

EFFECT OF MERIDIONAL VARIATION OF MOIST STATIC ENERGY ON MIXED ROSSBY-GRAVITY WAVES

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1. Introduction

Wang and Rui [1990] used numerical solutions for the coupled moist Kelvin-Rossby wave on an equatorial beta plane to show that when there is a meridional gradient in SST, there is a substantial modification of the nature of the equatorially trapped waves. Gadgil and Srinivasan [1990] and Srinivasan et al. [1993] have shown that meridional variation in moist static energy will result in a meridional variation in the moist gravity wave speed. In this paper we study the effect of meridional variation in moist gravity wave speed on the equatorially trapped waves using a simple linear model of the tropical atmosphere.

2. Governing Equations

The 2-layer model for equatorially trapped waves (Wang and Rui, 1990) is

$$\begin{aligned} u_t + Uu_x - fv &= -\phi_x \\ v_t + Uv_x + fu &= -\phi_y \\ \phi_t + U\phi_x + c^2(y)[u_x + v_y] &= 0 \end{aligned}$$

Following Wang (1988), the specific humidity is assumed to vary with the sea surface temperature (SST), which we assume to be linear with latitude. The moist gravity wave speed c then is written as

$$c^2 = c_0^2[1 + \Gamma y]$$

here c_0 is the moist gravity wave speed at the Equator. The form for zonal and time dependence is $\exp[i(Kx - Ut - \omega t)]$. The latitudinal velocity is then governed by the Hermite differential equation, with the maximum value displaced in latitude toward the cooler side by an amount proportional to Γ . The eigenvalue equation is

$$[\omega^3 - (\tilde{K}^2 + 2n + 1)\tilde{\omega} - \tilde{K}][\omega^2 - \tilde{K}^2] = -\Gamma^2 \tilde{\omega}^3 [\omega^2 - 3\tilde{K}^2]$$

where n is the order of the Hermite polynomial of the $V(y)$ solution.

For gamma equal to zero, the problem reverts to the adiabatic case for mixed Rossby-gravity waves, which was treated by Lindzen (1967). The effect of SST variation with latitude is to increase the order of the eigenvalue equation from three to seven. With a gradient of moist static energy, four additional roots appear. Two of these roots are real, corresponding to stable oscillations, and the other two are a pair of complex roots, indicating a mode which grows in time.

Figure 1 shows root 1 for wavenumber K from 0 to 10 for modes $n = 1$ through 10. The adiabatic case ($\Gamma = 0$) is indicated by the solid lines and is included for $n = 0$. The moist cases $\Gamma = 0.5$ and 1.0 are indicated by the dashed and dotted lines. For n equal to one or greater, the moist static energy gradient Γ makes little difference for this root.

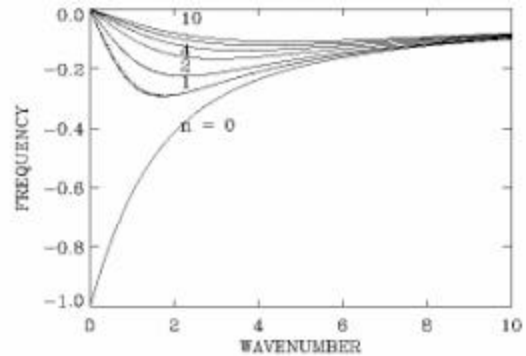


Figure 1: Eigenvalue 1 as function of wavenumber for various n . Solid line indicates $\Gamma = 0$, dashed line 0.5 and dotted line 1.0.

Figure 2 shows root 2 in like manner. For this root there is little effect of Γ for n not equal to 0 also. Figure 3 shows root 3. The effect of Γ is small except for the $n = 0$ case and small wavenumber K .

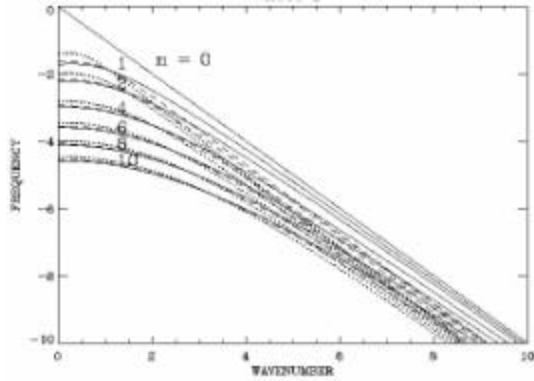


Figure 2: Eigenvalue 2 as function of wavenumber for various n . Line code same as for fig. 1.

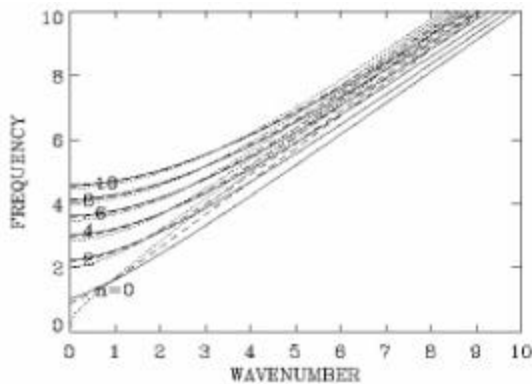


Figure 3: Eigenvalue 3 as function of wavenumber for various n . Line code same as for fig. 1.

The case of $n = 0$ for roots 1 and 2 is shown by fig. 4. The effects of gradient of moist static energy for this case are quite interesting.

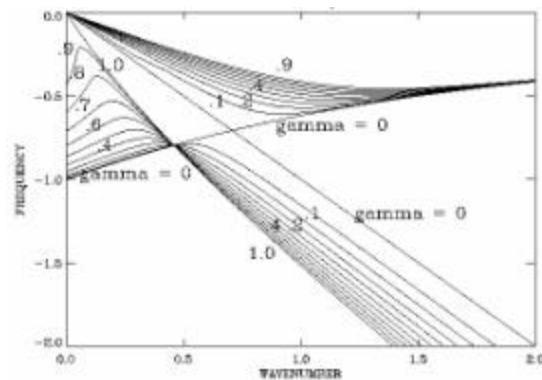


Figure 4. Frequencies of modes 1 and 2 a functions of wavenumber and moist static energy gradient Γ .

The four new roots which appear have imaginary parts, thus describe modes which grow with time for Γ not zero. Figure 5 shows the

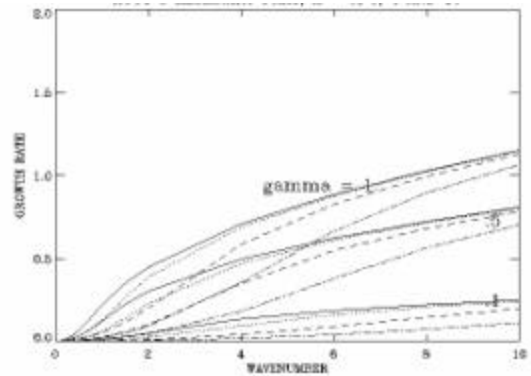


Figure 5. Growth rate of mode due to moist static energy gradient in latitude for $G = 0.1, 0.5$ and 1.0 . Line code: $n = 0$, solid; $n = 1$, dotted; $n = 4$, dashed; $n = 10$, dash-dot.

The presence of a meridional gradient in moist energy causes the amplification of westward migrating mixed Rossby-gravity waves with wavelength in the range of 8000 to 9000 km and a period on the order of 4 days (10 days) when the mean zonal wind is easterly (westerly) with a speed of 10 m/sec. The predictions of this model are consistent with observed mixed Rossby-gravity waves.

6. References

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