SOM Results

Early results from SOM analysis of annually, semi-annually and seasonally averaged 700 mb geopotential height, temperature and relative humidity are promising. In general, geopotential height maps show expected features (e.g., the climatological low in the Amundsen Sea region) and significant changes on interannual and seasonal timescales. Further work will attempt to relate the SOM-based climatology to the ice-core data (glaciochemistry and accumulation rate) and ENSO.

4. SUMMARY

ANNs provide a potentially powerful tool for repairing and extending the AWS surface meteorological record in West Antarctica. They also offer a new way to study the synoptic climatology and investigate how it relates to the ice core proxy record. Results from the AWS work have been encouraging in many respects and a full analysis of these records will be even more rewarding. Early results from the SOM-based work are promising as well. The potential benefits to the ice coring community towards improved interpretations of paleoclimate proxies are but one reason of many to continue.

5. REFERENCES


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Genthon, C. and A. Braun, 1995: ECMWF Analyses and Predictions of the Surface Climate of Greenland and Antarctica. J. Climate, 8, 2324-2332.


